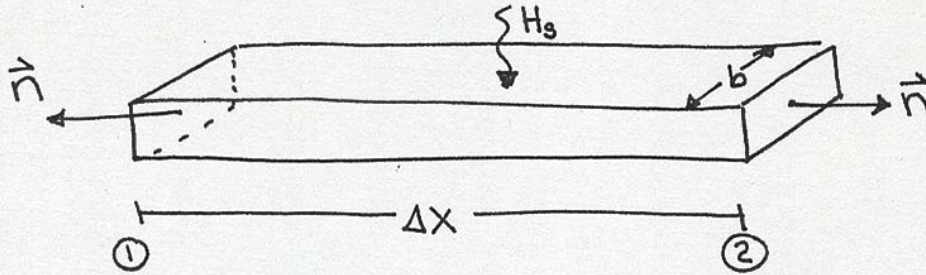


ANSWER 1.9

CONTROL-VOLUME APPROACH :

SELECT A SHORT LENGTH OF RIVER, ΔX , AND EVALUATE THE CONTROL VOLUME (INTEGRAL) FORM OF THE CONSERVATION EQUATION. FOR CONSERVATION OF HEAT ENERGY, REPLACE $C = \rho C_p T$ IN EQ. 4

$$(A) \frac{d}{dt} \int_{CV} \rho C_p T dV = - \int_{CS} \rho C_p T \vec{V} \cdot \vec{n} dA + \int_{CS} D_n \frac{dC}{dn} dA + H_s \Delta X b$$



AS THE PROBLEM STATEMENT DOES NOT INDICATE ANY UNSTEADINESS, WE ASSUME STEADY FLOW, I.E. $\frac{d}{dt} = 0$.

EVALUATING THE FLUX TERMS IN (A)

$$(B) \quad 0 = -\rho C_p U b h (T_2 - T_1) + \rho C_p D b h \left(\frac{\partial T}{\partial x} \Big|_2 - \frac{\partial T}{\partial x} \Big|_1 \right) + H_s \Delta x t$$

USING A TAYLOR EXPANSION, ASSUMING T IS CONTINUOUS IN x

$$T_2 = T_1 + \left(\frac{\partial T}{\partial x} \right) \Delta x$$

$$\frac{\partial T}{\partial x} \Big|_2 = \frac{\partial T}{\partial x} \Big|_1 + \frac{d}{dx} \left(\frac{\partial T}{\partial x} \right) \Delta x$$

PLUG THESE EXPANSIONS IN (B), AND DIVIDE OUT THE COMMON TERMS, $\Delta x b$

$$(c) \quad 0 = \rho C_p h \left(-U \frac{\partial T}{\partial x} + D \frac{\partial^2 T}{\partial x^2} \right) + H_s$$

FROM WHICH, ONE COULD SOLVE FOR $\partial T / \partial x$.

IT IS USEFUL TO CONSIDER THE RELATIVE IMPORTANCE OF THE ADVECTIVE AND DIFFUSIVE FLUXES. HERE, SPECIFICALLY THE RELATIVE MAGNITUDES OF $U \partial T / \partial x$ AND $D \partial^2 T / \partial x^2$. THE SCALE OF EACH TERM CAN BE ESTIMATED FOR THIS SYSTEM. CONSIDER THE CONTROL VOLUME LENGTH, Δx , AS AN APPROPRIATE LENGTH-SCALE, THEN

$$U \frac{\partial T}{\partial x} \sim U \frac{\Delta T}{\Delta x}$$

$$D \frac{\partial^2 T}{\partial x^2} \sim D \frac{\Delta T}{\Delta x^2},$$

WHERE ΔT IS THE TEMPERATURE CHANGE ACROSS Δx .

THE RELATIVE MAGNITUDE OF THESE TERMS IS THEN,

$$\frac{\text{ADVECTIVE FLUX}}{\text{DIFFUSIVE FLUX}} = \frac{U \frac{\Delta T}{\Delta X}}{D \Delta T / \Delta X^2} = \boxed{\frac{U \Delta X}{D}}$$

THIS DIMENSIONLESS PARAMETER IS CALLED THE PÉCLET NUMBER. IT IS DISCUSSED IN DETAIL IN CHAPTER 5.

IF $\frac{U \Delta X}{D} \gg 1$, THEN ADVECTIVE FLUXES DOMINATE DIFFUSIVE FLUXES, AND WE CAN DROP THE TERM $D \frac{\partial^2 T}{\partial X^2} \ll U \frac{\partial T}{\partial X}$.

IF $\frac{U \Delta X}{D} \ll 1$, DIFFUSIVE FLUXES ($D \frac{\partial^2 T}{\partial X^2}$) ARE MUCH LARGER THAN ADVECTIVE FLUXES ($U \frac{\partial T}{\partial X}$), AND WE CAN DROP $U \frac{\partial T}{\partial X}$.

SINCE ΔX IS NOT SPECIFICALLY DEFINED, WE ASK, EG, FOR WHAT LENGTH-SCALE WILL ADVECTION DOMINATE TRANSPORT?

$$\frac{U \Delta X}{D} \gg 1 \quad \text{IFF} \quad \Delta X \gg \frac{D}{U} = \frac{0.1 \text{ m}^2 \text{ s}^{-1}}{1 \text{ m s}^{-1}} = 0.1 \text{ m}$$

\therefore OVER ANY LENGTH-SCALE, $\Delta X \gg 0.1 \text{ m}$, WE MAY NEGLECT THE IMPACT OF DIFFUSION IN (A) FOR THIS SYSTEM. ~~As~~ THE PROBLEM ASKS FOR A DESCRIPTION OF $\partial T / \partial X$ ALONG A RIVER CHANNEL. IN SUCH A SYSTEM, THE LENGTH SCALES OF INTEREST ARE MUCH LARGER THAN 10 cm !, AND ARE MORE LIKE 100-m TO km 's. THEREFORE, FOR THIS SYSTEM,

WE CAN SAFELY DROP THE DIFFUSIVE TRANSPORT TERM. THEN, (c) REDUCES TO

$$(d) \quad 0 = -\rho C_p h U \frac{\partial T}{\partial x} + H_s$$

FROM WHICH,

$$(e) \quad \frac{\partial T}{\partial x} = \frac{H_s}{\rho C_p h U} = \frac{J s^{-1} m^{-2}}{(kg m^{-3})(J kg^{-1} K^{-1})(m)(m s^{-1})} = \frac{K}{m}$$

UNITS CHECK OUT.

USING THE STATED PARAMETERS

$$(f) \quad \frac{\partial T}{\partial x} = \frac{800 W m^{-2}}{(1000 \frac{kg}{m^3})(4200 \frac{J}{kg K})(1m)(1 \frac{m}{s})} = 2 \times 10^{-4} \frac{K}{m} = 0.2 \frac{C}{km}$$

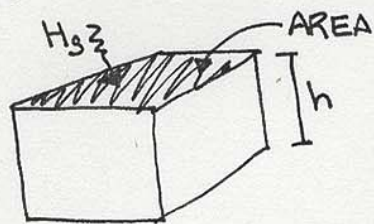
DIFFERENTIAL APPROACH :

THE CONSERVATION OF MASS EQUATION CAN BE APPLIED TO THE TRANSPORT OF HEAT ENERGY BY NOTING THE CONCENTRATION OF HEAT ENERGY, $C [J/m^3] = \rho C_p T$. THEN, THE DIFFERENTIAL FORM OF THE CONSERVATION EQUATION IS, FOR INCOMPRESSIBLE FLOW

$$(g) \quad \frac{d}{dt}(\rho C_p T) + u \frac{d}{dx}(\rho C_p T) + v \frac{d}{dy}(\rho C_p T) + w \frac{d}{dz}(\rho C_p T) =$$

$$\frac{d}{dx} D_x \frac{d}{dx}(\rho C_p T) + \frac{d}{dy} D_y \frac{d}{dy}(\rho C_p T) + \frac{d}{dz} D_z \frac{d}{dz}(\rho C_p T) \pm S$$

- IF WE NEGLECT $\rho = f(T)$, THEN $\rho \neq f(x, y, z, t)$.
- THE PROBLEM STATEMENT GIVES US $V = W = 0$, AND ISOTROPIC, HOMOGENEOUS $D = D_x = D_y = D_z \neq f(x, y, z)$
- IF WE ASSUME THE SYSTEM IS UNIFORM (WELL-MIXED) IN $y \approx z$, THEN $T \neq f(y, z)$
- THE SOURCE TERM IS GIVEN AS A SURFACE FLUX, $H_s = [J s^{-1} m^{-2}]$. SINCE THE EQUATION DEALS IN VOLUME CONCENTRATION, WE MUST DIVIDE BY DEPTH TO PUT THE SOURCE TERM IN CONSISTENT UNITS.



$$H_s = \frac{J/m^2}{s} = \frac{\text{ENERGY PER SURFACE AREA}}{s}$$

$$\frac{H_s}{h} = \frac{J/m^3}{s} = \frac{\text{ENERGY PER VOLUME}}{s}$$

APPLYING THE ABOVE POINTS, (g) REDUCES TO

$$(h) \quad \rho C_p \left[\underset{1.}{\frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x}} \right] = D \underset{2.}{\frac{\partial^2 T}{\partial x^2}} + \frac{H_s}{h}$$

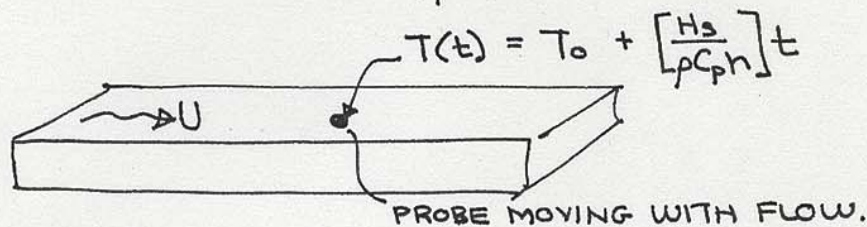
FOR TYPICAL LENGTH SCALES OF INTEREST ALONG A RIVER CHANNEL, $\Delta x \sim 100m$ TO km^5 , IT IS EASY TO SHOW THAT THE DIFFUSION TERM (2) IS SMALL COMPARED TO (1) THE ADVECTION TERM. THUS WE WILL DROP $D \frac{\partial^2 T}{\partial x^2} \ll u \frac{\partial T}{\partial x}$

SEE SCALING ARGUMENTS GIVEN ABOVE

FINALLY, NOTE THAT THE BRACKETED TERM IN (h) IS THE TOTAL DERIVATIVE

$$(i) \quad \frac{DT}{Dt} = \frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} = \frac{H_s}{\rho c_p h} = \left[\frac{^\circ C}{T} \right]$$

THIS EQUATION MAY BE READ IN THE LAGRANGIAN CONTEXT AS, FOLLOWING A PARTICULAR FLUID PARTICLE, WE WOULD OBSERVE ITS TEMPERATURE TO INCREASE AT THE RATE $\left[\frac{H_s}{\rho c_p h} \right] ^\circ C/S$.



IF THE FLOW / THERMAL CONDITIONS ARE STEADY, $\partial T / \partial t = 0$, THEN (i) ALSO PROVIDES A SIMPLE DESCRIPTION OF SPATIAL GRADIENT.

$$(j) \quad \frac{\partial T}{\partial x} = \frac{H_s}{u \rho c_p h} = \left[\frac{^\circ C}{L} \right]$$