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http://dx.doi.org/10.1073/pnas.1400643111

National Academy of Sciences (U.S.)

Final published version

Sun Apr 07 06:07:17 EDT 2019

http://hdl.handle.net/1721.1/91287

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The origin of methanethiol in midocean ridge hydrothermal fluids

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Edited by David M. Karl, University of Hawaii, Honolulu, HI, and approved March 4, 2014 (received for review January 12, 2014)

Simple alkyl thiols such as methanethiol (CH$_3$SH) are widely speculated to form in seafloor hot spring fluids. Putative CH$_3$SH synthesis by abiotic (nonbiological) reduction of inorganic carbon (CO$_2$ or CO) has been invoked as an initiation reaction for the emergence of protometabolism and microbial life in primordial hydrothermal settings. Thiols are also presumptive ligands for hydrothermal trace metals and potential fuels for associated microbial communities. In an effort to constrain sources and sinks of CH$_3$SH in seafloor hydrothermal systems, we determined for the first time its abundance in diverse hydrothermal fluids emanating from ultramafic, mafic, and sediment-covered midocean ridge settings. Our data demonstrate that the distribution of CH$_3$SH is inconsistent with metastable equilibrium with inorganic carbon, indicating that production by abiotic carbon reduction is more limited than previously proposed. CH$_3$SH concentrations are uniformly low ($\sim 10^{-8}$ M) in high-temperature fluids ($>200$ °C) from all unsedimented systems and, in many cases, suggestive of metastable equilibrium with CH$_4$ instead. Associated low-temperature fluids ($<200$ °C) formed by admixing of seawater, however, are invariably enriched in CH$_3$SH (up to $\sim 10^{-4}$ M) along with NH$_3$ and low-molecular-weight hydrocarbons relative to high-temperature source fluids, resembling our observations from a sediment-hosted system. This strongly implicates thermogenic interactions between upwelling fluids and microbial biomass or associated dissolved organic matter during subsurface mixing in crustal aquifers. Widespread thermal degradation of subsurface organic matter may be an important source of organic production in unsedimented hydrothermal systems and may influence microbial metabolic strategies in cooler near-seafloor and plume habitats.

Significance

Simple alkyl thiols such as methanethiol are widely speculated to spontaneously form in seafloor hot spring fluids and are implicated in facilitating the emergence of protometabolism and microbial life in early Earth hydrothermal systems, the complexation of hydrothermally derived metals, and as fuels for microbial ecosystems. Existing models suggest that methanethiol forms by nonbiological reduction of hydrothermal inorganic carbon (CO$_2$ or CO). We demonstrate that methanethiol is actively produced in low-temperature mixing zones of hydrothermal systems, but our data suggest it is the thermal destruction of preexisting organic matter (likely subsurface microbial biomass) that is responsible. Formation of organosulfur compounds and other degradation products during subsafefloor mixing may influence the biogeochemistry of low-temperature hydrothermal fluids inhabited by microbial life.

Author contributions: E.P.R. and J.S.S. designed research; E.P.R., J.M.M., and J.S.S. performed research; E.P.R., J.M.M., and J.S.S. analyzed data; and E.P.R. and J.S.S. wrote the paper.

The authors declare no conflict of interest.

This article is a PNAS Direct Submission.

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This article contains supporting information online at www.pnas.org/lookup/suppl/doi:10.1073/pnas.1400643111/-/DCSupplemental.
organisms) carbon compounds to hydrothermal fluids are poorly constrained at present (7, 26).

The fundamental physicochemical processes governing the formation of hydrothermal fluids and the abundances of presumptive inorganic substrates for biotic synthesis (CO$_2$/CO, H$_2$, and H$_2$S) have likely persisted through geologic time (3); hence, diverse modern analogs present an opportunity to assess the potential for abiotic synthesis of CH$_3$SH on early Earth. Therefore, to clarify its production and origin, we investigated the distribution of CH$_3$SH in 38 hot spring fluids from diverse geologic environments throughout the global midocean ridge system that span a range of temperatures and redox states. Isobaric gas-tight fluid samplers (27) were used to collect the hydrothermal fluids, and the citations for each system included here provide the most relevant or recent reported fluid compositions. Samples were taken from four unsedimented hydrothermal vent fields hosted in basaltic rock: Lucky Strike (28) and TAG (29) on the Mid-Atlantic Ridge (MAR), 9°50’N (30) on the East Pacific Rise, and Piccard (31, 32) on the Mid-Cayman Rise (MCR). Three diverse unsedimented vent fields where serpentinization of ultramafic rock is postulated to influence fluid compositions were sampled: Von Damm (33) on the MCR, and Lost City (24) and Rainbow (34) on the MAR. Fluids were also collected from the sediment-hosted Guaymas Basin vent field in the Gulf of California (35, 36), where extensive hydrothermal alteration of organic-rich sediment is occurring. Our results indicate that production of abiotic CH$_3$SH from inorganic carbon is not occurring to a significant extent and that abiotic degradation of preexisting organic matter may instead be the dominant source. The widespread production we observe during crustal mixing of hydrothermal fluids has numerous implications for the biogeochemistry of seafloor hydrothermal systems, which we discuss here.

**Results and Discussion**

Measured CH$_3$SH concentrations vary from $\sim 10^{-9}$ to $10^{-6}$ M in unsedimented hydrothermal systems (Dataset S1). In general, low-temperature (<200 °C) fluids formed by subsurface mixing of high-temperature endmember fluids with seawater (see Endmember and Mixed Vent Fluid Compositions) are most enriched. For example, CH$_3$SH concentrations in mixed fluids at Piccard are above $10^{-6}$ M, whereas associated endmember sources are $\sim 10^{-8}$ M. Cooler endmember fluids at the sediment-hosted Guaymas Basin vent field contain over $10^{-5}$ M CH$_3$SH—the highest concentrations observed—whereas nearby hotter endmember fluids in the same system have abundances similar to those associated with endmember fluids from unsedimented systems ($\sim 10^{-8}$ M). Concentrations are lowest in endmember fluids (94–96 °C) of the Lost City hydrothermal field (1.4–1.9 $\times 10^{-9}$ M).

**Thermodynamic Evaluation of CH$_3$SH Abundances.** The abiogenic production of CH$_3$SH in hydrothermal solutions is typically described (13, 22) by the overall reaction of CO$_2$, H$_2$, and H$_2$S according to the relationship

$$\text{CO}_2^{(aq)} + 2\text{H}_2\text{S}^{(aq)} + 3\text{H}_2^{(aq)} = \text{CH}_3\text{SH}^{(aq)} + 2\text{H}_2\text{O}. \quad [1]$$

Although CO is also considered a possible substrate for reduction (12, 13, 22), it does not provide an alternate pathway to CH$_3$SH that is independent of CO$_2$ reduction in fluids emanating from unsedimented systems. CO is predicted to be present at very low abundances due to rapid equilibrium with CO$_2$ and H$_2$ (37). Equilibrium between CO and CO is confirmed by calculated chemical affinities near zero for the high-temperature vent fluids from unsedimented systems presented in this study (see Assessment of Metastable Equilibrium Using Chemical Affinities and Widespread CO$_2$/H$_2$/CO Equilibrium in Unsedimented Systems and Eqs. S1 and S2). The stoichiometry of reaction 1 indicates that for a given temperature and pressure, CH$_3$SH abundance at metastable equilibrium should be highly sensitive to variations in dissolved H$_2$ due to the third-power H$_2$ dependence of the associated mass action expression. Using measured concentrations of CO$_2$, H$_2$, and H$_2$S and thermodynamic data (22) for reaction 1 at measured vent temperatures and pressures, CH$_3$SH abundances in unsedimented hydrothermal fluids are predicted to vary by over ten orders of magnitude for metastable chemical equilibrium (see Fig. 1 and Thermodynamic Prediction of Metastable CH$_3$SH Abundances). Aqueous H$_2$ concentrations that vary by almost three orders of magnitude notwithstanding, observed concentrations of CH$_3$SH are relatively uniform ($\sim 10^{-8}$ M) in high-temperature endmember fluids regardless of the geologic setting (Fig. 1 and Dataset S1). Although CH$_3$SH concentrations below predicted values might suggest kinetic inhibition of reaction 1 in high-H$_2$ fluids (e.g., at Rainbow, Piccard, and Von Damm), this does not explain the strongly enriched nature of very high-H$_2$-poor fluids relative to predictions (e.g., TAG and Lucky Strike). It is possible that endmember fluids may cool during ascent from higher temperature and pressure subsurface reaction zones, where equilibrium according to reaction 1 might regulate CH$_3$SH abundances. Such a scenario is unlikely, however, because predicted metastable equilibrium concentrations according to reaction 1 decrease strongly with increasing temperature (22), which would result in the apparent enrichments in low-H$_2$ fluids becoming even more pronounced. Collectively, these observations suggest that CH$_3$SH, CO$_2$, H$_2$S, and H$_2$O do not attain a state of metastable equilibrium at the diverse conditions encountered by fluids in modern hydrothermal systems. This implies that abiogenic synthesis of CH$_3$SH from inorganic carbon is unlikely and incapable of sustaining metastable equilibrium abundances.

The relative uniformity of endmember abundances shown in Fig. 1, despite widely differing H$_2$ abundances, suggests that CO$_2$
reduction is not responsible for the production of CH$_3$SH in high-temperature endmember fluids. Aside from CO$_2$, CH$_4$ is invariably the next largest stable pool of dissolved carbon in hydrothermal solutions, with all other aqueous single carbon species representing metastable states (7, 26). Indeed, models that propose reaction 1 as the source of CH$_3$SH in hydrothermal fluids (5, 13, 22) inherently assume that metastable CH$_3$SH is kinetically inhibited from destruction by further reduction to CH$_4$. For many high-temperature endmember fluids presented here, however, calculated chemical affinities (see Assessment of Metastable Equilibrium Using Chemical Affinities and Eq. S1) show CH$_3$SH is indeed at or close to metastable equilibrium with CH$_4$ according to the reaction

$$\text{CH}_3\text{SH}_{(aq)} + \text{H}_2_{(aq)} = \text{CH}_4_{(aq)} + \text{H}_2\text{S}_{(aq)}. \quad [2]$$

For example, affinities for reaction 2 for all Rainbow endmember fluids (−3.8 to +0.5 kJ/mol) are within typical uncertainties of equilibrium at conditions of venting (see Assessment of Metastable Equilibrium Using Chemical Affinities). At Piccard (+6.4 to +14.6 kJ/mol) and Lucky Strike (−4.7 to −12.8 kJ/mol), endmembers are also close to equilibrium in several cases, as are some vents at 9°50′N (e.g., Tica, +5.5 kJ/mol). Thus, in contrast to reaction 1, measured CH$_3$SH concentrations in endmember fluids are more consistent with predicted values according to reaction 2. Although we cannot exclude that the reverse of reaction 2 is occurring, there is to date no evidence to suggest that aqueous CH$_4$ can react with H$_2$S under hydrothermal conditions. On the contrary, CH$_3$SH was observed to react to form small quantities of CH$_4$ at 100 °C in the thioester synthesis experiments of Huber and Wächterhäuser (8). Thermodynamic data (22) indicate that reaction 2 would maintain CH$_3$SH at low levels with respect to CH$_4$ for most hydrothermal fluid compositions, with an inverse dependence on H$_2$ abundance. This is incompatible with the notion of greater abiotic CH$_3$SH production with increasing H$_2$ abundance, as invoked in scenarios for prebiotic hydrothermal thioester production (8, 12, 13). Regardless of whether or not CH$_3$SH forms by inorganic carbon reduction (reaction 1) or other biogenic or thermogenic processes during hydrothermal fluid circulation, the data presented here strongly suggest that metastable equilibrium with CH$_4$ at high temperatures is sufficiently fast that it regulates CH$_3$SH abundances in endmember fluids according to reaction 2.

**Thermogenic CH$_3$SH Production.** The highest CH$_3$SH concentrations were observed in endmember vent fluids from the sediment-covered Guaymas Basin rift zone, where the influence of hydrothermal alteration of immature organic matter and biomass is readily apparent (Fig. 2). At Guaymas Basin, basaltic dikes and sills intrude into 0.5-km-thick organic-rich diatomaceous ooze overlaying the ridge axis, resulting in rapid and widespread hydrothermal alteration of immature sedimentary organic matter and expulsion of hydrothermal petroleum at the seafloor (36, 38). In addition to abundant NH$_4^+$ and dissolved CO$_2$, multiple classes of thermogenic organic compounds are added to circulating fluids during this process (35, 36, 39). Alkyl thiols are considered to form predominantly at low thermal maturities during petroleum generation in slowly subsiding sedimentary basins (40), and their production in this setting is therefore not surprising. Cyclic polysulfide organosulfur compounds (thiolanes, thianes, and thiepanes) have previously been reported in fragments of an active smoker chimney from Guaymas Basin, indicating organosulfur production during hydrothermal petroleum generation (38). Production of CH$_3$SH during hydrothermal alteration of sedimentary organic matter could reflect the removal of organosulfur moieties from macromolecular organic structures or the secondary reaction of thermogenic products such as CO. Indeed, fluids with abundant CH$_3$SH at Guaymas also have excess CO relative to equilibrium with CO$_2$ and H$_2$ (see Dataset S1 and Widespread CO$_2$–H$_2$–CO Equilibrium in Unsedimented Systems).

The abundance of CH$_3$SH in fluids at Guaymas Basin is characterized by a bimodal distribution, with cooler endmember fluids being most enriched and hotter endmember fluids having similarly low abundances to fluids in unsedimented systems. CH$_3$SH-depleted Rebecca’s Roost and Toadstool fluids (288–299 °C; Fig. 2) have substantially higher C$_1$/C$_2$–C$_3$ ratios (124–132 versus 55–60) than the CH$_3$SH-rich cooler fluids (172–251 °C), suggesting higher thermal maturity in the hotter fluids and the conversion of longer-chain alkanes to shorter chains (39). These observations suggest that abundant CH$_3$SH is produced predominantly during early hydrothermal alteration of immature organic matter, consistent with observations from conventional
petroleum-producing systems (40). In a similar manner to C2x hydrocarbons, but unlike NH\textsubscript{4}, CH\textsubscript{3}SH may have decomposed (e.g., by reaction 2) at the higher thermal stress of Rebecca’s Roost and Toadstool fluids before venting.

Although sedimentary organic matter is lacking at unsedimented spreading centers, the potential still exists for thermogenic production of thiols due to the presence of a putative subsurface biosphere and vent-associated biomass (26, 41–43). In low-temperature hydrothermal fluids (<200 °C) from numerous unsedimented settings, there is substantial evidence to support CH\textsubscript{3}SH production during mixing of endmember fluids with seawater within hydrothermal upflow zones or crustal aquifers. Linear mixing relationships between measured Cl and Mg concentrations for fluids sampled at Rainbow, Von Damm, and Piccard indicate that a common source fluid at each vent field feeds all of the overlying vents (Fig. 3A). That many of the replicate fluid samples collected at the low-temperature vents are characterized by nearly identical Mg concentrations implies that mixing of Mg-depleted endmember hydrothermal fluids with Mg-rich seawater occurred in the subsurface before sampling because stochastic admixing of ambient seawater during fluid collection at vent orifices would likely result in highly variable Mg concentrations (see Endmember and Mixed Vent Fluid Compositions). In contrast to the conservative behavior that characterizes Cl abundances during mixing, CH\textsubscript{3}SH concentrations in mixed fluids are all substantially greater than expected for conservative dilution of the associated high-temperature endmember (Fig. 3B and Fig. S1). Such enrichments require production of CH\textsubscript{3}SH during subsurface mixing and associated cooling. Although we cannot completely exclude that some component of the CH\textsubscript{3}SH in mixed fluids is derived from CO\textsubscript{2} reduction or abiotic formation from other metastable intermediates, CH\textsubscript{3}SH abundances are far below values predicted for metastable equilibrium with respect to CO\textsubscript{2}, H\textsubscript{2}S, H\textsubscript{2}, and H\textsubscript{2}O (Fig. 1). It is difficult to argue that reaction 1 should only proceed at the lower temperatures and H\textsubscript{2} abundances of mixed fluids, given the complete lack of evidence that inorganic carbon reduction is responsible for CH\textsubscript{3}SH production in endmember fluids, where substantially faster reaction rates would be expected due to higher temperatures and reactant concentrations.

Elevated concentrations of other organically derived aqueous species in these low-temperature mixed fluids provide compelling support for thermogenic CH\textsubscript{3}SH production during mixing. As is evident at Guaymas, short-chain hydrocarbons and NH\textsubscript{4} are typically produced simultaneously during hydrothermal alteration of organic matter (35, 36, 39, 44). High CH\textsubscript{3}SH concentrations in mixed fluids at Piccard, Von Damm, Rainbow, and 9°50’N are associated with excess NH\textsubscript{4} (Fig. 3C and Fig. S1) and, in some cases, low-molecular-weight hydrocarbon (methane, ethane, and propane) enrichments relative to conservative dilution of precursor endmember fluids. For example, the low-temperature Hot Chimlet #1 vent at Piccard is enriched by more than 25% in NH\textsubscript{4} and CH\textsubscript{3}SH relative to conservative dilution of high-temperature endmember concentrations, and the low-temperature Ecurie vent at Rainbow is also enriched in both species (Dataset S1). Ethane and propane are substantially enriched in mixed fluids at Piccard, with concentrations as high as 90 nmol/kg ethane and 60 nmol/kg propane in the mixed fluids Hot Chimlet #1 and Hot Chimlet #2, whereas associated endmember fluids have much lower C\textsubscript{2x} hydrocarbon concentrations [8–20 nmol/kg ethane and propane (<10 nmol/kg) below detection].

Abiotic N\textsubscript{2} reduction is unlikely to be responsible for the NH\textsubscript{4} enrichment in mixed fluids, given that it is only thought to occur under high-temperature reaction zone conditions (42, 44). Admixing and abiotic reduction of seawater nitrate (NO\textsubscript{3}–) may be an alternate possible source of NH\textsubscript{4} to mixed fluids because

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Fig. 3. Plots of measured concentrations of chloride (Cl; A), methanethiol (CH\textsubscript{3}SH; B), and ammonium (NH\textsubscript{4}; C) versus Mg in fluid samples from three unsedimented hydrothermal vent fields. Mg is used as an index of mixing between seawater and vent fluid (see Endmember and Mixed Vent Fluid Compositions for further details). Low-Mg samples from high-temperature endmember fluids (open symbols) and higher-Mg samples from low-temperature fluids (half filled/solid symbols) at each field all lie on common mixing lines with respect to Cl, indicating single common source fluids at each site. At all vent fields, the higher Mg, low-temperature fluids formed by predominantly subsurface mixing with seawater (dashed lines) are enriched in CH\textsubscript{3}SH and NH\textsubscript{4} (and low-molecular-weight alkanes; see text) relative to conservative dilution of endmember fluids (solid lines) with seawater. Uncertainties (2σ) not shown are smaller than the data symbols.
(\sim 150 \, ^\circ \text{C}) \) mixed fluids from other mafic-hosted hydrothermal systems based on N-isotope measurements (42). Thermal degradation of dissolved organic nitrogen (DON) from admixed ambient deep ocean seawater is also an unlikely source of \( \text{NH}_3 \). Some mixed fluids are enriched in \( \text{NH}_3 \) above conservative endmember dilution by up to 4 \( \mu \text{M} \) (Fig. 3C), exceeding what might reasonably be expected based on complete N release from typical deep-ocean DON concentrations \(< 3 \mu \text{M} \) (47)).

Although microbially mediated nitrogen transformations can increase the abundance of \( \text{NH}_3 \) in very low-temperature/diffuse fluids (42), most large \( \text{NH}_3 \) and \( \text{CH}_3\text{SH} \) enrichments observed during this study (e.g., Hot Chimlet \#1) are in mixed fluids hotter than the 122 \( \, ^\circ \text{C} \) limit for microbial life (48), suggesting that microbial activity is not responsible for production of either \( \text{NH}_3 \) or \( \text{CH}_3\text{SH} \).

It is widely proposed that the permeable and porous upper oceanic crust (Layer 2A)—where mixing predominantly occurs—harbors microbial communities that constitute a deep biosphere (41–43). Given that \( \text{CH}_3\text{SH} \) in low-temperature mixed fluids is consistently associated with thermogenic indicators (Fig. 3 and Fig. S1) in systems free of sedimentary influence, we propose that entrainment and thermal alteration of microbial biomass, and/or associated dissolved organic matter (DOM), are responsible for production of \( \text{CH}_3\text{SH} \) within subsurface zones of mixing between high-temperature fluids and seawater. Previous observations of organic compounds derived from microbial biomass pyrolysis in low-temperature fluids support this possibility. Brault et al. (49) first described the presence of nonvolatile hydrocarbons in a diffuse \( \sim 15 \, ^\circ \text{C} \) and a high-temperature \( \sim 250 \, ^\circ \text{C} \) vent from 13\,°N, East Pacific Rise, and found strong enrichments of microbial lipid residues in the cooler vent. More recent work also suggests that dissolved organic carbon (DOC) is enriched in low-temperature fluids relative to both endmember precursors and ambient seawater, for example, and correlates with high temperature (50). To address whether these observations could reflect active microbial processes, our results suggest that higher-temperature portions of subsurface mixing zones generate dissolved organic compounds through abiotic thermal degradation. Fluids in the temperature range presented here (e.g., 126–191 \( ^\circ \text{C} \)) have rarely been reported but provide key insights into processes occurring within crustal hydrothermal aquifers below the known 122 \( ^\circ \text{C} \) (48) temperature limit of life.

The ubiquitously and uniformly low levels of \( \text{CH}_3\text{SH} \) \( \sim \times 10^{-10} \) \( \text{M} \) in endmember fluids from unseminated systems, despite the broad range of temperature, salinity, and \( \text{H}_2 \) concentrations, may be a remnant of pyrolysis of trace organic matter at high temperatures. Brault et al. (49) noted minor quantities of thermally mature hopanes in a hot fluid \( \sim 250 \, ^\circ \text{C} \) from 13\,°N, suggesting that entrainment and rapid pyrolysis of biomass to a high extent of maturity occurs during fluid venting at the seafloor. Endmember fluids are also known to be depleted in DOC relative to ambient seawater (50), implying thermal decomposition. A thermogenic carbon source would provide an explanation for the presence of \( \text{CH}_3\text{SH} \) in fluids with no thermodynamic drive to create it from \( \text{CO}_2 \). In a similar manner to the hotter endmember fluids at Guaymas Basin, consumption of \( \text{CH}_3\text{SH} \) by reaction 2 in high-temperature fluids could therefore represent the high maturity stage of much more limited organic matter pyrolysis during hydrothermal circulation or venting that produces insignificant changes in the abundance of major species like \( \text{CH}_4 \). Lost City endmember fluids, despite quite low vent temperatures (94–96 \( ^\circ \text{C} \), Dataset S1), are also consistent with this explanation. Several lines of evidence from the inorganic (S) and organic (S2) composition of endmember fluids point to substantial conductive cooling and much higher temperatures (likely in excess of 250 \( ^\circ \text{C} \)) in the subsurface reaction zone. Collectively, our data therefore imply that the distribution of methanethiol in seafloor hydrothermal fluids is largely controlled by thermal maturation of preexisting biological organic matter, with endmember and mixed vent fluids representing higher and lower thermalmatures, respectively.

**Implications.** Despite the diversity of geologic settings and potential catalytic minerals present in hydrothermal reaction zones, our results show no evidence for abiotic methanethiol synthesis from the inorganic precursors \( \text{CO}_2 \), \( \text{H}_2 \), and \( \text{H}_2\text{S} \) in modern hydrothermal fluids. This suggests that analogous hydrothermal systems on early Earth may not represent an abundant source of abiotic \( \text{CH}_3\text{SH} \) necessary for thromboer production (4, 8, 12). The production of thermogenic organic compounds in crustal mixing zones, however, has numerous interesting biogeochemical implications for modern seafloor hydrothermal systems. Not only does widespread pyrolysis of subsurface organic matter provide further indirect support for a putative deep biosphere in unseminated hydrothermal aquifers, it implies that such carbon may be recycled and returned to cooler near-surface environments by subsurface mixing processes. Production of methylated organic compounds by subsurface pyrolysis raises the possibility that metallohydrothermal trace metals (e.g., methylotrophy) in hydrothermal systems traditionally considered to have limited available organic compounds. A predominance of thermogenic products in low-temperature fluids could support larger populations of organotrophic microbes in these mixing zones relative to those immediately surrounding high-temperature vent structures, for example. Given that thiol functional groups are hypothesized to play a significant role in the complexity and delivery of hydrothermal trace metals (e.g., \( \text{Fe} \) and \( \text{Cu} \)) to the deep ocean (17–19), thermogenic production of organosulfur compounds with a high affinity for metals may constitute a key mechanism for this process in unseminated hydrothermal systems.

**Materials and Methods**

All fluid samples were collected using isobaric gas-tight (IGT) samples (27) during cruises to the Mid-Atlantic Ridge, Guaymas Basin, and East Pacific Rise in 2008 and Mid-Cayman Rise in 2012, using either ROV Jason or HOV Alvin. In most cases a minimum of two IGT samples were taken from each vent. Reported vent temperatures are the maximum measured in real time during fluid collection (27).

Dissolved \( \text{CH}_3\text{SH} \) concentrations were determined at sea by sampler recovery by purge-and-trap gas chromatography (GC) with flame ionization detection (FID). FID, unlike sulfur-specific detection, is insensitive to the extremely high \( \text{H}_2\text{S} \) concentrations in vent fluids. Gas-tight fluid aliquots (<4 mL) were acidified with ~1 mL of 25 wt % phosphoric acid, and \( \text{CH}_3\text{SH} \) was sparged with \( \text{He} \) gas (30 mL/min for 10 min) and cryofocused on an n-octane–coated silica trap (~78 \( ^\circ \text{C} \)), then thermally desorbed (145 \( ^\circ \text{C} \)) directly onto a Carbograph 1SC packed GC column (30 mL/min \( \text{He} \), 40 \( ^\circ \text{C} \)). To limit potential losses of gaseous \( \text{CH}_3\text{SH} \) during sparging (53, 54), deactivated glass and polytetrafluoroethylene tubing were used wherever possible in the purge-and-trap system. Sparging was assumed to be quantitative given the long sparge time and volatility of \( \text{CH}_3\text{SH} \), and resparging tests on samples revealed no significant evidence of incomplete removal. Before calibration and between samples, the trap was heated at >145 \( ^\circ \text{C} \) with \( \text{He} \) flowing (30 mL/min) for a minimum of 10 min to remove residual \( \text{CH}_3\text{SH} \) from the tubing, and any residual \( \text{CH}_3\text{SH} \) remaining in the trap (typically <1%) from previous analyses. In almost all cases, resulting \( \text{CH}_3\text{SH} \) concentrations from separate discrete samples of the same vent yield mixing lines within bottom seawater (no detectable \( \text{CH}_3\text{SH} \)) and a hydrothermal endmember or mixed fluid composition when plotted against Mg (see Endmember and Mixed Vent Fluid Compositions for further details). This indicates not only that methanethiol is conservative with respect to accidental seawater entrainment during sample collection but that any losses or additions of \( \text{CH}_3\text{SH} \) due to the analytical method are not significant. Reported uncertainties (2s) for \( \text{CH}_3\text{SH} \) (Dataset S1) are the larger of either the error of reproducibility or the uncertainty of the commercial gas standard used (±5%). \( \text{H}_2 \) and \( \text{CO}_2 \) were analyzed at sea by a headspace extraction GC technique, using thermal conductivity and helium ionization detection, respectively (37, 39). \( \text{H}_2\text{S} \) (total dissolved, \( \Sigma \text{H}_2\text{S} \)) was determined either gravimetrically by precipitation as Ag\(_2\text{S}\) (55) or at sea by electrochemical (28, 34) or iodometric titration. \( \Sigma \text{H}_2\text{S} \) was determined at sea by electrode (28, 34, 55). Aliquots for dissolved inorganic carbon (DIC, abbreviated as \( \text{CO}_2 \), \( \text{H}_2\text{O} \), and \( \text{C}_2 \)) hydrocarbons were stored in evacuated glass serum vials (poisoned with HgCl\(_2\)) for headspace gas GC analysis (28, 55). \( \text{C} \) was determined by ion chromatography (IC, ±5% (53)) or electrochemical detection (±3% (28, 34, 55)). \( \Sigma \text{H}_2\text{S} \) was determined by either flow injection analysis (57) for unseminated systems or IC (Guaymas Basin), with reported errors representing the larger of either error of reproducibility (2s) or the
typical ±5% reproducibility of prepared standards. Analytical uncertainties (2s) are ±0.05 for pH (25 °C), ±10% for H₂S, CO₂, H₂, and C₂H, hydrocarbon concentrations; and ±5% for Mg, SO₂, CO₂H, and CH₄ concentrations (28, 39, 55).

ACKNOWLEDGMENTS. We thank the captains and crews of the R/V Atlantis and R/V Roger Revelle and the crews of the HOV Alvin and the ROV Jason for their indispensable skills in sample collection. We are grateful to N. Pester and W. Seyfried for providing Mg and H₂S data for TAG and Lost City. This research was supported by National Science Foundation OCE-0702677, OCE-0548929, and OCE-1061663, and NASA Grant NNX-327 09A875G (to J.S.S.); and by funding from the Woods Hole Oceanographic Institution Deep Ocean Exploration Institute, InterRidge, and the Deutsche Forschungsgemeinschaft Research Center/Cluster of Excellence MARUM “The Ocean in the Earth System” (E.P.R.).


28. Reeves et al. (2014) April 15, 2014 | vol. 111 | no. 15 | 5479