

**MICROBIAL CYCLING OF MARINE HIGH MOLECULAR WEIGHT
DISSOLVED ORGANIC MATTER**

By

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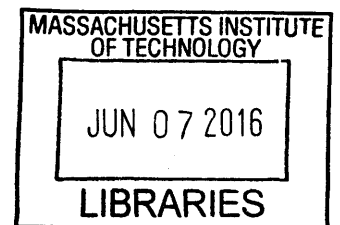
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Microorganisms play a central role mediating biogeochemical cycles in the ocean. Marine dissolved organic matter (DOM) – a reservoir of organic solutes and colloids derived from plankton is a major source of carbon, nutrients, and energy to microbial communities. The biological transformation and remineralization of DOM sustains marine productivity by linking the microbial food web to higher trophic levels (the microbial loop) and exerts important controls over the cycles of carbon and bio-essential elements, such as nitrogen and phosphorus, in the sea. Yet insight into the underlying metabolism and reactions driving the degradation of DOM is limited partly because its exact molecular composition is difficult to constrain and appropriate microbial model systems known to decompose marine DOM are lacking. This thesis identifies marine microorganisms that can serve as model systems to study the metabolic pathways and biochemical reactions that control an important ecosystem function, DOM turnover. To accomplish this goal, bacterial isolates were obtained by enriching seawater in dilution-to-extinction culturing experiments with a natural source of DOM, specifically, the high molecular weight (HMW) fraction (>1 kDa nominal molecular weight) obtained by ultrafiltration. Because it is relatively easy to concentrate and it is fairly uniform in its chemical composition across the global ocean and other aquatic environments, HMW DOM has the potential to serve as a model growth substrate to study the biological breakdown of DOM. The phylogeny, genomes, and growth characteristics of the organisms identified through this work indicate that HMW DOM contains bioavailable substrates that may support widespread microbial populations in coastal and open-ocean environments. The availability of ecologically relevant isolates in culture can now serve to test hypothesis emerging from cultivation-independent studies pertaining the potential role of microbial groups in the decomposition of organic matter in the sea. Detailed studies of the biochemical changes exerted on DOM by selected bacterial strains will provide new insight into the processes driving the aerobic microbial food chain in the upper ocean.

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WHEN YOU LOOK INTO THE ABYSS, THE ABYSS ALSO LOOKS INTO YOU.
Friedrich Nietzsche

NVLLIAS IN VERBA
“SEE FOR YOURSELF” (OR QUESTION AUTHORITY)
Royal Society of London

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CHAPTER ONE

Introduction

The oceans harbor one of the largest reservoirs of organic carbon on the planet (660 Pg; Hansell, Carlson, Repeta, & Schlitzer, 2009; Hansell & Carlson, 1998; Williams & Druffel, 1987) comparable in magnitude to the organic carbon stored in upper bioturbated surface layer (10-20 cm) of marine sediments (~150 Pg; Emerson & Hedges, 1988), terrestrial soils (1576 Pg ; Eswaran, Van Den Berg, & Reich, 1993), and carbon dioxide content in our atmosphere [850 Pg (400 ppm); (Keeling, 2015)]. The majority of the organic carbon in the ocean (>97%) is in the form of dissolved organic matter (DOM), an operationally-defined term typically used to describe colloids and organic solutes in water that pass freely through a filter with pores measuring around 200 nm in diameter (Benner, 2002). Marine DOM is thought to be the result of an assortment of food web processes that produce and consume organic matter originally derived from photosynthesis and the progressive concentration of organic compounds particularly resistant to degradation (Carlson, 2002). Early studies comparing the respiratory activity of different size classes of marine biota pointed out that heterotrophic microorganisms such as bacteria, rather than zooplankton – as was originally suspected – were the primary sink of the ocean’s photosynthetically-derived organic matter (Pomeroy, 1974) and estimated that as much as 60% of primary production cycles through these microorganisms (Williams, 1981) most of which is converted into DOM. Bacteria are in fact the dominant chemoorganoheterotrophic microorganisms in the sunlit regions of the ocean (Baumann, Baumann, Mandel, & Allen, 1972; Ducklow & Hill, 1985; Fuhrman, Sleeter, Carlson, & Proctor, 1989) fueling their metabolism with carbon and energy obtained from oxidation-reduction reactions of organic compounds, though it is now known that many marine heterotrophic bacteria can also harvest light – and

abundant energy source in clear surface waters – to complement their metabolism (Béjà, Spudich, Spudich, Leclerc, & DeLong, 2001). The ecological role of bacterioplankton in aquatic ecosystems is central in the “microbial loop” as summarized by Azam et al. (1983). In this trophic cascade, energy and nutrients stored in DOM are shunted back to the food web via DOM-consuming bacteria which are preyed upon by protozoans and other bacterivores, which are in turn eaten by zooplankton and higher trophic levels. Additionally, the microbial decomposition of DOM contributes significantly to the regeneration of inorganic forms of nitrogen and phosphorus that ultimately support marine productivity (Karl, 2002). These properties make the link between DOM and microorganisms a central component of Earth’s carbon cycle (Hansell & Carlson, 2001).

The biogeochemical nature of DOM

Even though DOM is fundamental to multiple ecological processes in the ocean, its molecular composition and biogeochemistry remain poorly characterized (Benner, 2002) making it difficult to investigate the underlying biological processes that contribute to its production and consumption. Depending on its characteristic residence time, DOM is typically categorized as labile, semi-labile, or refractory. Labile DOM cycles in timescales less than one day, and thus represents a very small fraction of bulk DOM at any given time (Carlson & Ducklow, 1996; Cherrier, Bauer, & Druffel, 1996; Ducklow & Carlson, 1992; Fuhrman et al., 1986). Bacterial production assays based on ^3H -thymidine or ^3H -leucine incorporation typically span a few hours and thus are thought to reflect the microbial cycling of labile DOM (Fuhrman & Azam, 1980, 1982; Kirchman, K’nees, & Hodson, 1985). Most DOM however, is thought to be either refractory or semi-labile. Refractory organic matter is thought to have a residence time of

thousands of years as suggested by the depletion of radiocarbon in the DOC pool in the deep ocean. Williams & Druffel (1987) found that open-ocean surface waters accumulate DOC enriched in radiocarbon ($\Delta^{14}\text{C}$ of -146‰) relative to deep waters ($\Delta^{14}\text{C}$ -540‰) reflective of the active cycling of newly fixed bomb radiocarbon but also indicative of the existence of a major old and refractory organic carbon reservoir that cycles through the ocean's conveyor over the course of thousands of years (Druffel, Williams, & Suzuki, 1989). In turn, the timescales associated with the turnover of semi-labile DOM – weeks, months or a few years – are derived from models that use labile and refractory DOM residence times as end members to constrain the turnover time of the remaining DOM inventory (Carlson & Ducklow, 1995; Davis & Benner, 2007; Druffel, Williams, Bauer, & Ertel, 1992; Luo, Friedrichs, Doney, Church, & Ducklow, 2010; Ogawa & Tanoue, 2003). These results warranted the need to chemically characterize DOM in detail to understand why it accumulated in the ocean and explain the predicted residence times of the different pools of DOM.

Considerable progress has been made in the last decades to dissect and describe the chemical and molecular nature of DOM (Nebbioso & Piccolo, 2013) although most marine organic matter cannot be easily characterized due to its small molecular size, the difficulty to isolate polar organic material from dilute saline solutions, or because of its chemical and physical state (Hedges, 1992; Benner, 2002). Considerable efforts aimed at describing DOM have focused on characterizing the high-molecular-weight (HMW) fraction – typically made up of compounds with a nominal molecular weight >1kDa and less than 30–100 kDa (Benner, 2002). Tangential flow ultrafiltration is a chemically-unbiased technique used to concentrate HMW DOM from thousands of liters of seawater to obtain sufficient quantities for analysis (Aluwihare, Repeta, & Chen, 1997; McCarthy, Hedges, & Benner, 1993; McCarthy, Hedges, &

Benner, 1996). About 30% of the DOM standing stock in open-ocean surface waters can be concentrated by ultrafiltration while the rest is too small to be sampled by this method and is thus defined as low-molecular-weight (LMW) DOM (Fig. 1.1). However, it is unclear if this material is representative of the bulk DOM pool (Benner, Biddanda, Black, & McCarthy, 1997; Carlson, Brann, Mague, & Mayer, 1985). Nonetheless, the isolation of HMW DOM from seawater has enabled molecular characterization studies which in turn have begun to shed light on the properties that may be implicated in its lability and biogeochemistry as well as its ecological role in the microbial loop.

Radiocarbon measurements of HMW DOM have provided direct evidence supporting semi-labile turnover times predicted by lability models. For example, HMW DOM from open-ocean surface waters and its constituting neutral sugars are highly enriched in bomb radiocarbon ($\Delta^{14}\text{C}$ of 10‰ and a $\Delta^{14}\text{C}$ 47-67‰, respectively) relative to the bulk DOC pool ($\Delta^{14}\text{C}$ of -146‰). These values are similar to the radiocarbon content of dissolved inorganic carbon (DIC) in surface waters ($\Delta^{14}\text{C}$ 72±7‰; Repeta & Aluwihare, 2006) and are thus indicative of recent production. These observations suggest that subcomponents of marine DOM such as HMW polysaccharides cycle faster (1-25 years) than its mean residence time consistent with the hypothesis that DOM is composed of several carbon reservoirs that turnover at different rates. Investigating the biological and chemical basis of DOM lability is critical to our understanding of the impact of this global reservoir of exchangeable carbon on Earth's carbon cycle and climate.

Atomic ratio analysis of the principal bio-elements of organic matter can be used as an indicator of its degradation state in the environment. For example, marine planktonic biomass has an average carbon, nitrogen, and phosphorus stoichiometry (C:N:P) of 106:16:1 (Redfield,

Ketchum, & Richards, 1963; Redfield, 1934). In contrast, HMW DOM is depleted in nitrogen and phosphorus relative to carbon (C:N:P=298:18:1; Fig. 1.2; Benner, 2002) regardless of the depth at which it is sampled (Benner, Pakulski, McCarthy, Hedges, & Hatcher, 1992; Mccarthy et al., 1996). The major organic nitrogen and phosphorus constituents of HMW DOM are also different than from plankton biomass or particulate organic matter. Nuclear magnetic resonance analysis has revealed the major organic constituents of organic matter. HMW DOM from open-ocean surface waters contains on average 40-60 nmol N per $\mu\text{mol C}$ primarily in the form amino polysaccharides (Aluwihare, Repeta, Pantoja, & Johnson, 2005; Mccarthy et al., 1996) and only 3-4 nmol P per $\mu\text{mol C}$ as part of phosphate esters and phosphonate molecules (Clark & Ingall, 1998; Kolowitz, Ingall, & Benner, 2001) wherase newly produced biomass is dominated by nitrogen in the form of amino acids (Cowie & Hedges, 1992) and phosphorus in the form of phosphate esters (Kolowitz, Ingall, & Benner, 2001). These observations suggest that selective microbial decomposition processes, particularly regenerating labile forms of organic nitrogen and phosphorus, contribute to the accumulation of carbon-rich HMW DOM in the water column. However, without direct evidence constraining the source of HMW DOM alternative hypothesis explaining the observed elemental stoichiometry must be considered.

The ^{13}C -NMR spectrum of HMW DOM from marine surface waters indicates it is largely composed of carbohydrates (50%), humic substances (25%), and a lesser amount of lipids and proteins (Aluwihare et al., 1997; Benner et al., 1992). Remarkably, the carbohydrate NMR-based molecular signature of HMW DOM is highly conserved across the global ocean and is even similar to HMW DOM in some freshwater environments (Aluwihare & Repeta, 1999; Aluwihare et al., 1997; Mccarthy et al., 1996). In fact, HMW DOM largely consists of a conserved family of closely related heteropolysaccharides with a predictable monosaccharide composition

including arabinose, rhamnose, fucose, xylose, mannose, galactose, and glucose (Aluwihare et al., 1997) which suggests that widespread food web processes, if not the same organisms, produce this material. In addition to its complex sugar composition, HMW DOM polysaccharides are substituted with acetyl (Aluwihare et al., 1997), amino (Aluwihare et al., 2005), deoxy, methyl (Panagiotopoulos, Repeta, & Johnson, 2007; Quan & Repeta, 2007), and phosphono sugars (Daniel Repeta, personal communication) with moderate branching indicating these polysaccharides serve a very different function than, for example, cellular glucose polysaccharides used as carbon and energy reserve like glycogen and starch. It is possible that such chemical moieties, like the acetyl and amino functional groups, represent available substrates for heterotrophic bacteria. However, the chemical substitutions may render HMW DOM polysaccharides recalcitrant, for example, by impeding binding by carbohydrate-degrading enzymes. This would explain in part the semi-labile nature of HMW DOM. Identifying bacterioplankton groups that breakdown and metabolize HMW DOM is expected to enable testing this hypothesis and to shed light on the biochemical processes underlying DOM cycling.

The role of bacterioplankton in the cycling of DOM in the sea

The emergent role of bacterioplankton as a main player in the ocean's carbon cycle and novel insights on the bulk composition of organic matter set the stage to investigate in more detail individual components and processes of DOM cycling. The microbial decomposition of marine DOM in some ways resembles the gradual decomposition of organic carbon in terrestrial soils, which consists of the stepwise breakdown of a easily degradable, intermediate, and resistant fractions of material (Wolf & Wagner, 2005). Decomposition of soil organic matter is driven primarily by the activities of bacteria and fungi (Hopkins & Gregorich, 2005) albeit fungi

are not prevalent in the open-ocean. Heterotrophic soil microorganisms fall under two main groups; those that respond primarily to the addition of fresh carbon substrates (r-selected biomass) and those that derive their energy mainly from the decomposition of older, more recalcitrant forms of organic carbon (autochthonous or K-selected biomass) (Hopkins & Gregorich, 2005). Similar ecological divisions related to substrate lability and resource availability are often used to describe marine bacterioplankton. Copiotrophic bacteria, often considered r-strategists, are better adapted to nutrient-rich environments and display boom-and-bust population dynamics, while oligotrophic bacteria, or K-strategists, are associated with dilute and stable environments and steady growth (Koch, 2001). Moreover, genome properties of cultured bacteria, such as the number of regulatory networks and substrate transporters, seem to be consistent with these microbial life-styles (Lauro et al., 2009; Singer et al., 2011). Despite having this ecological information, it remains difficult to associate specific microbial groups with the decomposition of the different pools of marine DOM, particularly in the slow cycling of the less labile reservoirs.

Seawater incubation experiments, often referred to as microcosms, have been practical for studying short term responses of microbial communities to organic matter perturbations. Carlson et al. (2004) performed field studies in the Sargasso Sea consisting of microcosm incubations combining filtered seawater and microbial inocula from different depths to show that the DOM that accumulates periodically in surface waters becomes bioavailable in the presence of mesopelagic microbial communities whereas it appears to be refractory to surface microbial communities. The consumption of surface DOC exported to mesopelagic depths has also been observed with the use of biogeochemical tracers such as chlorofluorocarbon-based estimates of water mass ventilation age (Carlson et al., 2010), oxygen mass balance (Emerson, Quay, Stump,

Wilbur, & Schudlich, 1995), and DOC concentration gradients in the North Atlantic (Carlson, Ducklow, & Michaels, 1994; Hansell & Carlson, 2001). These field observations illustrate the semi-labile nature of DOM and highlight the importance of the composition of microbial heterotrophic communities in the cycling of organic carbon. However, it remains to be described which microbial groups are the most metabolically active and relevant to the degradation of this carbon reservoir.

The application of cultivation-independent techniques has been useful to identify bacterioplankton groups relevant to the degradation of operationally-defined fractions of organic matter. One approach has been to obtain profiles of the SSU rRNA gene pool from a microbial sample *in situ* or in controlled experiments. For example, 16S rRNA gene clone libraries obtained from environmental samples indicate that a large proportion of bacterial communities associated to particulate organic matter (POM) are closely related to the Cytophaga-Flavobacter and *Planctomycetes* taxa (DeLong, Franks, & Alldredge, 1993; Fontanez, Eppley, Samo, Karl, & DeLong, 2015). These microbial assemblages are thought to be responsible for the remineralization of sinking particles (Cho & Azam, 1988; Smith, Simon, Alldredge, & Azam, 1992) and play a central role in regulating the efficiency of the biological pump. Cottrell & Kirchman (2000) used microautoradiography combined with fluorescence *in situ* hybridization techniques (based on 16S rRNA gene) to investigate which microorganisms became enriched in seawater treatments with HMW or LMW substrates. The Cytophaga-Flavobacter group appeared to be associated with consumption of relevant marine HMW substrates like chitin and N-acetyl glucosamine while *Alphaproteobacteria* were more successful in competing for amino acids. Covert & Moran (2001) also observed considerable differences in the composition of bacterial

communities that developed in HMW or LMW DOM treatments in estuaries using terminal restriction fragment length polymorphism analysis of 16S rRNA gene sequences.

Most recently, more powerful cultivation-independent approaches like metagenomics and community gene expression analysis, or metatranscriptomics, have been employed to reveal the DOM degradation activities of microbes *in situ* or in microcosm experiments. In addition to surveying the SSU rRNA composition of microbial samples in response to experimental treatments or environmental stimuli, these techniques capture the metabolic potential of microbial communities which is particularly useful to infer transformations of matter. For example, McCarren et al. (2010) applied a metatranscriptomic analysis to an experimental microcosm of Hawaii surface water supplemented with natural HMW DOM. The data revealed an apparent succession of bacterial taxa including *Alteromonadales* and methylotrophic bacteria that were hypothesized to operate in syntrophy whereby the first group produced or released one-carbon compounds that supported the metabolism of the latter. In another cultivation-independent study, Teeling et al. (2012) employed metagenomics and proteomics to resolve microbial processes over the course of a bacterioplankton bloom in the North Sea. Similar to the McCarren et al. (2010) study, Teeling et al. (2012) observed a succession of taxa thought to be associated with changes in the organic matter field as the bloom progressed including an overrepresentation of *Bacteroidetes* populations encoding carbohydrate-active enzymes involved in the degradation of polysaccharides at the end of the bloom. These results are consistent with genomic studies of other cultured and non-cultured *Bacteroidetes* which predict these microorganisms are adapted to degrade biopolymers, including polysaccharides and proteins (González et al., 2008; Woyke et al., 2009).

Recent field studies investigating the activity of specific members of bacterioplankton in response to carbon amendments indicate that many taxonomic groups are more specialized than previously thought leading to the emerging hypothesis that the organic matter field may be responsible for structuring microbial communities (Dinasquet, Kragh, Schrøter, Søndergaard, & Riemann, 2013; Dixon, Sargeant, Nightingale, & Colin Murrell, 2013; Gifford, Sharma, Booth, & Moran, 2012; Gómez-Consarnau, Lindh, Gasol, & Pinhassi, 2012; Nelson & Carlson, 2012; Sarmiento & Gasol, 2012). The widespread occurrence of genome streamlining (Dupont et al., 2012; Giovannoni, Cameron Thrash, & Temperton, 2014; Grote et al., 2012; Woyke et al., 2009) also indicates that marine microbes occupy narrow ecological niches that originate in stable environments though small-genome size is not the only strategy to succeed in a specialized niche. For example, recently, single-cell genomics have revealed uncultured marine bacterioplankton from the phylum *Verrucomicrobia* with an estimated genome size of 3.2 to 5.7 Mb that may be key players in the degradation of polysaccharides due to the high content of carbohydrate-degrading enzymes (Martinez-Garcia et al., 2012). Nevertheless, even if compositional changes in the DOM pool result in alterations of the bacterioplankton community structure, the transformative potential of organic carbon may remain the same given the functional redundancy between taxa (Allison & Martiny, 2008). Moreover, trophic factors and physical conditions are thought to be more important drivers of community structure than resource availability such as the quality or quantity of DOM (Mou, Sun, Edwards, Hodson, & Moran, 2008). Hypothetically though, a highly stable environment like the North Pacific Subtropical Gyre, where bacterioplankton communities have evolved continuously exposed to a predominantly local source of organic matter, may exhibit a naturally orchestrated aerobic microbial decomposition chain wherein different bacterioplankton populations participate in the

slow decomposition of less labile DOM. This mechanism may be similar to the open-ocean bacterioplankton transcriptional cascades observed by Ottesen et al. (2014) associated with the tight coupling between primary producers and heterotrophic consumption processes, though these processes may be more representative of the cycling of labile substrates.

Even though the majority of microcosm experiments investigating the heterotrophic activities of bacterioplankton have only resolved broader group responses and employed model organic carbon substrates or operationally-defined size fractions of organic matter, their results indicate that many marine microbes may occupy specialized niches carved in adaptation to different sources of organic matter in the environment and the types and quantities of substrates available for growth. Cultivation-independent techniques like genomics – despite known methodological biases – have also provided numerous observations consistent with this hypothesis. However, more appropriate microbial model systems, i.e., cultivable and ecologically relevant microorganisms, are necessary to corroborate these findings.

Microbial model systems relevant to semi-labile DOM cycling

One of the main challenges in marine microbial ecology is the ability to cultivate marine microorganisms to study them in detail in pure culture. In contrast to traditional cultivation methods which employ pre-defined media enriched with organic and inorganic nutrients, the dilution-culturing approach (Button, Schut, Quang, Martin, & Robertson, 1993) makes an effort to simulate the environmental conditions of the microbial inocula by using a sterile form of natural, minimally altered water as growth medium to isolate strains from a highly dilute sample (Fig. 1.3). Modifications of the dilution-to-extinction method (Connon & Giovannoni, 2002) made it possible to cultivate diverse marine microorganisms including some of the most

abundant heterotrophic bacterioplankton clades like SAR11 (Rappé, Connon, Vergin, & Giovannoni, 2002) and Oligotrophic Marine Bacteria (OMG) (Cho & Giovannoni, 2004). Representative cultured members of the SAR11 clade have been explored to investigate the role of this bacterioplankton clade in the marine environment. Field studies, genome analysis and pure culture experiments indicate that this diverse marine clade is best suited to metabolize simple carbon compounds – e.g., acetate, one-carbon compounds, and amino acids – and are thought to play a major role in carbon cycling (Grote et al., 2012; Malmstrom, Kiene, Cottrell, & Kirchman, 2004; Schwalbach, Tripp, Steindler, Smith, & Giovannoni, 2010; Sun et al., 2011a). However, other abundant groups within the *Proteobacteria* such as the SAR86 clade remain uncultured (Dupont et al., 2012) reflecting our lack of understanding of the ecology and physiology of marine microbes.

Even though traditional and new cultivation techniques have generated large and diverse culture collections, most culture studies typically determine the carbon and energy requirements of new marine isolates by testing simple model substrates such as LMW organic acids and simple sugars (e.g., Baumann et al., 1972; Sun et al., 2011). For the purposes of investigating the implication of microbial heterotrophic processes in the marine carbon cycle, such an approach is incomplete because it ignores the gradual decomposition of heterogeneous organic matter components derived from particulates (e.g., sinking phytoplankton and marine snow) to HMW polymers, and ultimately to LMW oligomers and monomers amenable for cell uptake and growth. Though this linear representation of organic matter degradation oversimplifies the process it highlights the need for more ecologically-relevant microbial model systems and experimentation with naturally-occurring growth substrates to study in detail the cycling of DOM.

Thesis goal and working hypotheses

One of the primary objectives of this thesis is to develop a screening system to identify and cultivate ecologically-relevant marine microorganisms that are centrally involved in DOM decomposition in the sea. The central focus in this project is on marine HMW DOM, an abundant source of carbon that is relatively easy to concentrate from seawater without chemical biases. Because its molecular composition and spectroscopic properties are well conserved across the global ocean, HMW DOM is predicted to support widespread heterotrophic microbial populations. The second component is a robust microbiological screening and cultivation method aimed at isolating microorganisms with the phenotype of interest. The dilution-to-extinction technique has proven successful in cultivating previously uncultured bacterioplankton groups. In this thesis, a modification of this technique consisting of amending dilution-to-extinction cultures with HMW DOM is predicted to select for microorganisms adapted to degrade this naturally-occurring substrate to obtain carbon and energy for growth (Fig. 1.4). Having such microbial systems available in the laboratory is expected to enable the use of physiological, genomic, and genetic experimentation to test hypotheses similar to those generated by cultivation-independent studies and to reveal details about the microbial processes that contribute to HMW DOM turnover including:

1. The identity of microorganisms actively involved in HMW DOM turnover.
2. The types of HMW DOM carbon constituents available for bacterial growth.
3. The genes and enzymes catalyzing the degradation of identifiable HMW DOM constituents.
4. The exact biochemical transformations of these substrates.

The information contained in microbial genomes can be used to infer enzymatic capabilities and metabolic properties enabling HMW DOM degradation. Simple growth assays can help determine the extent to which different microorganisms degrade HMW DOM and the types of organic constituents they can metabolize. Additionally, knowledge of the chemical constitution of HMW DOM can help select appropriate analytical techniques to detect the expected biochemical transformations of HMW DOM substrates. For example, biotransformation assays consisting of batch microbial cultures enriched with HMW DOM can be used to identify decomposition byproducts or alterations in HMW DOM using analytical instruments such as high-performance liquid chromatography, gas chromatography, nuclear magnetic resonance, or mass spectrometry, or simply thin-layer chromatography. Moreover, the combination of genomic and physiological information can help design genetic tests to pin point the genes and biochemical pathways that encode the ability to decompose HMW DOM. Microbiological assays may thus provide a powerful framework to elucidate the decomposition reactions underlying DOM turnover, information that can help answer longstanding questions in biological and chemical oceanography.

One observation that has puzzled scientists studying marine DOM for over a decade, is its significant enrichment in carbohydrates, materials typically thought of as quite bioavailable. Paradoxically, microorganisms are well known to degrade complex carbohydrates such as the cellulose and pectin that provide plants structure and rigidity, and the chitin that makes up the exoskeletons of crustaceans. Another intriguing and important question is whether marine DOM will act as a sink or source of carbon during global environmental change (Jiao et al., 2014). Carbon locked away in the ocean by living organisms in marine sediments and in the water column represents a considerable and climate-sensitive, sink of anthropogenic CO₂ (Nellemann

et al., 2009). In the open ocean, the microbial carbon pump (MCP), a process by which Bacteria and Archaea produce refractory DOC, is thought to sequester atmospheric CO₂ for geological timescales (Jiao et al., 2014, 2010) and its magnitude may be responsive to changes in sea-surface temperature, water column stratification, and production of particulate organic carbon (Legendre et al., 2015). However, we still lack a mechanistic understanding of the biological processes that drive the opposite direction of carbon flow, the aerobic microbial remineralization of DOC to carbon dioxide in marine surface waters, including the physical and environmental parameters that influence these processes. The longstanding debate of whether the ocean is net autotrophic or net heterotrophic (Duarte, Regaudie-de-Gioux, Arrieta, Delgado-Huertas, & Agustí, 2013; Ducklow & Doney, 2013; Williams, Quay, Westberry, & Behrenfeld, 2013) clearly illustrates the difficulty in constraining the implication of microbial DOM cycling in the planet's carbon budget.

The goal of this thesis is to use cultivation-based techniques to elucidate the biochemical and genetic determinants underlying the decomposition of marine HMW DOM. Throughout this work, new and unexpected findings regarding the degradation properties of microbial isolates from coastal and open-ocean ecosystems have expanded our knowledge of how different functional bacterioplankton groups may participate in HMW DOM cycling and highlighted the potential of microbial-mediated carbon transformations to alter the environment and Earth's climate. Knowledge such as this is expected to contribute to a more mechanistic understanding of the aerobic decomposition of semi-labile DOM in the sea. The data chapters encompassing this work are outlined below.

Chapter summaries

Chapter 2: High-molecular-weight dissolved organic matter enrichments select for methylotrophs in dilution to extinction cultures.

In this chapter, I put in practice the modified dilution-to-extinction cultivation strategy designed to isolate microbial degraders of HMW DOM. The study was conducted using the endemic microbial community from a coastal marine environment (Nahant Bay, MA) and HMW DOM collected from oligotrophic surface waters in the open-ocean near Hawaii. Even though the microbial inoculum and HMW DOM samples were obtained from different environments the strategy was successful in identifying several groups of heterotrophic bacteria that could use HMW DOM for growth supporting the hypothesis that this abundant substrate is well conserved across the global ocean and can be degraded by predictable and widespread functional groups of bacteria. The work encompassing this chapter was originally published by Sosa, Gifford, Repeta, & DeLong (2015) and represents the foundation for the work developed in the subsequent chapters.

Chapter 3: Assessment of the degradation capabilities of open-ocean bacterial isolates enriched with marine high-molecular-weight dissolved organic matter

Because microbial heterotrophic communities may be better adapted to degrade co-occurring sources of organic matter, in Chapter 3 I re-applied the modified screening methodology to identify microbial degraders of HMW DOM from an open-ocean environment. For this purpose, a series of cultivation experiments were conducted with microbial inocula and HMW DOM samples collected in the open-ocean near Hawaii. The results of this work were similar to Chapter 2 in that several bacterial groups were stimulated by HMW DOM amendments but different in that most isolates had a reduced ability to utilize the substrate for

growth. Nevertheless, genomic analysis revealed some clues regarding the enzymatic degradation potential of these cultures to transform HMW DOM substrates, information that is expected to guide future investigations.

Chapter 4: Production of hydrocarbon gases during bacterial degradation of covalently-bound phosphonates associated with marine high-molecular-weight dissolved organic matter

In Chapter 4, I delve deeper into the characterization of two such bacterial isolates from the open-ocean able to utilize known organic constituents of marine HMW DOM polysaccharides, namely, phosphonates – molecules with a direct carbon-phosphorus (C-P) bond rather than an oxygen ester (R-O-P) bond. In this chapter, pure culture studies in combination with analytical techniques used to detect the degradation byproducts of HMW DOM demonstrate the usefulness of the microbial model system approach to dissect the underlying biochemical transformations that drive the cycling of DOM in the ocean.

Chapter 5: Summary and future directions

Finally, chapter 5 summarizes the findings of this thesis, discusses the limitations of the microbial model system approach, and suggests ways in which the information gathered throughout this work can help guide future experiments with candidate microbial decomposers of HMW DOM identified in chapters 2 and 3.

Figures

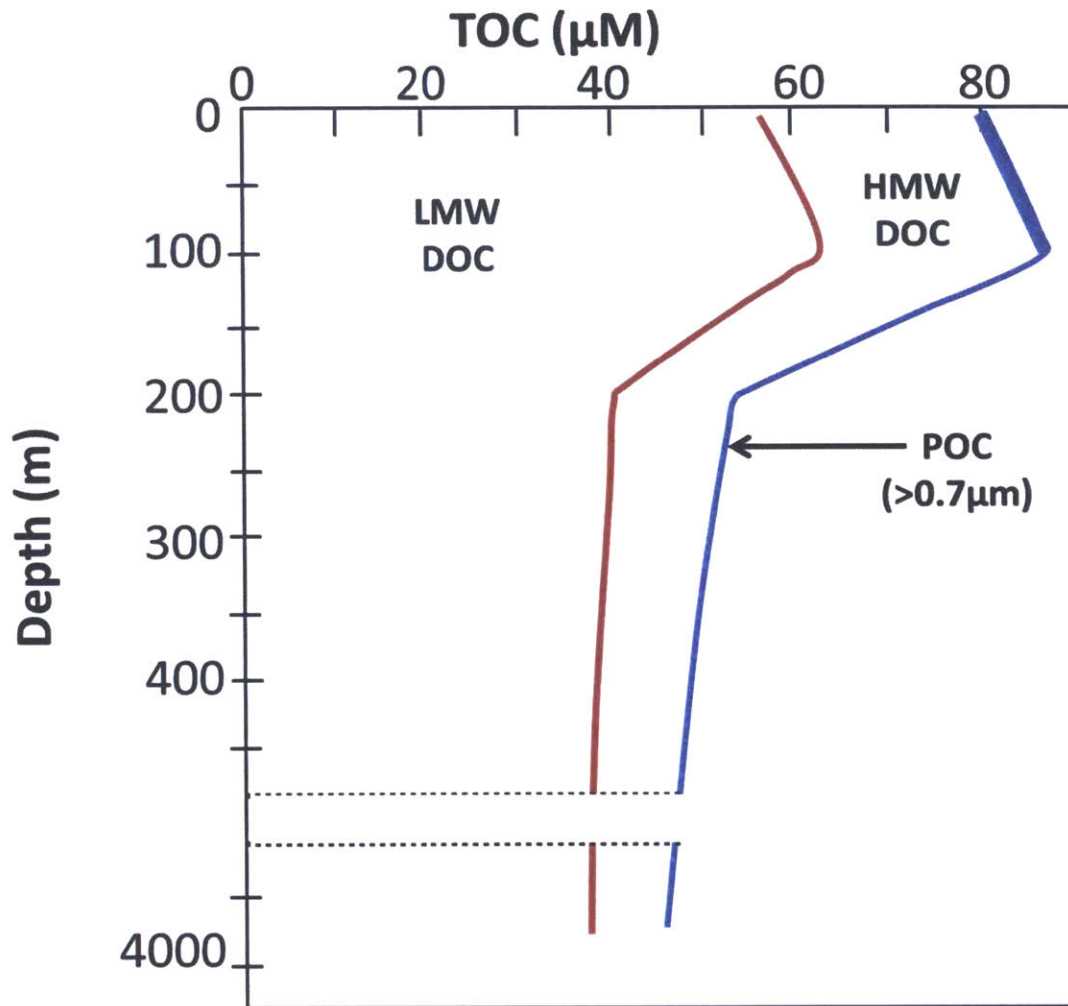


Figure 1.1. Size distribution and abundance depth profile of total organic carbon (TOC) characteristic of open ocean regions. Low-molecular-weight dissolved organic carbon (LMW DOC) dominates the TOC inventory (area left to the red line). In the surface regions high-molecular-weight (HMW) DOC accounts for up to ~29% of TOC and diminishes to ~18% at depth (area between the red and blue lines). Particulate organic carbon is approximately ~1% of TOC. Profiles are based on data from Benner et al. (1997).

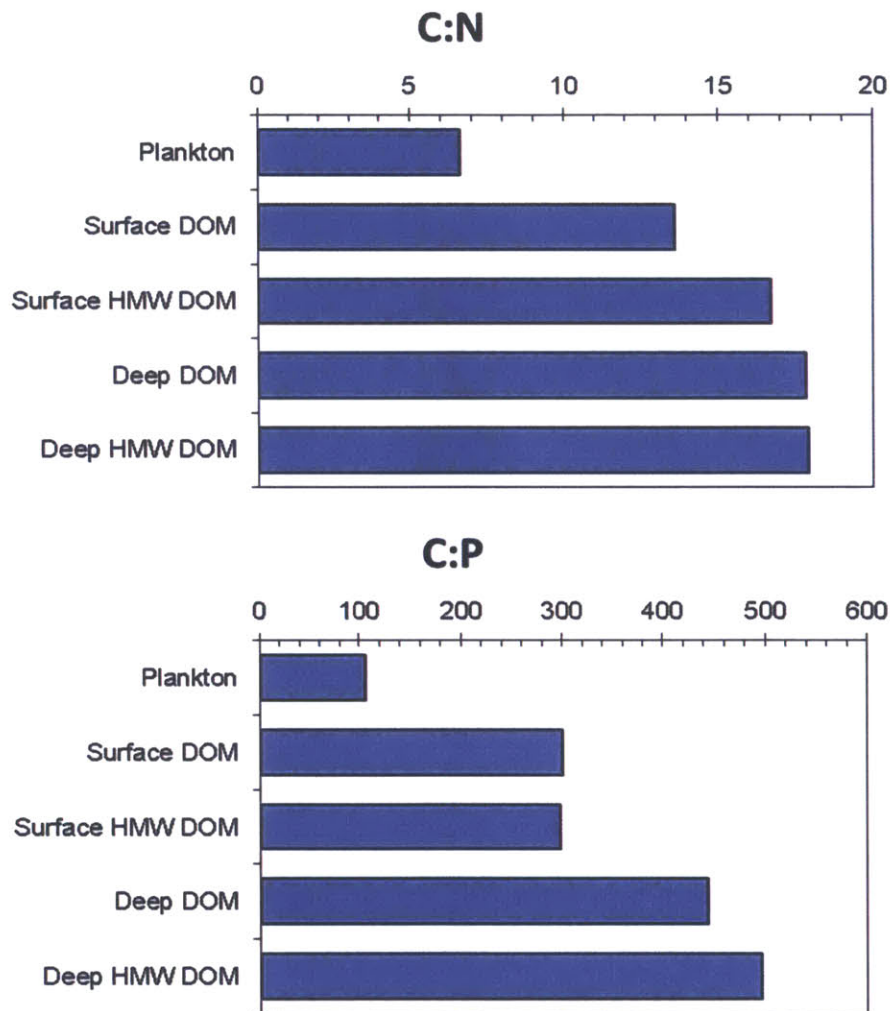


Figure 1.2. Elemental stoichiometry of marine organic matter. High-molecular-weight (HMW) dissolved organic matter (DOM) has carbon to nitrogen atomic ratios (C:N) ~2.5 times higher than plankton biomass, and slightly higher ratios than the total DOM pool. The carbon to phosphorus atomic ratios (C:P) are similarly ~3 times higher than for plankton. Both C:N and C:P of HMW DOM increase with depth. Elemental stoichiometry is based on data summarized by Benner (2002).

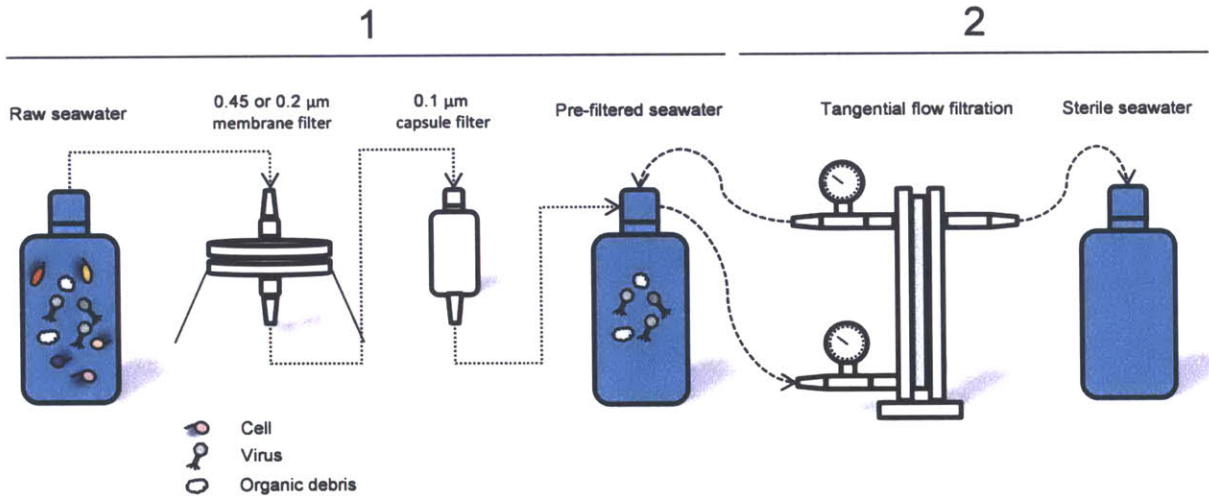


Figure 1.3. Tangential flow filtration (TFF) methodology for the sterilization of natural seawater media used in dilution-to-extinction experiments and culturing. Seawater is collected in acid-clean polycarbonate carboys and stored preferably in darkness and chilled before sterilization. The first step in the process is pre-filtration. Normally, two pre-filters are used to remove particulates and cells from seawater. The first pre-filter used for this work was a 0.22 μm Supor membrane filter (PALL Corp.) to remove most cells. The second pre-filter was 0.1 μm double membrane capsule filter to ensure that even small cells were removed. The second step in the protocol is to use TFF with a 30 kDa membrane cassette to remove any remaining cells as well as viruses and a fraction of high-molecular-weight dissolved organic matter, like organic debris from cells and particulates. The permeate of the TFF membrane is collected in sterile polycarbonate bottles and stored at 4°C on darkness until used for cultivation.

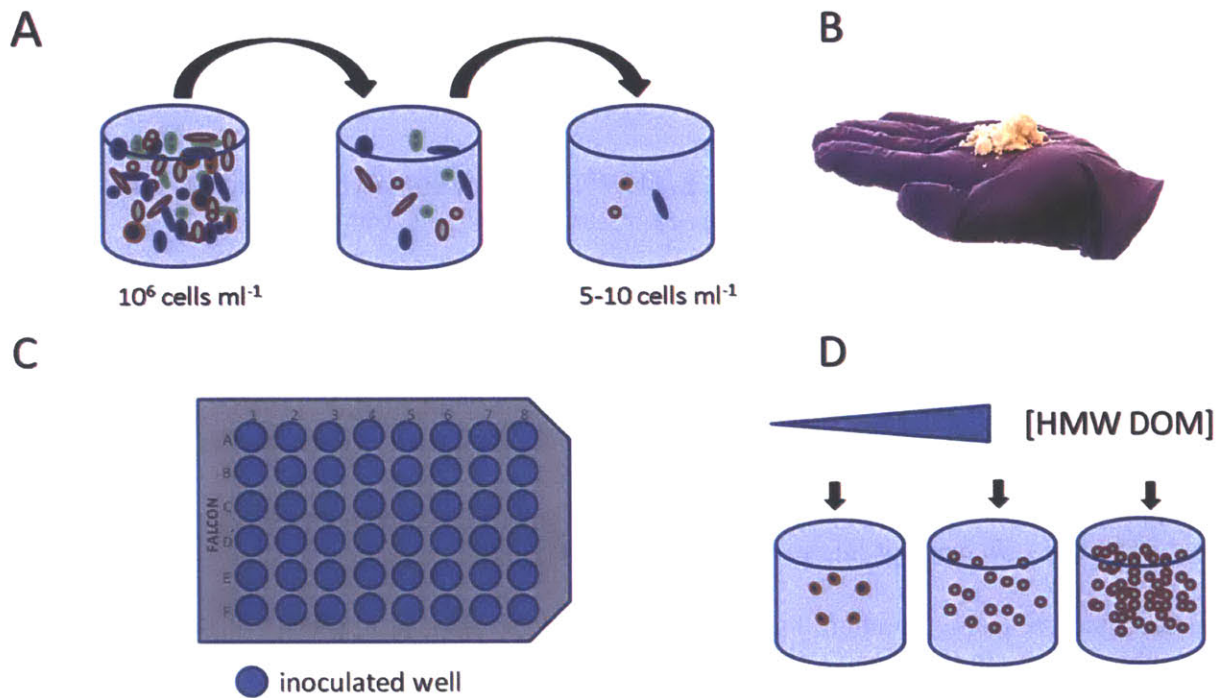


Figure 1.4. Enrichment of dilution-to-extinction cultures with high-molecular-weight (HMW) dissolved organic matter (DOM). A. The microbial community of a natural seawater sample is diluted with sterile natural seawater (Fig. 1.3) to a theoretical cell density of $5-10 \text{ cells ml}^{-1}$. B. HMW DOM collected by ultrafiltration is added to the dilution-to-extinction sample to increase the dissolved organic carbon content 5-10X over the background carbon ($\sim 60-80 \mu\text{M}$), though organic forms of nitrogen and phosphorus are also present in HMW DOM. C. One milliliter aliquots of the amended dilution-to-extinction samples are arrayed into 48-well polystyrene cultivation plates (Falcon™ plates are commercially available through Fisher Scientific). D. The addition of HMW DOM is hypothesized to stimulate the growth of microorganisms capable of degrading HMW DOM substrates.

References

- Allison, S. D., & Martiny, J. B. H. (2008). Resistance, resilience, and redundancy in microbial communities, *Proceedings of the National Academy of Sciences of the United States of America*, *105*, 11512–11519.
- Aluwihare, L. I., & Repeta, D. J. (1999). A comparison of the chemical characteristics of oceanic DOM and extracellular DOM produced by marine algae. *Marine Ecology Progress Series*, *186*, 105–117.
- Aluwihare, L. I., Repeta, D. J., & Chen, R. F. (1997). A major biopolymeric component to dissolved organic carbon in surface sea water. *Nature* *387*, 166–169.
- Aluwihare, L. I., Repeta, D. J., Pantoja, S., & Johnson, C. G. (2005). Two chemically distinct pools of organic nitrogen accumulate in the ocean. *Science*, *308*, 1007–1010.
- Azam, F., Fenchel, T., Field, J., Gray, J., Meyer-Reil, L., & Thingstad, F. (1983). The ecological role of water-column microbes in the sea. *Marine Ecology Progress Series*, *10*, 257–263.
- Baumann, L., Baumann, P., Mandel, M., & Allen, R. D. (1972). Taxonomy of aerobic marine eubacteria. *Journal of Bacteriology*, *110*(1), 402–429.
- Béjà, O., Spudich, E. N., Spudich, J. L., Leclerc, M., & DeLong, E. F. (2001). Proteorhodopsin phototrophy in the ocean. *Nature*, *411*, 786–789.
- Benner, R. (2002). Chemical Composition and Reactivity. In Hansell, D. A. and Carlson, C. A. (Eds.). *Biogeochemistry of Marine Dissolved Organic Matter* (pp. 59–90). Academic Press, San Diego.
- Benner, R., Biddanda, B., Black, B., & McCarthy, M. (1997). Abundance, size distribution, and stable carbon and nitrogen isotopic compositions of marine organic matter isolated by tangential-flow ultrafiltration. *Marine Chemistry*, *57*, 243–263.
- Benner, R., Pakulski, J. D., McCarthy, M., Hedges, J. I., & Hatcher, P. G. (1992). Bulk chemical characteristics of dissolved organic matter in the ocean. *Science*, *255*, 1561–1564.
- Button, D. K., Schut, F., Quang, P., Martin, R., & Robertson, B. R. (1993). Viability and isolation of marine bacteria by dilution culture: Theory, procedures, and initial results. *Applied and Environmental Microbiology*, *59*(3), 881–891.
- Carlson, C. A. (2002). Production and removal processes. In Hansell, D. A. and Carlson, C. A. (Eds.). *Biogeochemistry of Marine Dissolved Organic Matter* (pp. 139–151). Academic Press, San Diego.

- Carlson, C. A., Giovannoni, S. J., Hansell, D. A., Goldberg, S. J., Parsons, R., & Vergin, K. (2004). Interactions among dissolved organic carbon, microbial processes, and community structure in the mesopelagic zone of the northwestern Sargasso Sea. *Limnology and Oceanography*, *49*(4), 1073–1083.
- Carlson, C. A., & Ducklow, H. W. (1995). Dissolved organic carbon in the upper ocean of the central equatorial Pacific Ocean, 1992: Daily and finescale vertical variations. *Deep Sea Research Part II: Topical Studies in Oceanography*, *42*(2), 639–656.
- Carlson, C. A., & Ducklow, H. W. (1996). Growth of bacterioplankton and consumption of dissolved organic carbon in the Sargasso Sea. *Aquatic Microbial Ecology*, *10*, 69–85.
- Carlson, C. A., Ducklow, H. W., & Michaels, A. F. (1994). Annual flux of dissolved organic carbon from the euphotic zone in the northwestern Sargasso Sea. *Nature*, *371*, 405–408.
- Carlson, C. A., Hansell, D. A., Nelson, N. B., Siegel, D. a., Smethie, W. M., Khatiwala, S., ... Halewood, E. (2010). Dissolved organic carbon export and subsequent remineralization in the mesopelagic and bathypelagic realms of the North Atlantic basin. *Deep-Sea Research Part II: Topical Studies in Oceanography*, *57*(16), 1433–1445.
- Carlson, D., Brann, M., Mague, T., & Mayer, L. (1985). Molecular weight distribution of dissolved organic materials in seawater determined by ultrafiltration: a re-examination. *Marine Chemistry*, *16*, 155–171.
- Cherrier, J., Bauer, J., & Druffel, E. R. M. (1996). Utilization and turnover of labile dissolved organic matter by bacterial heterotrophs in eastern North Pacific surface waters. *Marine Ecology Progress Series*, *139*, 305–307.
- Cho, B. C., & Azam, F. (1988). Major role of bacteria in biogeochemical fluxes in the ocean's interior. *Nature* *332*, 441–443.
- Cho, J. C., & Giovannoni, S. J. (2004). Cultivation and growth characteristics of a diverse group of oligotrophic marine *Gammaproteobacteria*. *Applied and Environmental Microbiology*, *70*(1), 432–440.
- Clark, L. L., & Ingall, E. D. (1998). Marine phosphorus is selectively remineralized. *Nature*, *393*, 426.
- Connon, S. A. & Giovannoni, S. J. (2002). High-throughput methods for culturing microorganisms in very-low-nutrient media yield diverse new marine isolates high-throughput methods for culturing microorganisms in very-low-nutrient media yield diverse new marine isolates. *Applied and Environmental Microbiology*, *68*(8), 3878–3885.
- Cottrell, M. T., & Kirchman, D. L. (2000). Natural assemblages of marine *Proteobacteria* and members of the Cytophaga-Flavobacter clustler consuming low- and high- molecular-

- weight dissolved organic matter. *Applied and Environmental Microbiology*, 66(4), 1692–1697.
- Covert, J. S., & Moran, M. A. (2001). Molecular characterization of estuarine bacterial communities that use high- and low-molecular weight fractions of dissolved organic carbon. *Aquatic Microbial Ecology*, 25, 127–139.
- Cowie, G. L., & Hedges, J. I. (1992). Sources and reactivities of amino acids in a coastal marine environment. *Limnology and Oceanography*, 37(4), 703–724.
- Davis, J., & Benner, R. (2007). Quantitative estimates of labile and semi-labile dissolved organic carbon in the western Arctic Ocean: A molecular approach. *Limnology and Oceanography*, 52(6), 2434–2444.
- Delong, E. F., Franks, D. G., & Alldredge, A. L. (1993). Phylogenetic diversity of aggregate-attached vs. free-living marine bacterial assemblages. *Limnology and Oceanography*, 38(5), 924–934.
- Dinasquet, J., Kragh, T., Schrøter, M.-L., Søndergaard, M., & Riemann, L. (2013). Functional and compositional succession of bacterioplankton in response to a gradient in bioavailable dissolved organic carbon. *Environmental Microbiology*, 15, 2616–2628.
- Dixon, J. L., Sargeant, S., Nightingale, P. D., & Colin Murrell, J. (2013). Gradients in microbial methanol uptake: productive coastal upwelling waters to oligotrophic gyres in the Atlantic Ocean. *The ISME Journal*, 7(3), 568–80.
- Druffel, E. R. M., Williams, P. M., Bauer, E., & Ertel, J. R. (1992). Cycling of dissolved and particulate organic matter in the open ocean. *Journal of Geophysical Research*, 97, 15639–15659.
- Druffel, E. R. M., Williams, P. M., & Suzuki, Y. (1989). Concentrations and radiocarbon signatures of dissolved organic matter in the Pacific Ocean. *Geophysical Research Letters*, 16(9), 991–994.
- Duarte, C. M., Regaudie-de-Gioux, A., Arrieta, J. M., Delgado-Huertas, A., & Agustí, S. (2013). The oligotrophic ocean is heterotrophic. *Annual Review of Marine Science*, 5, 551–569.
- Ducklow, H. W., & Carlson, C. A. (1992). Oceanic Bacterial Production. *Advances in Microbial Ecology*, 12(1986), 113–181.
- Ducklow, H. W., & Doney, S. C. (2013). What is the metabolic state of the oligotrophic ocean? A debate. *Annual Review of Marine Science*, 5, 525–533.
- Ducklow, H. W., & Hill, S. M. (1985). The growth of heterotrophic bacteria in the surface waters of warm core rings. *Limnology and Oceanography*, 30, 239–259.

- Dupont, C. L., Rusch, D. B., Yooseph, S., Lombardo, M.-J., Alexander Richter, R., Valas, R., ... Craig Venter, J. (2012). Genomic insights to SAR86, an abundant and uncultivated marine bacterial lineage. *The ISME Journal*, 6(6), 1186–1199.
- Emerson, S., & Hedges, J. I. (1988). Processes controlling the organic carbon content of open ocean sediments. *Paleoceanography*, 3(5), 621.
- Emerson, S., Quay, P. D., Stump, C., Wilbur, D., & Schudlich, R. (1995). Chemical tracers of productivity and respiration in the subtropical Pacific Ocean. *Journal of Geophysical Research*, 100, 15873–15887
- Eswaran, H., Van Den Berg, E., & Reich, P. (1993). Organic carbon in soils of the world. *Soil Science Society of America Journal*, 57, 192–194.
- Fontanez, K. M., Eppley, J. M., Samo, T. J., Karl, D. M., & DeLong, E. F. (2015). Microbial community structure and function on sinking particles in the North Pacific Subtropical Gyre. *Frontiers in Microbiology*, 6, 469.
- Fuhrman, J. A., & Azam, F. (1980). Bacterioplankton secondary production estimates for coastal waters of British-Columbia, Canada, Antarctica, and California, USA. *Applied and Environmental Microbiology*, 39(6), 1085–1095.
- Fuhrman, J. A., & Azam, F. (1982). Thymidine incorporation as a measure of heterotrophic bacterioplankton production in marine surface waters: Evaluation and field results. *Marine Biology*, 66, 109–120.
- Fuhrman, J. a, Ducklow, H. W., Kirchman, D. L., Hudak, J., McManus, G. B., & Kramer, J. (1986). Does adenine incorporation into nucleic acids measure total microbial production? *Limnology and Oceanography*, 31(3), 627–636.
- Fuhrman, J. A., Sleeter, T. D., Carlson, C. A., & Proctor, L. M. (1989). Dominance of bacterial biomass in the Sargasso Sea and its ecological implications. *Marine Ecology Progress Series*, 57, 207–217.
- Gifford, S. M., Sharma, S., Booth, M., & Moran, M. A. (2012). Expression patterns reveal niche diversification in a marine microbial assemblage, 7(2), 281–298.
- Giovannoni, S. J., Cameron Thrash, J., & Temperton, B. (2014). Implications of streamlining theory for microbial ecology. *The ISME Journal*, 8(8), 1553–1565.
- Gómez-Consarnau, L., Lindh, M. V., Gasol, J. M., & Pinhassi, J. (2012). Structuring of bacterioplankton communities by specific dissolved organic carbon compounds. *Environmental Microbiology*, 14, 2361–2378.
- González, J. M., Fernández-Gómez, B., Fernández-Guerra, A., Gómez-Consarnau, L., Sánchez, O., Coll-Lladó, M., ... Pedrós-Alió, C. (2008). Genome analysis of the proteorhodopsin-

- containing marine bacterium *Polaribacter* sp. MED152 (Flavobacteria). *Proceedings of the National Academy of Sciences of the USA*, 105(25), 8724–8729.
- Grote, J., Cameron Thrash, J., Huggett, M. J., Landry, Z. C., Carini, P., Giovannoni, S. J., & Rappé, M. S. (2012). Streamlining and core genome conservation among highly divergent members of the SAR11 clade. *mBio*, 3.
- Hansell, D. A., & Carlson, C. A. (2001). Marine Dissolved Organic Matter and the Carbon Cycle. *Society*, 14(4), 41–49.
- Hansell, D. A., & Carlson, C. A. (1998). Deep-ocean gradients in the concentration of dissolved organic carbon. *Nature*, 395, 263–266.
- Hansell, D. A., Carlson, C. A., Repeta, D. J., & Schlitzer, R. (2009). Dissolved organic matter in the ocean: A controversy stimulates new insights. *Oceanography*, 22(4), 202–211.
- Hedges, J. I., Keil, R. G., & Benner, R. (1997). What happens to terrestrial organic matter in the ocean? *Organic Geochemistry*, 27(5), 195–212.
- Hopkins, D. W., & Gregorich, E. G. (2005). Carbon as a substrate for soil organisms. In R. D. Bardgett, M. B. Usher, & D. W. Hopkins (Eds.), *Biological diversity and function in soils* (pp. 57–79). Cambridge University Press.
- Jiao, N., Herndl, G. J., Hansell, D. a., Benner, R., Kattner, G., Wilhelm, S. W., ... Azam, F. (2010). Microbial production of recalcitrant dissolved organic matter: long-term carbon storage in the global ocean. *Nature Reviews Microbiology*, 8(8), 593–599.
- Jiao, N., Robinson, C., Azam, F., Thomas, H., Baltar, F., Dang, H., ... Zhang, R. (2014). Mechanisms of microbial carbon sequestration in the ocean – future research directions. *Biogeosciences*, 11, 5285–5306.
- Karl, D. M. (2002). Nutrient dynamics in the deep blue sea. *Trends in Microbiology*, 10(9), 410–418.
- Keeling, R. (2015). *Measurement Note: An Adjustment to the Record*. Retrieved December 6, 2015, from <https://scripps.ucsd.edu/programs/keelingcurve/>
- Kirchman, D., K'nees, E., & Hodson, R. (1985). Leucine incorporation and its potential as a measure of protein synthesis by bacteria in natural aquatic systems. *Applied and Environmental Microbiology*, 49(3), 599–607.
- Koch, A. L. (2001). Oligotrophs versus copiotrophs. *BioEssays*, 23, 657–661.
- Kolowitz, L. C., Ingall, E. D., & Benner, R. (2001). Composition and cycling of marine organic phosphorus. *Limnology and Oceanography*, 46(2), 309–320.

- Lauro, F. M., McDougald, D., Thomas, T., Williams, T. J., Egan, S., Rice, S., ... Cavicchioli, R. (2009). The genomic basis of trophic strategy in marine bacteria. *Proceedings of the National Academy of Sciences of the USA*, *106*, 15527–15533.
- Legendre, L., Rivkin, R. B., Weinbauer, M. G., Guidi, L., & Uitz, J. (2015). The microbial carbon pump concept: Potential biogeochemical significance in the globally changing ocean. *Progress in Oceanography*, *134*, 432–450.
- Luo, Y. W., Friedrichs, M. a M., Doney, S. C., Church, M. J., & Ducklow, H. W. (2010). Oceanic heterotrophic bacterial nutrition by semilabile DOM as revealed by data assimilative modeling. *Aquatic Microbial Ecology*, *60*, 273–287.
- Malmstrom, R. R., Kiene, R. P., Cottrell, M. T., & Kirchman, D. L. (2004). Contribution of SAR11 bacteria to dissolved dimethylsulfoniopropionate and amino acid uptake in the North Atlantic Ocean. *Applied and Environmental Microbiology*, *70*(7), 4129–4135.
- Martinez-Garcia, M., Brazel, D. M., Swan, B. K., Arnosti, C., Chain, P. S. G., Reitenga, K. G., ... Stepanauskas, R. (2012). Capturing single cell genomes of active polysaccharide degraders: An unexpected contribution of verrucomicrobia. *PLoS ONE*, *7*(4), 1–11.
- McCarren, J., Becker, J. W., Repeta, D. J., Shi, Y., Young, C. R., Malmstrom, R. R., ... DeLong, E. F. (2010). Microbial community transcriptomes reveal microbes and metabolic pathways associated with dissolved organic matter turnover in the sea. *Proceedings of the National Academy of Sciences of the USA* *107*, 16420–16427.
- McCarthy, M. D., Hedges, J. I., & Benner, R. (1993). The chemical composition of dissolved organic matter in seawater. *Chemical Geology*, *107*, 503–507.
- McCarthy, M., Hedges, J., & Benner, R. (1996). Major biochemical composition of dissolved high molecular weight organic matter in seawater. *Marine Chemistry* *55*, 281–297.
- Mou, X., Sun, S., Edwards, R. A., Hodson, R. E., & Moran, M. A. (2008). Bacterial carbon processing by generalist species in the coastal ocean. *Nature* *451*.
- Nebbioso, A., & Piccolo, A. (2013). Molecular characterization of dissolved organic matter (DOM): A critical review. *Analytical and Bioanalytical Chemistry*, *405*, 109–124.
- Nellemann, C., Corcoran, E., Duarte, C. M., Valdés, L., De Young, C., Fonseca, L., & Grimsditch, G. (2009). *Blue carbon: A Rapid Response Assessment*. (p. 71 pp.). United Nations Environment Programme, GRID-Arendal. Retrieved from http://www.grida.no/files/publications/blue-carbon/BlueCarbon_screen.pdf
- Nelson, C. E., & Carlson, C. a. (2012). Tracking differential incorporation of dissolved organic carbon types among diverse lineages of Sargasso Sea bacterioplankton. *Environmental Microbiology*, *14*, 1500–1516.

- Ogawa, H., & Tanoue, E. (2003). Dissolved organic matter in oceanic waters. *Journal of Oceanography*, 59, 129–147.
- Ottesen, E. A., Young, C. R., Gifford, S. M., Eppley, J. M., Marin, R., Schuster, S. C., ... DeLong, E. F. (2014). Multispecies diel transcriptional oscillations in open ocean heterotrophic bacterial assemblages. *Science*, 345, 207–12.
- Panagiotopoulos, C., Repeta, D. J., & Johnson, C. G. (2007). Characterization of methyl sugars, 3-deoxysugars and methyl deoxysugars in marine high molecular weight dissolved organic matter. *Organic Geochemistry*, 38, 884–896.
- Pomeroy, L. (1974). The ocean's food web, a changing paradigm. *BioScience* 24, 499-504.
- Quan, T. M., & Repeta, D. J. (2007). Periodate oxidation of marine high molecular weight dissolved organic matter: Evidence for a major contribution from 6-deoxy- and methyl sugars. *Marine Chemistry*, 105, 183–193.
- Rappé, M. S., Connon, S. a, Vergin, K. L., & Giovannoni, S. J. (2002). Cultivation of the ubiquitous SAR11 marine bacterioplankton clade. *Nature*, 418(August), 630–633.
- Redfield, A C., Ketchum, B. H., & Richards, F. A. (1963). The influence of organisms on the composition of sea water. In Hill, M. N. (Ed.) *The Sea*, vol. 2. (pp. 26–27). Interscience Publishers, New York.
- Redfield, A. C. (1934). On the proportions of organic derivatives in sea water and their relation to the composition of plankton. In *James Johnstone Memorial Volume* (pp. 176-192). University Press of Liverpool, Liverpool, U. K.
- Repeta, D. J., & Aluwihare, L. I. (2006). Radiocarbon analysis of neutral sugars in high-molecular-weight dissolved organic carbon: Implications for organic carbon cycling. *Limnology and Oceanography*, 51(2), 1045–1053.
- Repeta, D. J., Ferron, S., Sosa, O. A., DeLong, E. F., & Karl, D. M. (2016). Title of paper. Unpublished manuscript.
- Sarmiento, H., & Gasol, J. M. (2012). Use of phytoplankton-derived dissolved organic carbon by different types of bacterioplankton. *Environmental Microbiology*, 14, 2348–2360.
- Schwalbach, M. S., Tripp, H. J., Steindler, L., Smith, D. P., & Giovannoni, S. J. (2010). The presence of the glycolysis operon in SAR11 genomes is positively correlated with ocean productivity. *Environmental Microbiology*, 12, 490–500.
- Singer, E., Webb, E. A., Nelson, W. C., Heidelberg, J. F., Ivanova, N., Pati, A., & Edwards, K. J. (2011). Genomic potential of *Marinobacter aquaeolei*, a biogeochemical “Opportunitroph.” *Applied and Environmental Microbiology*, 77(8), 2763–2771.

- Smith, D. C., Simon, M., Alldredge, A. L., & Azam, F. (1992). *Nature*, 359, 139–142.
- Sosa, O. A, Gifford, S. M., Repeta, D. J., & DeLong, E. F. (2015). High molecular weight dissolved organic matter enrichment selects for methylotrophs in dilution to extinction cultures. *The ISME Journal*, 9, 2725–2739.
- Sun, J., Steindler, L., Thrash, J. C., Halsey, K. H., Smith, D. P., Carter, A. E., ... Giovannoni, S. J. (2011). One carbon metabolism in SAR11 pelagic marine bacteria. *PLoS ONE*, 6(8), e23973.
- Teeling, H., Fuchs, B. M., Becher, D., Klockow, C., Gardebrecht, a., Bennke, C. M., ... Amann, R. (2012). Substrate-controlled succession of marine bacterioplankton populations induced by a phytoplankton bloom. *Science*, 336, 608–611.
- Williams, P. J. L. B., Quay, P. D., Westberry, T. K., & Behrenfeld, M. J. (2013). The oligotrophic ocean is autotrophic. *Annual Review of Marine Science*, 5, 535–549.
- Williams, P. J. leB. (1981). Microbial contribution to overall marine plankton · metabolism: direct measurements of respiration. *Oceanologica Acta*, 4, 359–364.
- Williams, P. M., & Druffel, E. R. M. (1987). Radiocarbon in dissolved organic matter in the central North Pacific Ocean. *Nature*, 330, 246–248.
- Wolf, D., & Wagner, G. (2005). Carbon transformations and soil organic matter formation. In D. Sylvia, J. Fuhrmann, P. Hartel, & D. Zuberer (Eds.), *Principles and applications of soil microbiology* (pp. 285–232). Pearson Prentice Hall.
- Woyke, T., Xie, G., Copeland, A., González, J. M., Han, C., Kiss, H., ... Stepanauskas, R. (2009). Assembling the marine metagenome, one cell at a time. *PLoS ONE*, 4(4).

WAVES ARE NOT MEASURED IN FEET AND INCHES, BUT IN INCREMENTS OF
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CHAPTER TWO

High-molecular-weight dissolved organic matter enrichment selects for methylotrophs in dilution-to-extinction cultures

Abstract

The role of bacterioplankton in the cycling of marine dissolved organic matter (DOM) is central to the carbon and energy balance in the ocean, yet there are few model organisms available to investigate the genes, metabolic pathways, and biochemical mechanisms involved in the degradation of this globally important carbon pool. To obtain microbial isolates capable of degrading semi-labile DOM for growth, we conducted dilution-to-extinction cultivation experiments using seawater enriched with high molecular weight (HMW) DOM. In total, 93 isolates were obtained. Amendments using HMWDOM to increase the dissolved organic carbon concentration 4x (280 μM) or 10x (700 μM) the ocean surface water concentrations yielded positive growth in 4–6% of replicate dilutions, whereas <1% scored positive for growth in non-DOM-amended controls. The majority (71%) of isolates displayed a distinct increase in cell yields when grown in increasing concentrations of HMW DOM. Whole-genome sequencing was used to screen the culture collection for purity and to determine the phylogenetic identity of the isolates. Eleven percent of the isolates belonged to the gammaproteobacteria including Alteromonadales, the SAR92 clade, and *Vibrio*. Surprisingly, 85% of isolates belonged to the methylotrophic OM43 clade of betaproteobacteria, bacteria thought to metabolically specialize in degrading C1 compounds. Growth of these isolates on methanol confirmed their methylotrophic phenotype. Our results indicate that dilution to extinction cultivation enriched with natural sources of organic substrates has a potential to reveal the previously unsuspected relationships between naturally occurring organic nutrients and the microorganisms that consume them. The ISME Journal advance online publication, 15 May 2015; doi:10.1038/ismej.2015.68

Introduction

Marine dissolved organic matter (DOM) supports a considerable fraction of the carbon, energy and nutrient requirements of bacterioplankton in the ocean, yet the biological processes that control DOM turnover are poorly understood. This is due in part to the chemical complexity of naturally occurring DOM and the current level of molecular characterization, along with a lack of information on the variety of bacteria and metabolic strategies involved in DOM degradation. Marine DOM characterization has been largely limited to defining its bulk chemical properties. Marine DOM can be operationally separated into two different fractions based on size: low molecular weight (<1 kDa) and high molecular weight (HMW: >1 kDa). The HMW DOM fraction is isolated from seawater by ultrafiltration and typically accounts for 25–40% of the total dissolved organic carbon (DOC) in marine waters (Benner, 2002). The elemental composition of HMW DOM (carbon:nitrogen:phosphorus=298:18:1; Benner, 2002) contrasts with marine plankton (Redfield et al., 1963) in that HMW DOM is depleted in nitrogen and phosphorus relative to carbon. HMW DOM from open ocean- surface waters contains 40–60 nmol nitrogen per μmol carbon primarily in the form of amides (Aluwihare et al., 2005; McCarthy, Pratum, Hedges, & Benner, 1997) and only 3–4 nmol phosphorus per μmol carbon as part of phosphate esters and phosphonate molecules (Clark & Ingall, 1998; Kolowitz et al., 2001). The ^{13}C -NMR spectrum of HMW DOM from marine surface waters indicates it is largely composed of carbohydrates (50%), humic substances (25%) and a lesser amount of lipids and proteins (Aluwihare et al., 1997; Benner et al., 1992; Benner, 2002). Radiocarbon analysis offers insight of the fate and reactivity of DOM. HMW DOM from the North Pacific has a $\Delta^{14}\text{C}$ of 10‰ and its constituting neutral monosaccharides derived from acid hydrolysis have $\Delta^{14}\text{C}$ values of 47–67‰; this is similar to the $\Delta^{14}\text{C}$ of dissolved inorganic carbon in surface waters ($72\pm 7\%$;

Repeta & Aluwihare, 2006) and indicates recent production. In contrast, total DOC is radio-carbon depleted ($\Delta^{14}\text{C}$ of -146‰ and -540‰ in surface and deep North Pacific waters, respectively) likely owing to the presence of refractory carbon (Williams & Druffel, 1987). This data suggests that subcomponents of HMW DOM from surface waters, including polysaccharides, cycle faster (1–25 years; Repeta & Aluwihare, 2006) than the total pool of DOC (6000 years; Williams & Druffel, 1987). These chemical characteristics, along with field observations of semi-labile DOM utilization in natural seawater (Carlson et al., 2004b; McCarren et al., 2010b), indicate HMW DOM can fuel microbial metabolism on ecologically relevant timescales. Difficulties in identifying the specific chemical moieties supporting such growth, however, limits current understanding of DOM transformation, remineralization and ultimate impact on the flux of carbon and nutrients through the marine food web. Bacteria and Archaea are the dominant degraders of DOM (Carlson, 2002), but the specific taxa, genes and molecular mechanisms most responsible for DOM metabolism remain relatively obscure. Recently, several culture-independent studies have begun to identify the microbial community members that respond to DOM enrichment, though the majority of these studies have focused on DOM isolated from phytoplankton cultures (Beier, Rivers, Moran, & Obernosterer, 2014; Landa et al., 2013; Nelson & Carlson, 2012; Poretsky, Sun, Mou, & Moran, 2010; Sarmiento & Gasol, 2012; Sharma et al., 2014) which is presumed to be more labile than HMW DOM collected in oligotrophic environments. This is likely because of the presence of labile proteins and amino acids (Sarmiento et al., 2013) and homopolysaccharides (Meon & Kirchman, 2001) in phytoplankton DOM. Aluwihare & Repeta (1999) and Aluwihare et al. (1997) observed rapid drawdown of phytoplankton DOM and evolution of polysaccharides from homopolysaccharides (energy storage products) to heteropolysaccharides with a composition like HMW DOM. In

contrast, there have been fewer culture-independent studies focusing on the cycling of HMW DOM (McCarren et al., 2010b; Sharma et al., 2014). In all of these studies, DOM additions resulted in the enrichment of select members of the microbial community and in several cases a temporal succession of DOM-degrading bacteria (McCarren et al., 2010b; Sharma et al., 2014). Although these “omic”-based studies have generated hypothesis as to the functional and metabolic roles of DOM-degrading bacteria, appropriate model systems upon which to test such hypothesis and dissect the genes, metabolic pathways and chemical transformation reactions driving DOM degradation are still lacking. The goal of this study was to isolate HMW DOM metabolizing bacteria from marine environments to develop model systems for investigating the biochemical and metabolic pathways driving DOM transformations in the ocean. Dilution-to-extinction cultivation has greatly expanded the collection of marine bacterial isolates (Button et al., 1993; Connon & Giovannoni, 2002b) and has proved instrumental in isolating organisms resistant to traditional cultivation techniques, including members of the SAR11 and SAR116 clades (Rappé et al., 2002; Stingl, Desiderio, Cho, Vergin, & Giovannoni, 2007), oligotrophic marine gammaproteobacteria (Cho & Giovannoni, 2004), members of the SUP05 clade of gammaproteobacterial sulfur oxidizers (Marshall & Morris, 2012) and psychropiezophilic Alphaproteobacteria (Eloe et al., 2011). Field studies, genome analysis and pure cultures experiments enabled by the cultivation of bacterial clades like SAR11 have expanded our understanding of the ecology and physiology of marine bacteria and their role in the ocean carbon cycle (Grote et al., 2012; Malmstrom et al., 2004; Schwalbach et al., 2010; Sun et al., 2011b). In this study we expanded on the technique by enriching dilution-to-extinction cultures with HMW DOM collected directly from ocean surface waters, leveraging the high-throughput

nature of dilution to extinction cultivation to screen a microbial assemblage for HMW DOM-utilizing organisms.

Materials and Methods

Concentration of HMW DOM from seawater

DOM collection was conducted using the method described by (Quan & Repeta, 2007). Briefly, the HMW fraction of DOM (>1 kDa) was concentrated from 24,000 L of surface seawater (15 m) collected 2 km offshore of the Island of Hawaii, Hawaii at the National Energy Laboratory Hawaii Authority (NELHA) in February of 2006. The HMW DOM concentrated seawater was then filtered through a 30 kDa ultrafiltration membrane to remove cell debris and viral particles, diafiltered to remove salts, and freeze-dried. A total of 9.5 g of freeze-dried HMW DOM was obtained that was 32% C (-21.8 ‰), 2.7% N (6.5‰) with a C/N 13.9 representing 18% of the DOC in the original raw seawater.

Molecular characterization of HMW DOM

DOM was characterized by ¹H nuclear magnetic resonance spectroscopy to determine major constituents. Spectra were recorded on a Bruker Avance DPX 400 MHz spectrometer fitted with an inverse broadband 5mm probe. Approximately 1-2 mg of sample was dissolved in 100% D₂O. Chemical shifts were referenced to residual HOD at δ 4.80 ppm. Acid hydrolysis (2.8 M trifluoroacetic acid, heated at 120°C for four hours under nitrogen) of HMWDOM yields a suite of neutral sugars (arabinose, fucose, galactose, glucose, mannose, rhamnose, xylose) that were separated by reverse phase high pressure liquid chromatography (Ascentis C-18 column; 150x1 mm, 3 μm, eluted at 120 μl min⁻¹ with 10/90 (v/v) acetonitrile/water) and quantified at 307 nm as their aminobenzoate ethyl ester (ABEE) derivatives (Baik & Cheong, 2007). Our

chromatographic analysis does not separate xylose and arabinose which are reported as the sum xylose+arabinose. Analyses of unhydrolyzed HMW DOM does not yield any ABEE products that interfere with our monosaccharide analyses.

DOC measurements

Samples for DOC concentration analysis were transferred into combusted glass vials and acidified with 25% phosphoric acid solution before sealing with acid-washed Teflon septa. Sample concentrations were determined using the high-temperature combustion method on a Shimadzu TOC-VCSH with platinized alumina catalyst alongside potassium hydrogen phthalate standards and consensus reference materials (CRM) provided by the DOC-CRM program.

Preparation of seawater media

The water for extinction culturing and subsequent cultivation media was collected off Bermuda (33.2497° N 65.7103° W) during the KN207-01 cruise aboard the R/V Knorr on May 3rd, 2012 via the ship's flow through system with a 0.2 µm filter. After collection and storage, the Sargasso seawater was sterilized by tangential flow filtration (TFF) as described by Becker, Brandon, & Rappé (2007) with minor modifications. The seawater was pre-filtered through a 142 mm-diameter 0.2 µm pore-size Supor membrane (Pall Corp., Ann Arbor, MI, USA) and a 0.1 µm pore-size Supor capsule filter (Pall Corp.), and collected in an acid-cleaned polycarbonate carboy. Pre-filtered water circulated through a Pellicon 2 Mini tangential flow ultrafiltration system (Millipore Corp., Billerica, MA, USA) consisting of a 30 kDa cassette of regenerated cellulose (Millipore Corp.). The tangential flow filtrate was collected in autoclaved, acid-clean polycarbonate carboys and stored at 4°C. Seawater remained sterile at room temperature after the addition of nutrients and sterile rich media. The final TFF seawater medium had a total organic carbon concentration of 70 µM.

Extinction culturing with HMW DOM additions

The inoculum was collected in Nahant Bay, MA (42° 25.195 N, 70°54.463 W) on August 23rd, 2012. Bacterioplankton cell densities were obtained by direct cell counts of DNA-stained cells using SYBR green I (Molecular Probes, Life Technologies, Carlsbad, CA, USA) approximately 1 hour after sample collection. The inoculum was diluted into sterile seawater to an estimated cell density of 3 cells mL⁻¹ in a final volume of 5 L. The dilute inoculum was split into three 1 L aliquots, one left untreated (70 μM DOC) and two which were supplemented with HMW DOM to 210 and 630 μM HMW DOC to obtain media with a final 280 (4x) and 700 (10x) μM DOC, respectively. One mL aliquots from each treatment were then distributed into non-tissue culture treated polystyrene 48-well plates (BD Biosciences, Franklin Lakes, NJ, USA). Each experimental treatment consisted of 840 potential extinction cultures in 20 48-well plates. Every plate included 6 negative controls consisting of wells filled with the corresponding treatment media made up with sterile seawater. Plates were covered in foil to reduce evaporation and incubated at 24°C in the dark for 4-5 weeks. In order to ensure that potential extinction cultures remained viable and capable of growth in the corresponding cultivation media, after the 4-5 week incubation the samples were re-diluted 10,000-fold into new plates containing fresh TFF seawater and the appropriate DOM concentrations. These were then incubated for 6-7 weeks under the same conditions as before.

Detection of culture growth

Extinction cultures were detected using a guava easyCyte 8HT flow cytometry system (Millipore Corp.) by transferring 50 μL from each cultivation plate well to a 96-well microtiter plate containing 10 μL of 500-fold diluted SYBR green I and 140 μL of sterile seawater. Samples were stained for at least 30 minutes before flow cytometry. Each well was analyzed for 15 s at a

flow rate of $0.59 \mu\text{L s}^{-1}$ with a blue laser (488 nm excitation) to detect green fluorescence. Wells were scored positive for cell growth when the concentration was $\geq 10^5$ cells ml^{-1} .

HMW-DOM dose response screen of the culture collection

Cultures scoring positive for growth were examined for increased cell yields under different HMW DOM concentrations. To limit cell growth to consumption of HMW DOM added, background DOC in the TFF seawater was photo-oxidized via exposure to high-intensity ultraviolet light for 4 hours, reducing DOC concentrations by 30-40%. During this process, approximately an eighth of the volume was lost due to evaporation, and this was replenished with ultrapure water. The UV oxidized TFF seawater ($25 \mu\text{M DOC}$) was then supplemented with 210, 1330, and 6930 $\mu\text{M DOC}$ using HMW DOM to obtain $\sim 4\text{x}$, $\sim 20\text{x}$, or $\sim 100\text{x}$ DOC media respectively in similitude to media prepared for extinction culturing. Non DOM-amended UV oxidized seawater served as control. No inorganic nutrients or vitamins were added to the media. Each treatment media was distributed into 1-mL aliquots in 48-well cultivation plates. Wells were inoculated with 10- μL sub-samples of the dilution to extinction cultures and incubated at 24°C in the dark. Subsamples were taken every 5-7 days for cell density enumeration using the SYBR green I flow cytometry assay described above.

SSU rRNA sequencing

400-800 μl from the 93 cultures scoring positive for growth were transferred to a $0.2 \mu\text{m}$ Supor-membrane 96-well filter plate (Pall Corp.) and vacuum filtered. The cells were resuspend from the filter with two separate aliquots of 125 μl of lysis buffer (40mM EDTA, 50mM Tris at pH 8.3, 0.73M sucrose 1.15 mg/ml lysozyme (Sigma-Aldrich, St. Louis, MO, USA), 200 ug/ml RNase (Qiagen) and transferred to a 96 deep-well plate, which was then placed at -80°C . After freezing, the plate was thawed, Proteinase K and SDS were added to final concentrations of 0.65

mg ml⁻¹ and 10% SDS, respectively, and allowed to incubate for 2 hr at 55°C. The lysate was then transferred to new 2ml deep well plate and DNA extracted using a DNeasy 96 well Blood and Tissue kit (Qiagen).

The SSU rRNA gene was PCR amplified using universal Eubacterial primers pA *E.coli* 8-28F (5'- AGA GTT TGA TCC TGG CTC AG -3') and *E. coli* 1510-1492R (5'- GGT TAC CTT GTT ACG ACT T -3') in 50 µl reactions consisting of 1 µM each forward and reverse primer, 1 µl of FailSafe Enzyme Mix (Epicentre, Madison, WI, USA), 25 µl of 2x FailSafe PCR PreMix E (Epicentre) water, and 20 µl DNA template. After amplification, the PCR amplicons were run on 1% agarose gel and bands running at ~1500-1600 nt were excised and PCR purified with a 96 well gel extraction kit (Qiagen) followed by an additional purification using the QIAquick 96-well PCR purification kit (Qiagen) and eluted in 80 µl elution buffer. Sanger sequencing was conducted following the BigDye v3.1 sequencing protocol (Applied Biosystems, Foster City, CA, USA) on a ABI 3730 DNA analyzer (Applied Biosystems) using the SSU rRNA universal primers described above (pA *E. coli* 8-28, *E. coli* 1510-1492) and with bacterial primers 519F 5'-CAGCMGCCGCGGTAATWC-3' and 800R 5'-TACCAGGGTATCTAATCC-3'.

Whole genome sequencing

70 cultures were selected for whole genome sequencing based on their growth yields and response to HMW DOM substrate additions. To obtain sufficient amounts of DNA for the NexterXT library preparation, 10 µl culture was inoculated into acid-clean polycarbonate bottles containing 200 mL TFF seawater amended with HMW DOM to 3x DOC. After incubating at 22°C in the dark for 4-5 weeks, the cells were harvested on 0.1 µm pore-size, 25 mm diameter Durapore membranes (Millipore Corp.) using Swinnex filter-holders (Millipore Corp.) and

peristaltic pumping. Filters were stored frozen (-80°C) in 600 µL of tissue lysis buffer (Qiagen). DNA was extracted from thawed filters using a DNeasy Blood and Tissue Kit (Qiagen) following the manufacturer instructions. DNA samples were prepared for sequencing and barcoded using the Nextera XT DNA 96-sample preparation kit (Illumina Co.), and sequenced with one 250x250 nt paired end MiSeq run (Illumina Co.). Sequences are deposited in the NCBI sequence read archive (SRA) under study SRP045600.

Bioinformatic and phylogenetic analysis

FastQ files from the MiSeq run were imported into the CLC Genomics Workbench (CLC bio, Denmark). Paired reads were joined and then assembled into contigs using CLC's De Novo assembler with automatic word and bubble sizes, a minimum contig length of 200, insertion and deletion costs set to 3, mismatch cost set to 2, length fraction set to 0.5, and the similarity fraction set to 0.8. Only contigs ≥ 1000 nt in length were further examined. Contigs containing small and large subunit rRNA sequences were identified and annotated by a BLASTn sequence similarity search against the SILVA database (<http://www.arb-silva.de>, version 111).

To ensure that low abundance reads with a divergent SSU rRNA were not missed during the assembly process the unprocessed miSeq reads were trimmed and paired-end joined using trimomatic (<http://www.usadellab.org/cms/?page=trimmomatic>) and PandaSeq (<https://github.com/neufeld/pandaseq/wiki/PANDAseq-Assembler>), respectively. The trimmed and joined reads were compared to the assembled SSU rRNA for each sample via a BLASTn sequence similarity search, with hits having an e-value < 0.001 , a read length > 50 nt, and 45% identity over the BLAST alignment considered significant. The unassembled, but trimmed and joined reads, were also taxonomically binned based on a LAST (Frith, Hamada, & Horton, 2010; Kielbasa, Wan, Sato, Horton, & Frith, 2011) sequence similarity search against NCBI's RefSeq

database (version 64) with a score penalty for frameshift of 500 and an initial match multiplicity of 10. Top hits with a score ≥ 40 and alignment length ≥ 75 nt were considered significant and used for the taxonomic annotation.

Partial SSU rRNA gene sequences obtained via Sanger sequencing were quality-trimmed, manually inspected, and assembled with Sequencher version 5.1 (GeneCodes, Ann Arbor, MI, USA). SSU rRNA gene sequences from Sanger or WGS were aligned using the SILVA incremental aligner online tool (Pruesse, Peplies, & Glockner, 2012) and curated in ARB (Ludwig, 2004) using the All-Species Living Tree project release 115 (Munoz et al., 2011). Full-length SSU rRNA gene sequences obtained from WGS were used to build phylogenetic trees and were deposited in NCBI GenBank under accession numbers KP770034 through KP770106. For the betaproteobacteria OM43 clade strains a tree was inferred from 1350 alignment positions using the RAxML maximum likelihood method (Stamatakis, 2014) and neighbor joining (Saitou & Nei, 1987). For *Gammaproteobacteria* isolates, a tree was constructed using 1293 unambiguous alignment positions using RAxML. RAxML 8.0.24 was implemented on Cipres Science Gateway (Miller, Pfeiffer, & Schwartz, 2010) and neighbor joining was implemented on MEGA5 (Tamura et al., 2011). RAxML trees were curated in EMBL's interactive Tree of Life tool (Letunic & Bork, 2007).

Results and Discussion

Molecular characterization of HMW DOM

HMW DOM was characterized by ^1H nuclear magnetic resonance spectroscopy to determine major constituents. The spectrum looked similar to previously published ^1H NMR spectra from samples collected at the Natural Energy Laboratory study site (Repeta & Aluwihare, 2006) with

strong resonances at 5.5 ppm from carbohydrate anomeric (O-CH-O), 4.0 ppm from carbohydrate (HCOH), 2.7 ppm from N-acetyl N-methyl aminosugar (H₃C-N-C(O)CH₃), 2.0 ppm from N-acetyl aminosugars N-C(O)CH₃, and 1.3 ppm from 6-deoxysugar protons. The spectral region between 0.9-3.0 ppm also included a broad unresolved baseline attributed to co-occurring humic substances. Monosaccharide analysis of hydrolysable sugars showed the presence of similar amounts of glucose, galactose, mannose, rhamnose, fucose, and xylose+arabinose. This distribution is typical of HMW DOM hydrolysable sugars in open-ocean surface seawater (Aluwihare et al., 1997). The sum of hydrolysable neutral sugars represents 10% of total HMW DOC.

Experimental overview

The culturing method applied here is a modification of the “high-throughput” culturing (HTC) technique described by Connon & Giovannoni (2002) to isolate marine bacteria in seawater media using dilution to extinction. Our main modification consisted of supplementing filter-sterilized natural seawater with HMW DOM (> 1 kDa) isolated by ultrafiltration from oligotrophic ocean surface waters (Fig. 2.1). To limit the amount of background DOM in the natural seawater and to increase the probability of obtaining isolates responding to the HMW DOM additions we used relatively nutrient deplete, Sargasso Sea seawater sterilized by tangential flow filtration (TFF) to prepare dilution to extinction samples.

The inoculum was collected from Nahant Bay in the coastal North Atlantic and had an initial cell concentration of 1.5×10^6 ($\pm 2.2 \times 10^5$ s.d.) cells ml⁻¹. The inoculum was diluted into TFF seawater to a final concentration of 3 cells ml⁻¹ and then divided into three treatments: a no DOM addition (70µM dissolved organic carbon [DOC] background) and two HMW DOM additions (4x DOC [280µM], and 10x DOC [700µM]). One ml aliquots were placed in 48-well

cultivation plates and incubated at 24°C in the dark. After 4 weeks of incubation, the cultures were diluted 10,000 fold into fresh TFF media containing the appropriate DOM concentrations to ensure viability before the screening.

Growth screen

After 6-7 weeks of incubation, the cultures were screened for growth by flow cytometry and any well with $>10^5$ cells ml^{-1} was scored growth positive. Out of 2520 potential extinction cultures, 93 (3.7%) scored positive. Only 1 of the 840 non-DOM amended inoculated wells (0.1%) scored growth positive (Table 2.1 and Table S1). By contrast, the HMW DOM amendments had significantly more wells scoring growth positive (Table 2.1; $p < 0.05$, bootstrap comparison with 10^6 iterations; Table S1), increasing the percent recovery to 4.7% and 6.0% for the 4x and 10x DOC treatments, respectively. Cell concentrations also tended to increase with higher DOM dose (Fig. 2.2), with upper concentrations reaching $>10^6$ cells ml^{-1} (max. 2.4×10^6), the majority of which occurred in the 10x DOC treatment. No growth was detected in any of the non-inoculated control wells (Table 2.1 and Table S1). The single culture obtained from the non-DOM amended treatment showed little sustained growth, and could not be maintained in successive transfers of the culture collection grown on 4x DOC.

HMW DOM dose response

To verify that the isolates had a discernible growth response to the HMW DOM, the 93 cultures were tested for increased cell yields with DOM additions. The TFF seawater was first UV oxidized to reduce background DOC concentrations from 70 μM DOC to 25 μM DOC. HMW DOM was then added at 4x (280 μM), 20x (1400 μM), or 100x (7000 μM) open-ocean DOC concentrations. Sixty six of the cultures showed an increase in cell yields with DOM addition, reaching $>5 \times 10^5$ cells ml^{-1} in the DOM amended treatments and having no growth in the non-

DOM amended controls (Fig. S1). Furthermore, 36 cultures showed a direct proportional increase in cell yields with increasing HMW DOM concentrations (100x > 20x > 4x).

Identification and purity screen

In an initial phylogenetic screen of the isolates, 88 small subunit (SSU) rRNA gene sequences were determined via PCR amplification and subsequent Sanger sequencing. Ten of these sequences clustered within the *Gammaproteobacteria Alteromonadaceae* (7 cultures including three members of the SAR92 clade and one of the OM60 clade), *Pseudoalteromonadaceae* (1 culture), and *Vibrio* (2 cultures). The remaining 78 sequences all clustered within the OM43 clade of *Methylophilaceae* betaproteobacteria. There was no significant difference in the relative recovery of any specific taxonomic group in the 4x versus 10x DOC treatments.

The quality of the SSU rRNA Sanger sequencing provided an initial assessment of culture purity, with positive PCR amplifications yielding poor quality Sanger sequences being potentially indicative of mixed cultures. We applied an additional approach to assess the purity of cultures and identify their phylogenetic origins using “tagmentation” sequencing library preparation and Illumina paired-end sequencing for whole genome shotgun (WGS) sequencing (Fig 1 and see methods). Sixty-eight cultures were successfully sequenced, with an average of 120,000 paired reads (200 X 200 bases) per sample.

The purity of the cultures was assessed using the WGS data in three ways. In the first approach, the reads were *de novo* assembled into contigs which were then searched via BLASTn against the SILVA rRNA database to identify SSU and LSU rRNA sequences. Samples having multiple rRNA genes binning to different organisms were considered mixed. Ninety percent of the 68 cultures had a single copy small and large subunit rRNA gene binning to one organism

with high identity (Table 2.2), either an OM43 clade betaproteobacterium (82%) or miscellaneous Gammaproteobacteria (9%). Five samples contained multiple rRNA genes hitting two different organisms. These were primarily an OM43 clade betaproteobacteria with a SAR92 clade Gammaproteobacteria, although one culture (NB0016) contained an OM43 clade betaproteobacteria and a flavobacterium.

To verify that rRNAs did not escape detection in the course of sequence assembly we searched for SSU rRNA sequences in the unassembled reads. The quality controlled (trimmed and paired-end joined) unassembled reads were searched against each sample's assembled SSU rRNA sequence. Samples having a large portion of reads matching at low identity were taken as an indication of a mixed culture. Four of the five same samples that yielded multiple, divergent rRNA reads after assembly (Table 2.2) had high levels of low identity rRNA matching reads (5-17%; Fig. S2 and S3).

Finally, a metagenomic approach was taken, taxonomically binning the unassembled reads for each sample based on the top hit from a LAST sequence similarity search against NCBI's RefSeq database. For 54 of the 60 cultures with betaproteobacterial SSU rRNA sequences the majority of reads binned to betaproteobacterial reference genomes (Fig. 2.3). The remaining six cultures had a substantial enrichment in reads binning to non-betaproteobacterial reference genomes (17% to 53% of total reads). Five of these cultures were also identified as mixed using the SSU rRNA analyses above but the sixth (NB0010) was only confidently detected with the metagenomic approach because the unassembled SSU rRNA analysis revealed only a marginal proportion of low identity SSU rRNA reads (1.4%). In summary, these three screening approaches produced congruent assessments of culture identity and purity, with only six out of 68 of WGS sequenced cultures indicating the presence of mixed populations.

Isolate phylogeny

The WGS assembled SSU rRNA sequences matched excellently with the Sanger reads and provided phylogenetic placement of the culture collection. Of the 58 WGS SSU rRNA sequences (length = 1542 nt) that had overlapping regions with quality Sanger sequences (489 to 1444 bp, quality >97%), only 5 sequences had mismatches between the Sanger and WGS data (NB0034 [10 mismatches], NB0054 [11], NB0057 [32], NB0068 [12], and NB0072 [194]), with the remaining 53 cultures matching at 100% identity between the two sequencing methods.

The full length WGS assembled rRNA sequences were used to determine the phylogenetic placement of cultures. The OM43 clade cultures consisted of 16 unique groups of sequences that clustered with sequences obtained from coastal metagenomes and clone libraries (Fig. 2.4). The majority of methylotroph isolates were >99% identical to one another at the SSU rRNA level with the greatest sequence divergence occurring between isolate NB0027 and NB0070 (96.6% SSU rRNA identities, Fig. 2.4). The *Gammaproteobacteria* sequences were more phylogenetically diverse, spanning several Alteromonadales clades and *Vibrio* (Fig. S4). The SSU rRNA sequence of NB0015 clustered with the SAR92 clade isolate HTCC2207 (SSU 97% identity). Four additional SAR92 clade SSU rRNA sequences were identified mixed with OM43 clade cultures and were 98-100% similar to the NB0015 sequence. The remaining Alteromonadales sequences clustered with *Teredinibacter turnerae* (NB0011 and NB0077), *Alteromonas macleodii* (NB0094), and *Psychrosphaera saromensis* (NB0030). *Vibrio* sequences (NB0059, NB0072, and NB0080) were closely related to marine strains *Vibrio chagasii*, *Vibrio lentus*, and *Vibrio pomeroyi* (Fig. S4).

Growth responses to differing media and C substrates

Two isolates representative of the most commonly identified bacterial groups, OM43 clade betaproteobacterium NB0046 and SAR92 clade gammaproteobacterium NB0015, were selected for further characterization. Growth was undetectable for the OM43 betaproteobacterium NB0046 on rich media (Marine Broth 2216), methanol amended agar plates (Janvier, Frehel, Grimont, & Gasser, 1985), or TFF seawater amended with glucose and succinate (100 μM each). In non-carbon amended TFF oligotrophic seawater, NB0046 will typically reach cell concentrations of 1×10^5 cells ml^{-1} , although we have found this to be variable, with some cultures showing zero growth in the no carbon amendments, while others can reach almost 10^6 cells ml^{-1} . We also tested cultures in media prepared with UV oxidized seawater. UV oxidation lowers the total amount of naturally occurring DOM in the seawater, but the oxidation process can also produce a broad suite of low molecular weight, highly oxidized organic compounds that could serve as substrates or inhibitors to our isolates (Moran & Zepp, 1997). However, in our experiments, UV oxidation of the seawater medium consistently reduced growth yields in non-carbon amended cultures to much lower cell concentrations ($\leq 10^3$ cells ml^{-1}).

The addition of methanol to seawater media significantly increased OM43 isolate cell yields (Fig. 2.5), with maximum cell concentrations achieved between 10-100 μM methanol; higher methanol concentrations (>1 mM) tended to inhibit, rather than increase cell yields, consistent with previous observations of OM43 clade cultures (Halsey et al., 2011). The highest cell yields were achieved with the HMW DOM additions, where isolate NB0046 reached $>5 \times 10^6$ cells ml^{-1} in HMW DOM-containing media at 50x ambient DOC concentrations (~ 4 mM DOC, Fig. 2.5). However, normalizing the observed cell yields to carbon added to the media indicates

that methanol supports higher growth yields (2.5×10^7 to 2.1×10^8 cells $\mu\text{mol}^{-1}\text{C}$) than HMW DOC treatments (1.6×10^6 to 6.4×10^6 cells $\mu\text{mol}^{-1}\text{C}$).

SAR92 clade gammaproteobacterium NB0015 grew in rich media (Marine Broth 2216), although only at dilute concentrations (1:10 and 1:100 strength). No growth was detected on full strength media, marine agar, seawater amended with methanol, or seawater amended with glucose and succinate (100 μM each). In non-carbon amended TFF oligotrophic seawater, NB0015 reached 1×10^5 cell ml^{-1} . Supplementing TFF oligotrophic seawater with HMW DOM (4x and 8x DOC) increased NB0015 growth to $2\text{-}5 \times 10^5$ cells ml^{-1} (Fig. 2.5). The cell yield normalized to carbon added as HMW DOC was 7.1×10^5 cells $\mu\text{mol}^{-1}\text{C}$, an order of magnitude lower than for the OM43 clade isolate NB0046. In contrast, cultures of NB0015 reached on average a maximum cell concentration of 1.5×10^6 cell ml^{-1} when grown in an artificial seawater medium (Kester, Duedall, Connors, & Pytkowicz, 1967) supplemented with a mix of simple carbon compounds (D-glucose [55 μM], succinate [85 μM], pyruvate [114 μM], glycerol [109 μM], N-acetyl D-glucosamine [45 μM], and ethanol [434 μM]; Fig. S5), consistent with the previous isolation of SAR92 clade strains (Cho & Giovannoni, 2004).

Conclusion

Dilution to extinction cultivation has proven a useful approach for obtaining highly abundant but difficult to cultivate bacteria. It enables separation of slower growing cells from faster, more easily cultured microorganisms, and it also allows the growth of cells in media that more closely mimics the native environment. While several studies have supplemented natural seawater media with defined carbon substrates in efforts to improve cell yields (Cho & Giovannoni, 2004) we applied the alternative approach of enriching seawater media with HMW

DOM collected directly from the environment. We postulated that the native, complex mixture of marine HMW DOM could increase the probability of obtaining relevant DOM degrading microorganisms. Such isolates could then be used as tools to help define the chemical characteristics of marine HMW DOM and the biological processes that degrade it. A similar tactic was successfully applied by Hotalle-Schmelzer, Zwirrmann, Krüger, & Grossart (2010) who enriched bacterioplankton dilution cultures with the humic fraction of lake-derived DOM.

In our study, dilution to extinction cultures enriched with HMW DOM increased culturing efficiency and yielded over 90 organisms with DOM degrading capabilities. Our non-DOM amended cultures had lower culturing efficiency (0.1%) than previous dilution culturing studies, which typically range from 3 to 25% (Cho & Giovannoni, 2004; Connon & Giovannoni, 2002; Song, Oh, & Cho, 2009). The reduced culturing efficiency in our experiments may be explained by the deliberate use of oligotrophic Sargasso Sea surface water to prepare the media, which likely has low levels of vitamins (e.g., <0.1 pM vitamin B₁₂ [Menzel & Spaeth, 1962], and 20-25 pM thiamin [Carini et al., 2014]), inorganic nutrients (often below detection limits, 0.05 $\mu\text{mol kg}^{-1}$ for nitrate + nitrite and 0.03 $\mu\text{mol kg}^{-1}$ for phosphate; <http://batsftp.bios.edu/BATS/bottle>), and DOC concentrations (70 μM ; as measured for this study) in the basal media than are typical of coastal seawater. While this oligotrophic medium may have prohibited the enrichment of some coastal microorganisms, its low DOC and nutrient background (inorganic N and P were not added) provided greater sensitivity for detecting DOM addition effects without requiring extensive manipulation of the natural media. The fact that culturing efficiencies significantly increased in DOM amended treatments suggested that DOM addition provides, at least partially, growth factor(s) or organic sources of N and P that were missing in oligotrophic seawater.

Establishing the purity of the culture collection is essential for linking the metabolism of DOM to specific taxa, genes and biochemical mechanisms. Determining culture purity, however, can be challenging, especially when a culture contains a secondary isolate at very low abundance (Shrestha, Nevin, Shrestha, & Lovley, 2013). While SSU rRNA gene sequencing or screens such as restriction fragment length polymorphisms (RFLP) are most often used to identify cultures, they suffer from primer design and preferential amplification biases which can decrease sensitivity for detecting mixed cultures. The WGS screen used in this study does not suffer from oligonucleotide primer mismatch and excludes most other PCR biases thereby providing a more sensitive and universal method for determining culture purity. While PCR-based SSU rRNA gene sequencing revealed no mixed cultures in our study, WGS sequencing showed that 6 of the 68 screened cultures were mixed as congruently determined by three different bioinformatic analyses: 1) number of contigs containing SSU rRNA genes, 2) unassembled read's sequence similarity to assembled SSU rRNA sequences, and 3) unassembled read taxon binning based on sequence similarity search against NCBI's RefSeq database.

The ability of WGS sequencing to identify mixed cultures is directly dependent on the relative abundance of the secondary culture in relation to the sequencing depth, as well as on the phylogenetic relatedness of the co-isolates. In terms of resolving mixed cultures of closely related isolates, we found cultures with no other indication of being mixed had the majority of unassembled SSU rRNA reads at $\geq 98\%$ identity to the assembled SSU rRNA sequence (Fig. A2), suggesting that two different populations with $<2\%$ SSU rRNA divergence would be difficult to detect. We note that this would include many of the betaproteobacterial isolates obtained here. The metagenomic approach of identifying mixed cultures via taxonomic binning will be less sensitive than the SSU rRNA bioinformatic approaches to resolving closely related

populations due to database limitations. The breadth of genes examined using the metagenomic approach, however, should provide greater sensitivity for detecting low abundance but phylogenetically diverse secondary populations. For the betaproteobacterial genomes examined here, 5 to 10% of unassembled reads had significant matches to non-betaproteobacterial genomes likely due to factors such as short reads lengths, horizontal gene transfer events, and limited reference genomes in the database. This suggests a lower bound for detecting co-isolates at around 10% total abundance, though we note differences in genome size will also affect this sensitivity. Alternative approaches to increase power for detecting closely related co-isolates at low abundance would likely require more sensitive bioinformatic analysis, such as the statistical frequency of genome single nucleotide polymorphisms (Shrestha et al., 2013) focusing on less conserved genes such as those typically used in multi-locus sequence typing.

De Novo assembly of the WGS data produced contigs containing full length SSU and LSU sequences that were as accurate, and often had 100% identities, with overlapping Sanger reads. Phylogenetic placement of these WGS SSU rRNA sequences revealed the culture collection was dominated by members of the OM43 clade of betaproteobacteria.

The OM43 clade, initially described by Rappe, Kemp, & Giovannoni (1997), was first isolated from the Oregon coast by Cannon & Giovannoni (2002). Their distribution is generally limited to the coastal zone, where they can compose 1 to 3% of the bacterioplankton community (Suzuki et al., 2004), though their abundance has been shown to significantly increase during phytoplankton blooms (Morris, Longnecker, & Giovannoni, 2006; Rich, Pham, Eppley, Shi, & DeLong, 2011). Metatranscriptomic and metaproteomic data suggest they are active components of coastal bacterioplankton communities (Gifford, Sharma, & Moran, 2014; Sowell et al., 2011). The OM43 isolates obtained in this study were closely related to these previously characterized

coastal strains and sequences (Fig. 2.4). The majority of our isolates clustered most closely with the Hawaiian strain HIMB624 (Huggett, Hayakawa, & Rappé, 2012) and related SSU rRNA phylotypes identified in diverse coastal locations. Three of our isolates formed an outgroup with the Oregon coastal isolate HTCC2181 as well as sequences obtained primarily from higher latitudes. Though closely related, the presence of SSU rRNA microdiversity within our isolate collection suggests the prevalence of OM43 clade members in the dilution to extinction cultures was not due to the enrichment of a single clonal population.

Genome sequences of OM43 betaproteobacteria are missing the E1 subunit of the α -ketoglutarate dehydrogenase complex of the TCA cycle, a trait thought to be indicative of an obligatory methylotrophic lifestyle (Halsey, Carter, & Giovannoni, 2012; Huggett et al., 2012). A preliminary genome analysis of NB0046 shows that it is also missing the E1 α -ketoglutarate dehydrogenase subunit. Furthermore, NB0046 shares similar characteristics with other OM43 lineage genomes (Giovannoni et al., 2008; Huggett et al., 2012), including a small genome size (<1400 genes), the presence of an alternative methanol dehydrogenase (XoxF) instead of the canonical Mxa or Mdh methanol dehydrogenase (Chistoserdova, 2011), genes for formaldehyde oxidation via tetrahydrofolate but not via tetrahydromethanopterin, and carbon assimilation through the RUMP cycle instead of the serine cycle. These observations suggest the methylotrophs isolated on HMW DOM have a similar ‘obligate’ methylotrophic lifestyle as other OM43 clade members, although a more detailed and thorough analysis of the 60 OM43 genomes sequenced here will be required to confirm this.

Given their reliance on single carbon substrates, the predominance of C1 compound specialists in our cultures amended with HMW DOM was unexpected. Our results, however, are predicated in a previous study by McCarren et al. (2010), who observed that HMW DOM added

to oligotrophic ocean bacterioplankton microcosms increased the relative abundance and transcriptional activity of methylotrophs, particularly Methylophaga species belonging to the Gammaproteobacteria. McCarren et al. (2010) postulated that HMW DOM polysaccharides containing a large fraction of methylated sugars (Quan & Repeta, 2007) were the source of single carbon compounds, including methanol, which provided the methylotrophs their growth substrate. We further propose that the prevalence of pure cultures of methylotrophs in our HMW DOM-enriched dilution-to-extinction samples is due to their ability to cleave C1 groups from the methylated sugars in the DOM polymer. Additionally, low levels of abiotic oxidation of the HMW DOM during collection and storage may possibly degrade the methylated sugars, releasing low molecular weight compounds that can be directly consumed by methylotrophs. In either case, the data indicated that HMW DOM can serve as a sole source of carbon and energy for these methylotrophs.

The order of magnitude difference in growth yields (cell yield normalized to C added) between methanol (6.8×10^7 to 5.3×10^8 cells $\mu\text{mol}^{-1}\text{C}$) and HMW DOM (1.6×10^6 to 6.4×10^6 cells $\mu\text{mol}^{-1}\text{C}$) suggests that growth on HMW DOM is less efficient than on methanol. Similar low growth yields (7.1×10^5 cells $\mu\text{mol}^{-1}\text{C}$) were observed for SAR92 clade strain NB0015 in cultures amended with HMW DOM. Alternatively, the low growth yields observed may indicate that only a small fraction of HMW DOM may be available to a single organism with limited metabolic potential like isolates NB0046 and NB0015 despite the total amount of DOC added to cultures. The efficiency at which DOC is converted into biomass, typically <20% in the open-ocean (Carlson & Ducklow, 1996) has implications on the magnitude of carbon that cycles through bacterioplankton. Measurements of DOC drawdown and respiration, though not obtained in HMW DOM dose-response experiments, will be necessary to determine bacterial

growth efficiencies on HMW DOC and to estimate the bacterial carbon demand this organic carbon pool can support.

The second major taxonomic group found in the HMW DOM enriched cultures was the Gammaproteobacteria. This included isolates belonging to the *Alteromonadales* and *Vibrionales*. The *Vibrio* (NB0059, NB0072, and NB0080), alteromonad (NB0094), and pseudoalteromonad (NB0030) isolates did not exhibit a proportional increase in cell yields in response to added HMW DOM in dose-response experiments even though these isolates were initially obtained from the HMW DOM enriched dilution to extinction cultures. Among the Gammaproteobacteria only strains NB0011 and NB0077 and SAR92 clade isolates increased their growth proportionally in the presence of HMW DOM. The SAR92 clade isolate NB0015 exhibited the strongest growth response to HMW DOM although cell yields were low ($2-4 \times 10^5$ cells ml⁻¹) compared to the OM43 clade isolates ($>5 \times 10^6$ cells ml⁻¹). However, NB0015 growth yields reached $>1 \times 10^6$ cells ml⁻¹ when cultured in defined media with a mixture of simple carbon substrates indicating that this isolate may require a variety of carbon compounds and nutrients not available in HMW DOM to supplement its metabolism. The SAR92 clade was the most frequently recovered Gammaproteobacteria group in the dilution to extinction cultures including four SAR92 clade-related SSU rRNA sequences found mixed with cultures of OM43 clade strains, more than any other taxa. Including these additional sequences, the SAR92 clade was the second most common group identified in our dilution to extinction cultivation experiment.

The growth of SAR92 clade strains in our HMW DOM-amended samples may partially be explained by its numerical abundance in coastal bacterioplankton assemblages (Stingl et al., 2007) and their ease of recovery by dilution to extinction cultivation (Cho & Giovannoni, 2004). These factors alone, however, do not solely account for the prevalence of SAR92 in our cultures

since SAR92 was only identified in the HMW DOM enriched dilution to extinction cultures and not in the non-DOM amended cultures. We postulate that the carbohydrate-rich component of HMW DOM in particular may have stimulated the growth of SAR92 bacteria. Some of the closest relatives of the SAR92 clade (90-93% SSU rRNA identity) include cultured isolates with carbohydrate degrading capabilities. For example, *Microbulbifer hydrolyticus* (González, Mayer, Moran, Hodson, & Whitman, 1997), *Saccharophagus degradans* 2-40 (Andrykovitch & Marx, 1988; Weiner et al., 2008) and *Simiduia agarivorans* SA1 (Lin, Shieh, Chen, & Tang, 2013; Shieh, Liu, Lin, Jean, & Chen, 2008) can breakdown several recalcitrant polysaccharides, including agar, alginate, cellulose, or chitin. Another relative, *Teredinibacter turnerae* T7902, a bacterium associated with wood-boring bivalves, is capable of digesting cellulose (Distel, Morrill, MacLaren-Toussaint, Franks, & Waterbury, 2002). It is plausible that SAR92 clade strains may directly degrade HMW DOM polysaccharides in contrast to the OM43 clade methylotrophs which may utilize C1 compounds that decorate the polysaccharides. The availability of model laboratory organisms able to grow on HMW DOM now allows us to test these hypotheses. Future work chemically characterizing the HMW DOM before and after microbial degradation, as well as examining the transcriptional and proteomic responses of the isolates during DOM metabolism will help determine the bonded nature of the carbon sustaining the cultures.

In summary, dilution to extinction cultures enriched with HMW DOM extended the power of the dilution cultivation technique by not only growing cells in a media closely mimicking their native environment, but also by stimulating growth using naturally derived organic substrates. This approach is critical for obtaining model DOM degrading isolates as both the carbon substrates and the organisms acting upon them are unknown. Using this technique we

found that organisms ranging from obligate methylotrophs to less fastidious microorganisms were able to grow using oligotrophic ocean HMW DOM as a substrate, suggesting there are multiple metabolic strategies involved in the degradation of HMW DOM. We note, however, that only a fraction of the total DOM added to our cultures was remineralized, suggesting that there were other growth limiting factors, or potential requirement for syntrophic microbial partners to further degrade the DOM polymers (McCarren et al., 2010). Co-culture experiments will likely be essential to further elaborate the biological, physiological and biochemical details of consortial DOM degradation processes in the sea.

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Tables

Table 2.1. HMW DOM-enriched dilution to extinction cultivation experimental design and growth screen results. Forty-eight well plates were filled with TFF oligotrophic seawater amended with HMW DOM at 4x (280 μM) DOC, 10x (700 μM) DOC or non-amended and then inoculated with the diluted coastal bacterioplankton community. The 'controls' consisted of six wells on each treatment plate, as well as an entire microtiter plate, that contained the same media and DOM enrichment, but were not inoculated. Positive growth ('cultures detected') was defined as a well having a density of 10^5 or more cells per ml after 6–7 weeks of incubation and one round of redilution.

		noDOM	4X DOM	10X DOM	total
wells	inoculated	840	840	840	2520
	controls	54	54	54	162
cultures detected	inoculated	1	40	52	93
	controls	0	0	0	0

Figures

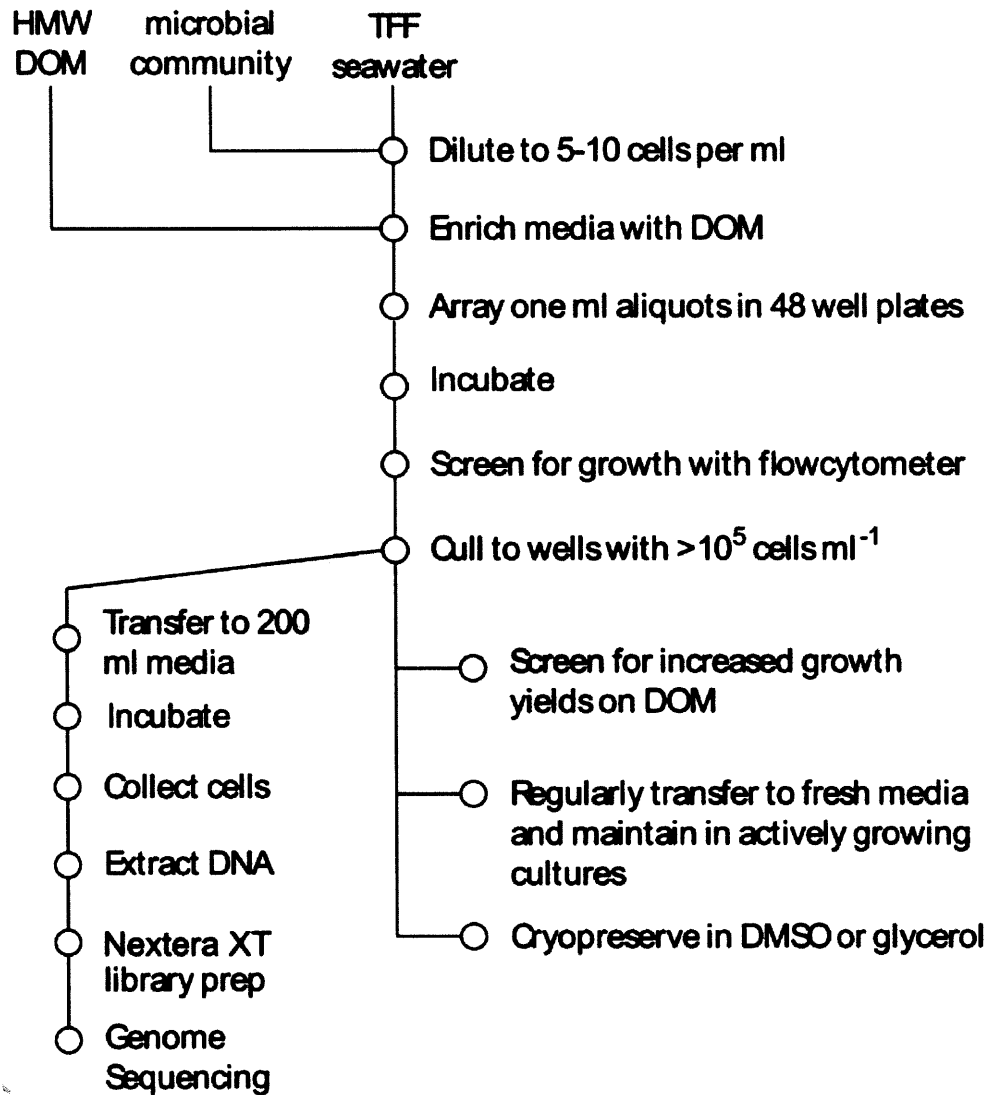


Figure 2.1. Procedure for the setup, incubation, screening and downstream analysis of dilution to extinction cultures enriched with high molecular weight dissolved organic matter (HMW DOM), including purity screening and phylogenetic identification via whole-genome sequencing. TFF, tangential-flow filtered.

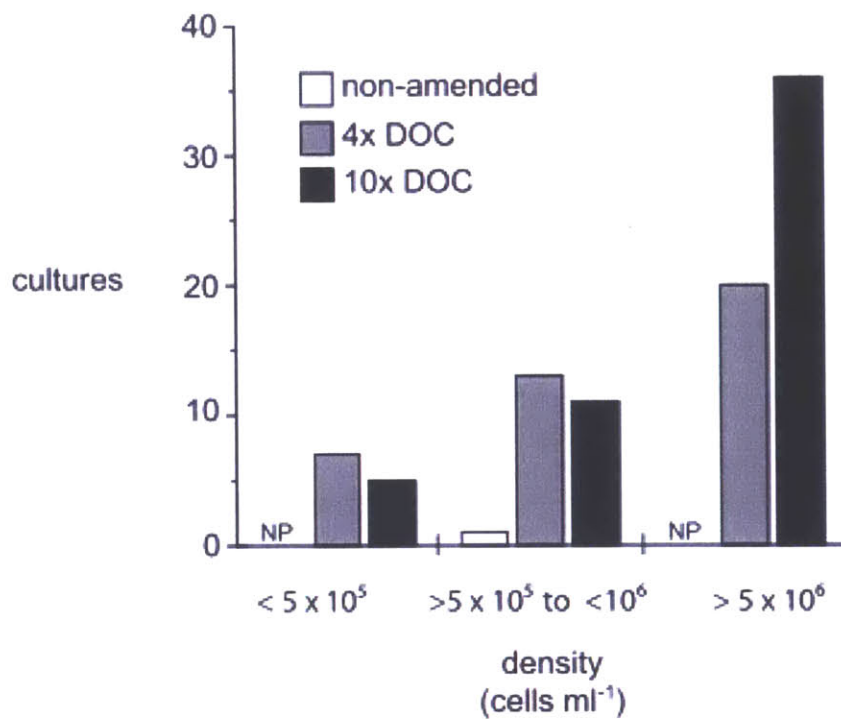


Figure 2.2. Effect of HMW DOM enrichment on dilution to extinction cultures' final cell density. Only wells that scored positive for growth (≥ 10 cells per ml) are included. NP, no positive cultures detected; 4x DOC, $\sim 280 \mu\text{M}$ DOC; 10x DOC, $\sim 700 \mu\text{M}$ DOC.

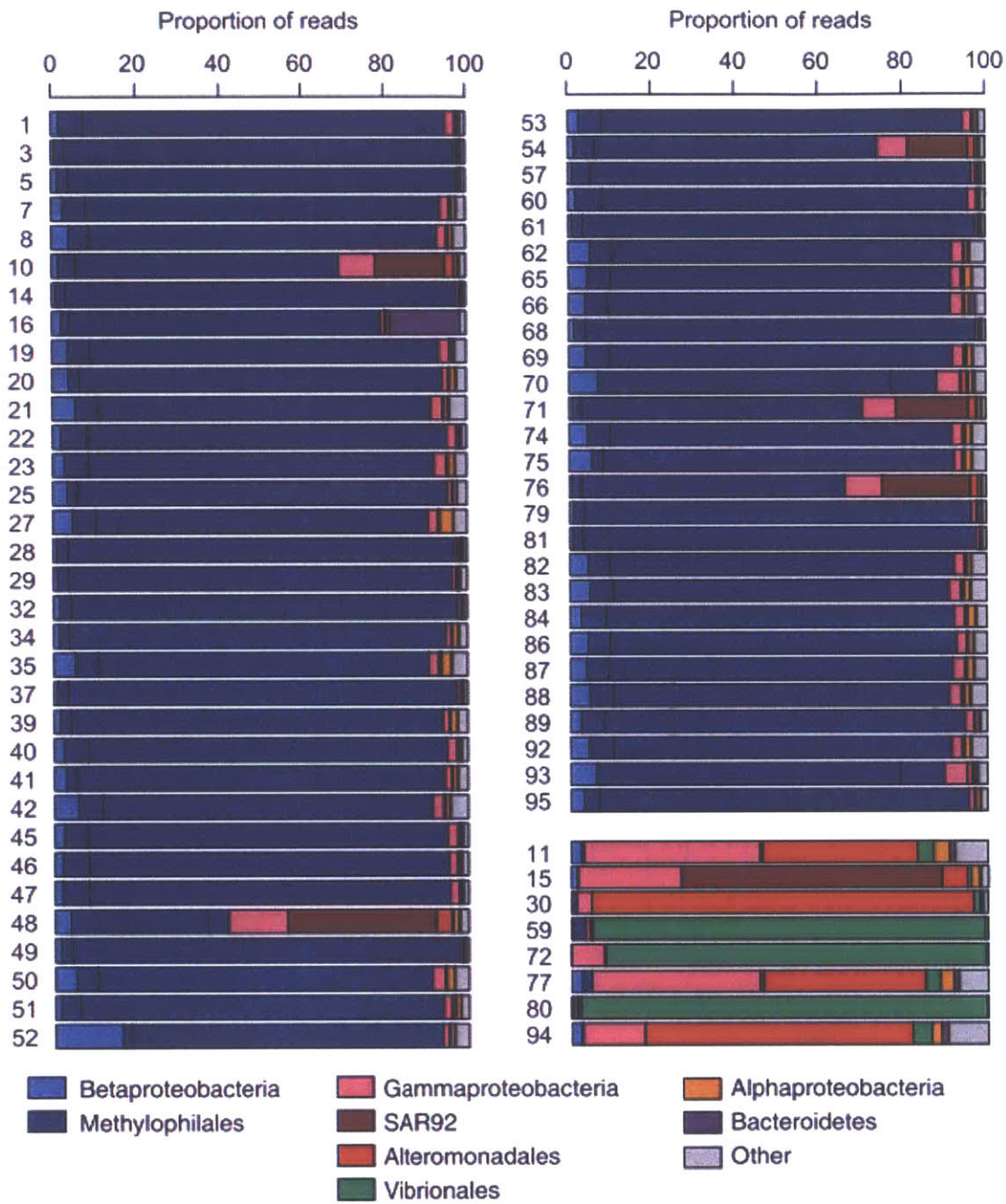


Figure 2.3. Taxonomic binning of unassembled reads based on the top hit of a LAST sequence similarity search against NCBI's REFSeq database. The culture identification number (NB00##) is located on the left of the sample bar.

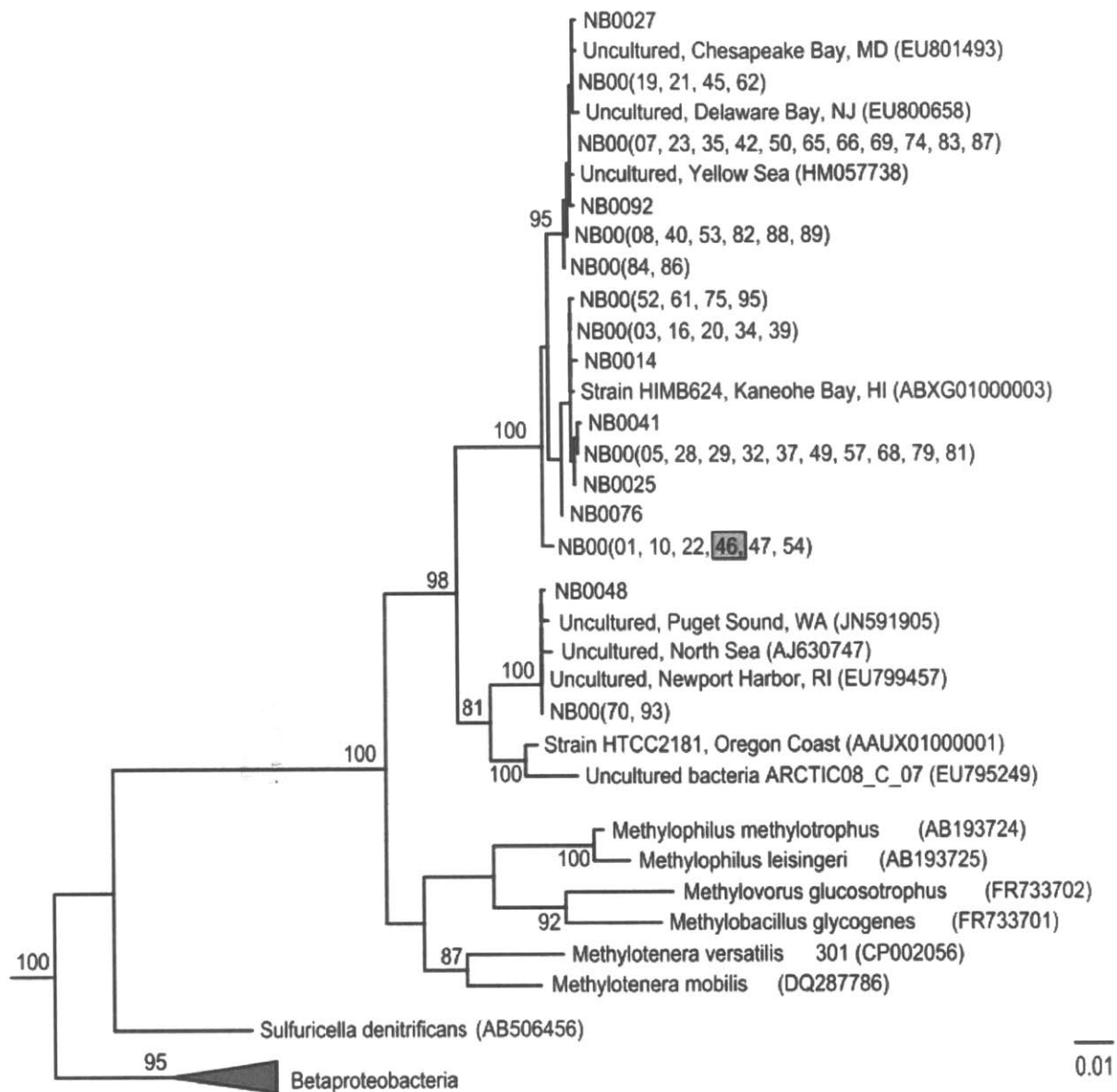


Figure 2.4. Phylogenetic relationships based on the SSU rRNA gene extracted from the whole-genome sequences of Nahant Bay isolates (NB00##) belonging to the OM43 clade of *Betaproteobacteria*. The grey box highlights strain NB0046 used in additional growth experiments (Fig. 2.5). The tree was inferred from 1350 alignment positions from sequences curated in the ARB software using the RAxML (maximum likelihood) method. RAxML bootstrap values (1000 replicates) are shown (480%) on nodes. Neighbor-joining inference method also produced bootstrap values 480% for these nodes. The scale bar indicates substitutions per site. *Gammaproteobacteria* and *Alphaproteobacteria* sequences (not shown) were used as outgroups.

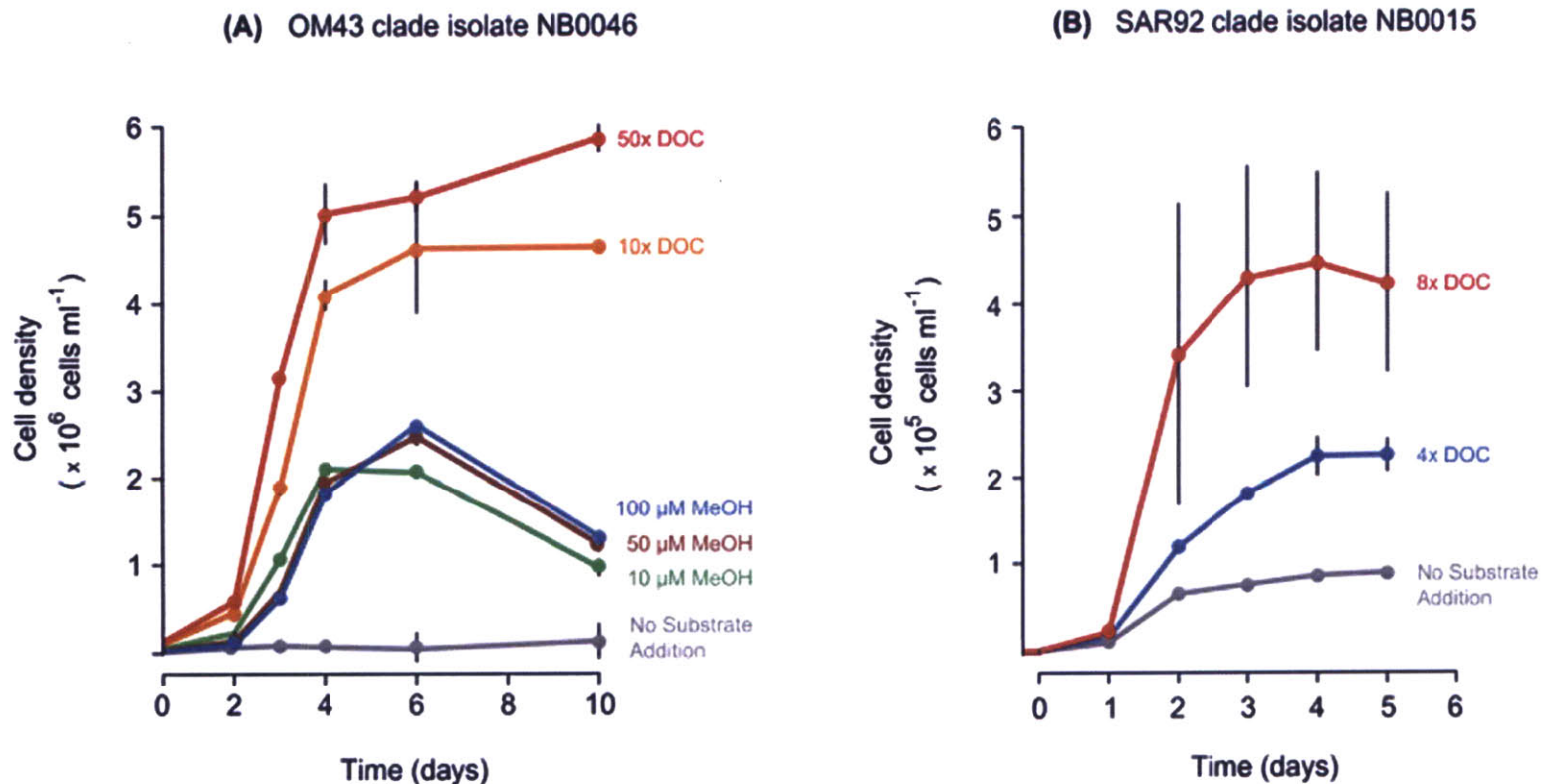


Figure 2. 5. Growth profiles of select cultures amended with HMW DOM. Cells from either (a) OM43 clade betaproteobacterium isolate NB0046 or (b) SAR92 clade gammaproteobacterium isolate NB0015 were inoculated into TFF oligotrophic seawater amended with HMW DOM and inorganic nutrients (OM43 culture: 400 μM NH_4 , 30 μM PO_4 ; SAR92 culture: 200 μM NH_4 , 200 μM NO_3 , 25 μM PO_4). Additional methanol (MeOH) treatments were included for the OM43 clade isolate. Points are the mean cell concentrations of four (OM43) or three replicates (SAR92), with error bars representing the s.d. (where not visible, error bars are smaller than the symbols).

References

- Aluwihare, L. I., & Repeta, D. J. (1999). A comparison of the chemical characteristics of oceanic DOM and extracellular DOM produced by marine algae. *Marine Ecology Progress Series*, 186, 105–117.
- Aluwihare, L. I., Repeta, D. J., & Chen, R. F. (1997). A major biopolymeric component to dissolved organic carbon in surface sea water. *Nature* 387, 166–169.
- Aluwihare, L. I., Repeta, D. J., Pantoja, S., & Johnson, C. G. (2005). Two chemically distinct pools of organic nitrogen accumulate in the ocean. *Science*, 308, 1007–1010.
- Andrykovitch, G., & Marx, I. (1988). Isolation of a new polysaccharide-digesting bacterium from a salt marsh. *Applied and Environmental Microbiology*, 54(4), 1061–1062.
- Baik, Y. S., & Cheong, W. J. (2007). Development of SPE for recovery of polysaccharides and its application to the determination of monosaccharides composition of the polysaccharide sample of a lactobacillus KLB 58. *Journal of Separation Science*, 30, 1509–1515.
- Becker, J. W., Brandon, M. L., & Rappé, M. S. (2007). Cultivating Microorganisms from Dilute Aquatic Environments: Melding Traditional Methodology with New Cultivation Techniques and Molecular Methods. In C. J. Hurst, R. Crawford, J. Garland, D. Lipson, A. Mills, & L. Stetzenbach (Eds.). *Manual of Environmental Microbiology* (3rd ed., pp. 339–406). ASM Press, Washington D. C.
- Beier, S., Rivers, A. R., Moran, M. A., & Obernosterer, I. (2014). The transcriptional response of prokaryotes to phytoplankton-derived dissolved organic matter in seawater. *Environmental Microbiology*, 17(10), 3466–3480.
- Benner, R. (2002). Chemical Composition and Reactivity. In Hansell, D. A. and Carlson, C. A. (Eds.). *Biogeochemistry of Marine Dissolved Organic Matter* (pp. 59–90). Academic Press, San Diego.
- Benner, R., Pakulski, J. D., McCarthy, M., Hedges, J. I., & Hatcher, P. G. (1992). *Science*, 255, 1561–1564.
- Button, D. K., Schut, F., Quang, P., Martin, R., & Robertson, B. R. (1993). Viability and isolation of marine bacteria by dilution culture: Theory, procedures, and initial results. *Applied and Environmental Microbiology*, 59(3), 881–891.
- Carini, P., Campbell, E. O., Morré, J., Sañudo-Wilhelmy, S. a, Thrash, J. C., Bennett, S. E., ... Giovannoni, S. J. (2014). Discovery of a SAR11 growth requirement for thiamin's pyrimidine precursor and its distribution in the Sargasso Sea. *The ISME Journal*, 8, 1727–38.

- Carlson, C. A. (2002). Production and Removal Processes. In Hansell D. A. and Carlson, C. A. (Eds.). *Biogeochemistry of Marine Dissolved Organic Matter* (pp. 139–151). Academic Press, San Diego.
- Carlson, C. a, Giovannoni, S. J., Hansell, D. a, Goldberg, S. J., Parsons, R., & Vergin, K. (2004). Interactions among dissolved organic carbon, microbial processes, and community structure in the mesopelagic zone of the northwestern Sargasso Sea. *Limnology and Oceanography*, *49*, 1073–1083.
- Carlson, C. A., & Ducklow, H. W. (1996). Growth of bacterioplankton and consumption of dissolved organic carbon in the Sargasso Sea. *Aquatic Microbial Ecology*, *10*, 69–85.
- Chistoserdova, L. (2011). Modularity of methylotrophy, revisited. *Environmental Microbiology*, *13*, 2603–2622.
- Cho, J. C., & Giovannoni, S. J. (2004). Cultivation and growth characteristics of a diverse group of oligotrophic marine *Gammaproteobacteria*. *Applied and Environmental Microbiology*, *70*(1), 432–440.
- Clark, L. L., & Ingall, E. D. (1998). Marine phosphorus is selectively remineralized. *Nature*, *393*, 426.
- Connon, S. A., & Giovannoni, S. J. (2002). High-throughput methods for culturing microorganisms in very-low-nutrient media yield diverse new marine isolates. *Applied and Environmental Microbiology*, *68*(8), 3878–3885.
- Distel, D. L., Morrill, W., MacLaren-Toussaint, N., Franks, D., & Waterbury, J. (2002). *Teredinibacter turnerae* gen. nov., sp. nov., a dinitrogen-fixing, cellulolytic, endosymbiotic-proteobacterium isolated from the gills of wood-boring molluscs (*Bivalvia: Teredinidae*). *International Journal of Systematic and Evolutionary Microbiology*, *52*, 2261–2269.
- Eloe, E. A, Malfatti, F., Gutierrez, J., Hardy, K., Schmidt, W. E., Pogliano, K., ... Bartlett, D. H. (2011). Isolation and characterization of a psychropiezophilic alphaproteobacterium. *Applied and Environmental Microbiology*, *77*(22), 8145–8153.
- Frith, M. C., Hamada, M., & Horton, P. (2010). Parameters for accurate genome alignment. *BMC Bioinformatics*, *11*, 80.
- Gifford, S. M., Sharma, S., & Moran, M. A. (2014). Linking activity and function to ecosystem dynamics in a coastal bacterioplankton community. *Frontiers in Microbiology*, *5*, 1–12.
- Giovannoni, S. J., Hayakawa, D. H., Tripp, H. J., Stingl, U., Givan, S. a., Cho, J. C., ... Rappé, M. S. (2008). The small genome of an abundant coastal ocean methylotroph. *Environmental Microbiology*, *10*, 1771–1782.

- González, J. M., Mayer, F., Moran, M. a, Hodson, R. E., & Whitman, W. B. (1997). *Microbulbifer hydrolyticus* gen. nov., sp. nov., and *Marinobacterium georgiense* gen. nov., sp. nov., two marine bacteria from a lignin-rich pulp mill waste enrichment community. *International Journal of Systematic Bacteriology*, 47(706), 369–376.
- Grote, J., Cameron Thrash, J., Huggett, M. J., Landry, Z. C., Carini, P., Giovannoni, S. J., & Rappé, M. S. (2012). Streamlining and core genome conservation among highly divergent members of the SAR11 clade. *mBio*, 3(5), e00252–12.
- Halsey, K. H., Carter, A. E., & Giovannoni, S. J. (2012). Synergistic metabolism of a broad range of C1 compounds in the marine methylotrophic bacterium HTCC2181. *Environmental Microbiology*, 14, 630–640.
- Huggett, M. J., Hayakawa, D. H., & Rappé, M. S. (2012). Genome sequence of strain HIMB624, a cultured representative from the OM43 clade of marine *Betaproteobacteria*. *Standards in Genomic Sciences*, 6, 11–20.
- Hutalle-Schmelzer, K. M. L., Zwirnmann, E., Krüger, A., & Grossart, H. P. (2010). Enrichment and cultivation of pelagic bacteria from a humic lake using phenol and humic matter additions. *FEMS Microbiology Ecology*, 72, 58–73.
- Janvier, M., Frehel, C., Grimont, F., & Gasser, F. (1985). *Methylophaga marina* gen. nov., sp. nov. and *Methylophaga thalassica* sp. nov., marine methylotrophs. *International Journal of Systematic Bacteriology*, 35(2), 131–139.
- Kester, D. R., Duedall, I. W., Connors, D. N., & Pytkowicz, R. M. (1967). Preparation of artificial seawater. *Limnology and Oceanography*, 12(1), 176–179.
- Kielbasa, S. M., Wan, R., Sato, K., Horton, P., & Frith, M. C. (2011). Adaptive seeds tame genomic sequence comparison. *Genome Research*, 21, 487–93.
- Kolowith, L. C., Ingall, E. D., & Benner, R. (2001). Composition and cycling of marine organic phosphorus. *Limnology and Oceanography*, 46(2), 309–320.
- Landa, M., Cottrell, M. T., Kirchman, D. L., Kaiser, K., Medeiros, P. M., Tremblay, L., ... Obernosterer, I. (2013). Phylogenetic and structural response of heterotrophic bacteria to dissolved organic matter of different chemical composition in a continuous culture study. *Environmental Microbiology*, 16(6):1668–1681.
- Letunic, I., & Bork, P. (2007). Interactive tree of life (iTOL): An online tool for phylogenetic tree display and annotation. *Bioinformatics*, 23, 127–128.
- Lin, S. Y., Shieh, Y., Chen, J., & Tang, S. (2013). Complete genome sequence of *Simiduia agarivorans* SA1 T, a marine bacterium able to degrade a variety of polysaccharides. *Genome Announcements*, 1(1), 1–2.

- Ludwig, W. (2004). ARB: a software environment for sequence data. *Nucleic Acids Research*, 32(4), 1363–1371.
- Malmstrom, R. R., Kiene, R. P., Cottrell, M. T., & Kirchman, D. L. (2004). Contribution of SAR11 both dissolved dimethylsulfoniopropionate and amino acid uptake in the North Atlantic Ocean contribution of SAR11 bacteria to dissolved dimethylsulfoniopropionate and amino acid uptake in the North Atlantic Ocean. *Applied and Environmental Microbiology*, 70(7), 4129–4135.
- Marshall, K. T., & Morris, R. M. (2012). Isolation of an aerobic sulfur oxidizer from the SUP05/Arctic96BD-19 clade. *The ISME Journal*, 7(2), 452–455.
- McCarren, J., Becker, J. W., Repeta, D. J., Shi, Y., Young, C. R., Malmstrom, R. R., ... DeLong, E. F. (2010). Microbial community transcriptomes reveal microbes and metabolic pathways associated with dissolved organic matter turnover in the sea. *Proceedings of the National Academy of Sciences of the USA*, 107, 16420–7.
- McCarthy, M. D., Pratum, T., Hedges, J. I., & Benner, R. (1997). Chemical composition of dissolved organic nitrogen in the ocean. *Nature*, 390, 150–154.
- Menzel, D. W., & Spaeth, J. P. (1962). Occurrence of vitamin B12 in the Sargasso Sea. *Limnology and Oceanography*, 7, 151–154.
- Meon, B., & Kirchman, D. L. (2001). Dynamics and molecular composition of dissolved organic material during experimental phytoplankton blooms. *Marine Chemistry*, 75, 185–199.
- Miller, M. a, Pfeiffer, W., & Schwartz, T. (2010). Creating the CIPRES Science Gateway for inference of large phylogenetic trees. *2010 Gateway Computing Environments Workshop, GCE 2010*, 1–8.
- Moran, M. A., & Zepp, R. G. (1997). Role of photoreactions in the formation of biologically labile compounds from dissolved organic matter. *Limnology and Oceanography*, 42(6), 1307–1316.
- Morris, R. M., Longnecker, K., & Giovannoni, S. J. (2006). Pirellula and OM43 are among the dominant lineages identified in an Oregon coast diatom bloom. *Environmental Microbiology*, 8, 1361–1370.
- Munoz, R., Yarza, P., Ludwig, W., Euzéby, J., Amann, R., Schleifer, K.-H., ... Rosselló-Móra, R. (2011). Release LTPs104 of the All-Species Living Tree. *Systematic and Applied Microbiology*, 34, 169–70.
- Nelson, C. E., & Carlson, C. A. (2012). Tracking differential incorporation of dissolved organic carbon types among diverse lineages of Sargasso Sea bacterioplankton. *Environmental Microbiology*, 14, 1500–1516.

- Poretzky, R. S., Sun, S., Mou, X., & Moran, M. A. (2010). Transporter genes expressed by coastal bacterioplankton in response to dissolved organic carbon. *Environmental Microbiology*, *12*, 616–627.
- Pruesse, E., Peplies, J., & Glockner, F. O. (2012). SINA: Accurate high-throughput multiple sequence alignment of ribosomal RNA genes. *Bioinformatics*, *28*(14), 1823–1829.
- Quan, T. M., & Repeta, D. J. (2007). Periodate oxidation of marine high molecular weight dissolved organic matter: Evidence for a major contribution from 6-deoxy- and methyl sugars. *Marine Chemistry*, *105*, 183–193.
- Rappé, M. S., Connon, S. A., Vergin, K. L., & Giovannoni, S. J. (2002). Cultivation of the ubiquitous SAR11 marine bacterioplankton clade. *Nature*, *418*, 630–633.
- Rappé, M. S., Kemp, P. F., & Giovannoni, S. J. (1997). Phylogenetic diversity of marine coastal picoplankton 16S rRNA genes cloned from the continental shelf off Cape Hatteras, North Carolina. *Limnology and Oceanography*, *42*, 811–826.
- Redfield, A. C., Ketchum, B. H., & Richards, F. A. (1963). The influence of organisms on the composition of sea water. In Hill, M. N. (Ed.) *The Sea*, vol. 2. (pp. 26–27). Interscience Publishers, New York.
- Repeta, D. J., & Aluwihare, L. I. (2006). Radiocarbon analysis of neutral sugars in high-molecular-weight dissolved organic carbon: Implications for organic carbon cycling. *Limnology and Oceanography*, *51*(2), 1045–1053.
- Rich, V. I., Pham, V. D., Eppley, J., Shi, Y., & DeLong, E. F. (2011). Time-series analyses of Monterey Bay coastal microbial picoplankton using a “genome proxy” microarray. *Environmental Microbiology*, *13*, 116–134.
- Saitou, N., & Nei, M. (1987). The neighbor-joining method: A new method for reconstructing phylogenetic trees. *Molecular Biology and Evolution*, *4*(4), 406–425.
- Sarmiento, H., & Gasol, J. M. (2012). Use of phytoplankton-derived dissolved organic carbon by different types of bacterioplankton. *Environmental Microbiology*, *14*, 2348–2360.
- Sarmiento, H., Romera-Castillo, C., Lindh, M., Pinhassi, J., Sala, M. M., Gasol, J. M., ... Taylor, G. T. (2013). Phytoplankton species-specific release of dissolved free amino acids and their selective consumption by bacteria. *Limnology and Oceanography*, *58*(3), 1123–1135.
- Schwalbach, M. S., Tripp, H. J., Steindler, L., Smith, D. P., & Giovannoni, S. J. (2010). The presence of the glycolysis operon in SAR11 genomes is positively correlated with ocean productivity. *Environmental Microbiology*, *12*, 490–500.
- Sharma, A. K., Becker, J. W., Ottesen, E. a, Bryant, J. a, Duhamel, S., Karl, D. M., ... DeLong, E. F. (2014). Distinct dissolved organic matter sources induce rapid transcriptional

responses in coexisting populations of *Prochlorococcus*, *Pelagibacter* and the OM60 clade. *Environmental Microbiology*, 16, 2815–30.

- Shieh, W. Y., Liu, T. Y., Lin, S. Y., Jean, W. D., & Chen, J. S. (2008). *Simiduia agarivorans* gen. nov., sp. nov., a marine, agarolytic bacterium isolated from shallow coastal water from Keelung, Taiwan. *International Journal of Systematic and Evolutionary Microbiology*, 58, 895–900.
- Shrestha, P. M., Nevin, K. P., Shrestha, M., & Lovley, D. R. (2013). When is a microbial culture “Pure”? Persistent cryptic contaminant escapes detection even with deep genome sequencing. *mBio*, 4(2), e00591–12.
- Song, J., Oh, H. M., & Cho, J. C. (2009). Improved culturability of SAR11 strains in dilution-to-extinction culturing from the East Sea, West Pacific Ocean. *FEMS Microbiology Letters*, 295, 141–147.
- Sowell, S. M., Abraham, P. E., Shah, M., Verberkmoes, N. C., Smith, D. P., Barofsky, D. F., & Giovannoni, S. J. (2011). Environmental proteomics of microbial plankton in a highly productive coastal upwelling system. *The ISME Journal*, 5(5), 856–865.
- Stamatakis, A. (2014). RAxML version 8: A tool for phylogenetic analysis and post-analysis of large phylogenies. *Bioinformatics*, 30, 1312–1313.
- Stingl, U., Desiderio, R. a., Cho, J. C., Vergin, K. L., & Giovannoni, S. J. (2007). The SAR92 clade: An abundant coastal clade of culturable marine bacteria possessing proteorhodopsin. *Applied and Environmental Microbiology*, 73(7), 2290–2296.
- Sun, J., Steindler, L., Thrash, J. C., Halsey, K. H., Smith, D. P., Carter, A. E., ... Giovannoni, S. J. (2011). One carbon metabolism in SAR11 pelagic marine bacteria. *PLoS ONE*, 6(8).
- Suzuki, M. T., Preston, C. M., Béjà, O., de la Torre, J. R., Steward, G. F., & DeLong, E. F. (2004). Phylogenetic screening of ribosomal RNA gene-containing clones in bacterial artificial chromosome (BAC) libraries from different depths in Monterey Bay. *Microbial Ecology*, 48, 773–775.
- Tamura, K., Peterson, D., Peterson, N., Stecher, G., Nei, M., & Kumar, S. (2011). MEGA5: Molecular evolutionary genetics analysis using maximum likelihood, evolutionary distance, and maximum parsimony methods. *Molecular Biology and Evolution*, 28(10), 2731–2739.
- Weiner, R. M., Taylor, L. E., Henrissat, B., Hauser, L., Land, M., Coutinho, P. M., ... Hutcheson, S. (2008). Complete genome sequence of the complex carbohydrate-degrading marine bacterium, *Saccharophagus degradans* strain 2-40T. *PLoS Genetics*, 4(5), e1000087.
- Williams, P. M., & Druffel, E. R. M. (1987). Radiocarbon in dissolved organic matter in the central North Pacific Ocean. *Nature*, 330, 246–248.

Supplemental material

Table S1. Cell counts of dilution to extinction wells after growth screen. Wells with >100,000 cell ml⁻¹ were scored positive. Sterile wells were prepared on separate plates than control wells, which were prepared on the treatment plates (6 per plate).

10³ Cells ml⁻¹	Sterile wells	Control wells	No DOC	4X DOC	10X DOC
<10	48	360	953	908	876
10-50	0	0	4	11	30
50-100	0	0	2	1	2
100-500	0	0	0	7	5
500-1000	0	0	1	13	11
>1000	0	0	0	20	36
Total wells	48	360	960	960	960

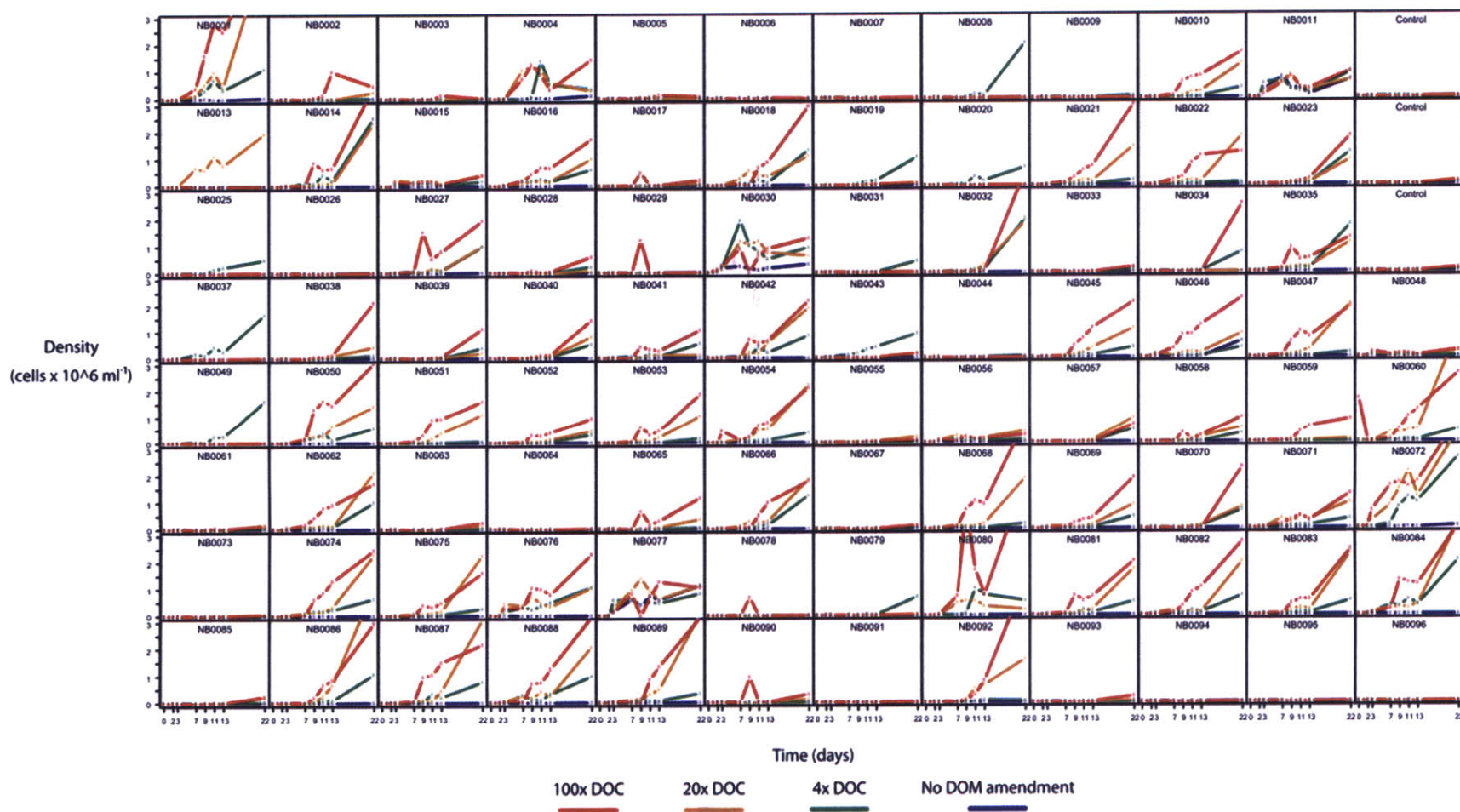
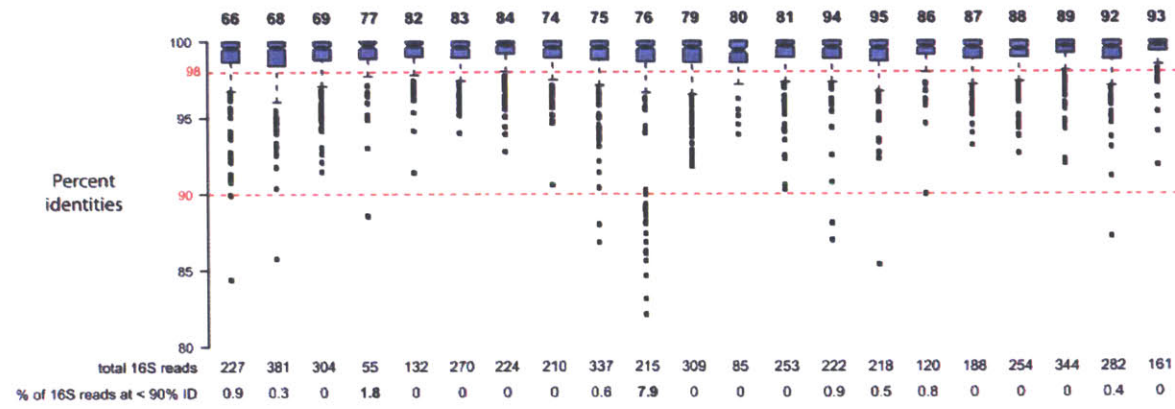
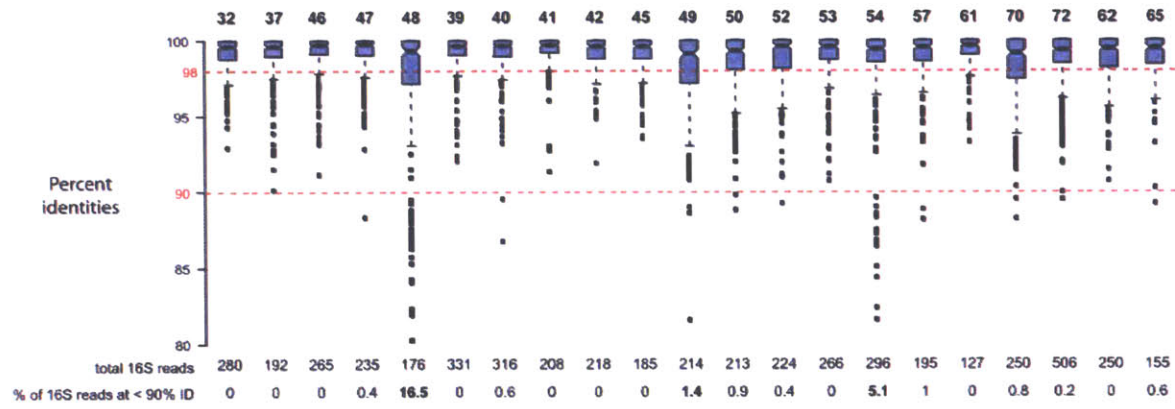
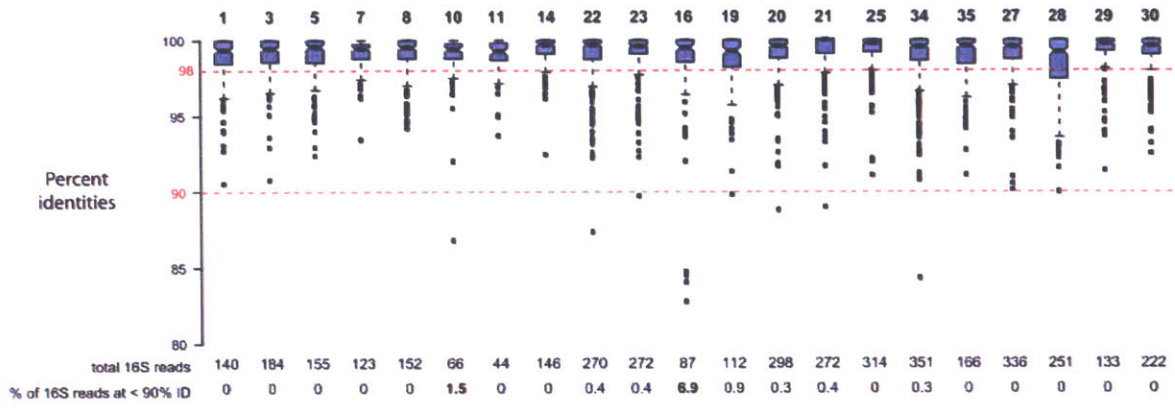


Figure S1. Growth yields of the entire HMW DOM-enriched culture collection under different HMW DOM concentrations (100x: 6930 μ M DOC, 20x: 1330 μ M DOC, 4x: 210 μ M DOC, and no DOM amendment) in TFF UV oxidized seawater.

(Opposite page)

Figure S2. Distribution of the cultures' unassembled SSU rRNA reads' percent identity to the assembled SSU rRNA sequence. Quality controlled unassembled reads >50 nt in length were compared to the assembled SSU sequences via a BLASTn search, with reads hitting the SSU rRNA with a bit score > 50 considered significant hits. The distribution of percent IDs from the BLASTn search is shown as a boxplot, with the bottom and top of the box representing the 1st and 3rd quartiles, and the inner notch marking the median. Outliers (percent IDs > 1.5 times the interquartile range) are shown as individual points. The total reads meeting the BLASTn significance criteria, as well as the percent of those reads at less than 90% identity to the assembled sequence, are shown at the bottom. The culture identification number (NB00##) is shown above its respective plot.



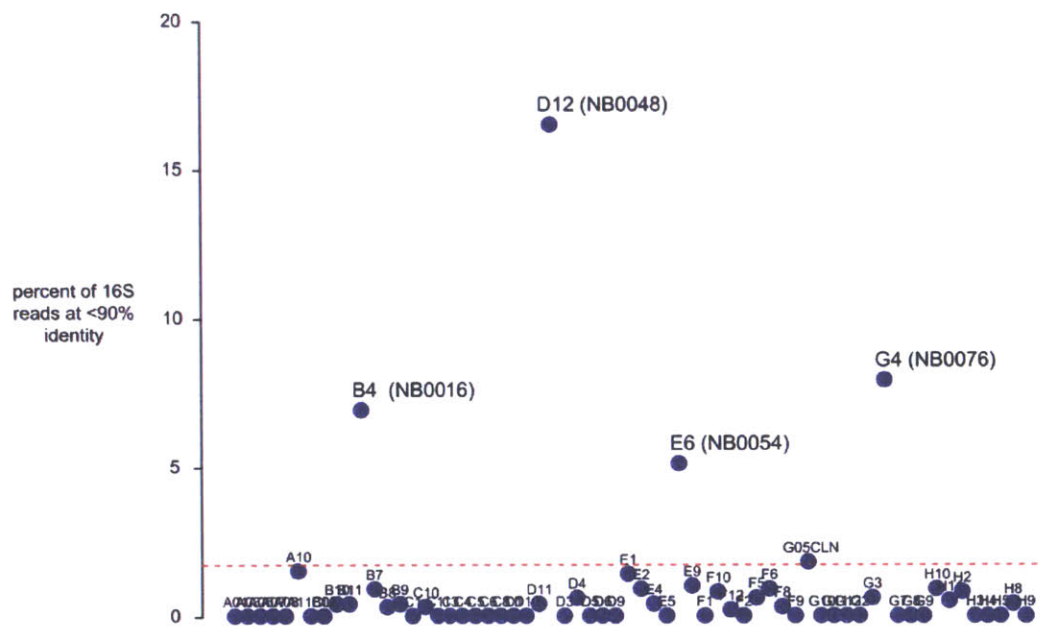


Figure S3. Distribution of unassembled SSU rRNA reads at less than 90% identity to the assembled SSU rRNA gene sequence. See Figure S2 legend for annotation criteria.

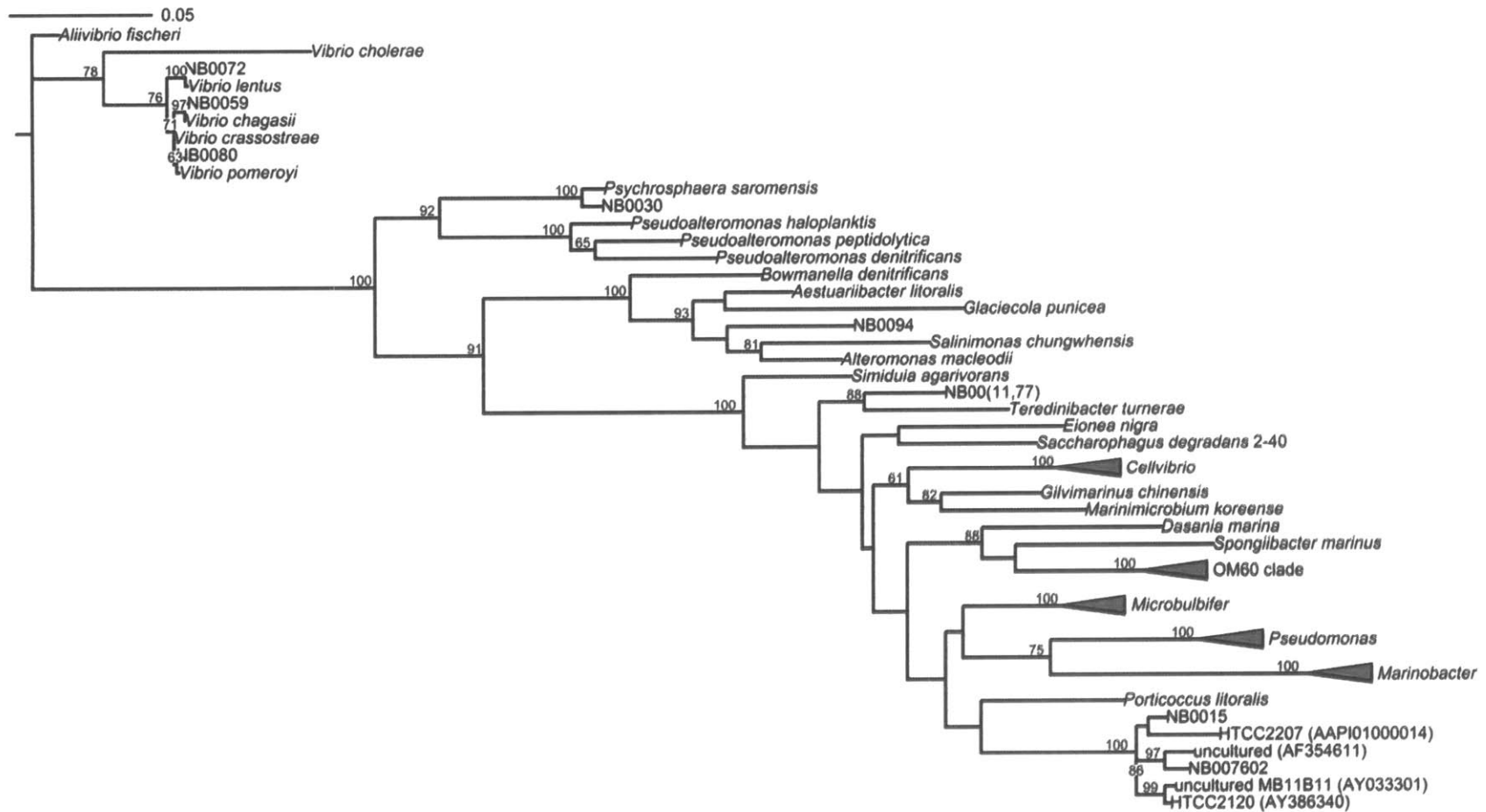


Figure S4. Phylogenetic relationships based on the SSU rRNA gene extracted from Nahant Bay (NB00##) gammaproteobacteria isolates. The tree was inferred from 1293 unambiguous alignment positions from sequences curated in the ARB software (reference) using the RAxML (maximum likelihood) method. RAxML bootstrap values (1000 replicates) are shown (>60%) on nodes. The tree is mid-point rooted between the most distantly related taxa (*Vibrio cholera* and *Marinobacter szutsaonensis*, not shown). The scale bar indicates substitutions per site.

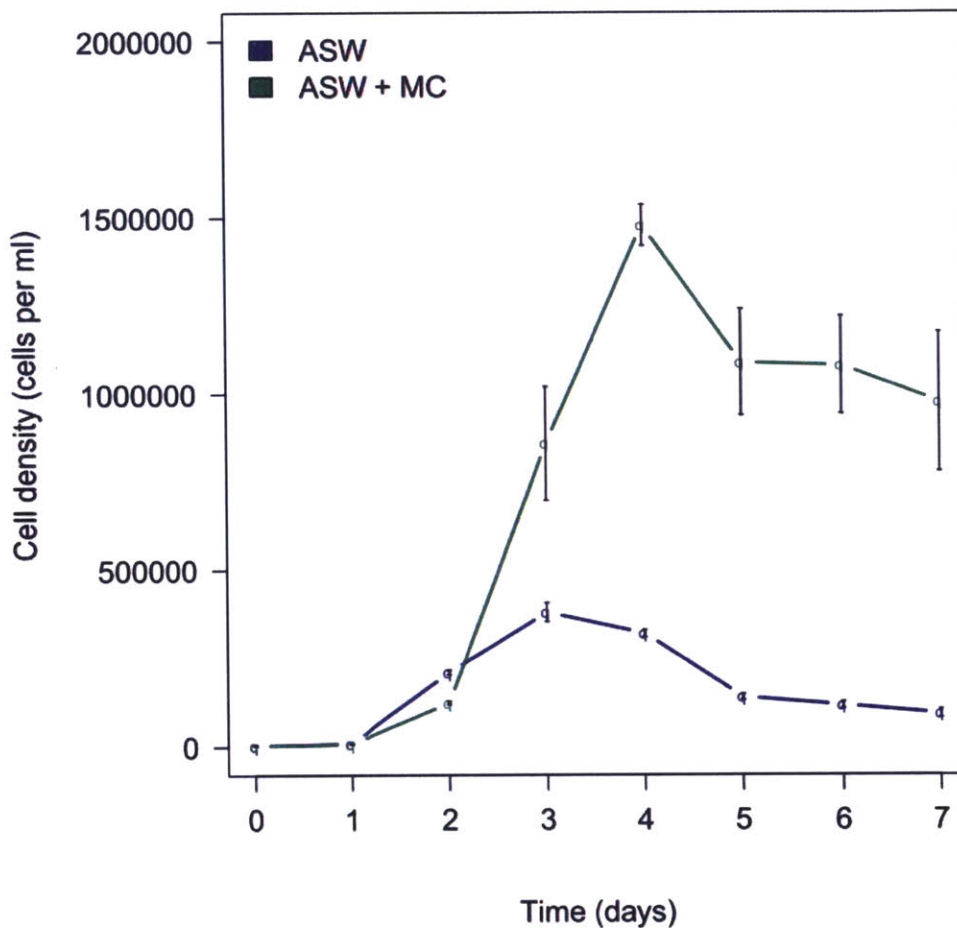


Figure S5. Growth profiles of SAR92 clade isolate NB0015 in defined media. Cells of NB0015 were grown in artificial seawater (ASW) with and without mixed carbons (MC) including D-glucose (55 μM), succinate (85 μM), pyruvate (114 μM), glycerol (109 μM), N-acetyl D-glucosamine (45 μM), and ethanol (434 μM). Points are the mean cell concentrations of three replicates, with error bars representing the standard deviation (where not visible, error bars are smaller than the symbols).

CHAPTER THREE

Assessment of the degradation capabilities of open-ocean bacterial isolates enriched with marine high-molecular-weight dissolved organic matter

Abstract

In the previous chapter, I described how the experimental additions of high-molecular-weight (HMW) dissolved organic matter (DOM) to dilution-to-extinction cultures resulted in the enrichment of bacterial groups able to degrade this natural substrate to obtain carbon and energy for growth. In this chapter I apply the same methodology to cultivate bacterial groups from open-ocean planktonic communities and investigate their biochemical and metabolic potential to degrade HMW DOM through growth experiments and genome analysis. Two DOM samples, ultra-filtered HMW DOM and a purified sub-fraction (F1) enriched in HMW polysaccharides obtained by chromatography, were used to supplement dilution-to-extinction cultures with dissolved organic carbon (DOC) above ambient concentrations. The inocula consisted of seawater collected at Station ALOHA (22° 45'N, 158° 00'W) from the deep chlorophyll maximum (DCM) and from mesopelagic depths (>200 m). The addition of HMW DOM resulted in a significantly higher number of cultures recovered from mesopelagic samples whereas DCM samples did not respond significantly to the treatments. Genomic analysis indicated that these isolates have the genetic pathways required to break down and metabolize known DOM constituents such as carbohydrates and phosphonates. However, cultures exhibited relatively small changes in cell yields in the presence of HMW DOM when DOC was increased from ~80 μM ambient concentrations to over 1200 μM in treatments. Taken together these results indicate that as monocultures these isolates have limited access to HMW DOM though they seem sensitive to increases in DOC concentrations. This may indicate that in the open-ocean, the breakdown of HMW DOM is heavily dependent on microbial consortia and abiotic degradation processes. Future work employing co-cultures rather than monocultures may help to further identify important microbial interactions mediating DOM cycling.

Introduction

The subtropical gyres of the Pacific and Atlantic Oceans represent some of the most extensive biomes on the planet, and as such have important influence on the global carbon cycle. These open-ocean regions are characteristically oligotrophic ecosystems. Their distance from large land masses, the central gyre hydrography, and strong stratification of the water column limits the availability of nutrients – compared to coastal and upwelling systems – that sustain primary production. Such conditions favor the so-called “regeneration loop” (Sarmiento & Gruber, 2006) in which the organic matter produced in the surface is efficiently recycled to renew nutrient (inorganic nitrogen and phosphorus) supplies in great part through the “microbial loop” (Azam, 1998; Azam et al., 1983).

Microbial heterotrophic communities are thought to play a major role regulating the turnover of organ matter. In addition to establishing a link between microorganisms and higher trophic levels through the microbial loop, DOC remineralization maintains the balance between the organic carbon transported downward by the biological pump and the upward transport of dissolved inorganic carbon (Sarmiento & Gruber, 2006). The organic matter that escapes biological remineralization in the surface is exported to the dark ocean via the biological pump by lateral and downward transport of DOM and by sinking of POM (Ducklow, Steinberg, & Buesseler, 2001). In the North Atlantic Subtropical Gyre for example, organic matter is exported during the spring after winter convective mixing returning to regenerated production thereafter whereas in the North Pacific Subtropical Gyre the water column is permanently stratified favoring regenerated production (Brix, Gruber, Karl, & Bates, 2006).

In addition to the strong stratification in the gyres, the accumulation of degradable DOM in the surface ocean is thought to be the result of a malfunctioning microbial loop wherein DOC

cannot be consumed as quickly as it is produced (Thingstad, Hagström, & Rassoulzadegan, 1997). This is because microbial heterotrophic communities in the sunlit regions of the ocean are thought to be highly dependent for their growth and metabolism on the production of organic carbon by photosynthetic organisms with which they also compete for nutrients. Accordingly, estimates of bacterial activity are usually best predicted by phytoplankton primary productivity (Cole, Findlay, & Pace, 1988) and diel patterns are often observed in parameters like cellular volume and bacterial abundance relating to phytoplankton activity (Kuipers, Van Noort, Vosjan, & Herndl, 2000). Even at the transcriptional level diverse bacterial heterotrophic clades appear to be synchronized to autotrophic processes (Aylward et al., 2015; Ottesen et al., 2014). On the other hand, in the dark ocean, where the available light no longer supports photosynthesis, microorganisms can obtain carbon, energy, and nutrients from sinking POM (Cho & Azam, 1988). However, the particulate organic carbon (POC) flux alone cannot explain the observed respiration rates in the dark ocean; to close the budget DOC exported to these depths ($\sim 2.0 \text{ mol C m}^{-2} \text{ yr}^{-1}$) must be remineralized by microbial heterotrophic communities (Carlson et al., 1994; Emerson, Quay, Karl, Winn, & Tupas, 1997). The majority ($\sim 70\%$) of the respiration in the dark ocean, $3\text{--}4 \text{ mol C m}^{-2} \text{ yr}^{-1}$, is thought to occur within the mesopelagic (150-1000 m) fueled primarily by POC and of which exported DOC can contribute between 10-20% (Aristegui, Agustí, Middelburg, & Duarte, 2005; Carlson et al., 2010, 1994). Illustrating the significance of DOC export, open-ocean microcosm experiments in the North Atlantic Subtropical Gyre found that growth of mesopelagic microbial communities was stimulated in the presence of filtered surface water which harbors the surplus of semi-labile DOM that accumulates periodically (Carlson et al., 2004a). However, the transformation and degradation processes by which semi-labile DOM supports microbial metabolism remain largely unexplored.

Ultra-filtered marine HMW DOM may prove a useful and appropriate substrate to study microbial cycling of semi-labile DOM. HMW DOM is operationally defined as the fraction of DOM that can be typically concentrated by 1kDa ultrafiltration membranes and constitutes about a third of the standing stock of DOM in marine surface waters (Benner et al., 1997; McCarthy et al., 1996). Notably, this HMW fraction is made up primarily of carbohydrates (~50% in surface waters) with a fairly homogenous monosaccharide composition across major ocean basins indicating this material has a widespread and common source throughout the global ocean (Aluwihare et al., 1997; McCarthy et al., 1996). Radiocarbon analysis of the monosaccharides making up HMW DOM carbohydrates indicates this fraction has an estimated residence time of <3 years (Repeta & Aluwihare, 2006) consistent with the residence time of semi-labile DOM (100 to 1000 days) obtained from ecosystem models (Luo et al., 2010). Interestingly, the carbohydrate content of HMW DOM has been observed to decrease with depth indicating that polysaccharides are a major bio-reactive component (Benner et al., 1992; McCarthy et al., 1993).

Despite several lines of evidence that demonstrate the importance of microorganisms in the cycling of DOM, there is limited information on what types of microbial groups and biochemical pathways mediate this basic ecological function. The work described in this chapter aims to identify microbial model systems that may serve to investigate in detail the processes that drive biological semi-labile DOM turnover. For this purpose, HMW DOM, a widespread and naturally occurring carbon-rich and nutrient-poor substrate representative of semi-labile DOM, is supplemented to dilution-to-extinction cultures to stimulate the growth of HMW DOM-degrading microorganisms. The experiments were conducted with seawater collected in the field site of the Hawaii Ocean Time series (HOT), Station ALOHA, a region representative of general conditions in the North Pacific Subtropical Gyre. The cultivation media for dilution-to extinction

experiments consisted of natural seawater collected in Hawaii waters sterilized by tangential flow filtration as described in Chapters 1 and 2. Microbial samples for cultivation experiments were obtained from two depths, one within the euphotic zone specifically targeting the DCM (~100 m), and one below it in the upper mesopelagic (below 200 m). The DCM at Station ALOHA is situated below the surface mixed layer year-round and coincides roughly with the depth of 1% surface photosynthetically available radiation (Laws, Letelier, & Karl, 2014; Letelier, Karl, Abbott, & Bidigare, 2004). In this region, *Prochlorococcus* and picoeukaryotes dominate the photosynthetic biomass (Campbell, Liu, Nolla, & Vaultot, 1997) and heterotrophic populations may rely more heavily on newly produced organic matter to obtain carbon and energy for growth. In contrast, in mesopelagic depths where light levels no longer support photosynthesis, heterotrophic bacteria may be better adapted to degrade semi-labile HMW DOM. Thus, I expected to identify a higher number of microbial degraders of HMW DOM in cultivation experiments prepared from mesopelagic samples than from DCM samples. It is possible, however, that heterotrophic microorganisms from both depths support a fraction of their carbon demand with HMW DOM.

Materials and Methods

Collection of HMW DOM and isolation of neutral polysaccharides

Ultrafiltration

Large-volume seawater samples were drawn from the 20m intake pipe at the Natural Energy Laboratory in Kona, Hawaii (19.727°W, 156.063°N) in February 2013. The samples were filtered to remove bacteria and small particles using a cleaned (10% HCl) Suporflow dual-stage (0.8 and 0.2 microns) Gelman polyether sulfone (PES) cartridge filter (Chisolm-Pall) fitted to an

Advanta stainless-steel housing. The high molecular weight dissolved organic matter fraction (HMWDOM) was collected using two spiral wound ultrafiltration systems consisting of a stainless-steel centripetal pump and membrane housings and a fluorinated high-density polyethylene reservoir. Each system was plumbed with Teflon tubing and PVDF valves, and fitted with two ultrafiltration membranes (Osmonics GE-H4040C; Separation Engineering) that nominally retain organic matter with a relative molecular weight >1 kDa ($> 99\%$ rejection of vitamin B-12). Membranes were cleaned using isopropanol, detergent (0.01% micro), HCl (0.01 M), and NaOH (0.01 M); stored in sodium azide (0.55 mM); and rinsed with water immediately before use. Approximately 32,241 liters of seawater were concentrated over a period of 14 days. Each day, samples (~ 1150 liter per system) were concentrated to 20 liter, filtered through a 0.2 micron PES cartridge filter (Propor demi-cap) into a fluorinated HDPE carboy and stored in a -20°C freezer. The sample was cooled, but did not freeze during storage. The following day, this concentrate was combined with the new day's concentrate, reduced in volume to 20 liter, and again stored at -20°C . This process was repeated daily. At the end of the sampling period, the pooled samples were frozen, and returned to Woods Hole for further processing. Dissolved organic carbon measurements showed 22% of TOC was recovered in the concentrate. In Woods Hole, samples were thawed, filtered through a 0.2 micron PES filter, then filtered again through a stirred cell system fitted with a 30 kD cellulose fiber (ultracell) filter to remove viruses. Samples were desalted by a smaller ultrafiltration system fitted with Osmonics GE-E4010 filters. Twenty liters of sample was concentrated to 2 liter, to which 2 liter of ultra-pure water was added. The diluted sample was again reduced in volume to 2 liter, and the process repeated a total of 12 times, until the no visible precipitate was observed with the addition of sample to 10 mg ml^{-1} silver nitrate. Desalted samples were lyophilized to a fluffy white powder.

Isolation of HMW DOM neutral polysaccharides

A 2.2 cm i.d. x 30 cm glass chromatography column was slurry packed with 16g of Biorex-5 (BioRad Corp.) anion exchange resin (chloride form) suspended in ultrapure water (Q-water). The water drained to the resin bed. Sodium hydroxide (1.92 g) was dissolved in 88 mls of Q-water and the solution used to wash the column to convert the resin to the hydroxide form. The column was washed with 80 mL Q-water, and the pH checked to ensure the final wash was neutral. Approximately 0.5 g of freeze dried HMWDOM (~36% by weight C) was dissolved in 5 mL of warm Q-water, and the solution applied to the column. The column was drained to the bed. The first fraction (F1) was eluted by washing the column with 80 mLs of Q water. The fraction was freeze dried to yield ~ 150 mg of a pure white, cotton-like material. ¹H and ¹³C NMR of F1 showed the preparation to be carbohydrate. ³¹P NMR showed the presence of phosphonates (20% total P), phosphate esters (70% of total P) and pyrophosphate (10% of total P).

Overview of dilution-to-extinction experiments at Station ALOHA

Two dilution-to-extinction cultivation experiments were prepared with seawater collected in the open-ocean near Hawaii (Station ALOHA, 22° 45'N, 158° 00'W). The experiments consisted of treatments enriched with HMW DOM and control media with no DOC added. The cultivation media consisted of natural Hawaii seawater collected offshore, sterilized by tangential flow ultrafiltration (30 kDa) as described in Chapter 2 (Sosa, Gifford, Repeta & DeLong, 2015) and amended with inorganic nitrogen and phosphorus to alleviate nutrient limitation and promote HMW DOM carbon utilization. Trace metals and vitamins were not added to the media. Details on the preparation of each experiment including HMW DOM and nutrient amendments are summarized in Table 3.1.

Experiment I

Seawater inoculum samples for dilution-to-extinction were collected from the DCM (110 m-depth, based on in situ fluorometry) and from the Mesopelagic (200 m-depth) during the Hawaii Ocean Time series (HOT) cruise 250 (KM 13-05) on March 9, 2013. Inoculum samples were initially diluted ~1000-fold onboard with sterile seawater prepared by tangential flow ultrafiltration (30 kDa cut-off). Total cell counts were obtained from formaldehyde-fixed samples by epi-fluorescence microscopy for the raw and diluted inoculum seawater. Samples were kept onboard at room temperature in darkness for 14 hours until arrival to port and transport to the laboratory where the dilution-to-extinction was completed. Dilution-to-extinction samples incubated for an additional 24 hours during transport from Hawaii to MIT before they were amended with HMW DOM and arrayed into 48-well cultivation (1 ml well⁻¹). Three DOC treatments were prepared with dilution-to-extinction samples from each depth. The DOC treatments consisted of a control with no DOC added, and two experimental treatments; the first was supplemented with 2.5x DOC (ca. 200 μM) and the second with 7x DOC (ca. 560 μM) using a HMW DOM concentrate (4592 μM DOC liquid sample; not freeze-dried) obtained by ultrafiltration. In addition, dilution-to-extinction samples were supplemented with 400 μM nitrate and 30 μM phosphate. For both the DCM and Mesopelagic dilution-to-extinction cultures, no DOC control treatments were arrayed into 15 cultivation plates (720 wells total) and 2.5x and 7x DOC treatments were arrayed into 30 plates each (1440 wells total). Non-inoculated media prepared with the same DOC concentrations were also arrayed into 5 plates each (240 well total). Plate samples were incubated for a period of 1-3 months at 22°C in darkness. To assess the effect of HMW DOM treatments on cell growth, wells were subsampled to enumerate SYBR Green I-stained cells on a GUAVA easyCyte 8HT (Millipore) flow cytometry system. Flow cytometry

data analysis was performed in R with flowCore and prada packages. Dilution-to-extinction wells with $>10^5$ cells ml^{-1} were considered positive hits.

Experiment II

Seawater inoculum samples for dilution-to-extinction were collected from the fluorometric DCM (95 m-depth) and from the Mesopelagic (250 m-depth) during the Hawaii Ocean Experiment - Budget of Energy I research cruise (KM 14-09) on March 23, 2014 (<http://hahana.soest.hawaii.edu/hoeboe/hoeboe.html>). In order to calculate the dilutions necessary to yield of 1-5 cells ml^{-1} the inoculum samples were analyzed by flow cytometry. Briefly, triplicate 1 ml sub-samples were fixed with 15 μl of 16% PFA for 10 minutes and stained with 5 μl of Sybr Green I (200-fold diluted stock) for 20 minutes in the dark. A total of 50 μl of each replicate were analyzed through an Influx cell sorter (BD Biosciences, San Jose, CA) to determine the total number of fluorescent events. Samples incubated onboard for 14 hours before reaching the laboratory where final dilution-to-extinction samples were prepared. The DCM inoculum was diluted to 3 cell ml^{-1} in a final volume of 6 liters of sterile seawater. Similarly, the mesopelagic inoculum was diluted to an estimated 2 cells ml^{-1} . Each 6-liter dilution-to-extinction sample was amended with inorganic phosphate (7 μM), nitrate (60 μM), and ammonium (4 μM). The 6-liter dilution-to-extinction samples were split into four 1-liter aliquots to receive the HMW DOM amendments (Table 3.1). Each dilution-to-extinction treatment was arrayed into 15 cultivation plates resulting in total of 720 wells. Each control or treatment media was also arrayed into 3 plates (144 wells). DCM samples were set to incubate at 26°C under a diurnal cycle and mesopelagic samples incubated at 25 °C in darkness for 10-11 weeks. To determine the effect of HMW DOM treatments on cell growth, wells were subsampled to enumerate SYBR Green I-stained cells on a GUAVA easyCyte PLUS (Millipore)

flow cytometry system. Flow cytometry data analysis was performed in R with flowCore and prada packages. Dilution-to-extinction wells with $>10^4$ cells ml^{-1} were considered positive hits.

Identification of cultures by whole genome shotgun sequencing

DNA purification and genome sequencing

The cultures identified in dilution-to-extinction samples were routinely maintained in the laboratory by passaging samples into sterile seawater. Sub-samples of selected cultures were transferred into 0.2 to 1.5 liters of sterile seawater to obtain sufficient genomic DNA for whole genome shotgun (WGS) sequencing. Culture growth was determined by enumerating total cells by flow cytometry. Upon reaching stationary phase cells were harvested by filtration onto 0.22 μm Supor membranes (Pall) and stored at -80°C in sucrose lysis buffer. Genomic DNA was purified using Qiagen's Blood and Tissue Kit (Experiment I) or through an in-house automated protocol using Proteinase K and Lysozyme additions using an MSM robot (Experiment II). Samples expected to have low DNA yield based on total cell counts were analyzed in a Fragment Analyzer (Advanced Analytical, Heidelberg, Germany) to determine sample quality. All samples of purified genomic DNA were quantified by the PicoGreen assay (Life Technologies). DNA concentrations were normalized before sequencing. DNA samples were prepared for sequencing using the Nextera XT 96 DNA sample preparation kit (Illumina, San Diego, CA, USA) to obtain indexed paired-end libraries. The libraries were sequenced with a 2x250 nt paired-end MiSeq run (Experiment I) and with a 2x300 nt paired-end MiSeq run (Experiment II).

Phylogenetic identification of cultures

FastQ files from the MiSeq runs were imported into the CLC Genomics Workbench (CLC bio, Aarhus, Denmark). Reads were assembled into contigs using CLC's de novo assembler with automatic word and bubble sizes, a minimum contig length of 200, insertion and

deletion costs set to 3, mismatch cost set to 2, length fraction set to 0.5 and the similarity fraction set to 0.8. Only contigs \geq 1000 nt in length were further examined. Small subunit (SSU) rRNA sequences were identified in assembled contigs and annotated by a LAST (Frith et al., 2010; Kielbasa et al., 2011) sequence similarity search against the SILVA non-redundant SSU database (version 119). The SSU rRNA gene sequences identified were aligned using the SILVA Incremental Aligner (SINA) version 1.2.11 (<http://www.arb-silva.de/aligner/>; Pruesse, Peplies, & Glockner, 2012) and imported into ARB (Ludwig et al., 2004) for alignment with the SILVA “All-Species Living Tree” (LTPs115) SSU reference tree (Munoz et al., 2011). A total of 240 sequences including closely related reference sequences from SILVA were selected for phylogenetic comparisons. The sequence alignment was filtered in ARB to exclude gaps, missing data ambiguous nucleotide positions, and lower case positions (nucleotides outside of an *E. coli* reference sequence). The filtered alignment contained 1172 comparable nucleotide positions across all sequences. Phylogenetic analysis was implemented using the Neighbor-Joining method (Saitou & Nei, 1987) with 1000-replicate bootstrap confidence limits (Felsenstein, 1985) in MEGA6 (Tamura et al., 2013). The optimal tree sum of branch length = 3.36318869. The evolutionary distances were computed using the Kimura 2-parameter method (Kimura, 1980) and are in the units of the number of base substitutions per site. The rate variation among sites was modeled with a gamma distribution (shape parameter = 2).

Assessment of the effect of HMW DOM treatments on the phylogeny of isolates recovered

To determine if the control and HMW DOM treatments had a notable effect on the phylogeny of isolates recovered in Experiment II, SSU rRNA genes were binned into phylo-types based on a 98% similarity threshold with BLASTClust (<http://toolkit.tuebingen.mpg.de/blastclust>). In addition, a chi-squared goodness of fit test was

applied to determine if the occurrence of phylo-types (irrespective of depth) was independent from the treatments (P -value <0.05). To evaluate the correlation between the occurrence of phylo-types and HMW DOM treatments a principal component analysis was performed in R statistical software using `prcomp` from the package `stats`.

DOM-dose response tests

Isolates were examined for higher growth yields under different HMW DOM treatments (increasing DOC concentrations). The isolates from Experiment I (cruise HOT 250) were tested in four different DOC treatments. The HMW DOM sample used as DOC supplement consisted of a 1-30 kDa fraction concentrated in seawater (4592 μM DOC) by tangential flow filtration. The seawater base media and control treatment consisted of UV-oxidized seawater with DOC reduced from 80 to 25 μM and amended with sodium nitrate (400 μM) and phosphoric acid (30 μM). Trace metals and vitamins were not added to the base media. The DOC treatments supplemented with HMW DOM concentrate were 3X DOC (230 μM), 15X DOC (1148 μM), and 30X DOC (2296 μM). The media was arrayed into 48-well plates and inoculated with 10 μl of culture maintained in HMW DOM media. Duplicate sets of cultures were prepared for each isolate. Samples were incubated at 22°C in darkness. Growth was determined by flow cytometry analysis of DNA-stained cells as described previously.

Selected isolates from Experiment II (cruise HOE-BOE I) were tested with seawater media with increasing HMW DOM concentrations. The control base media consisted of Hawaii sterile seawater from 250 m-depth amended with phosphoric acid (7 μM), sodium nitrate (60 μM), and ammonium chloride (4 μM). Trace metals and vitamins were not added to the base media. Five HMW DOM treatments were prepared with the base media, the first three (4X, 8X, 12X DOC) amended with freeze-dried HMW DOM, and two additional media (8X and 16X

DOC) amended with the freeze-dried HMW DOM sub-fraction F1 consisting of neutral polysaccharides. The 4X, 8X, 12X and 16X DOC treatments indicate the estimated DOC concentrations above the typical concentration in open-ocean Hawaii waters (~80 μM DOC) which resulted from the addition of 7.2, 14.4, 21.6, and 28.8 $\mu\text{g ml}^{-1}$ of HMW DOM sample to seawater, respectively. Samples were inoculated with <100 cells ml^{-1} of each isolate in triplicate and incubated at 25°C in darkness. Growth was determined by flow cytometry analysis of DNA-stained cells as described previously.

Carbon substrate utilization and growth analysis

To identify the types of carbon substrates Hawaii isolates could utilize for growth, representative isolates (HI0003, HI0035, HI0037, HI0049, HI0053, HI0054, HI0113, HI00120, and HI0133) were assayed for increases in cell yields in Biolog multi-carbon plates (PM1 and PM2A) containing a total of 190 different organic compounds. The exact formulation of the Biolog plates is proprietary but is similar to the recipes published by Bochner, Gadzinski, & Panomitros (2001) which utilize 25 mM carbon compound concentrations and 2 and 5 mM inorganic phosphorus and nitrogen concentrations, respectively. Because these nutrient concentrations are orders of magnitude greater than the concentrations found in open-ocean waters the dried contents of Biolog plates were re-suspended in 100 μl of sterile water and used a ~10,000X stock solutions to amend seawater media. The isolates tested were pre-grown in sterile Hawaii seawater amended with 1 μM phosphoric acid and ammonium chloride, as well as trace metals and vitamins (see Appendix I Table A3 for details on media supplements). Upon reaching stationary phase cultures were arrayed into plates with 1X-concentrated Biolog substrate. Samples incubated for 3-7 days at 26°C under a 12 hour diurnal light cycle. Cell growth was measured by detection of DNA-stained (SYBR Green I) cells on an Attune cytometer system

(Applied Biosystems, Foster City, CA, USA). To determine which Biolog substrates supported isolate growth, the end-point (2-4 day incubations) cell yields were analyzed using the Hampel (1974) outlier identifier method as modified by Leys, Ley, Klein, Bernard, & Licata (2013) which relies on the Median Absolute Deviation (MAD) to determine upper (median+2×MAD) and lower (median−2×MAD) confidence intervals of the data. MAD was implemented in R with function `mad` and default options. The Standard Boxplot Rule outlier detection method, which uses the first and third quartiles (Q1 and Q3) of the data to identify the upper (Q3+1.5×IQD) and lower (Q1−1.5×IQD) confidence intervals (where IQD is the interquartile difference) was implemented for comparison. Substrates deemed to support growth had cell yields greater than the upper MAD-based thresholds. The assay was also able to detect detrimental treatments (negative growth) indicated by cell yields less than the lower MAD-based threshold. In addition, the relative growth effect of each Biolog substrate was evaluated by normalizing the cell yield to the median across all treatments (190 substrates) and \log_2 -transforming this ratio to produce a doubling scale (\log_2 fold change). Positive growth agreed well with cell yields resulting in a \log_2 fold change > 1 and negative growth with \log_2 fold change < -1 . The median was selected to analyze the data and normalize cell yields because its value closely approximated the cell yields in negative controls (samples with no Biolog carbon substrate present; not used for analysis) as most treatments did not result in considerable growth.

Functional genomics and metabolic analysis

Genome draft assemblies and functional annotation

For a subset of isolates with high quality Illumina data, paired-end reads were first assembled using MIRA version 4.9.3 (Chevreux, Wetter, & Suhai, 1999) with options for removing remaining Illumina adapters, low quality stretches, phiX174 standard, as well as

common Illumina artifacts. This generated a very conservative assembly which was used to call genes. As part of the MIRA assembly process, a set of cleaned/clipped reads was also generated. This set of clean reads was used as input to SPAdes version 3.5 (Bankevich et al., 2012) with default parameters and explored with different k-mer settings. To annotate the assembled genomes (including rRNA and tRNA genes, and CDS), the assembled contigs were analyzed in PROKKA (Seemann, 2014) implemented through the Galaxy platform (Cuccuru et al., 2014). PROKKA-predicted CDS were further annotated by comparison to the NCBI RefSeq database (release 69) and the Kyoto Encyclopedia of Genes and Genomes (KEGG, release 67.11; Kanehisa & Goto, 2000) using the LAST algorithm (Frith et al., 2010; Kiełbasa et al., 2011).

Functional and metabolic analysis of DOM degradation capabilities

The conservative genome drafts of open-ocean isolates were examined to identify genes and biochemical pathways that have the potential to participate in the degradation of known HMW DOM constituents including carbohydrates and phosphonates. The analysis included the genome of *Pseudomonas stutzeri* HI00D01 isolate from Experiment I, 55 genome assemblies obtained from Experiment II isolates, and the genome of SAR92 clade coastal isolate NB0015 (described in Chapter 2), for comparison. In addition, to identify metabolic and functional differences among genomes, PROKKA-predicted protein sequences of 28 isolates including *P. stutzeri* HI00D01 and the coastal strain NB0015 were analyzed on the KEGG Automatic Annotation Server (<http://www.genome.jp/tools/kaas/>) using the BBH (bi-directional best hit) method to assign KEGG orthologs (KO) and generate KEGG pathways. The representation of KEGG pathways and predicted KOs were compared between closely related isolates and across all isolates tested.

To assess the potential for the degradation of carbohydrates – a major component of HMW DOM – PROKKA-predicted protein sequences of all available genomes were compared to the Carbohydrate Active Enzyme (CAZy) database (Cantarel et al., 2009) using HMMER3.0 (Eddy, 2009) and hidden Markov models of CAZy signature domains (Yin et al., 2012). Differences in the average number of predicted CAZy family genes were assessed between DCM and Mesopelagic genomes (Mann-Whitney U test) and among phylogenetic clusters (one-way ANOVA). Post hoc analysis via Tukey's honest significance difference test was employed to identify phylogenetic groups in which the genomes contained significantly different number of predicted CAZy genes.

In addition, the C-P lyase pathway (Metcalf et al., 1993a, 1993b) was specifically targeted in the genome functional analysis because of the prevalence of phosphonates in marine HMW DOM (Kolowitz et al., 2001). To determine if the C-P lyase pathway (Metcalf & Wanner, 1993a, 1993b) was encoded in the genomes of open-ocean isolates a BLASTp analysis (Camacho et al., 2009) was implemented in Galaxy (Cock, Grüning, Paszkiewicz, & Pritchard, 2013) using as queries the nucleotide sequences of phosphonate degradation operons identified in several uncultivated marine bacteria enriched in phosphonate addition experiments to open-ocean surface seawater (Martínez, Ventouras, Wilson, Karl, & DeLong, 2013).

Results and Discussion

This chapter explores the hypothesis that, as was observed in a coastal ecosystem microbial community, the open-ocean harbors bacterial groups that contribute to the breakdown of HMW DOM. To test this hypothesis two cultivation experiments were prepared aimed to stimulate the growth of such bacterial groups by enriching seawater with HMW DOM above the

ambient DOC concentrations (~60-80 μM ; Fig. 3.1). In the following section I describe the effect that HMW DOM treatments had on the growth and number of cultures recovered from dilution-to-extinction samples. Later I present the phylogenetic classification of the isolates obtained and discuss the unexpected low cell yields of isolates grown in media supplemented with HMW DOM relative to unamended controls. Finally, I present some of the isolates' physiological and metabolic characteristics identified through growth experiments and genome analysis and discuss their potential implication in degrading specific constituents of HMW DOM.

Effect of HMW DOM additions on the recovery and cell yields of dilution-to-extinction cultures

Experiment 1

The first dilution-to-extinction experiment was prepared with seawater collected near Hawaii at Station ALOHA and transported to MIT for laboratory incubations. Because a prolonged handling time of seawater samples may alter the viability and composition of the microbial community, as a precaution the inoculum samples were diluted to extinction before transporting them to MIT. The cell concentration of inoculum seawater samples were $197,052.2 \pm 46,677.4$ cell ml^{-1} and $45,161.7 \pm 2,303.7$ cell ml^{-1} (mean \pm s.d., $n=3$) for DCM and Mesopelagic samples, respectively. The estimated final cell densities for dilution-to-extinction samples were 2-4 cells ml^{-1} . Inspection of dilution-to-extinction samples by epi-fluorescence microscopy did not reveal a drastic increase in cell numbers after transport, though small changes in cell density cannot be ruled given that such low cell numbers cannot be quantified accurately even by flow cytometry. Because the samples seemed unaltered they were deemed appropriate for cultivation. The complete flow cytometry growth screen results are presented in Appendix I Table A1.

For Mesopelagic samples, the 7X DOC and No DOC treatments resulted in >50% of wells populated with bacterial cultures $>10^5$ cell ml⁻¹ (positive hits), exceeding the number of expected cultures, typically 4-5% (Fig. 3.2). In contrast, the fraction of wells with grown cultures (8%) in the 2.5X DOC treatments was closer to the expected (Fig. 3.2). Dilution-to-extinction plates prepared with seawater samples from the DCM did not have as strong a response to HMW DOM treatments as the Mesopelagic samples (<1.2% positive hits, Fig. 3.2). In addition, the flow cytometry screen revealed a small fraction of sterile media control wells with positive hits (<1%).

In order to select for isolates with the strongest growth response to HMW DOM and to eliminate potential non-HMW DOM consuming microorganisms, a subset of five Mesopelagic plates from each treatment were serially diluted (1:100) into sterile media (final 10⁸-fold dilution) and incubated for 1 month (Fig. 3.1). A total of 71 of 210 wells recovered cultivars on either the last or second-to-last step of the dilution series. After passaging these cultures into sterile media, 53 (91%) Mesopelagic isolates from the 7X DOC treatment, 6 from the 2.5X DOC treatment, and 7 from the No DOC treatment were recovered.

Because the DCM dilution-to-extinction samples did not show an equally robust growth-response to HMW DOM treatments as the Mesopelagic samples did after one month of incubation, the plates were left incubating for an additional 1.5 months (Fig. 3.1). Re-screening of the DCM plates after the extended incubation period revealed wells with positive hits, cultures that reached cell densities $>10^5$ cell ml⁻¹ (Fig. 3.3). Samples from the No DOC treatment produced 18 (2.85%) new cultures and samples from the 2.5X DOC treatment resulted in 133 (10%) positive hits, the highest number of cultures recovered (Fig. 3.3). In contrast, plates with 7X DOC treatments did not produce any positive hits (Fig.3.3). To verify their viability, cultures

considered positive hits were serially diluted into sterile DOM media and incubated for 2 months under the same conditions (Fig. 3.1). Only 17 of 18 cultures from the No DOC treatment were recovered. Cultures from 2.5X DOC and 7X DOC treatments were not viable or recoverable. In total, 83 isolates were recovered from Experiment I: 66 from Mesopelagic samples and 17 from the DCM. These cultures were identified by whole genome shotgun sequencing (Table 3.2).

Experiment II

The cell concentrations of inoculum seawater were $498,653 \pm 14,536 \text{ ml}^{-1}$ and $108,235 \pm 2,626$ (mean \pm s.d., $n=3$) for DCM and Mesopelagic samples, respectively. These cell concentrations fell within the typical ranges observed at Station ALOHA (Campbell et al., 1997). The complete flow cytometry growth screen results are presented in Table A2 (Appendix I). Wells with positive hits ($>10^4 \text{ cell ml}^{-1}$) were detected across all samples including non-DOM amended controls. Flow cytometry failed to detect positive hits in sterile media preparations. Sterile controls only registered wells with $<5,000 \text{ events ml}^{-1}$, 98.4% of which had $<1,000 \text{ events ml}^{-1}$. Notably, HMW DOM supplements had a significant effect in the number of positive wells detected in Mesopelagic samples relative to controls (Fig. 3.4; $P < 0.05$, bootstrap comparisons with 10^4 iterations). The percentage of wells with positive hits in HMW DOM treated Mesopelagic samples was 1.67-5.56% compared to 0.69% in unamended controls (Appendix I Table A2). In contrast, DCM samples seemed to be unaffected by DOM additions (Fig. 3.4; $P < 0.05$, bootstrap comparisons with 10^4 iterations). To assess the viability of cultures identified in wells identified as positive hits, sub-samples ($\sim 0.5 \text{ ml}$) from each well were transferred into sterile seawater media (5-10 ml) with HMW DOM and incubated under the same conditions. After passaging cultures twice in sterile HMW DOM seawater, 16 cultures (40% of positive hits) from the DCM and 58 cultures (61.7% of positive hits) from the Mesopelagic were able to re-

grow (Fig. 3.4). In total, 74 cultures (55.2% of all positive hits) were deemed amenable to sub-culturing and were grown in large volume to obtain genomic DNA for whole genome shotgun sequencing and identification (Table 3.2).

Phylogenetic placement of cultures and patterns of bacterial groups enriched in HMW DOM treatments

WGS sequencing produced genomic data for a total of 157 cultures from which complete or partial SSU rRNA gene sequences were obtained (Table 3.2). The HMW DOM-treated open-ocean dilution-to-extinction samples resulted in the isolation of diverse bacterial groups within the *Alphaproteobacteria* and *Gammaproteobacteria*. The cultures were phylogenetically placed based on the SSU rRNA percent sequence similarity to closest neighbors in the SILVA database (Table 3.2; see also Fig. A1 and A2 in Appendix I). Based on the identity of the SSU rRNA gene sequences obtained from assembled Illumina reads all cultures consisted of a single isolate and there was no indication of mixed populations (Table 3.2; see also Appendix I Table A9 for a list of PROKKA gene identifications of SSU rRNA gene sequences of 60 draft genomes available under NCBI BioProject PRJNA305749). In Experiment I, single taxon was recovered from each DOC treatment. The 7X DOC treatment in Mesopelagic samples was dominated by the gamma-proteobacterium *Pseudomonas stutzeri* while the 2.5X DOC treatment selected for *Alphaproteobacteria* of the genus *Erythrobacter* and the No DOC treatment produced *Vibrio* isolates. The No DOC treatment of DCM samples recovered *Halioglobus* isolates (order *Cellvibrionales*). Viable cultures were not recovered from the DCM 2.5X DOC and 7X DOC treatments even though the flow cytometry screen detected wells with growth in those samples. In Experiment II, the types of isolates obtained in the DCM were similar to the Mesopelagic samples and could be grouped into 12 phylo-types based on 98% rRNA gene sequence similarity

clustering. These phylo-types corresponded well with the Neighbor-Joining phylogenetic tree branches (Appendix I Fig. A1 and A2). All *Alcanivorax* isolates were grouped into the phylo-type (>98% similarity) though DCM and Mesopelagic *Alcanivorax* isolates clustered separately in Neighbor-Joining trees due to differences in nucleotide the SSU rRNA gene sequence (Appendix I, Fig. A1). Similarly, some *Erythrobacter* SSU rRNA gene sequences were unique to the DCM while other sequences occurred both in DCM and Mesopelagic samples (Table 3.2). The relationship between phylo-types and DOC treatments was evaluated using principal component analysis (Fig. 3.4B). Notably, in Mesopelagic samples 64% of isolates were close relatives of *Oleiphilus messinensis* (*Oleiphilus* 1 and 2 phylo-types, Appendix I Fig. A1) and dominated the 10X DOC and 4X DOC F1 treatments (Fig. 3.4). The *Oceanibulbus/Sulfitobacter* phylo-type was also strongly associated with the 10X DOC treatment and with the 4X DOC treatment (Fig. 3.4B). Some taxa were unique or strongly associated with a particular DOC treatment, like the *Alteromonas* isolates (4 total) which only occurred in the 4X DOC F1 treatments and the *Erythrobacter* 1 phylo-type which was associated with the 10X DOC treatment (Fig. 3.4). In contrast, the *Alcanivorax* phylo-type occurred across all DOC treatments, including No DOC treatments, which suggests these isolates may have grown independently of HMW DOM additions. The remaining taxa were present in at least one DOC treatment and were absent from the No DOC controls though the total number of isolates recovered was too low to infer a strong association to a particular DOC treatment (Fig. 3.4).

DOM dose-response of selected Hawaii isolates

Experiment I isolates

Among the Mesopelagic isolates, the majority of *Pseudomonas stutzeri* cultures obtained from 7X DOC treatments increased their cell yields correspondingly with DOC concentration

(Fig. 3.5). The *Erythrobacter* cultures obtained from 2.5X DOC treatments also increased cell yields with DOC additions although not proportionally and seemed to have a maximum growth despite the higher DOC treatments (Fig. 3.5). In turn, the *Vibrio* isolates did not appear to be stimulated by HMW DOM additions, consistent with their isolation in No DOC controls samples (Fig. 3.6). Unexpectedly, the DCM *Halioglobus* cultures obtained from the No DOC treatment exhibited increased cell yields correspondingly with DOC additions (Fig. 3.6). The *Halioglobus* isolates were not detected by flow cytometry originally after the first month of incubation. It is possible that the prolonged incubation period allowed them to utilize semi-labile DOC although there growth might be induced by evaporation in the wells and the resulting concentration of salts and organic matter. Overall in this experiment, the growth profiles of closely related isolates exposed to HMW DOM was similar and consistent with the DOC treatment from which they were obtained. However, proportional increases in cell yields with DOC added were difficult to detect.

Experiment II isolates

The HMW DOM growth response of isolates HI0053 (*Pseudoalteromonas*), HI0054 (*Sulfitobacter*), HI0077 (*Erythrobacter*), HI0090 (*Alteromonas*), HI0096 (*Alcanivorax*), and HI0113 (*Oleibacter*) is presented in Fig. 3.7. These cultures did not show a significant growth response to DOC additions despite their isolation in HMW DOM treatments. For example, the Mesopelagic *Alteromonas* isolates (3 were tested including HI0090) obtained from the 4X DOC F1 treatment (HMW neutral polysaccharide fraction) did not to have a discernible growth response to increasing concentrations of HMW DOM despite the addition of up to 12X and 16X DOC (~960 and ~1280 μM , respectively). Instead, the cell yields in unamended media (No DOC) were slightly higher than the growth in media with DOC additions suggesting that the

addition of HMW DOM may be detrimental to cells. Similarly, the Mesopelagic *Alcanivorax* isolate HI0096 obtained from the 4X DOC F1 treatment observed higher growth yields in media without HMW DOM than in media with HMW DOM. On the other hand, some isolates including the HI0053, HI0063, HI0077, and HI0113 showed slightly higher cell yields in media supplemented with HMW DOM, though the growth difference with respect to control media was not statistically significant over the course of the incubation period (Fig. 3.7).

Carbon substrate utilization profiles of representative Hawaii isolates

The Biolog multi-carbon substrate assay was used to determine the growth requirements of 9 representative bacterial isolates obtained from Experiment II. The cell yields of inoculum cultures pre-grown in Hawaii seawater with inorganic nutrients, trace metals, and vitamins ranged from 1×10^4 to 1×10^5 cell ml⁻¹. Hawaii isolates closely related to known hydrocarbon degraders were compared for differential substrate utilization (Fig. 3.8). The *Alcanivorax* HI0003 and HI0035 were able to grow on capric acid and caproic acid (fatty acids), and sebacic acid (dicarboxylic acid) as well as on low-molecular-weight compounds like propionate, pyruvate, and methyl pyruvate (\log_2 fold change >1). Another isolate, *Oleibacter* HI0113, metabolized Tween 80 (polysorbate 80), capric acid, caproic acid, and the polysaccharides mannan (\log_2 fold-change >1), laminarin and pectin (\log_2 fold change ~0.9). Isolate *Thalassolituus* HI0120 observed similar carbon substrate preferences as HI0113. Growth was stimulated in the presence of Tween 80, L-proline, and capric acid (>2 \log_2 fold-change) as well as Tween 40, itaconic acid, L-homoserine, dextrin, acetic acid (>1 \log_2 fold-change), and Tween 20 (\log_2 fold change = 0.9). Finally, *Oleiphilus* HI0133 was stimulated by Tween compounds (polysorbate 20, 40, and 80), acetic acid, and the trisaccharide D-raffinose (\log_2 fold change >1; Fig. 3.8). Interestingly, the *Erythrobacter* isolate HI0037 was also strongly stimulated by Tween compounds (\log_2 fold

change > 2). In addition, HI0037 degraded two polysaccharides, laminarin (\log_2 fold change > 2) and pectin, and the polymer gelatin (\log_2 fold change >1), a derivative of collagen.

Two *Rhodobacteraceae* isolates, HI0049 (*Roseovarius*) and HI0054 (*Sulfitobacter*), were also tested (Fig. 3.9). HI0054 showed preferential utilization of α -amino acids (L-leucine, L-isoleucine, L-phenylalanine, L-methionine, L-proline, L-threonine, D-alanine), non-proteinogenic amino acids (γ -amino butyric acid, N-acetyl-L-glutamic acid, hydroxyl-L-proline, glycy-L-proline, glycy-L-aspartic acid, glycy-L-glutamic acid, δ -amino valeric acid), degradation products of amino acids (L-ornithine and putrescine), and other low-molecular-weight compounds including β -hydroxy butyric acid, α -keto-valeric acid, sebacic acid, thymidine, and D-lactic acid methyl ester (\log_2 fold-change >2). Similarly, isolate HI0049 had a strong preference for amino acids including L-valine, L-ornithine, L-proline, L-leucine, D-alanine, L-lysine, γ -amino butyric acid, δ -amino valeric acid, L-pyroglutamic acid, L-aspartic acid, and L-isoleucine (\log_2 fold-change >1). Carbohydrate substrates did not stimulate the growth of these *Rhodobacteraceae* isolates, with the exception of β -cyclodextrin to which HI0054 was responsive.

In contrast to the seemingly specialized metabolism of the hydrocarbon degraders and the *Rhodobacteraceae* isolates, the *Pseudoalteromonas* isolate HI0053 could be described as having a generalist metabolism. In total, HI0053 was stimulated positively by 44 substrates including carbohydrates and nucleosides such as maltose, D-trehalose, D-cellobiose, sucrose, maltotriose, laminarin, α -cyclodextrin, β -cyclodextrin, inosine, γ -cyclodextrin, adenosine, and 2-deoxy adenosine, and the amino acids L-proline and L-glutamic acid (\log_2 fold-change >3). In addition, HI0053 could metabolize N-acetyl-D-galactosamine, methyl pyruvate, L-arginine, acetic acid,

formic acid, D-mellobiose, propionic acid, succinic acid, citric acid, D-mannitol, uridine, and capric acid (\log_2 fold-change >2).

In addition to detecting growth substrates, the Biolog screen was able to detect potential toxic or detrimental compounds as observed by a relative decrease in cell numbers over the incubation period relative to the median value across all treatments. The *Alcanivorax* HI0003 seemed to be particularly sensitive to glycyl-L-glutamic acid (\log_2 fold change = -6) while *Alcanivorax* HI0035 was sensitive to a different set of compounds including D-galactose, succinic acid, D-saccharic acid, L-arabinose, N-acetyl-D-glucosamine, L-aspartic acid, D-alanine, D-mannose, D-trehalose, D-mannose, D-sorbitol, and α -ketobutyric acid (\log_2 fold change < -1). *Thalassolituus* HI0120 cell yields decreased in the presence of L-threonine and L-arginine (\log_2 fold change <-2). *Oleiphilus* HI0133 cells were also sensitive to amino acids including L-homoserine, L-histidine, and L-threonine (\log_2 fold change <-3). In contrast, *Oleibacter* HI0113 appeared to be uninhibited by Biolog substrates. The *Rhodobacteraceae* isolates were sensitive to different substrates. *Roseovarius* HI0049 was sensitive to gelatin, itaconic acid, α -hydroxybutyric acid, and D-tartaric acid (\log_2 fold change <-1). Similarly, cell yields of *Sulfitobacter* HI0054 cultured decreased in the presence of α -hydroxybutyric acid and 2-deoxy-D-ribose (\log_2 fold change <-1) but these were not deemed significant by the outlier test.

It is worth noting the polarized effect of substrates like amino acids serving, on the one hand, as growth substrates for bacterial groups like the *Rhodobacteraceae* and on the other, being potentially detrimental compounds for cultures of hydrocarbon degraders like *Oleiphilus* and *Thalassolituus*. It is possible that in the environment the dual role of such substrates helps

control the abundance or activity of bacterioplankton populations or that in turn, specific bacterial groups enable the growth of other taxa by clearing toxic compounds.

Genome draft assemblies

In total, 60 draft genomes were obtained from isolates, including one for the coastal isolate NB0015 (SAR92 clade), 4 for Hawaii open-ocean isolates from Experiment I, and 55 for Hawaii open-ocean isolates from Experiment II. The results of the assemblies are summarized in Appendix I Table A4. The genome coverage for Experiment II isolates was 24 ± 8 (mean \pm s.d.) Four representative genome drafts were obtained from Experiment I isolates including the DCM isolate *Halioglobus*, and the Mesopelagic *Vibrio*, *Erythrobacter*, and *Pseudomonas stutzeri* isolates. The raw reads obtained from these groups were pooled for assembly to obtain better coverage. The genome coverage for these assemblies was higher, 214 ± 102 (mean \pm s.d.), than in Experiment II. The highest N50 value obtained was for *Halioglobus* isolate HI00S01 (342,731 nucleotides) and the lowest was for *Oleiphilus* isolate HI0085 (2,070 nucleotides). The GC content was lowest in *Oleiphilus* isolates, ranging from 44-48%, and highest for the *Alphaprotobacteria* isolates, between 61-64%. The combined assemblies included 54,103,639 Illumina reads. A summary of the gene annotations obtained through analysis with PROKKA as well as LAST analysis against KEGG and RefSeq are summarized in Appendix Table A5 including the number of CDSs with hypothetical functions and number of SSU and LSU genes. The draft genome assemblies with annotations are available under NCBI BioProject PRJNA305749.

Genome functional and metabolic analysis

The analysis of draft genome assemblies of several Hawaii isolates revealed examples of HMW DOM degradation strategies consistent with several known chemical constituents of

HMW DOM, including polysaccharides and phosphonates. This analysis also provided evidence of the chemical class of carbon substrates preferentially used for growth by the isolates that could be compared to the Biolog multi-carbon growth assays. PROKKA CDSs were assigned to KEGG pathways, compared to the carbohydrate active enzymes (CAZy) database (Cantarel et al., 2009; <http://www.cazy.org/>), and aligned against reference protein sequences comprising the phosphonate degradation pathway.

Enrichment of KEGG pathways

The KEGG ortholog (KO) assignments of PROKKA-predicted protein sequences served to assess the enrichment of KEGG pathways among isolates. A total of 28 isolates, 16 from the DCM, 11 from Mesopelagic depths, including *P. stutzeri* HI00D01, and one coastal isolate, NB0015 were compared. These isolates represented all the Hawaii phylo-types identified by clustering, except the *Alteromonas* for which conservative draft genomes were not obtained. Several of the isolates compared were also tested with the Biolog multi-carbon plates. The enrichment of KEGG pathways varied among different bacterial groups (KO count mean \pm s.d., n=28) of which ABC transporters (58.5 ± 32.0), two-component systems (64.1 ± 13.3), flagellar assembly (19.4 ± 11.2), and lipopolysaccharide biosynthesis (12.8 ± 7.7) were the most differentially represented. For example, the *Rhodobacteracea* (*Roseovarius*, *Oceanibulbus*, and *Sulfitobacter* isolates) had the highest number of ABC transporters (107.3 ± 4.5 ; n=6) while the *Pseudoalteromonas* HI0053 and *P. stutzeri* HI00D01 had the most two-component system proteins (92 and 104, respectively). Two-component systems are signal transduction pathways that regulate the physiology of the cell in response to changes in the environment or within the cell. For example, the PhoR/PhoB regulon is a two-component system protein in which PhoR, a membrane-associated protein kinase, phosphorylates PhoB in response to changes in phosphate

levels (McGrath, Chin, & Quinn, 2013). This triggers a signal cascade that regulates gene expression of genes associated with phosphorus uptake and metabolism. The genome of *Pelagibacter ubique*, strain HTCC1062, a member of the oligotrophic SAR11 clade, contains only 4 two-component systems three of which are predicted to help regulate nutrient and osmotic stress (Giovannoni et al., 2005) indicating of the importance of these physiological parameters for bacteria. In contrast, the so-called opportunistic bacteria like *Marinobacter aquaeleoi* and several *Pseudomonas* species (Singer et al., 2011), as well as copiotrophic bacteria (Lauro et al., 2009) are enriched (~1% of the CDSs) in two-component regulators and signal transduction pathways reflecting the ability of these bacteria to quickly respond to environmental stimuli like pulses of nutrients. All isolates except *Alcanivorax* HI0003 and *Roseovarius* HI0049 encoded a complete flagellar apparatus (flagellar assembly KEGG pathway). On the other hand, all the *Erythrobacter* isolates analyzed lacked lipopolysaccharide biosynthesis genes as is expected of the *Sphingomonadales* which replace them with glycosphingolipids (Kawasaki et al., 1994). In some cases, the presence of genes encoding ABC transporters or degradation pathways associated with specific chemical classes of compounds could be compared to the Biolog carbon substrate utilization profiles of isolates. For example, The KEGG pathways of two different oil-degrading isolates, *Oleiphilus* HI0133 and the *Alcanivorax* HI0003, were compared. Biolog multi-carbon assays confirmed that these isolates preferentially used lipid-containing compounds for growth (Fig. 3.8). The *Alcanivorax* HI0003 could only metabolize simple fatty acids like capric (decanoic) and caproic (hexanoic acid), and sebacic (decanedioic) acids while *Oleiphilus* HI0133 preferred Tween compounds (polysorbate 20, 40, and 80) which in addition to the hydrocarbon tail contain a hydrophilic polyoxyethylene group. Interestingly, the genome of HI0133 encoded 4 ABC transporter sub-unit homologs (*mlacDEF*) associated with

phospholipids, molecules that resemble the Tween compounds, while HI0003 only contained a copy of *miaC*. In addition, HI0133 contained ABC transporters for various substrates including L-amino acids, proline, putrescine, and osmoprotectants though significant growth on these substrates was not observed. Another distinguishing feature of the *Oleiphilus* HI0133 genome was that it encoded a complete flagellar apparatus and several bacterial chemotaxis genes, 9 more than HI0003. In comparing the KEGG pathways of four closely related *Oleiphilus* isolates (HI0067, HI00118, HI0132, and HI0133) the most differences occurred in the number of two-component system proteins, ABC transporters, porphyrin and chlorophyll metabolism, nitrogen metabolism, and phosphonate/phosphinate metabolism. One distinguishing feature of *Oleiphilus* HI0067 was that its genome encoded complete ABC transporters for alkylphosphonates (Metcalf & Wanner, 1993a, 1993b), and several single sub-units of ABC transporters associated with oligosaccharides, and one complete ribose/D-xylose ABC transporter while the other *Oleiphilus* genomes did not. *Oleiphilus* HI0133 was able to grow on D-raffinose, a trisaccharide, but no oligosaccharide or monosaccharide ABC transporters were detected by the KEGG analysis. On the other hand, the NarX, NarL two-component system proteins and the NarG, NarH, NarJ, NarI subunits of nitrate reductase were only present in the genome of HI0132.

Two closely related *Rhodobacteraceae* isolates were also compared, *Roseovarius* HI0049 and *Sulfitobacter* HI0054. These isolates showed preferential utilization of amino acids for growth. In accordance with this result, HI0049 and HI0054 had 16 and 17 predicted ABC transporters sub-unit homologs for amino acids and 5 and 10 for oligopeptides and dipeptides, respectively. For comparison, *Oleiphilus* HI0133 and *P. stutzeri* HI00D01 had 13 amino acid ABC transporters sub-units each but lacked ABC transporters for the oligopeptides and

dipeptides. Differences in motility capabilities were also evident. *Sulfitobacter* HI0054, with 23 flagellar assembly genes, seems to be motile while *Roseovarius* HI0049, with only 1 predicted gene is limited in this respect, though it may use other forms of motility. In addition to the amino acid ABC transporters, *Roseovarius* HI0049 and HI0054 were enriched in genes involved in the degradation of aromatic compounds (25 and 10, respectively), specifically in benzoate and catechol oxidation. Isolate *P. stutzeri* HI00D01, with 14 genes, was also enriched in this KEGG pathway though the ability of these isolates to metabolize aromatic compounds remains to be tested since this chemical class was not well represented in Biolog multi-carbon plates.

Carbohydrate degradation

Significant differences in the number of predicted CAZy genes were observed between the genomes of DCM (16) and Mesopelagic (39) cultures (Mann-Whitney U test, $H_0: \bar{X}=\bar{Y}$, P-value<0.05; Fig. 3.10). Mesopelagic genomes harbored more predicted carbohydrate-binding module genes (CBM, P-value=0.01998) and fewer auxiliary activities (AA, P-value=0.001129) genes than the DCM genomes (mean \pm s.d.; 3.1 ± 2.1 vs. 1.6 ± 1.6 CBM genes and 15.6 ± 5.3 vs. 23.2 ± 8.3 AA genes, respectively). The CBM genes recognize and bind carbohydrates and require associated enzymes such as glycoside hydrolases to cleave substrates (Cantarel et al., 2009). Several CBM families were present in Mesopelagic genomes but absent in the DCM genomes. These included CBM6, CBM23, CBM35, and CBM61 which characteristically bind carbohydrates substrates such as mannan, xylan, cellulose, glucans, and galactans. Among the Mesopelagic isolates the *Pseudolateromonas* HI0053 had the highest number of CBM genes 8 different predicted proteins (Fig. 3.10). Close inspection of the *Pseudolateromonas* HI0053 CBM genes revealed some characteristic features of secreted carbohydrase systems used to degrade polysaccharides. For example, one CBM protein annotated as a beta-xylosidase/alpha-L-

arabinofuranosidase (67.6% identity to *Cellvibrio japonicus* strain Ueda107) was clustered with genes encoding an alpha-glucosidase and a metal-dependent hydrolase, an ABC-transporter, enzymes involved in protein secretion, namely the sec-independent Twin-Arginine Translocation (TAT) system, and two GGDEF domain-containing proteins that likely regulate cyclic di-guanylate (Ausmees et al., 2001), a secondary messenger involved in signal transduction. All of these proteins likely make up a functional carbohydrase system as observed in the complex carbohydrate-degrading bacterium *Saccharophagus degradans* 2-40 (Weiner et al., 2008) For example, a carbohydrase system needs to be secreted to the cell wall, a function that can be mediated by the TAT system, and requires sensor proteins at the cell surface like the GGDEF domain-containing proteins to detect the presence of carbohydrate substrates and convey the information to secondary messengers like cyclic di-guanylate, which in turn help regulate the expression of the carbohydrase genes. Isolate HI0053 also contained CBM genes predicted to encode a chitinase, a glycogen branching enzyme, a collagen-lytic enzyme or a serine protease, a secreted glycosyl hydrolase, a putative hemolysin-type calcium-binding protein, and two hypothetical proteins. A one-way ANOVA was conducted to assess differences in CAZy gene class representation between phylogenetic groups (Appendix I Table A6). Significant differences in CAZy-predicted genes between phylogenetic groups were identified for all enzyme classes except for the polysaccharide lyases (Appendix I Table A6, Fig. 3.10). A Tukey's honest significance difference (HSD) post hoc test was applied to determine differences in the mean number of CAZy genes for each enzyme class among phylogenetic groups (P-value<0.05). The analysis revealed that, in addition to the Pseudoalteromonad HI0053, The *Oleiphilus* 1 and *Erythrobacter* 2 clusters also had a higher number of CBM genes compared to other clusters (Fig. 3.10). In particular, two *Oleiphilus* strains, HI0071 and HI0118, each with 7 CBM genes were

enriched. It was unclear though if the CBM predicted genes of *Oleiphilus* isolates encoded secreted carbohydrases as most hits lacked a functional annotation and adjacent genes did not fit the description of the additional proteins required for polysaccharide degradation.

CAZy Auxiliary Activities (AA) class consists of genes encoding redox enzymes that act in conjunction with other CAZy enzymes and include ligninolytic enzymes and lytic polysaccharide mono-oxygenases (Cantarel et al., 2009). A one-way ANOVA was conducted to compare the number of predicted CAZy AA genes among phylo-types (Appendix I Table A6). Post hoc tests indicated that the *Alcanivorax* and *Oceanibulbus/Sulfitobacter* phylo-types harbored a significantly different number of AA genes compared to other bacterial groups (Fig. 3.10, Tukey's HSD, P-value<0.05). Of all isolates though, *Roseovarius* HI0049 had the highest number of AA genes (Fig. 3.10). In addition, the AA1 and AA3 families were enriched in DCM genomes (Mann-Whitney U test, P-value<0.05) compared to Mesopelagic genomes. The AA1 enzymes are multicopper oxidases that use diphenols and related substances as donors with oxygen as the acceptor while AA3 enzymes are flavoproteins that belong to the glucose-methanol-choline (GMC) oxidoreductase family. Among the DCM isolates, *Erythrobacter* HI0037 and HI0038 contained the highest number of AA1 predicted genes (9 and 8, respectively) while the *Alcanivorax* HI0035 contained the highest number (24) of AA3 predicted genes.

No significant differences were observed on the mean number of homologs predicted as carbohydrate esterases (CE), glycoside hydrolases (GH), and glycosyl transferases (GT) or polysaccharides lyases (PL) enzyme classes between DCM and Mesopelagic genomes (Mann-Whitney U test, P-value<0.05; Fig. 3.10). However, within these CAZy enzyme classes several families were differentially enriched in DCM and Mesopelagic genomes (Mann-Whitney U test,

P-value<0.01). For example, among GH families, GH73 (lysozymes and peptidoglycan hydrolases) homologs were enriched in Mesopelagic genomes as it was present in all *Oleiphilus* isolates, while GH107 (sulfated fucan endo-1,4-fucanase) and GH114 (endo- α -1,4-polygalactosaminidase) were only present in *Alcanivorax* isolates, and GH28 (galacturonase) in DCM *Erythrobacter* isolates. Similarly, among the CE families CE4 (acetyl xylan esterases and chitin deacetylases) and CE16 (carbohydrate acetyl esterases) were most prevalent in Mesopelagic genomes. In particular, three *Oleiphilus* isolates (HI0065, HI0071, and HI0133) contained 10 predicted CE4 genes. These acetyl esterases may be involved in the degradation of acetyl ester moieties in HMW DOM polysaccharides – up to 7% of the total carbon based on ^1H NMR (Aluwihare et al., 1997). In contrast, the CE7 (acetyl xylan esterases and cephalosporin-C deacetylases) and CE10 (aryl and carboxyl esterases and cholinesterases) families were most prevalent in the *Erythrobacter* and *Alcanivorax* genomes isolated from the DCM. Finally, the polysaccharide lyase (PL) family PL9 comprised of pectate and thiopeptidoglycan lyases was absent from DCM genomes and only occurred in two *Alcanivorax* and twelve *Oleiphilus* Mesopelagic isolates.

Phosphonate degradation pathway

Previously, Martínez et al. (2013) identified three phosphonate degradation gene clusters expressed in open-ocean microcosms amended with free methyl phosphonate and glucose. These clusters were related to sequences found in *Vibrio nigripulchritudo* ATCC27043, the gammaproteobacterium IMCC1989, and a *Rhodobacterales*, respectively. In addition, the isolate *Pseudomonas stutzeri* strain HI0D01 harbors a complete phosphonate degradation pathway with a similar arrangement as the *E. coli* and *V. nigripulchritudo* operons (Appendix I Table A8). The presence of phosphonate degradation genes in other isolates obtained from HMW DOM

enrichments may be an indication of their ability to partially degrade HMW as the phosphonates are thought to be bound to polysaccharides. Details of the *P. stutzeri* HI0D01 phosphonate-degrading capabilities are presented in Chapter 4.

Surveying the RefSeq functional assignments of PROKKA-predicted genes revealed an enrichment of phosphonate degradation genes in 10 isolates closely related to members of the *Rhodobacteraceae* and *Oleiphilaceae* (Appendix I Table A8). With the exception of one isolate (*Oleiphilus* HI0009), all of these isolates were obtained from HMW DOM-treated dilution-to-extinction samples (4x DOC, 10x DOC, and 4x DOC F1) from Hawaii Experiment II. To confirm this observation, the protein sequences comprising the phosphonate degradation operons B3TF_MPn_1, B3TF_MPn_2, and B3TF_MPn_8 identified by Martinez et al. (2013) were used as queries (Appendix Table A7) in a BLASTp analysis against the PROKKA-predicted protein sequences of the open-ocean isolates. The PROKKA-predicted CDSs matching C-P lyase genes of Hawaii isolates are listed in Appendix I Table A8. The *Roseovarius* strain HI0049 harbored all the genes present in the *Rhodobacteraceae* operon B3TF_MPn_2. One contig was composed of 7 genes: a tetR family transcriptional regulator, *phnC*, *D*, *E1*, and *E2* encoding ABC-type phosphate/phosphonate transport system proteins, a putative acetyltransferase as in the B3TF_MPn2 cluster, and *phnM1*. A second contig CDSs for *phnF* (a GntR-like transcriptional regulator; Gebhard & Cook, 2008), *phnG*, *H*, *I*, a short hypothetical protein, *phnJ*, *K*, and *L*. Finally, a third contig contained copies of *phnN*, *rcsF* (a phosphoesterase perhaps analogous to *phnP*; Martínez et al., 2013), and *phnM2*]. The draft assembly of this isolate's genome could not resolve if all these genes are clustered in the region. Nevertheless, they constitute a complete set of genes required for phosphonate degradation with additional putative auxiliary functions.

The *Oceanibulbus* HI0021, HI0027, HI0076, HI0082 isolates and the *Sulfitobacter* HI0054 also contained phosphonate degradation genes similar to *Roseovarius* HI0049 and the *Rhodobacterales* cluster B3TF_MPn_2 (Appendix I Table A8). A common difference was the absence of *phnM1* in the *Oceanibulbus* and *Sulfitobacter* genomes, though they carry a second copy homologous to *phnM2*. In the cytoplasm, PhnI catalyzes the conversion of MPn and ATP to adenosine and adenine in the presence of PhnG, H, and L. PhnM is a phosphatase that hydrolyzes α -D-ribose-1- methylphosphonate-5-triphosphate to form 5-phospho- α -D-ribosyl 1-alkylphosphonate, the substrate of the C-P bond cleavage reaction catalyzed by PhnJ (Zhang & van der Donk, 2012). Also lacking in all but one *Sulfitobacter* isolate (HI0054) was a putative transferase, located between *phnE2* and *phnM1* in the *Rhodobacterales* B3TF_MPn_2 cluster, though its function in the C-P lyase pathway remains to be determined. In these isolates' genome assembly the phosphonate degradation genes were found contiguous to one another in the same contig and are thus likely to form complete and functional operons. Interestingly, the *Rhodobacteraceae*, C-P lyase operons were accompanied by sequences encoding a blue-light activated protein, a photosensitive response regulator kinase (Swartz et al., 2005), and a bacterial SH3 domain-containing protein, which is likely involved in intracellular and membrane-associated functions (Morton & Campbell, 1994; Musacchio, Gibson, Lehto, & Saraste, 1992). If these genes are associated with C-P lyase activity the first may be an additional response regulator of the pathway that acts independently of the transcriptional regulator PhnP, and the second may mediate the assembly of the C-P lyase complex or aid in signal transduction. Expression of these genes during phosphonate degradation can be explored further as they may provide alternative regulatory pathways independent from the Pho regulon (sensitive to phosphorus levels) as is characteristic of the C-P lyase operon in *E. coli* (White and Metcalf, 2007).

The remaining isolates carrying phosphonate degradation genes comprised four *Gammaproteobacteria* related to the genus *Oleiphilus* (isolates HI0009, HI0066, HI0067, and HI0125). Interestingly, these *Oleiphilus* strains clustered together in the SSU rRNA gene tree (Appendix I Fig. A1). The phosphonate degradation genes of these strains were most similar to the genes in the gammaproteobacterium IMCC1989 cluster B3TF_MPn_1 though the percent identity between them ranged from 29-52%, a similarity perhaps too low to attribute the same function confidently. The low identity between gene clusters may be explained by the longer phylogenetic distance separating different *Gammaproteobacteria* lineages compared to other *Proteobacteria* (Williams et al., 2010). The C-P lyase gene clusters of isolates HI0009 and HI0066 were resolved on the same contig and consisted of *phnFDCEGHIJKLMNP*. Additionally, sequences similar to a TetR family transcriptional regulator and an integrase were found in the vicinity of these clusters. The two additional *Oleiphilus* isolates (HI0067 and HI0125) had the same phosphonate degradation genes though they were split between two or three contigs.

The *Pseudoalteromonas* strain HI0053, did not harbor genes matching these phosphonate degradation operons. It did carry genes annotated as a 2-aminoethylphosphonate--pyruvate transaminase (PhnW), next to a GntR family transcriptional regulator (perhaps repressing 2-aminoethylphosphonate transport), and a putative phosphohydrolase that may be associated with phosphonate metabolism. Also present were a second putative transcriptional regulator of 2-aminoethylphosphonate degradation and a 2-aminoethylphosphonate-pyruvate transaminase next to a predicted transcriptional regulator CysB. In addition, HI0053 contained a phosphonoacetaldehyde hydrolase (PhnX) and three sequences similar to PhnA (putative phosphonacetate metabolism). While these sequences do not comprise the canonical alkyl-

phosphonate degradation pathway they may encode enzymes that degrade other types of phosphonates occurring in the marine environment (Martinez, Tyson, & Delong, 2010).

Conclusion

The cultivation experiments conducted as part of this work were motivated by several lines of evidence indicating that isolates originating from mesopelagic microbial communities are capable of degrading semi-labile DOM, and therefore appear not totally reliant on recently produced photosynthate exported from the surface. In these experiments, HMW DOM treatments did in fact recover a larger number of cultures from Mesopelagic samples than from DCM samples. This result may be biased given the dominance of *Prochlorococcus* in the euphotic zone ($28 \pm 6.3\%$ of total prokaryotic cells in the upper 100 m at Station ALOHA) and its rapid decline below 150 m (Campbell et al., 1997) which may lower the chances of isolating bacterial heterotrophs from the DCM using the dilution-to-extinction method. However, considering the Experiment II dilution-to-extinction samples' final cell concentrations (3 cells ml⁻¹ and 2 cells ml⁻¹ for DCM and Mesopelagic experiments, respectively), even if only 2 out of every 3 cells in the DCM are heterotrophic bacteria (assuming 1 out of 3 cells are *Prochlorococcus*) the percent recovery of cultures in DCM samples is still 2-4 times lower than the percent recovery of Mesopelagic samples (Appendix I, Table A2). In this case it would seem that mesopelagic microbial communities harbor bacteria better adapted to degrade semi-labile HMW DOM. This result is consistent with microcosm experiments performed by Carlson et al. (2004) in which the mesopelagic bacterioplankton communities were more responsive than surface communities to the semi-labile organic matter amendments from filtered surface waters. However, since none of the Mesopelagic or DCM isolates tested observed significant increases in cell yields in response

to HMW DOM additions (see DOM dose-response experiment results in Figures 3.3, 3.4, and 3.5) it remains unclear if HMW DOM is more readily available to heterotrophic populations from either depth. The only isolates with a significant DOM dose-response to HMW DOM were the Pseudomonads but these are not dominant members of the marine microbiota. How well the properties of the isolates identified reflect the general characteristics of the indigenous communities remains to be determined.

Even though the presence of HMW DOM did not elicit a robust growth response in the isolates tested, an undetectable fraction of HMW DOM may be bio-available. Thus, through physiological screens and genomics I investigated which types of carbon substrates and metabolisms could support the growth of these isolates. The isolation of hydrocarbon-degrading bacteria related to the genus *Alcanivorax* and *Oleiphilus* was unexpected given that HMW DOM is primarily composed of carbohydrates. The occurrence of *Alcanivorax* cultures in particular did not increase in HMW DOM treatments compared to control cultures as did other putative hydrocarbon degraders like *Oleiphilus* and *Erythrobacter* isolates suggesting that the former are not stimulated by HMW DOM. However, a small proportion of low-molecular-weight lipids (6% of total carbon) has been detected in ultrafiltered HMW DOM samples (Aluwihare et al., 1997) which may be sufficient to stimulate the growth of these isolates in dilution-to-extinction cultures. It cannot be ruled out that other HMW DOM constituents like polysaccharides can supply the carbon and energy necessary for growth of such isolates. It is worth noting that many isolates, including members of the *Alcanivorax* and *Oleiphilus*, were able to support their growth with acetate as observed in Biolog multi-carbon screens. Proton NMR of HMW DOM revealed a clear signature of acetylated polysaccharides (Aluwihare et al., 1997) of which N-acetyl amino polysaccharides, for example, can contribute ~26% of the total carbon in HMW DOM from

surface waters (Aluwihare et al., 2005). Bacteria like the *Oleiphilus* isolates investigated may be capable of degrading the acetate moieties from these biomolecules using carbohydrate acetyl esterases as identified in the CAZy gene analysis. In addition, the prevalence of these and other hydrocarbon-degrading bacteria in dilution-to-extinction cultures and in the marine environment may be associated with the production of higher alkanes by cyanobacteria, in particular *Prochlorococcus* (Lea-Smith et al., 2015), the dominant autotroph in subtropical gyres. On the other hand, the *Rhodobacteraceae* isolates identified in DOC treatments, primarily stimulated by amino acids, may be utilizing N available in HMW DOM from the degradation of proteins (21% of total N) or in N-acetyl amino polysaccharides (43% of total N) using cell surface-bound amidases (Aluwihare et al., 2005). These and similar hypotheses generated using preliminary physiological and genomic evidence can now be tested in the laboratory with the bacterial isolates and appropriate analytical techniques.

Even though HMW DOM treated samples produced a greater number of cultures, the DOM-growth response of open-ocean isolates was only discernible when UV-oxidized seawater was used as the base media. This low-DOC media was not used in DOM dose response tests with isolates from Experiment II which may explain the small increase in cell yields observed despite high DOC treatments (Fig. 3.7). Regardless, the addition of HMW DOM was expected to prompt a proportional increase in cell yields if open-ocean cultures were able to utilize organic carbon constituents for growth but this was not the case. These observations indicate that, in its native form, HMW DOM collected from the upper ocean is not easily accessible to marine microbial heterotrophic populations, at least under the timescales and incubation conditions of these experiments. Most likely, microbial partners may be required to degrade different HMW DOM

constituents including amino acids in proteins, the acetate and amide nitrogen moieties in polysaccharides, trace lipids as well as humic acids.

Even though HMW DOM appears to be a limited source of carbon to the monocultures investigated it may contain sufficient labile carbon or nutrients to stimulate their initial growth from a single cell in dilution-to-extinction samples to measureable concentrations of 10^5 - 10^6 cell ml^{-1} . To illustrate this, bacterial cultures reaching 10^6 cell ml^{-1} in 12X DOC treatments (960 nmols C ml^{-1}) only require 8.3 nmols C assuming a carbon content of marine bacteria of 20 fg cell^{-1} (Lee & Fuhrman, 1987) and a growth efficiency of <20% (Carlson & Ducklow, 1996) which is <1% of the total DOC that was available. On the other hand, trace metals, vitamins, or other growth requirements may be necessary for these bacterial cultures to degrade HMW DOM although the coastal isolates described in Chapter 2 did not receive such supplements and were able to increase their cell yields with respect to the control low-DOC media. It is possible that the open-ocean bacterial isolates tested may only consume nanomolar amounts of carbon from HMW DOM and that microbial partners are required to further access HMW DOM substrates.

Another possibility is that HMW DOM samples contain growth factors like vitamins that stimulated the growth of bacteria on labile DOC already present in the natural seawater media but did not further enhance their growth. Vitamin B-12, for example, is present at picomolar concentrations in seawater (Carini et al., 2014; Menzel & Spaeth, 1962) and has a molecular weight of 1.2 kDa indicating it can be easily concentrated by the 1 kDa ultrafiltration membranes used to collect HMW DOM. An alternative explanation to the observed effect that HMW DOM had on the recovery of dilution-to-extinction cultures may be that it serves as antioxidant. Natural polysaccharides including fucans (sulfated sugars) and agar are known to have antioxidant properties that protect organisms by scavenging reactive oxygen species (Kong

et al., 2010; Wang et al., 2013). Because HMW DOM is largely composed of polysaccharides, the DOC treatments may have effectively increased the antioxidant potential of the cultivation media, especially the 10X DOC and 4X DOC F1 (HMW neutral polysaccharides) treatments, increasing the viability of the few cells present in the dilution-to-extinction samples at the beginning of the incubation. However, the loss of some cultures while passaging cells into sterile DOM media and the lower cell yields of *Alteromonas* cultures with increased DOC indicates that some chemical constituents of HMW DOM may also be inhibitory or toxic to cells.

Investigating DOM-degrading capabilities encoded in the genomes of these isolates remains difficult because the molecular composition of HMW DOM is not well resolved. Because HMW DOM carbohydrates consist on average of five-sugar oligosaccharides with an average molecular mass of 1 kDa (Aluwihare et al., 1997), I hypothesized that isolates enriched in the DOC-treated samples would harbor carbohydrate-degrading complexes similar to microorganisms that specialize in the breakdown of structural plant polysaccharides (e.g., lignins and cellulose) including marine bacterial isolates like *Saccharophagus degradans* (Weiner et al., 2008) and *Cellvibrio japonicus* (DeBoy et al., 2008). Nevertheless, only one clear example of a secreted carbohydrase was identified in the mesopelagic *Pseudoalteromonas* isolate HI0053. Even though HI0053 and other isolates were also able to metabolize simple sugars, none of these were able to clearly increase their cell yields with the HMW polysaccharides obtained from HMW DOM indicating that carbohydrates were not bio-available. Perhaps even at the mesopelagic depth of 200 m, sinking POM is still the primary source of carbon and energy to copiotrophic heterotrophs. It may be necessary to attempt to obtain more oligotrophic isolates, or sample deeper in the water column to find more efficient HMW DOM degraders, capable of metabolizing semi-labile organic carbon constituents.

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I am indebted to several people who assisted in field work and analysis in the laboratory. Special thanks to Daniel Repeta, Scott Gifford, and Rene Boiteau who organized HMW DOM sampling in Hawaii. Thank you to Anna Romano and Tsultrim Palden for assistance DNA sequencing and John Eppley and Torben Nielsen for help with genome assemblies and annotations. Thanks to the crew of the RV Kilo Moana, the Hawaii Ocean Time (HOT) series team members and C-MORE community aboard research cruises HOT 250 and HOE-BOE I for their dedication to sea operations. Financial support for this work was provided by the National Science Foundation Center for Microbial Oceanography: Research and Education (award #EF0424599) and the Gordon and Betty Moore Foundation (grant #492.01, #3777, and #3298).

Table 3.1. DOC treatments and incubation conditions of HMW DOM-enriched dilution-to-extinction culturing open-ocean experiments. Experiment I and II were prepared with seawater samples collected at Station ALOHA on March 2013 and March 2014, respectively.

Experiment	Inoculum depth	Final dilution (cell ml ⁻¹)	Nutrient additions	DOC treatments	Estimated HMW DOM additions (μM DOC)	Incubation temp.	Incubation period
I	110 m (DCM)	5	400 μM NaNO ₃	Control	No addition	22°C	2.5 months
	200 m (Meso)	4	30 μM KH ₂ PO ₄	2.5X DOC	114.5		1 month
				7X DOM	459.2		
II	95 m (DCM)	3	60 μM NaNO ₃	Control	No addition	26°C	2 months
	250 m (Meso)	2	4 μM NH ₄ Cl 7 μM KH ₂ PO ₄	4X DOC	229.7		
				10X DOC	472.1		
				4X DOC F1	319		

Abbreviations: DOM, dissolved organic matter; HMW, high-molecular-weight; DCM, deep chlorophyll maximum; Meso, mesopelagic.

Table 3.2. Phylogenetic identification of Hawaii isolates via assembled SSU rRNA genes.

The dilution-to-extinction experiment and DOC treatment from which isolates originated are indicated as well as the number of cultures with the same SSU rRNA gene sequence similarity.

Isolate ID ^a	Best matching described taxa via SSU rRNA ^b	%ID ^c	Length ^d	No. cultures ^e	DOC treatment
Experiment I					
DCM					
S01	Halioglobus; uncultured bacterium	99.73	1492	17	No DOC
Mesopelagic					
D01	<i>Pseudomonas stutzeri</i> T13	99.93	1517	19	7X
D02	<i>Pseudomonas</i> sp. R-25343	99.92	1312	11	7X
D04	<i>Pseudomonas stutzeri</i> T13	99.87	1517	22	7X
D40	<i>Pseudomonas</i> sp. R-25343	99.86	1408	1	7X
D59	<i>Erythrobacter</i> ; uncultured bacterium (EF659424.1.1483)	99.86	1453	6	2.5X
D65	<i>Vibrio crassostreae</i> 9ZC13	98.83	1535	6	No DOC
D71	<i>Vibrio</i> ; uncultured marine microorganism (EU188141.1.1516)	99.5	1205	1	No DOC
Experiment II					
DCM					
9	<i>Oleiphilus</i> ; uncultured bacterium AD12-E7	97.95	1514	1	No DOC
21	<i>Oceanibulbus indolifex</i> HEL-45	99.52	1444	1	10X
27	<i>Oceanibulbus indolifex</i> HEL-45	99.51	1219	1	10X
19, 28	<i>Erythrobacter</i> ; uncultured bacterium AD84-H10	99.86	1468	2	10X
23, 40	<i>Oceanibulbus</i> ; marine bacterium SCRIPPS_101	98.05	1437	2	10X, 4X F1
20, 37, 38	<i>Erythrobacter</i> ; uncultured bacterium (EF659424.1.1483)	99.93	1453	3	10X, 4X F1
3, 7, 11, 13, 33, 35	<i>Alcanivorax</i> sp. Abu-1	99.93	1507	6	No DOC, 4X, 4X F1
Mesopelagic					
43 ^f	<i>Oleiphilus</i> ; uncultured gamma proteobacterium (JQ579692.1.1494)	98.19	1494	16	No DOC
49	<i>Roseovarius</i> ; uncultured alpha proteobacterium (HQ727624.1.1431)	99.79	1431	1	4X
53	<i>Pseudoalteromonas</i> ; uncultured bacterium (KF906589.1.1533)	99.47	1509	1	4X
54	<i>Sulfitobacter</i> sp. CC-AMSY-48	99.93	1429	1	4X
63	<i>Erythrobacter</i> ; uncultured bacterium HF0500_24B12	100.00	1469	1	4X
65 ^g	<i>Oleiphilus</i> ; uncultured bacterium AD12-E7	99.27	1514	15	10X
75	<i>Oleibacter</i> ; uncultured bacterium (AF382127.1.1491)	99.73	1491	1	10X
76	<i>Oceanibulbus indolifex</i> HEL-45	99.54	1300	1	10X
82	<i>Oceanibulbus indolifex</i> HEL-45	99.45	1444	1	10X
90	<i>Alteromonas</i> sp. KOPRI 11568	99.76	1276	1	4X F1
92	<i>Alteromonas macleodii</i> str. 'Black Sea 11'	98.53	819	1	4X F1
96	<i>Alcanivorax</i> sp. Mho1	99.93	1435	1	4X F1
97 ^g	<i>Oceanibulbus indolifex</i> HEL-45	99.51	1219	1	4X F1
107	uncultured <i>Psychrobacter</i> sp. (JX530017.1.1434)	82.57	1113	1	4X F1
109	<i>Alteromonas macleodii</i> str. 'Black Sea 11'	100.00	1187	1	4X F1
113	<i>Oleibacter</i> ; uncultured bacterium (AF382127.1.1491)	99.80	1491	1	4X F1
120	<i>Thalassolituus</i> ; uncultured gamma proteobacterium (FJ403082.1.1500)	97.87	1500	1	4X F1
66, 67	<i>Oleiphilus</i> ; uncultured bacterium AD12-E7	97.95	1514	2	10X
74, 77	<i>Erythrobacter</i> ; uncultured bacterium AD84-H10	99.86	1468	2	10X
44, 83	<i>Alcanivorax</i> ; unidentified (E31375.1.1516)	99.54	1516	2	No DOC, 10X
69, 85	<i>Oleiphilus</i> ; uncultured gamma proteobacterium (JQ579692.1.1494)	98.13	1494	2	10X
94, 125	<i>Oleiphilus</i> ; uncultured bacterium AD12-E7	97.89	1514	2	4X F1
112, 129	<i>Oceanibulbus</i> ; marine bacterium SCRIPPS_101	98.05	1437	2	4X F1

^aCulture identification number (HI0###).

^bTop hit of a LAST sequence similarity search against the SILVA 119 SSU database.

^cPercent matching identities to SILVA reference sequence

^dAlignment length

^eNumber of cultures with equal SSU rRNA gene percent matching identities

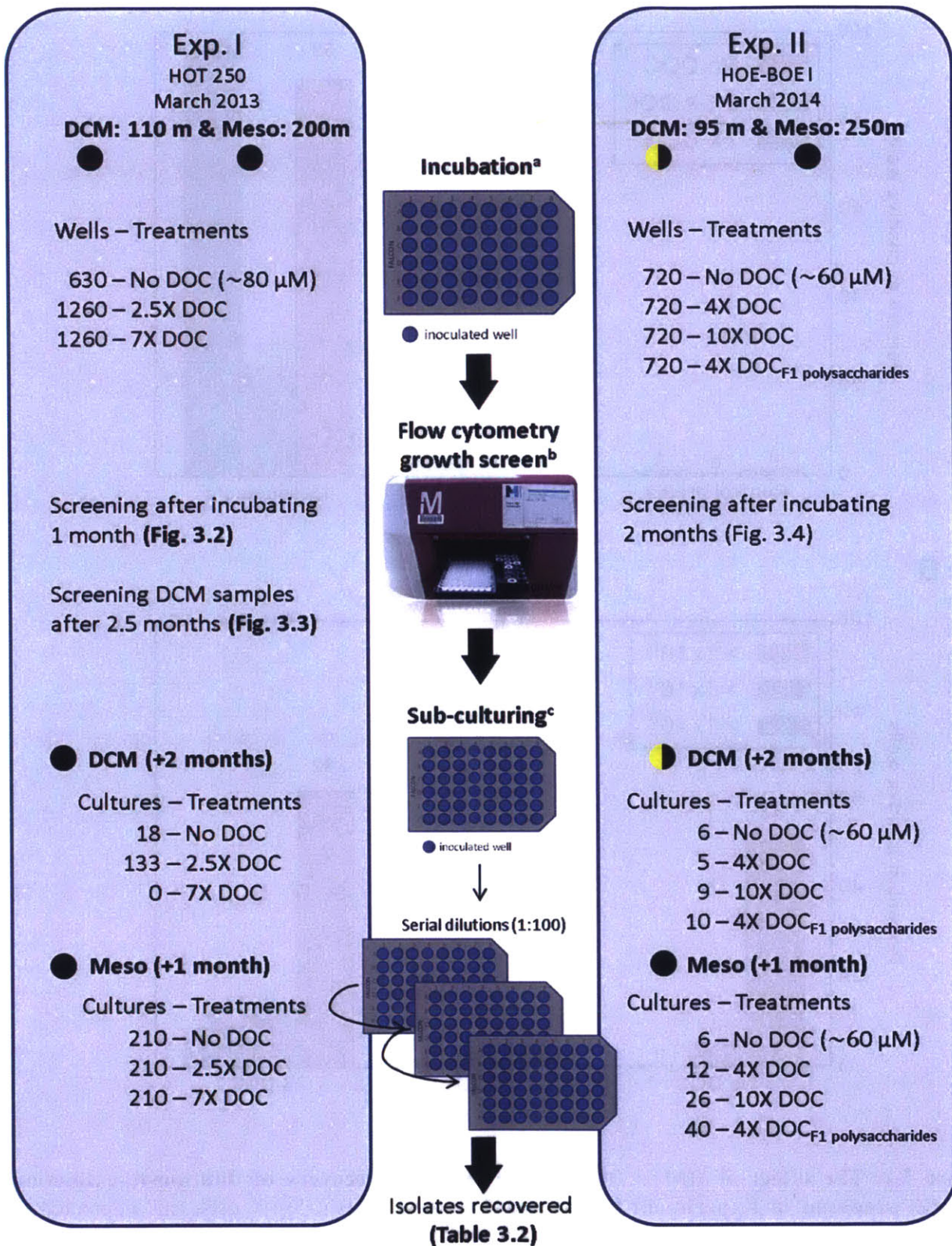
^fIsolate IDs (HI0###) matching HI0043: 050, 061, 068, 072, 078, 081, 086, 098, 100, 103, 111, 117, 123, 128, and 132

^gIsolate IDs (HI0###) matching HI0065: 071, 073, 079, 080, 087, 088, 104, 105, 108, 114, 118, 122, 130, 133

Figures

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Figure 3.1. HMW DOM enrichments of dilution-to-extinction cultures from Hawaii. Two dilution-to-extinction cultivation experiments (I and II) were prepared with HMW DOM enrichments of seawater samples collected at Station ALOHA from the deep chlorophyll maximum (DCM) and mesopelagic waters (Meso). ^a**Incubations.** The number of dilution-to-extinction wells prepared for each DOC treatment is indicated. Black circles indicated incubations in darkness. Half black/yellow circles indicate incubations under a 12-hour diurnal cycle). ^b**Flow cytometry growth screen.** The screening procedure is discussed in detail in the Materials and Methods. ^c**Sub-culturing.** Wells considered positive hits were sub-cultured by dilution into fresh media of the corresponding DOC treatment. The number of wells sub-cultured from each DOC treatment and the incubation conditions are indicated. The isolates recovered from the sub-culturing steps are indicated in **Table 3.2.**



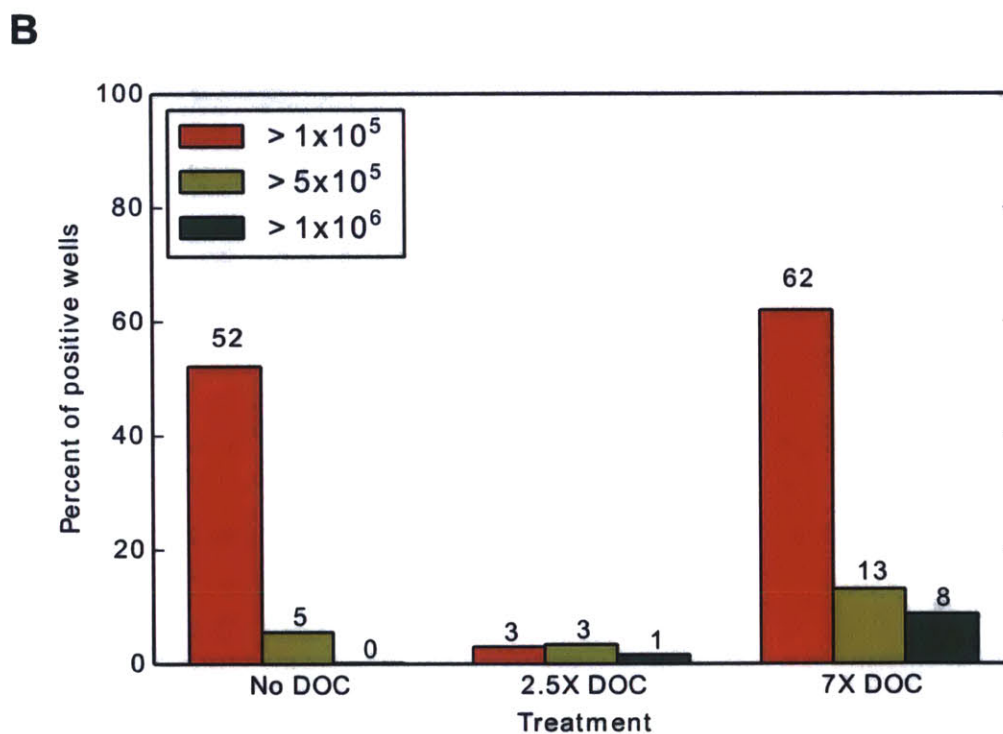
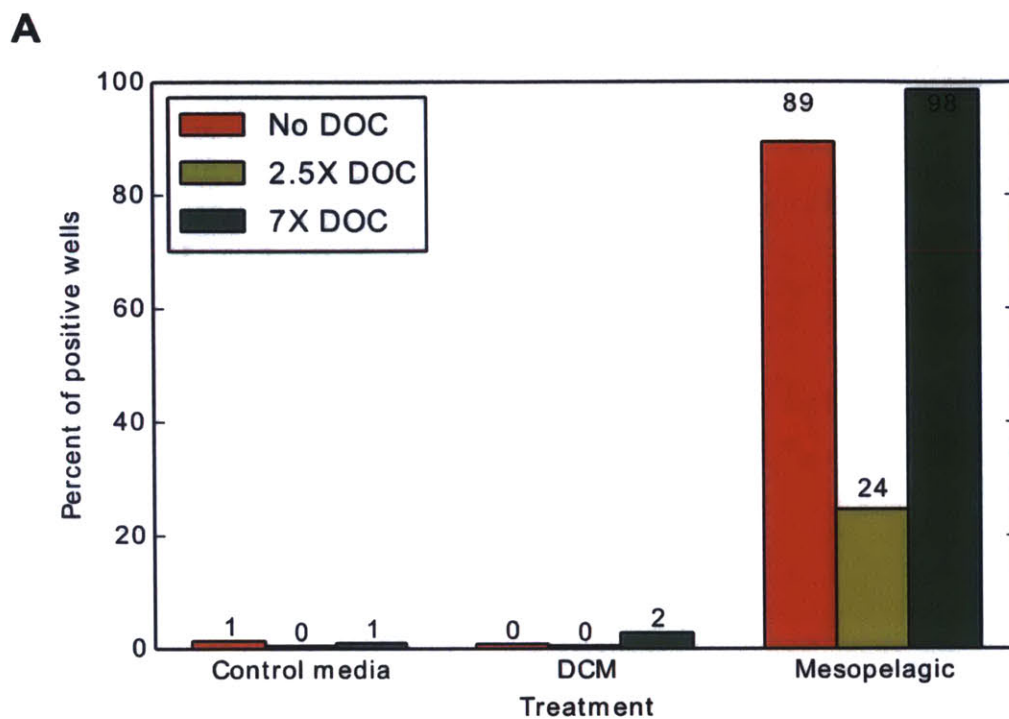


Figure 3.2. The effect of HMW DOM on growth and recovery of dilution-to-extinction cultures prepared in Experiment I. A) Percent positive wells ($>5 \times 10^4$ cells ml^{-1}) detected in each DOC treatment after one month of incubation. B) Mesopelagic well cell densities (cell ml^{-1}) are color coded: red, $>10^5$ to $<5 \times 10^5$; yellow, $>5 \times 10^5$ to $<10^6$; green, $>10^6$. No DOC, non-DOM amended samples; 2.5X DOC, $\sim 200 \mu\text{M C}$; 7X DOC, $\sim 560 \mu\text{M C}$.

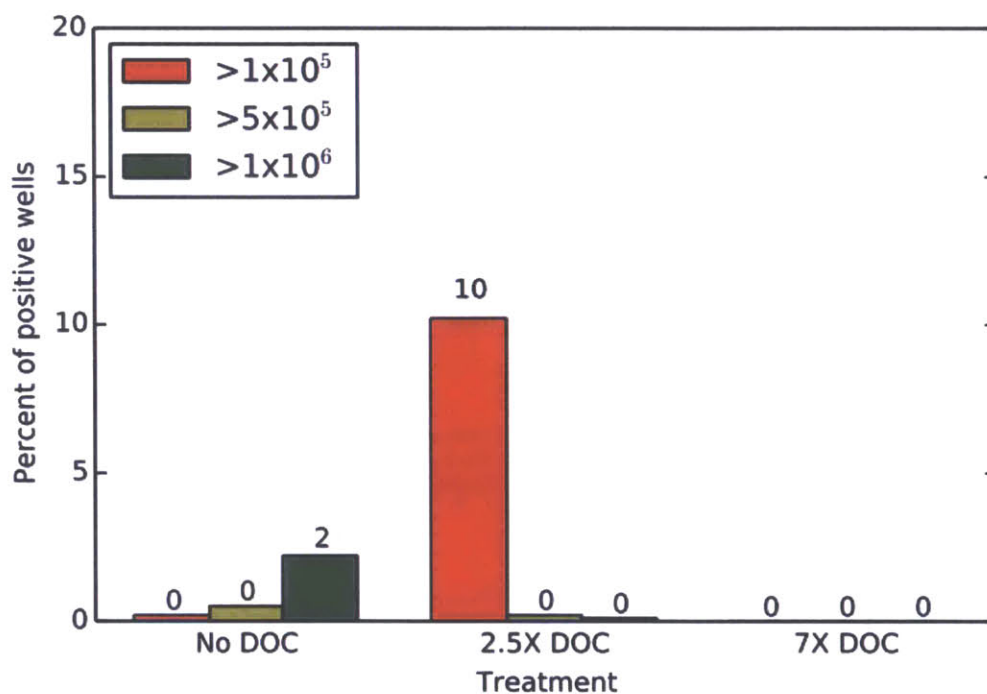


Figure 3.3. The effect of HMW DOM on growth and recovery of DCM dilution-to-extinction cultures after an extended incubation period. DCM cultures incubated for an additional 1.5 months after the first screen (Fig. 3.1). Positive wells are indicated as a fraction of the total number of wells screened for each treatment. Total number of wells screened: No DOC, 630 wells; 2.5X DOC, 1260 wells; 7X DOC, 630 wells. Well cell densities (cell ml^{-1}) are color coded: red, $>10^5$ to $<5 \times 10^5$; yellow, $>5 \times 10^5$ to $<10^6$; green, $>10^6$. DOC treatments: No DOC, non-DOM amended samples; 2.5X DOC, $\sim 200 \mu\text{M C}$; 7X DOC, $\sim 560 \mu\text{M C}$. Numbers on the top of each bar indicate the percent of positive wells (0 if $<1\%$).

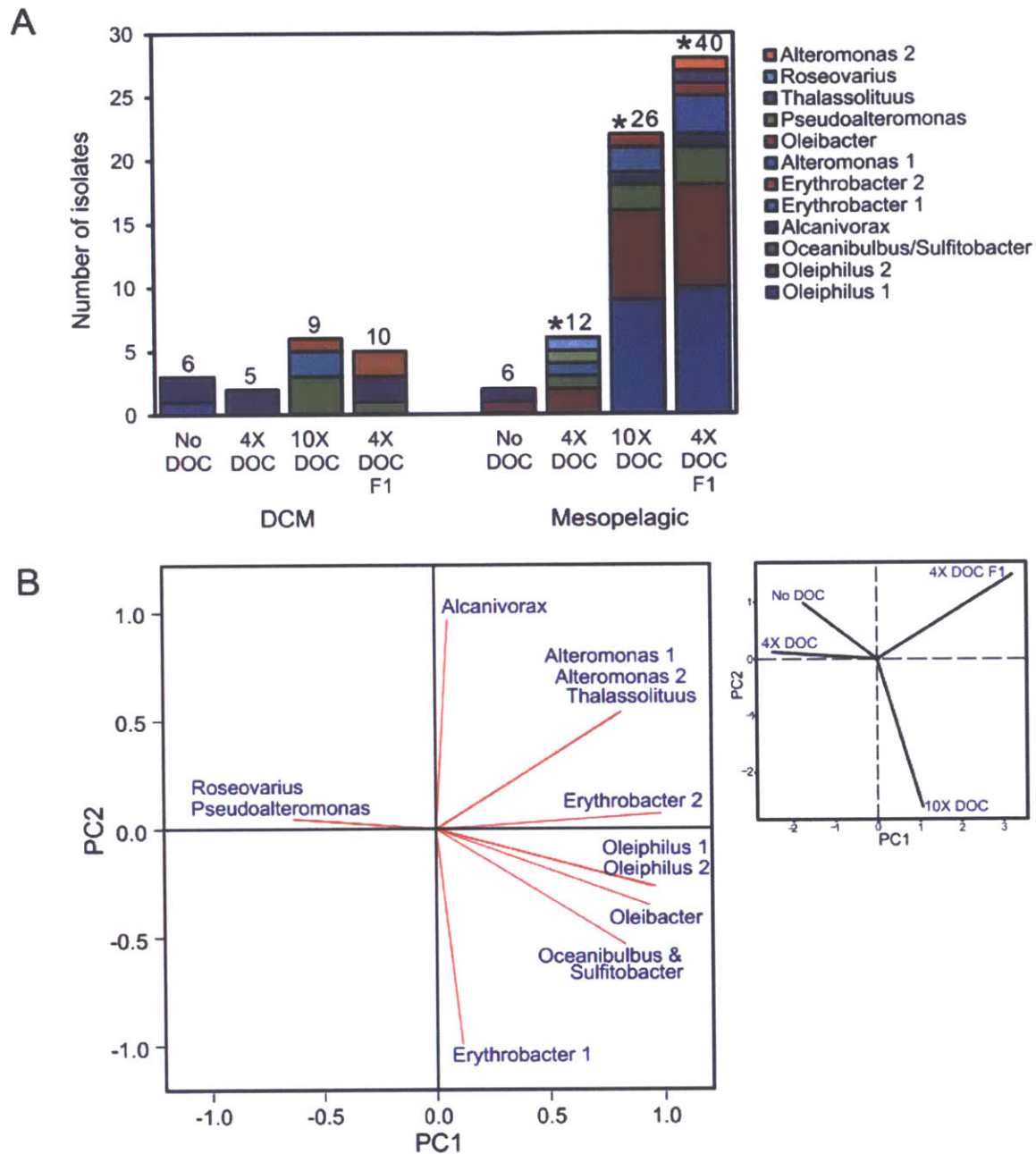


Figure 3.4. Bacterial phylogenetic groups identified in Experiment II. A) Occurrence of SSU rRNA gene clusters (98% sequence similarity) among DOC treatments. The number on the top of each column indicates the total number of positive wells ($>1 \times 10^4$ cells ml^{-1}) originally detected in the flow cytometry growth screen. Treatments marked with (*) were significantly different from the control ($P < 0.05$, bootstrap comparisons with 10^4 iterations). B) The occurrence of SSU rRNA gene clusters across treatments (independent of sample-depth) was assessed through a principal component analysis (PCA; PC1 = 59.79%, PC2 = 28.08% of variance, respectively). The inset indicates the PCA score received by each DOC treatment.

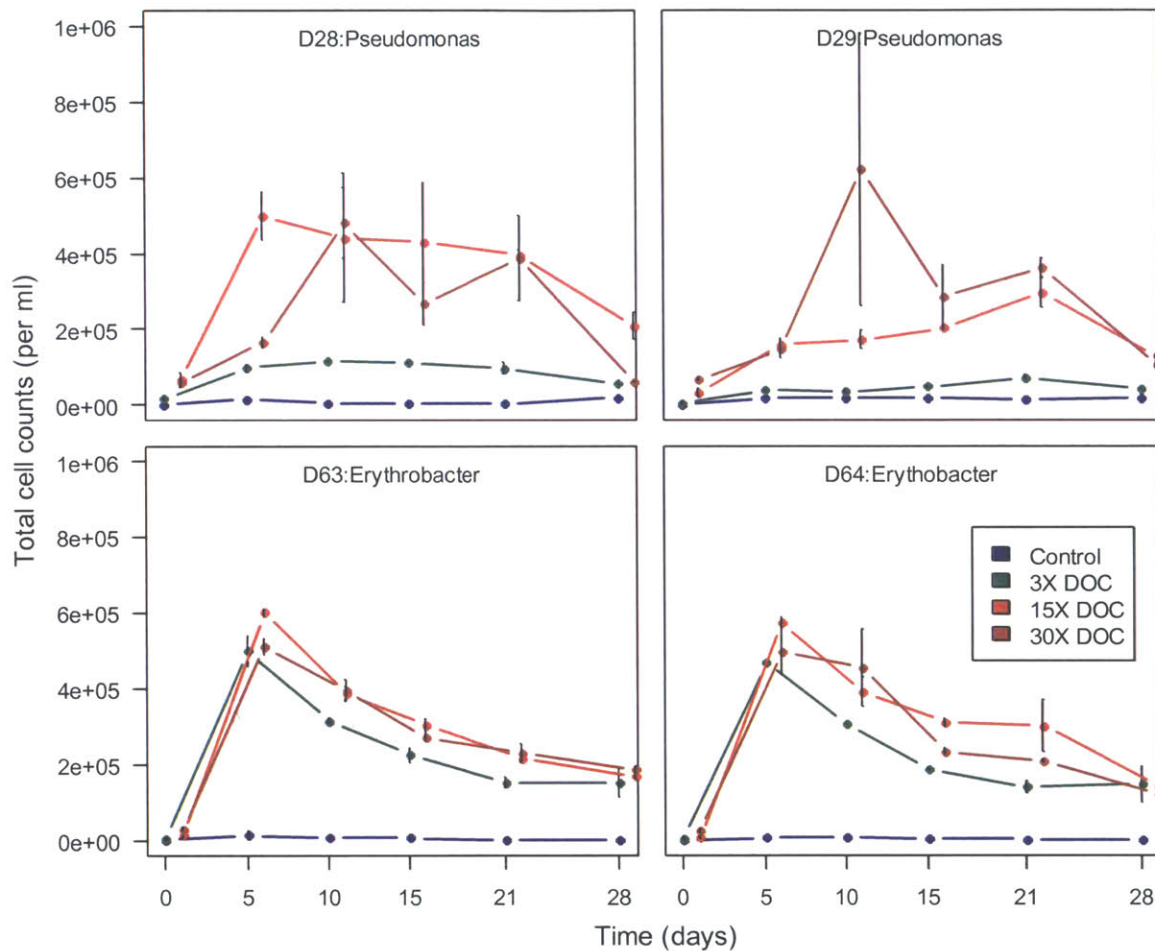


Figure 3.5. The growth response of Experiment I Hawaii *Pseudomonas* and *Erythrobacter* isolates to HMW DOM additions. *Pseudomonas* and *Erythrobacter* isolates were obtained from 7X and 2.5X DOC treatments of Mesopelagic samples, respectively. Controls consisted of UV-oxidized seawater (25 μM DOC). Treatments 3X, 15X, and 30X DOC indicate the estimated DOC concentration above natural Hawaii open-ocean waters ($\sim 80 \mu\text{M}$ DOC). Error bars indicate the difference between duplicate samples.

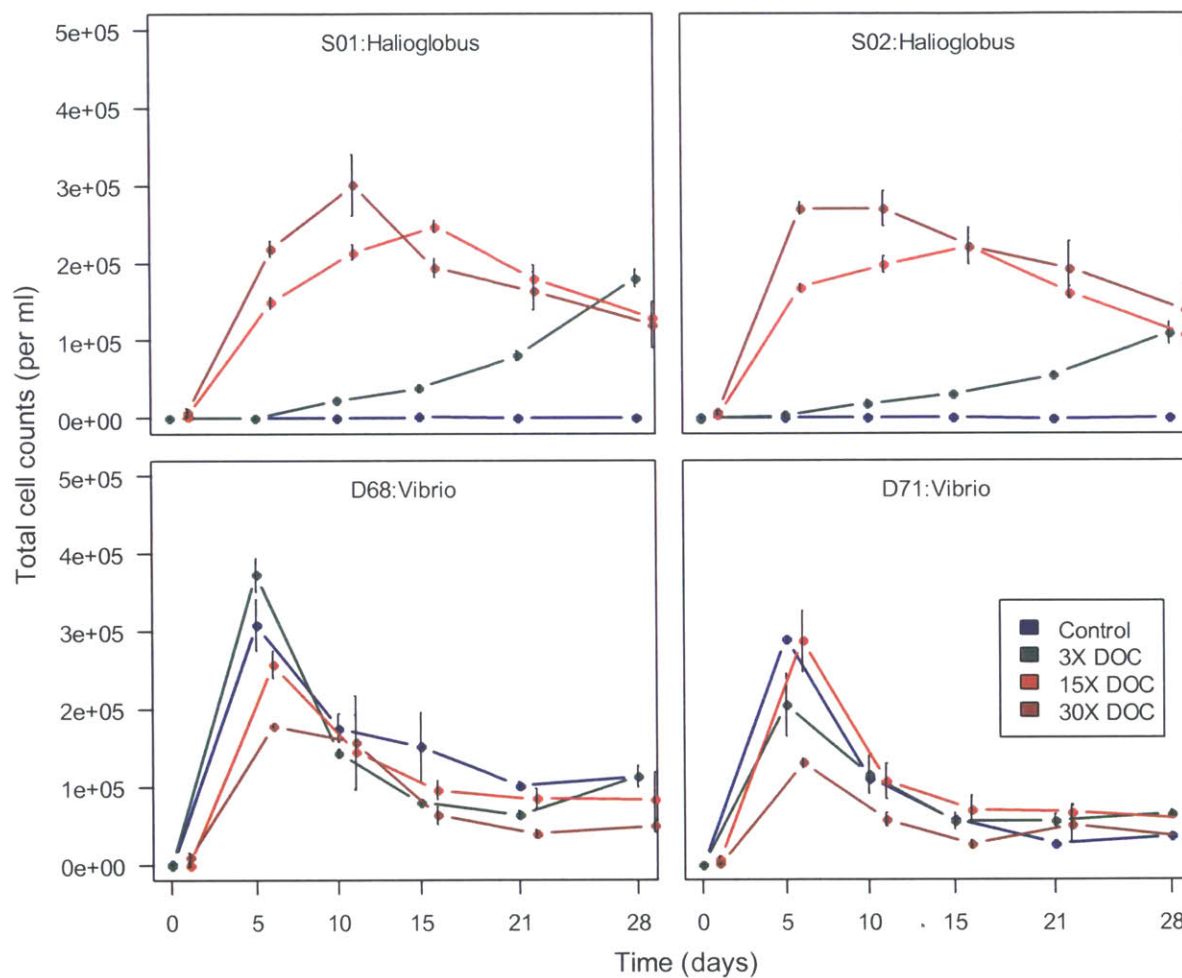


Figure 3.6. The growth response of Experiment I Hawaii *Halioglobus* and *Vibrio* isolates to HMW DOM additions. *Halioglobus* and *Vibrio* isolates were obtained from DCM and Mesopelagic No DOC treatments, respectively. Controls consisted of UV-oxidized seawater (25 μM DOC). Treatments 3X, 15X, and 30X DOC indicate the estimated DOC concentration above natural Hawaii open-ocean waters ($\sim 80 \mu\text{M}$ DOC). Error bars equal the difference between duplicate samples.

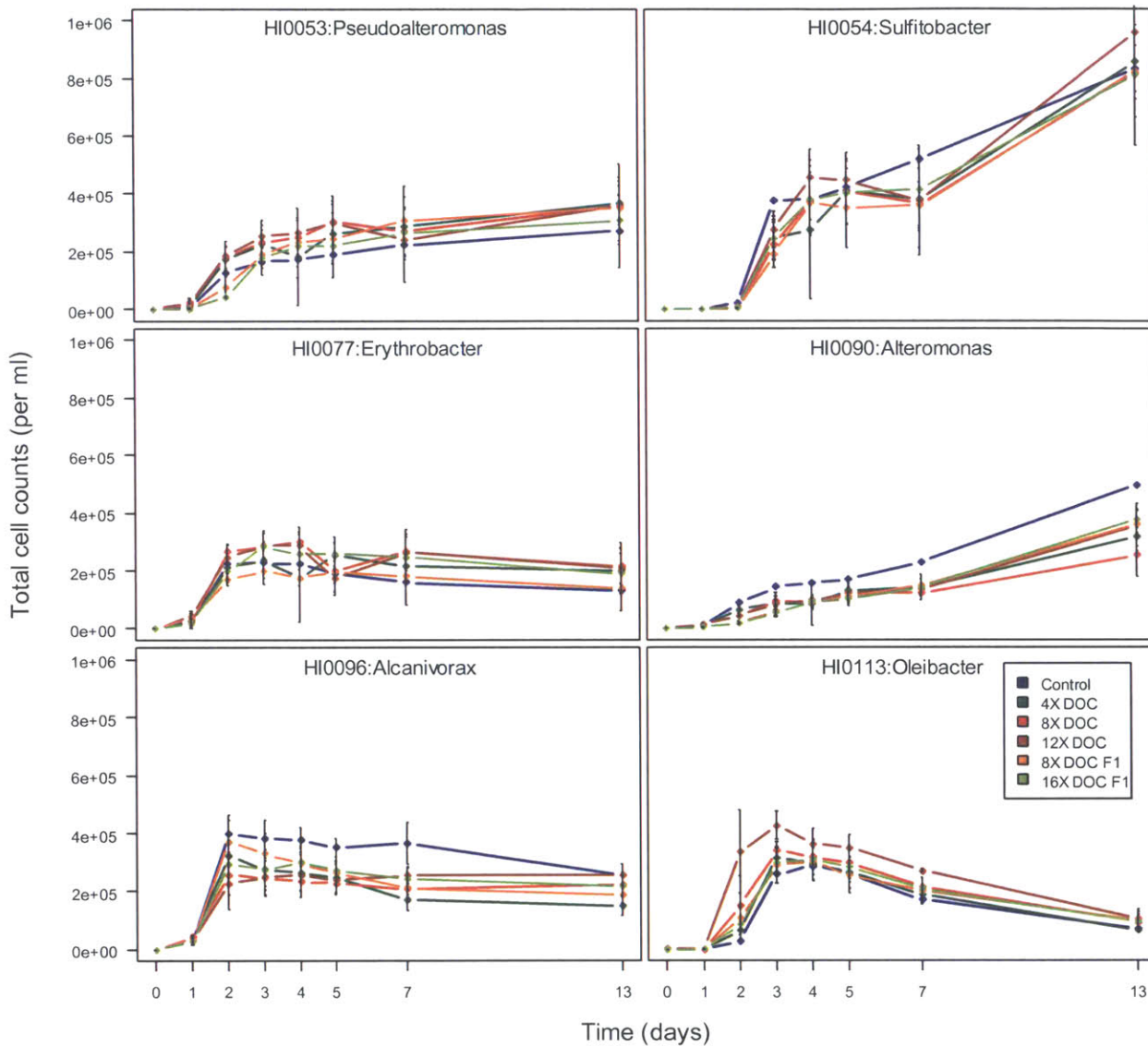


Figure 3.7. The growth response of Experiment II Hawaii isolates to increasing concentrations of HMW DOM. The control treatment consisted of unamended Hawaii seawater. 4X, 8X, 12X, and 16X indicate the DOC concentrations above the typical concentration in open-ocean Hawaii waters ($\sim 80 \mu\text{M}$ DOC). F1 refers to the neutral polysaccharide fraction of HMW DOM obtained by chromatography. Error bars equal one standard deviation.

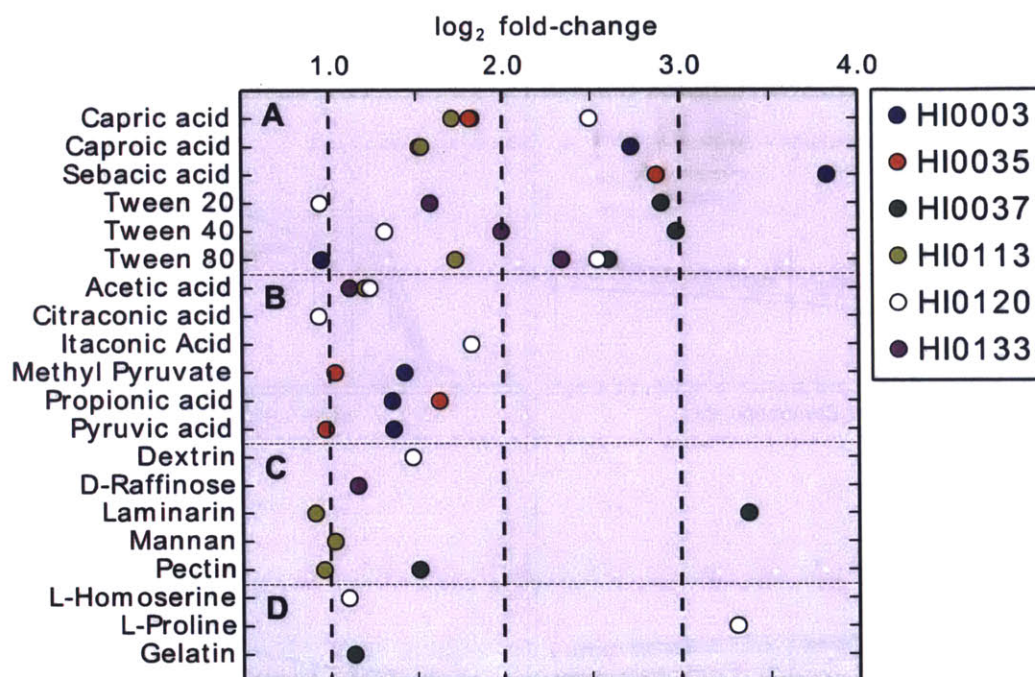
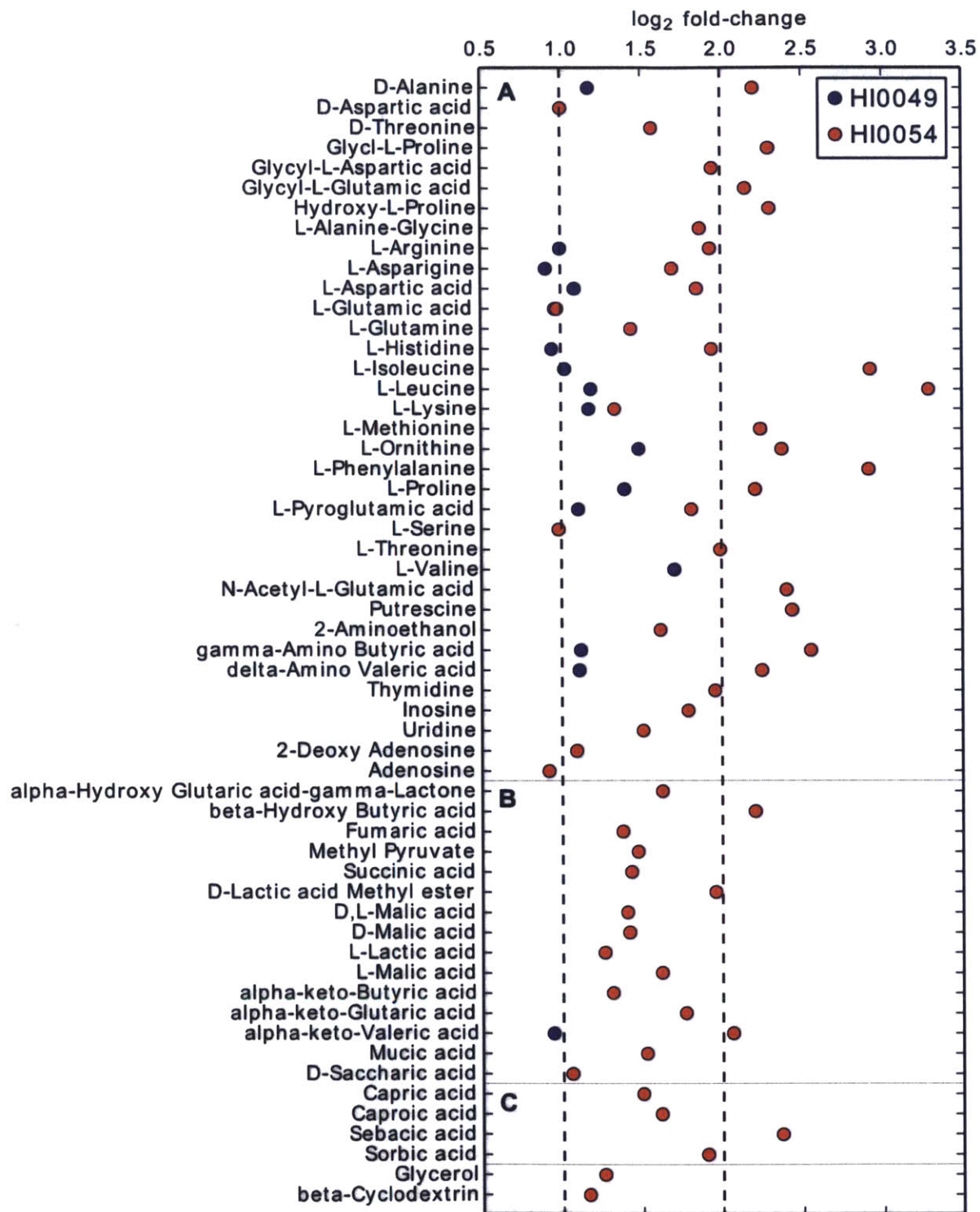


Figure 3.8. Carbon substrate utilization profiles of open-ocean putative hydrocarbon degraders. Sub-panels: (A) hydrocarbon chain-containing compounds, (B) low-molecular-weight organic acids, (C) carbohydrates, (D), amino acids and nitrogen-containing compounds. Biolog substrates for which the \log_2 fold change was >0.9 or <-1.0 are shown. A \log_2 fold change = 2 indicates cell yields $4\times$ the median cell yield for all substrates ($n=190$). Legend: HI0003 (*Alcanivorax*), HI0035 (*Alcanivorax*), HI0037 (*Erythrobacter*), HI0113 (*Oleibacter*), HI00120 (*Thalassolitus*), and HI0133 (*Oleiphilus*).

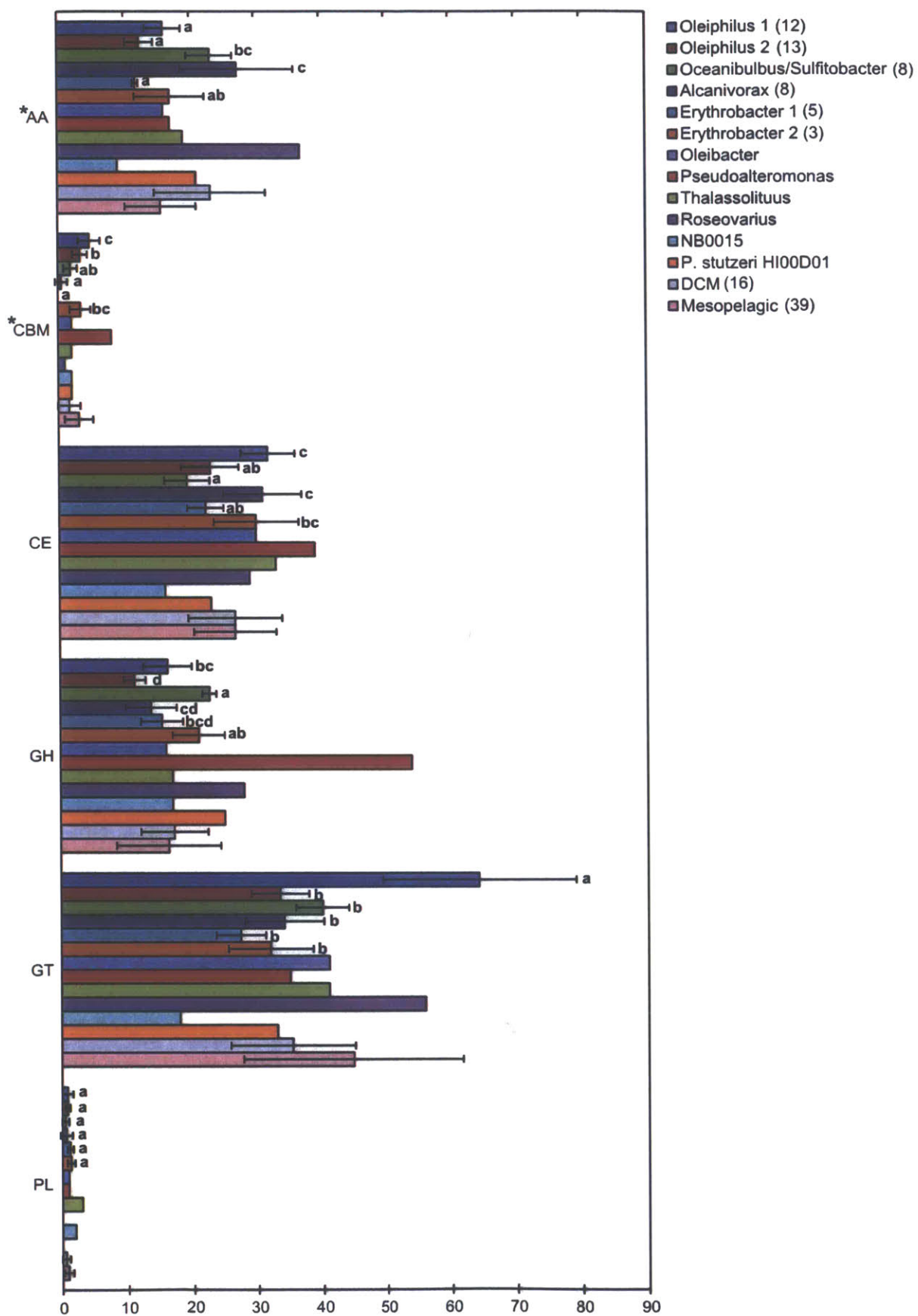
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Figure 3.9. Carbon substrate utilization profiles of Hawaii *Rhodobacteraceae* isolates. Sub-panels: (A) amino acids and nitrogen-containing compounds, (B) low-molecular-weight organic acids, (C) hydrocarbon chain-containing compounds. Biolog substrates for which the \log_2 fold change was >0.9 are shown. A \log_2 fold change = 1 indicates cell yields doubled the median cell yield for all substrates ($n=190$). Legend: HI0049 (*Roseovarius*) and HI0054 (*Sulfitobacter*).



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Figure 3.10. The occurrence of predicted CAZy genes in the genomes of Hawaii isolates. Bars indicate the average number of CAZy genes (per enzyme family) identified in each phylogenetic cluster (98% SSU rRNA sequence similarity) or in DCM and Mesopelagic genomes. The total number of genomes in each group is indicated in parenthesis. Groups with no number indicated consisted of a single representative. CAZy enzyme families: AA, auxiliary activities; CBM, carbohydrate-binding modules; CE, carbohydrate esterases; GH, glycoside hydrolases; GT, glycosyl transferases; PL, polysaccharide lyases. The average number of enzyme families marked with (*) was significantly different between DCM and Mesopelagic genomes (Mann-Whitney U test, P-value <0.05). A one-way ANOVA was conducted to compare the number of predicted CAZy genes of each class (AA, CBM, CE, GH, GT, and PL) among phylo-types (P-value<0.05). Bars marked with the same letter were not significantly different according to post hoc analysis using Tukey's honest significance difference test (P-value<0.05). Error bars are equal to the SEM of predicted CAZy genes for each phylo-type or isolate.



References

- Aluwihare, L. I., Repeta, D. J., & Chen, R. F. (1997). A major biopolymeric component to dissolved organic carbon in surface sea water. *Nature* 387, 166–169.
- Aluwihare, L. I., Repeta, D. J., Pantoja, S., & Johnson, C. G. (2005). Two chemically distinct pools of organic nitrogen accumulate in the ocean. *Science*, 308, 1007–1010. <http://doi.org/10.1126/science.1108925>
- Aristegui, J., Agustí, S., Middelburg, J. J., & Duarte, C. M. (2005). Respiration in the mesopelagic and bathypelagic zones of the oceans. In del Giorgio P. and Williams P. (Eds.). *Respiration in Aquatic Ecosystems*, (pp. 181–205). Oxford University Press, New York.
- Ausmees, N., Mayer, R., Weinhouse, H., Volman, G., Amikam, D., Benziman, M., & Lindberg, M. (2001). Genetic data indicate that proteins containing the GGDEF domain possess diguanylate cyclase activity. *FEMS Microbiology Letters*, 204, 163–167.
- Aylward, F. O., Eppley, J. M., Smith, J. M., Chavez, F. P., Scholin, C. a., & DeLong, E. F. (2015). Microbial community transcriptional networks are conserved in three domains at ocean basin scales. *Proceedings of the National Academy of Sciences of the USA*, 112(17), 5443–5448.
- Azam, F. (1998). Microbial Control of Oceanic Carbon Flux: The Plot Thickens. *Science*, 280(5364), 694–696.
- Azam, F., Fenchel, T., Field, J., Gray, J., Meyer-Reil, L., & Thingstad, F. (1983). The ecological role of water-column microbes in the sea. *Marine Ecology Progress Series*, 10, 257–263.
- Bankevich, A., Nurk, S., Antipov, D., Gurevich, A. a., Dvorkin, M., Kulikov, A. S., ... Pevzner, P. a. (2012). SPAdes: A New Genome Assembly Algorithm and Its Applications to Single-Cell Sequencing. *Journal of Computational Biology*, 19(5), 455–477.
- Benner, R., Biddanda, B., Black, B., & McCarthy, M. (1997). Abundance, size distribution, and stable carbon and nitrogen isotopic compositions of marine organic matter isolated by tangential-flow ultrafiltration. *Marine Chemistry*, 57, 243–263.
- Benner, R., Pakulski, J. D., McCarthy, M., Hedges, J. I., & Hatcher, P. G. (1992). Bulk chemical characteristics of dissolved organic matter in the ocean. *Science*, 255, 1561–1564.
- Bochner, B. R., Gadzinski, P., & Panomitros, E. (2001). Phenotype Microarrays for high-throughput phenotypic testing and assay of gene function. *Genome Research*, 11, 1246–1255.

- Brix, H., Gruber, N., Karl, D. M., & Bates, N. R. (2006). On the relationships between primary, net community, and export production in subtropical gyres. *Deep Sea Research Part II: Topical Studies in Oceanography*, 53, 698–717.
- Camacho, C., Coulouris, G., Avagyan, V., Ma, N., Papadopoulos, J., Bealer, K., & Madden, T. L. (2009). BLAST plus: Architecture and applications. *BMC Bioinformatics*, 10, 421.
- Campbell, L., Liu, H., Nolla, H. a., & Vaulot, D. (1997). Annual variability of phytoplankton and bacteria in the subtropical North Pacific Ocean at Station ALOHA during the 1991-1994 ENSO event. *Deep-Sea Research Part I: Oceanographic Research Papers*, 44(2), 167–192.
- Cantarel, B. I., Coutinho, P. M., Rancurel, C., Bernard, T., Lombard, V., & Henrissat, B. (2009). The Carbohydrate-Active EnZymes database (CAZy): An expert resource for glycogenomics. *Nucleic Acids Research*, 37, 233–238.
- Carini, P., Campbell, E. O., Morr e, J., Sa udo-Wilhelmy, S. a, Thrash, J. C., Bennett, S. E., ... Giovannoni, S. J. (2014). Discovery of a SAR11 growth requirement for thiamin’s pyrimidine precursor and its distribution in the Sargasso Sea. *The ISME Journal*, 8, 1727–38.
- Carlson, C. A, Giovannoni, S. J., Hansell, D. a, Goldberg, S. J., Parsons, R., & Vergin, K. (2004). Interactions among dissolved organic carbon, microbial processes, and community structure in the mesopelagic zone of the northwestern Sargasso Sea. *Limnology and Oceanography*, 49(4), 1073–1083.
- Carlson, C. A., & Ducklow, H. W. (1996). Growth of bacterioplankton and consumption of dissolved organic carbon in the Sargasso Sea. *Aquatic Microbial Ecology*, 10, 69–85.
- Carlson, C. A., Ducklow, H. W., & Michaels, A. F. (1994). Annual flux of dissolved organic carbon from the euphotic zone in the northwestern Sargasso Sea. *Nature*, 371, 405–408.
- Carlson, C. A., Hansell, D. a., Nelson, N. B., Siegel, D. a., Smethie, W. M., Khatiwala, S., ... Halewood, E. (2010). Dissolved organic carbon export and subsequent remineralization in the mesopelagic and bathypelagic realms of the North Atlantic basin. *Deep-Sea Research Part II: Topical Studies in Oceanography*, 57(16), 1433–1445.
- Chevreur, B., Wetter, T., & Suhai, S. (1999). Genome Sequence Assembly Using Trace Signals and Additional Sequence Information. *Computer Science and Biology: Proceedings of the German Conference on Bioinformatics (GCB)*, (1995), 45–56.
- Cho, B. C., & Azam, F. (1988). Major role of bacteria in biogeochemical fluxes in the ocean’s interior. *Nature*, 332, 441–443.

- Cock, P. J. a, Grüning, B. a, Paszkiewicz, K., & Pritchard, L. (2013). Galaxy tools and workflows for sequence analysis with applications in molecular plant pathology. *PeerJ*, *1*, e167.
- Cole, J., Findlay, S., & Pace, M. (1988). Bacterial production in fresh and saltwater ecosystems: a cross-system overview. *Marine Ecology Progress Series*, *43*, 1–10.
- Cuccuru, G., Orsini, M., Pinna, A., Sbardellati, A., Soranzo, N., Travaglione, A., ... Fotia, G. (2014). Orione, a web-based framework for NGS analysis in microbiology. *Bioinformatics*, *30*(13), 1928–1929.
- DeBoy, R. T., Mongodin, E. F., Fouts, D. E., Tailford, L. E., Khouri, H., Emerson, J. B., ... Nelson, K. E. (2008). Insights into plant cell wall degradation from the genome Sequence of the soil bacterium *Cellvibrio japonicus*. *Journal of Bacteriology*, *190*(15), 5455–5463.
- Ducklow, H. W., Steinberg, D. K., & Buesseler, K. O. (2001). Upper ocean carbon export and the biological pump. *Oceanography*, *14*(4), 50–58.
- Eddy, S. R. (2009). A new generation of homology search tools based on probabilistic inference. *Genome Informatics. International Conference on Genome Informatics*, *23*, 205–211.
- Emerson, S., Quay, P., Karl, D. M., Winn, C., & Tupas, L. M. (1997). Experimental determination of the organic carbon flux from open-ocean surface waters. *Nature*, *389*, 951–954.
- Felsenstein, J. (2010). Confidence limits on phylogenies: an approach using the bootstrap. *Evolution*, *39*(4), 783–791.
- Frith, M. C., Hamada, M., & Horton, P. (2010). Parameters for accurate genome alignment. *BMC Bioinformatics*, *11*, 80.
- Gebhard, S., & Cook, G. M. (2008). Differential regulation of high-affinity phosphate transport systems of *Mycobacterium smegmatis*: Identification of PhnF, a repressor of the phnDCE operon. *Journal of Bacteriology*, *190*(4), 1335–1343.
- Giovannoni, S. J., Tripp, H. J., Givan, S., Podar, M., Vergin, K. L., Baptista, D., ... Mathur, E. J. (2005). Genome streamlining in a cosmopolitan oceanic bacterium. *Science*, *309*, 1242–1245.
- Hampel, F. R. (1974). The influence curve and its role in robust estimation. *Journal of the American Statistical Association*, *69*(346), 383–393.
- Huggett, M. J., Hayakawa, D. H., & Rappé, M. S. (2012). Genome sequence of strain HIMB624, a cultured representative from the OM43 clade of marine *Betaproteobacteria*. *Standards in Genomic Sciences*, *6*, 11–20.

- Kanehisa, M., & Goto, S. (2000). KEGG: Kyoto Encyclopaedia of Genes and Genomes. *Nucleic Acids Research*, 28(1), 27–30.
- Kawasaki, S., Moriguchi, R., Sekiya, K., Nakai, T., Ono, E., Kume, K., & Kawahara, K. (1994). The cell envelope structure of the lipopolysaccharide-lacking gram-negative bacterium *Sphingomonas paucimobilis*. *Journal of Bacteriology*, 176(2), 284–290.
- Kielbasa, S. M., Wan, R., Sato, K., Horton, P., & Frith, M. C. (2011). Adaptive seeds tame genomic sequence comparison. *Genome Research*, 21, 487–93.
- Kimura, M. (1980). A simple method for estimating evolutionary rates of base substitutions through comparative studies of nucleotide sequences. *Journal of Molecular Evolution*, 16(1330), 111–120.
- Kong, F., Zhang, M., Liao, S., Yu, S., Chi, J., & Wei, Z. (2010). Antioxidant activity of polysaccharide-enriched fractions extracted from pulp tissue of litchi chinensis sonn. *Molecules*, 15, 2152–2165.
- Kuipers, B., Van Noort, G. J., Vosjan, J., & Herndl, G. J. (2000). Diel periodicity of bacterioplankton in the euphotic zone of the subtropical Atlantic Ocean. *Marine Ecology Progress Series*, 201, 13–25.
- Lauro, F. M., McDougald, D., Thomas, T., Williams, T. J., Egan, S., Rice, S., ... Cavicchioli, R. (2009). The genomic basis of trophic strategy in marine bacteria. *Proceedings of the National Academy of Sciences of the U.S.A*, 106, 15527–15533.
- Laws, E. a., Letelier, R. M., & Karl, D. M. (2014). Estimating the compensation irradiance in the ocean: The importance of accounting for non-photosynthetic uptake of inorganic carbon. *Deep-Sea Research Part I: Oceanographic Research Papers*, 93, 35–40.
- Lea-Smith, D. J., Biller, S. J., Davey, M. P., Cotton, C. a. R., Perez Sepulveda, B. M., Turchyn, A. V., ... Howe, C. J. (2015). Contribution of cyanobacterial alkane production to the ocean hydrocarbon cycle. *Proceedings of the National Academy of Sciences of the USA*, 13591–13596.
- Lee, S., & Fuhrman, J. A. (1987). Relationships between biovolume and biomass of naturally derived marine bacterioplankton. *Applied and Environmental Microbiology*, 53, 1298–1303.
- Letelier, R. M., Karl, D. M., Abbott, M. R., & Bidigare, R. R. (2004). Light driven seasonal patterns of chlorophyll and nitrate in the lower euphotic zone of the North Pacific Subtropical Gyre. *Limnology and Oceanography*, 49(2), 508–519.
- Leys, C., Ley, C., Klein, O., Bernard, P., & Licata, L. (2013). Detecting outliers: Do not use standard deviation around the mean, use absolute deviation around the median. *Journal of Experimental Social Psychology*, 49, 764–766.

- Ludwig, W. (2004). ARB: a software environment for sequence data. *Nucleic Acids Research*, 32(4), 1363–1371.
- Luo, Y. W., Friedrichs, M. a M., Doney, S. C., Church, M. J., & Ducklow, H. W. (2010). Oceanic heterotrophic bacterial nutrition by semilabile DOM as revealed by data assimilative modeling. *Aquatic Microbial Ecology*, 60, 273–287.
- Martinez, A., Tyson, G. W., & Delong, E. F. (2010). Widespread known and novel phosphonate utilization pathways in marine bacteria revealed by functional screening and metagenomic analyses. *Environmental Microbiology*, 12, 222–238.
- Martínez, A., Ventouras, L. A., Wilson, S. T., Karl, D. M., & DeLong, E. F. (2013). Metatranscriptomic and functional metagenomic analysis of methylphosphonate utilization by marine bacteria. *Frontiers in Microbiology*, 4, 1–18.
- McCarthy, M. D., Hedges, J. I., & Benner, R. (1993). The chemical composition of dissolved organic matter in seawater. *Chemical Geology*, 107, 503–507.
- McCarthy, M., Hedges, J., & Benner, R. (1996). Major biochemical composition of dissolved high molecular weight organic matter in seawater. *Marine Chemistry*, 5, 281–297.
- McGrath, J. W., Chin, J. P., & Quinn, J. P. (2013). Organophosphonates revealed: new insights into the microbial metabolism of ancient molecules. *Nature Reviews. Microbiology*, 11, 412–419.
- Menzel, D. W., & Spaeth, J. P. (1962). Occurrence of vitamin B12 in the Sargasso Sea. *Limnology and Oceanography*, 7, 151–154.
- Metcalf, W. W., & Wanner, B. L. (1993a). Evidence for a fourteen-gene, phnC to phnP locus for phosphonate metabolism in *Escherichia coli*. *Gene*, 129, 27–32.
- Metcalf, W. W., & Wanner, B. L. (1993b). Mutational analysis of an *Escherichia coli* fourteen-gene operon for phosphonate degradation, using TnpA' elements. *Journal of Bacteriology*, 175(11), 3430–3442.
- Morton, C. J., & Campbell, I. D. (1994). SH3 domains. Molecular “Velcro”. *Current Biology : CB*, 4(7), 615–617.
- Munoz, R., Yarza, P., Ludwig, W., Euzéby, J., Amann, R., Schleifer, K.-H., ... Rosselló-Móra, R. (2011). Release LTPs104 of the All-Species Living Tree. *Systematic and Applied Microbiology*, 34, 169–70.
- Musacchio, A., Gibson, T., Lehto, V.-P., & Saraste, M. (1992). SH3 — an abundant protein domain in search of a function. *FEBS Letters*, 307(1), 55–61.

- Ottesen, E. a, Young, C. R., Gifford, S. M., Eppley, J. M., Marin, R., Schuster, S. C., ... DeLong, E. F. (2014). Multispecies diel transcriptional oscillations in open ocean heterotrophic bacterial assemblages. *Science*, *345*, 207–12.
- Pruesse, E., Peplies, J., & Glockner, F. O. (2012). SINA: Accurate high-throughput multiple sequence alignment of ribosomal RNA genes. *Bioinformatics*, *28*(14), 1823–1829.
- Repeta, D. J., & Aluwihare, L. I. (2006). Radiocarbon analysis of neutral sugars in high-molecular-weight dissolved organic carbon: Implications for organic carbon cycling. *Limnology and Oceanography*, *51*(2), 1045–1053.
- Saitou, N., & Nei, M. (1987). The neighbor-joining method: A new method for reconstructing phylogenetic trees. *Molecular Biology and Evolution*, *4*(4), 406–425.
- Sarmiento, J. L., & Gruber, N. (2006). Organic Matter Production. *Ocean Biogeochemical Dynamic*. Chapter 4 (pp. 102–172). Princeton University Press, Princeton, NJ.
- Seemann, T. (2014). Prokka: rapid prokaryotic genome annotation. *Bioinformatics*, *30*(14), 2068–2069.
- Singer, E., Webb, E. a., Nelson, W. C., Heidelberg, J. F., Ivanova, N., Pati, A., & Edwards, K. J. (2011). Genomic potential of *Marinobacter aquaeolei*, a biogeochemical “Opportunitroph.” *Applied and Environmental Microbiology*, *77*(8), 2763–2771.
- Sosa, O. a, Gifford, S. M., Repeta, D. J., & DeLong, E. F. (2015). High molecular weight dissolved organic matter enrichment selects for methylophiles in dilution to extinction cultures. *The ISME Journal*, *9*, 2725–2739.
- Swartz, T. E., Tseng, T., Frederickson, M. A., Paris, G., Comerci, D. J., Rajashekara, G., ... Bogomolni, R. A. (2005). *Science*, *317*, 1090–1093.
- Thingstad, T. F., Hagström, A., & Rassoulzadegan, F. (1997). Accumulation of degradable DOC in surface waters: Is it caused by a malfunctioning microbial loop? *Limnology and Oceanography*, *42*, 398–404.
- Wang, H., Liu, Y. M., Qi, Z. M., Wang, S. Y., Liu, S. X., Li, X., ... Xia, X. C. (2013). An overview on natural polysaccharides with antioxidant properties. *Current Medicinal Chemistry*, *20*, 2899–2913.
- Weiner, R. M., Taylor, L. E., Henrissat, B., Hauser, L., Land, M., Coutinho, P. M., ... Hutcheson, S. (2008). Complete genome sequence of the complex carbohydrate-degrading marine bacterium, *Saccharophagus degradans* strain 2-40T. *PLoS Genetics*, *4*(5), e1000087.
- White, A. K., & Metcalf, W. W. (2007). Microbial metabolism of reduced phosphorus compounds. *Annual Review of Microbiology*, *61*, 379–400.

- Williams, K. P., Gillespie, J. J., Sobral, B. W. S., Nordberg, E. K., Snyder, E. E., Shallom, J. M., & Dickerman, a. W. (2010). Phylogeny of Gammaproteobacteria. *Journal of Bacteriology*, *192*(9), 2305–2314.
- Yin, Y., Mao, X., Yang, J., Chen, X., Mao, F., & Xu, Y. (2012). dbCAN: a web resource for automated carbohydrate-active enzyme annotation. *Nucleic Acids Research*, *40*, W445–51.

CHAPTER FOUR

Production of hydrocarbon gases during bacterial degradation of covalently-bound phosphonates associated with marine high-molecular-weight dissolved organic matter

Abstract

The decomposition of dissolved organic matter (DOM), predominantly a microbial-mediated function, is fundamental to the cycling of bio-elements like carbon, nitrogen, and phosphorus and to the sustainability of primary productivity. Yet the underlying mechanisms – the microbial taxa and the enzymatic and biochemical reactions carried on by them – driving this important ecological process remain greatly unexplored in the ocean environment. In this chapter I make use of two putative high-molecular-weight (HMW) DOM-degraders, the gamma-proteobacterium *Pseudomonas stutzeri* strain HI00D01 and the alpha-proteobacterium *Sulfitobacter* strain HI0054 (described in Chapter 3), to investigate the bio-availability of alkylphosphonates – organophosphorus molecules with a direct carbon-phosphorus (C-P) bond – associated with marine HMW DOM polysaccharides. In addition to providing a source of phosphorus to microorganisms in the oligotrophic ocean, the degradation of HMW DOM phosphonates like methyl phosphonate may be an important source of methane and other hydrocarbons to the environment. Using batch culture experiments, I demonstrate that both isolates can produce methane and ethylene in the presence of HMW DOM polysaccharides under aerobic conditions. To demonstrate the implication of the phosphonate degradation pathway – namely the C-P lyase complex – in the production of hydrocarbons, the same experiment was performed with a mutant strain of *P. stutzeri* HI00D01 unable to metabolize simple phosphonate compounds. As expected, the mutant strain did not produce hydrocarbons when exposed to the HMW polysaccharide amendment. These observations support the hypothesis that HMW DOM-bound methyl phosphonates may be a significant source of methane in the water column, and that bacterial degradation of these compounds likely contributes to the subsurface methane maxima observed throughout the world's oceans.

Introduction

Phosphorus is often a scarce nutrient in the ocean limiting marine productivity especially in the oligotrophic gyres (Karl et al., 2001; Moore et al., 2013). The organic phosphorus stored in DOM (DOP) dominates the total dissolved phosphorus inventory (~75-80%) in subtropical surface waters (Karl et al., 2001) and is of special importance for microbial communities when inorganic phosphorus is scarce or unavailable (Björkman & Karl, 2003). The high-molecular-weight (HMW) fraction of DOM contains two main forms of organic phosphorus: phosphate esters (75%) and phosphonates (25%), the latter comprising approximately 10% of total DOP in open-ocean surface waters (Clark, Ingall, & Benner, 1999; Kolowitz et al., 2001). Unlike phosphate esters, phosphonates are often undetectable in particulate organic matter (POM) and plankton biomass (Kolowitz et al., 2001; Paytan, Cade-Menun, McLaughlin, & Faul, 2003) and yet they seem to accumulate in HMW DOP. The mechanism that leads to the prevalence of phosphonates in DOM is unclear, although most likely this occurs because these are more resistant to microbial degradation than the more labile phosphate esters. Nevertheless, both phosphonates and phosphate esters in DOP seem to be utilized by microorganisms at similar rates given that their relative concentrations remain stable as HMW DOP decreases with depth (Clark et al., 1999; Clark & Ingall, 1998; Kolowitz et al., 2001).

This chapter investigates the microbial degradation of phosphonates, molecules that contain a chemically stable C-P bond highly resistant to hydrolysis and bacterial degradation. Several bacterial enzymatic pathways are known to degrade phosphonates including the phosphonopyruvate hydrolase pathway, the phosphonoacetate hydrolase pathway, the phosphonoacetaldehyde hydrolase (phosphonatase) pathway, and the carbon-phosphorus lyase

(C-P lyase) pathway and are summarized by (White & Metcalf, 2007). The most widespread is the C-P lyase pathway and is the focus of this chapter.

The C-P lyase is an enzyme complex encoded by a 14-gene operon in *Escherichia coli* (*phnCDEFGHIJKLMNOP*) that catalyzes the degradation of a broad range of alkylphosphonates, such as methyl phosphonate and 2-aminoethyl phosphonate, to make the phosphorus available for cell uptake (Metcalf & Wanner, 1993a, 1993b). Cleavage of the C-P phosphonate bond results in the concomitant release of the alkyl group which may or may not be used as additional carbon source (White & Metcalf, 2007). Marine metagenomes and sequenced bacterial genomes, mainly of the *Proteobacteria* lineage, are known to encode several modifications of the C-P lyase gene operon (Martinez, Ventouras, Wilson, Karl, & Delong, 2013; Villarreal-Chiu, 2012) indicating that phosphonates are bio-available in the ocean. In accordance with these observations, microcosm studies have indicate that bacterioplankton groups including *Rhodobacterales*, *Vibrio*, and *Alteromonadales* express the C-P lyase pathway under to alleviate phosphorus limitation, in this case induced by the addition of methyl phosphonate and excess glucose (Martínez et al., 2013), producing phosphate for cell uptake and releasing methane into the water column. The microbial degradation of methyl phosphonates has gathered special interest because it has been implicated in the production of the sub-surface methane maxima observed across the ocean (Karl, 2014; Karl et al., 2008). Marine surface waters typically observe methane supersaturation with respect to the atmospheric equilibrium concentration and are considered a source of methane to the atmosphere (Fig. 4.1; Holmes, Sansone, Rust, & Popp, 2000; Tilbrook & Karl, 1995). Holmes et al. (2000) calculate a mean flux rate of oceanic methane to the atmosphere of $1.6 \mu\text{mol m}^{-2} \text{d}^{-1}$ for the North Pacific which translates to 0.25 Tg y^{-1} considering the area encompassed by the North Pacific subtropical gyre, or about 2.5% of the

annual natural ocean and freshwater methane emissions (10 Tg y^{-1} ; Reeburgh, 2007), although the latter represents only 2% of the global methane budget (Reeburgh, 2007). In addition, the methane flux to the atmosphere can increase with the deepening of the mixed layer (high wind speeds and storms) bringing methane up from the sub-surface maxima (Holmes et al., 2000). Nevertheless, in the past the source of this methane remained obscure.

In the North Pacific, surface ocean methane concentrations (upper 300 meters) are consistently 37-67% (up to 90% in the North Atlantic) higher than the expected atmospheric equilibrium concentrations over the course of the year (Holmes et al., 2000). In addition, methane sampled in the upper 300 meters is isotopically heavier than the atmospheric methane sampled 5 meters above the surface (-45.6‰ and -47.35‰ , respectively). These two pieces of evidence indicate that the methane detected in the water column has an in situ biogenic origin albeit different than the well-known anaerobic methanogenesis which would result in more isotopically depleted methane. The methyl phosphonate associated with HMW DOM ($\sim 10\%$ of HMW DOP; Daniel Repeta, personal communications) may be sufficient to account for the methane observed in the upper 300 meters of Station ALOHA and in the sub-surface methane maxima concentrations. Assuming a HMW DOC concentration of $24 \mu\text{M}$ (30% of total DOC), a C:P atomic ratio of 300:1 (Benner, 2002), and that all methyl phosphonate is bio-available, HMW DOM would produce 8 nM methane, enough to account for the sub-surface methane maxima concentrations. However, the microorganisms and biochemical pathways responsible for HMW DOM phosphonate degradation, and the environmental conditions under which HMW DOM methyl phosphonate degradation can aggravate the greenhouse effect remain to be determined.

The bacterium *Pelagibacteriales* sp. strain HTCC7211, a representative of the widespread and abundant SAR11 clade, has been shown to produce methane from the degradation of methyl phosphonate and is thought to be an important contributor to oceanic methane (Carini et al., 2014). However, methyl phosphonate has not been reported to exist in free form in marine waters. HMW DOM is most likely the major source of methyl phosphonates and other alkylphosphonates in the water column (Repeta et al., unpublished) though direct evidence of bacterial utilization of HMW DOM-bound phosphonates via the C-P lyase pathway is lacking.

In this work I investigate the bioavailability of HMW DOM phosphonates using batch cultures of the bacteria *Pseudomonas stutzeri* strain HI00D01 and *Sulfitobacter* strain HI0054. In addition, I explore inorganic phosphate as a possible regulator of phosphonate degradation and test the direct involvement of the C-P lyase in degradation of HMW DOM phosphonates via a genetic mutagenesis approach. Because both isolates were obtained from HMW DOM-enriched cultures and their genomes encode complete operons of the C-P lyase pathway, I hypothesized that they would be able to degrade alkylphosphonates present in HMW DOM.

Materials and Methods

Collection of HMW DOM and isolation of neutral polysaccharides

HMW DOM sampling and isolation of the HMW neutral polysaccharide fraction (F1) by chromatography was done as describes in Chapter 3 Materials and Methods (Ultrafiltration and Isolation of HMW DOM neutral polysaccharides, pp. 98-100).

Culture purity of *Pseudomonas* and *Sulfitobacter* isolates

P. stutzeri isolate HI00D01 was obtained from Hawaii Experiment I dilution-to-extinction samples treated with HMW DOM (7x DOC treatments; see Tables 3.1 and 3.2 in Chapter 3). *P.*

stutzeri strain HI00D01 was purified by re-streaking single colonies in marine agar plates. Smooth and rough colony morphologies were observed. Sanger sequencing of the SSU rRNA gene indicated 100% identity between all morphotypes. Cryostocks of smooth and rough colonies were prepared in 20% glycerol and stored frozen at -80°C. A rough clone (RA1) was used for subsequent experiments. *Sulfitobacter* isolate HI0054 was obtained from Hawaii Experiment II dilution-to-extinction samples treated with HMW DOM (4x DOC; see Tables 3.1 and 3.2 in Chapter 3). HI0054 was also purified on marine agar plates to obtain single colonies for culture-based experiments. The SSU rRNA gene sequence obtained from draft genomes of HI00D01 and HI0054 indicated that both primary cultures consisted of a single taxon (Chapter 3 Table 3.1).

Growth of *P. stutzeri* strain HI00D01 on marine HMW DOM polysaccharides

P. stutzeri HI00D01 was tested for higher cell yields under increasing concentrations of HMW DOM neutral polysaccharides (F1). HMW DOM was collected by ultrafiltration at Station ALOHA (KM2007). HMW DOM F1 polysaccharides (sample 07182014) consisted of neutral sugars eluted in ultrapure water after anion exchange chromatography of bulk HMW DOM. HMW DOM F1 samples were filter-sterilized through a 0.22 μm Supor membrane and stored frozen at -20°C before amending culturing media. The culturing media consisted of MOPS minimal media (Appendix II Table A1) supplemented with 50 μM phosphate (K_2HPO_4). Five DOC treatments were prepared: 1) No DOC, the control treatment consisting of the base media with no HMW DOM added, 2) ~ 100 μM DOC (4 $\mu\text{g ml}^{-1}$ of HMW DOM F1), 3) ~ 500 μM DOC (20 $\mu\text{g ml}^{-1}$ of HMW DOM F1), 4) ~ 1 mM DOC (40 $\mu\text{g ml}^{-1}$ of HMW DOM F1), 5) ~ 2 mM DOC (80 $\mu\text{g ml}^{-1}$ of HMW DOM F1). Growth was determined daily from triplicate cultures

by enumerating DNA-stained cells (SYBR Green I) on an Attune cytometer system (Applied Biosystems, Foster City, CA, USA).

Transposon mutagenesis of *P. stutzeri* H00D01 and identification of strains unable to metabolize phosphonates

Transposon mutagenesis and library preparation

Single colonies of *P. stutzeri* strain H100D01 clone RA1 were grown on 5 ml of marine broth and incubated at 30°C and 255 rpm overnight. Cells were harvested by centrifugation at 2000Xg for 5 min, the supernatant was discarded, and cells were rinsed with 10 mL of chilled 10% glycerol solution. Centrifugation and wash steps were repeated after which cells were concentrated in 100 µl of 10% glycerol. The cell suspension was transformed with 2 µL of EZ-Tn5 <KAN-2> Tnp Transposome (20 ng µl⁻¹ stock, Epicentre Biotechnologies, Madison, WI) by electroporation in a 2 mm cuvette. Electroporation parameters were 12.5 kV cm⁻¹, 200 Ω, and 25 µF. To facilitate cell outgrowth and expression of the kanamycin antibiotic marker, immediately after electroporation 0.4 ml of marine broth were added to the cuvette to dilute cells and incubated at 30°C for 30 minutes. 50 µl aliquots of the dilute cell suspension were plated on large square dish plates with marine agar media containing kanamycin (50 µg ml⁻¹) to select for transposon insertion mutants. Plates were incubated at 30°C overnight. A robot Q-pix Genetix was used to pick colonies from selective media to inoculate a total of sixteen 96 well microtiter plates containing marine broth and kanamycin. Plates were incubated at 30°C for 2 days and cryopreserved with 20% glycerol.

Screening mutants for strains deficient in metabolizing phosphonates

To identify strains that can no longer use phosphonates, the transposon mutant library was screened in media supplemented with 2-aminoethyl phosphonate (99% purity; Sigma-

Aldrich 268674) as the sole phosphorus source and 0.2% glucose. The screening media consisted of Hawaii seawater amended with 1 mM sodium nitrate and ammonium chloride, and kanamycin ($50 \mu\text{g ml}^{-1}$). To assess mutant viability, the library was screened in parallel on marine agar plates with kanamycin ($50 \mu\text{g ml}^{-1}$). A total of 1344 transposon mutants were spotted onto the screening media using a Q-pix Genetix microbial colony picker (Molecular Devices, Sunnyvale, CA). Marine agar plates incubated overnight at 30°C while the phosphonate screening media incubated for two days to allow bacterial colonies to grow further. Samples were stored at 4°C to halt growth.

Candidate mutants deficient in phosphonate degradation

To ensure that mutants identified in the screen were not deficient in metabolizing glucose rather than phosphonates, growth was tested in agar media consisting of MOPS minimal media supplemented with glucose (0.1%), inorganic phosphate (1 mM K_2HPO_4) and kanamycin ($50 \mu\text{g ml}^{-1}$). To confirm which glucose-viable mutants were deficient in metabolizing phosphonates (phn^-) growth was tested on MOPS agar media with 0.2% glucose as the sole carbon source and 2-aminoethyl phosphonate (1 mM) or inorganic phosphate (1.32 mM) as the sole phosphorus source. To identify mutation sites, genomic DNA from candidate phn^- mutant strains was sequenced on a MiSeq system using a Nextera XT 96 DNA sample preparation kit (Illumina, San Diego, CA, USA) to obtain indexed paired-end libraries and a 2x250 nt paired-end MiSeq run. Paired reads were assembled into contigs using CLC Genomics Workbench (CLC bio, Aarhus, Denmark) de novo assembler with automatic word and bubble sizes, a minimum contig length of 200, insertion and deletion costs set to 3, mismatch cost set to 2, length fraction set to 0.5 and the similarity fraction set to 0.8. The transposon insertion site in each mutant was located in the assembled contigs using transposon-specific forward and reverse primer sequences as well as the

mosaic end sequences (5'-AGATGTGTATAAGAGACAG-3') marking the end of the transposon. To identify which genes had been disrupted by the transposon, the sequences flanking the transposon (300 nucleotides upstream and downstream) were compared to PROKKA-predicted genes in the draft genome of HI00D01 using BLASTn.

Evaluation of *P. stutzeri* HI00D01 growth on alkylphosphonates as a source of phosphorus

P. stutzeri HI00D01 and mutant strain 14-E11 (transposon insertion mutant) were tested for growth on several alkylphosphonates as sole phosphorus source in agarose media containing 0.2% glucose. The media was prepared with MOPS minimal media supplied without inorganic phosphate (M2101; Teknova, Hollister, CA; see Table A1 in Appendix II) in ultrapure water and 1.7% SeaKem LE agarose (Lonza, Rockland, ME), a high-purity agarose to minimized background growth. Methyl phosphonate (98%; Sigma-Aldrich 289868), aminomethyl phosphonate (98%; Sigma-Aldrich 324817), ethyl phosphonate (98%; Sigma-Aldrich 289876), 2-aminoethyl phosphonate (99%; Sigma-Aldrich 268674), 2-phosphonopropionate (98%; VWR AAAL11588-04), 3-phosphonopropionate (94%; Sigma-Aldrich 228559), 3-aminopropyl phosphonate (99%; Sigma-Aldrich 268615), and 2-phosphonobutyric acid (97%; VWR AAAL12036-04) were tested at a concentration of 100 mM. The test media for strain 14-E11 was supplemented with kanamycin (50 $\mu\text{g ml}^{-1}$) as selection marker. Phosphonates, glucose, and antibiotic were added to the media after autoclaving. Plates incubated in darkness at 27°C.

Detection of hydrocarbon gases in batch cultures amended with phosphonates or HMW DOM polysaccharides

Overview

Batch cultures of *P. stutzeri* HI00D01 and the *Sulfitobacter* HI0054 were prepared in sealed glass bottles to test for the accumulation of hydrocarbon gases predicted to be released as

byproducts during phosphonate degradation. To stimulate utilization of phosphonates, the cultures were grown in a phosphorus-limited medium with excess glucose (*P. stutzeri* HI00D01) or glycerol (*Sulfitobacter* HI0054). Dissolved methane, ethylene, and ethane concentrations were measured by gas chromatography using a gas stripping and cryo-trap concentration method, described in detail by Tilbrook & Karl (1995). Dissolved gases were purged by bubbling a known volume of water with ultra high purity helium. The volume of sample purged, determined by weight, was approximately 50 ml. During the gas stripping procedure, the gas stream passed through Drierite® and Ascarite® columns to remove water vapor and carbon dioxide, and then the gases were concentrated in a 80-100 mesh Porapak Q® trap cooled in liquid N₂. Once the gas extraction was completed (12 minutes), the trap was heated with boiling water (~90°C) and the concentrated gases were released and injected into the gas chromatograph (Agilent 7980A). The gases were separated in a capillary column and analyzed with a flame ionization detector (FID). The FID was calibrated by injecting different size loops of a gas mixture standard containing 10 ppm of methane and ethylene in pure N₂ (Scott-Marrin). The loops were injected into the purging chamber and concentrated into the Porapak Q® trap following the same procedure as for the water samples. Tilbrook & Karl (1995) reported a detection limit for methane analysis of approximately 0.05 nM with ±3% precision of repeated analyses of samples.

Initial screen

Cultures of *P. stutzeri* HI00D01 grown in the presence of phosphonates and excess glucose were screened for production of hydrocarbons. Treatments consisted of MOPS minimal media containing 0.2% glucose (11 mM) and supplemented with phosphonates [104 µM methyl phosphonate, aminoethyl phosphonate, 2-aminoethyl phosphonate, or 3-aminopropyl phosphonate], or HMW DOM F1 polysaccharides (40 µg ml⁻¹ or ~1 mM carbon) as phosphorus

source. Assuming a HMW DOM C:P ratio of 300:1 (Kolowitz et al., 2001), the addition of HMW DOM F1 polysaccharides contributed approximately 4-5 μM phosphorus to the media of which 20% (0.8-0.9 μM) is estimated to consist of phosphonates. HMW DOM contains 1 mol of methyl phosphonate per 4000 mols of carbon (HMW DOC:Methyl phosphonate = 4000:1; Daniel Repeta, personal communication). Therefore, the addition of HMW DOM polysaccharides (1 mM DOC) contained ~ 250 nM methyl phosphonate. Treatments were prepared with an initial inoculum of 4000 cells ml^{-1} (pre-grown in 0.2% glucose and no additional phosphorus) in 72-ml glass bottles with no head space and sealed with Teflon-lined butyl septa to collect trace gases. Samples incubated at 26°C under a 12 hour diurnal cycle for 3 days. The screen included a culture of the *P. stutzeri* HI00D01 phn^- mutant clone 14-E11 as negative control. The media was supplemented with kanamycin (50 $\mu\text{g ml}^{-1}$; Table 4.1) to ensure selection of the transposon insertion mutant. At the end of the incubation period, culture samples were injected and sparged with helium for 12 minutes. Gases were analyzed on a gas chromatograph (Agilent 7890A) to quantify methane, ethylene, and nitrous oxide. Sub-samples from each experiment were also drawn for flow cytometry analysis to determine the final cell yields.

Inhibition of hydrocarbon production from phosphonates by inorganic phosphate additions in batch cultures of *P. stutzeri* strain HI00D01

A second batch culture experiment with *P. stutzeri* HI00D01 was prepared in which inorganic phosphate was added to determine if phosphonate degradation was inhibited. In addition, the growth conditions were modified to prevent anoxia and ensure phosphorus limitation. To test this hypothesis, cultures of the wildtype and phn^- mutant strain were pre-grown in 0.85 liters of MOPS minimal media supplemented with 100 μM carbon (glucose) and

no additional phosphorus to induce phosphorus starvation. To ensure that cultures were phosphorus-limited, sub-samples of cultures in stationary phase were stimulated with inorganic phosphate and glucose. Addition of 1 μM , 10 μM phosphate, or 100 μM carbon (glucose) alone did not stimulate growth. The cultures required both phosphate and glucose in order to continue growing. Once such conditions were established, the 0.85-liter cultures were brought up to 1 liter and were amended with an additional ~ 100 μM carbon in the form of glucose. The resulting media thus contained 100-200 μM carbon. In addition, samples were bubbled with filtered air to replenish oxygen. To prepare treatments, phosphorus-limited cultures were aliquoted into 40 ml gas-tight glass bottles sealed with Teflon-lined butyl septa and no head space. Treatments consisted of duplicate samples amended with methyl phosphonate (1 μM ; Sigma-Aldrich 289868), ethyl phosphonate (1 μM ; Sigma-Aldrich 289876), or HMW DOM polysaccharides (0.095 mg ml^{-1} or ~ 3.2 mM carbon) with and without additional inorganic phosphate (1 μM) to test for inhibition of the phosphonate utilization. Assuming a HMW DOC:Methyl phosphonate ratio of 4000:1, the addition of HMW DOM polysaccharides resulted in ~ 800 nM methyl phosphonate. The control treatment consisted of samples with no added phosphorus source. Samples of the wildtype and mutant cultures were killed with mercuric chloride at the beginning of the incubation to determine the initial concentration of hydrocarbon gases. Additionally, single replicate samples from each treatment were prepared to estimate the total oxygen consumption during batch culture growth from changes in oxygen/argon (O_2/Ar) ratios measured by membrane-inlet mass spectrometry (MIMS) as described by Ferrón, Wilson, Martínez-García, Quay, & Karl (2015) and Kana et al. (1994) using an O_2 and Ar solubilities at salinity of 0 ppt and a temperature of 26°C (253.55 $\mu\text{mol L}^{-1}$ and 12.44 $\mu\text{mol L}^{-1}$, respectively). The initial

oxygen concentration was also determined by MIMS from killed cultures at the beginning of the incubation. The oxygen drawdown was obtained from equation 3.1

$$\Delta O_2 = \frac{\frac{O_2}{Ar_{final}}}{\frac{O_2}{Ar_{initial}}} \times [O_2]_{initial} \quad [3.1]$$

Treatment samples incubated for 1.75 days at 26°C under a diurnal 12 hour cycle. To determine culture growth, total cell counts were obtained by flow cytometry analysis of DNA-stained cells (SYBR Green I) at the beginning and end of the incubation period.

Degradation of two representative HMW DOM phosphonates (methyl phosphonate and 2-hydroxyethyl phosphonate) by *P. stutzeri* HI00D01

Methyl phosphonate and 2-hydroxyethyl phosphonate are the major phosphonate residues identified in HMW DOM polysaccharides (Repeta et al., unpublished; personal communication). In this experiment, three biological replicates of wildtype and *phn*⁻ mutant strain 14-E11 *P. stutzeri* HI00D01 cultures were pre-grown (starting from a single colony) under phosphorus-limited conditions. Cultures were first grown in 200 ml of MOPS minimal media containing 100 μM carbon (glucose) with no additional phosphate. Kanamycin (50 μg ml⁻¹) was added to mutant cultures. The cultures incubated for 5 days at 27°C and 200 rpm shaking after which samples were diluted two-fold with sterile medium and supplemented with an additional 200 μM carbon (glucose). Samples were bubbled with filtered air to replenish oxygen. The phosphorus-limited cultures were then aliquoted in 72-ml glass bottles containing phosphonates or HMW DOM polysaccharide amendments. Triplicate samples of the wildtype and mutant cultures were killed with mercuric chloride at the beginning of the incubation to obtain the initial gas concentrations. Treatments consisted of triplicate samples amended with methyl phosphonate

(1 μM ; Sigma-Aldrich 289868), 2-hydroxyethyl phosphonate (1 μM ; BOC 22987-21-9), or HMW DOM polysaccharides (0.295 mg ml⁻¹ or \sim 9.8 mM carbon). Assuming a HMW DOC:Methyl phosphonate ratio of 4000:1 (Daniel Repeta, personal communication), the addition of HMW DOM resulted in \sim 2.45 μM methyl phosphonate. The control treatment consisted of samples with no additional phosphorus source. Samples incubated at 26°C in darkness for 3.8 days after which they were killed with mercuric chloride and stored in darkness until analysis.

Test for production of hydrocarbon gases using batch cultures of *Sulfitobacter* strain HI0054 amended with phosphonates and HMW DOM polysaccharides

Strain HI0054 was pre-grown under phosphorus-limited conditions. The media consisted of Hawaii seawater sterilized by tangential flow ultrafiltration and amended with 200 μM carbon (glycerol), 40 μM ammonium chloride, 0.5 μM phosphate (K_2HPO_4), 1 μM ferric chloride, trace metals and vitamins (Appendix table). Upon reaching stationary phase, the culture was supplemented with an additional 20 μM ammonium chloride and 100 μM carbon (glycerol), and bubbled with filtered air to replenish oxygen. Treatments consisted of duplicate samples of phosphorus-limited culture amended with inorganic phosphate (2 μM), methyl phosphonate (2 μM), 2-hydroxyethyl phosphonate (2 μM), 2-phosphonopropionate (2 μM), or HMW DOM polysaccharides (2.6 mM carbon). Assuming a HMW DOC:Methyl phosphonate ratio of 4000:1, the addition of HMW DOM resulted in \sim 650 nM methyl phosphonate. Samples incubated at 26°C under a diurnal 12 hour cycle for 2 days. Total cell counts were obtained at the end of the incubation period by flow cytometry analysis of DNA-stained cells (SYBR Green I).

Metagenomic screening of the C-P lyase pathway genes

The C-P lyase operon genes (Appendix I Table A8) identified in the draft-genomes of Hawaii open-ocean *Rhodobacteraceae*, *Oceanospirillaceae*, and *Pseudomonas* (strain HI00D01)

isolates (Chapter 3) and in metagenomes obtained from methyl phosphonate-amended microcosms (Martínez et al., 2013; Appendix I Table A7) were searched by BLAST analysis in the Ocean Microbial Reference Gene Catalogue (OM-RGC) and associated metagenomes obtained during the *Tara* Oceans Expedition (Sunagawa et al., 2015). OM-RGC hits matching predicted homologs of phosphonate transporter genes *phnCDE*, the regulatory gene *phnF*, and the catalytic genes *phnGHIJKLM* required for phosphonate degradation (Metcalf & Wanner, 1993b) were identified with a bitscore ≥ 60 and percent identity ≥ 60 . Each of these 11 C-P lyase gene homologs was deemed present at a given *Tara* Ocean metagenome if the OM-RGC hits had a combined relative coverage $\geq 10^{-6}$. Because the catalogue contains 40 million genes, this cutoff value is $40\times$ greater than the relative coverage of an individual OM-RGC reference gene (1 in 40 million or 2.5×10^{-8}) and thus represents a strict assessment of their occurrence in the metagenome. Presence of the C-P lyase pathway was determined in a given metagenome if at least 7 of the 11 C-P lyase gene homologs surpassed this threshold. The *Tara* Ocean Expedition station geographic locations and C-P lyase metagenomic screen results were overlaid onto a global map indicating the statistical annual mean of inorganic phosphate concentration in surface waters (<5 meters) as obtained from the World Ocean Atlas 2013 available through the NOAA National Oceanographic Data Center (<https://www.nodc.noaa.gov/OC5/woa13/woa13data.html>). Data downloaded in comma separated values format was visualized using on ArcGIS® ArcMap® version 10.2.2.

Results and Discussion

Introduction of *Pseudomonas stutzeri* HI00D01 and *Sulfitobacter* HI0054 bacteria as suitable model systems to study the degradation of HMW DOM phosphonates

The bacterium *P. stutzeri* strain HI00D01 is a member of the genus *Pseudomonas* sensu stricto which is often found in terrestrial soils, the rhizosphere, and in groundwater, although several strains have been isolated from marine environments including *P. stutzeri*, *P. balearica*, and *P. xanthomarina* (Lalucat, Bennasar, Bosch, García-Valdés, & Palleroni, 2006). In many habitats *P. stutzeri* forms part of the community of microbial decomposers and is often actively involved in the cycling of nitrogen (denitrification, nitrogen fixation), phosphorus (phosphite and hypophosphite oxidation) and sulfur (thiosulfate turnover) (Lalucat et al., 2006). *P. stutzeri* HI00D01 was isolated from a depth of 200 m in Station ALOHA waters 100 km north of Hawaii by dilution-to-extinction culturing method, using a seawater based media supplemented with marine HMW DOM (Chapter 3). Strain HI00D01 primarily forms rough colonies on marine agar and can be grown in natural seawater, marine broth (DIFCO 2216), LB broth, and MOPS minimal media. Most importantly, HI00D01 is able to use naturally occurring marine HMW DOM as a growth substrate, as indicated by a small but measureable increase in growth yield with increasing concentrations of HMW DOM polysaccharides (Fig. 4.2). Inspection of the genome draft of *P. stutzeri* HI00D01 indicated that it contains a complete operon encoding the C-P lyase pathway (Metcalf & Wanner, 1993a, 1993b). The *P. stutzeri* HI00D01 operon nucleotide sequence has a 95% identity to the *phnCDEFGHIJKLMNP* operon of *P. stutzeri* isolate WM88 (Metcalf and Wolfe, 1998). The C-P lyase pathway has been shown to metabolize several alkylphosphonate compounds including methyl phosphonate and 2-aminoethyl phosphonate (White and Metcalf, 2004). These characteristics make *P. stutzeri* HI00D01 a suitable model system to study the decomposition of HMW DOM phosphonates.

In contrast to the genus *Pseudomonas*, the ecology of the genus *Sulfitobacter* remains largely unexplored. However, several representative members of the *Roseobacter* clade

of which the genus *Sulfitobacter* is part of have been well characterized (reviewd in Buchan, González, & Moran, 2005). Members of this lineage have developed important physiological adaptations that have enabled its success in the marine environment including the capability for aerobic anoxygenic photosynthesis, dimethyl sulfoniopropionate (DMSP) degradation, carbon monoxide oxidation, diverse photoheterotrophic strategies, as well as apparent association with diverse phytoplankton groups and the organic matter derived from them, such as carbohydrate exudates and glycolate (Luo & Moran, 2014). *Sulfitobacter* strain HI0054, along with several closely related isolates (>96% SSU rRNA gene sequence similarity), was obtained from HMW DOM-treated dilution-to-extinction cultures prepared with seawater collected from 250 m-depth at Station ALOHA near Hawaii (Appendix I Fig. A2). Strain HI0054 shares 97.8-98.4% SSU rRNA gene sequence similarity with *Sulfitobacter delicatus* strain KMM 3584 (isolated from the starfish *Stellaster equestris* collected at a depth of 100 m in the South China Sea; Ivanova et al., 2004), *Sulfitobacter dubius* strain KMM 3554 (isolated from the sea grass *Zostera marina* collected from a depth of 5-8 m in the Sea of Japan; Ivanova et al., 2004), *Oceanibulbus indolifex* strain HEL-45 (North Sea seawater isolate; Wagner-Döbler et al., 2004) and $\leq 97\%$ similarity with other Roseobacters. Physiological tests indicate that strain HI0054 seems to preferentially utilize amino acids and organic acids like β -hydroxybutyrate (Fig. 3.9 in Chapter 3). HI0054 grows robustly on natural seawater media amended with inorganic nutrients and glycerol as the main carbon source and forms well uniformly round, smooth colonies, creamy in color, when grown in marine agar. The observed growth on β -hydroxybutyrate may be associated with the production of poly- β -hydroxybutyrate since the three aforementioned reference cultured relatives accumulate this compound intracellularly. *Sulfitobacter* strain HI0054 may contribute to DMSP turnover as it carries a homolog of DMSP demethylase (*dmdA*;

80.6% identity with *Roseobacter denitrificans* OCh 114), the enzyme initiating DMSP breakdown. Most relevant to this work, HI0054 encodes a complete C-P lyase pathway for phosphonate degradation similar in arrangement and amino acid sequence (mean \pm s.d. sequence similarity = $72.2 \pm 9.8\%$; n=15) to the operon expressed by a *Rhodobacterales* population enriched in methylphosphonate-amended microcosms (Martínez et al., 2013; Appendix I Table A7). Using a genetically-modified *E. coli* strain, Martínez et al. (2013) showed that the *Rhodobacterales* CP-lyase operon could catalyze the degradation of several alkylphosphonates including methyl phosphonate. These observations along with the growth characteristics of strain HI0054 indicate that it may partially degrade HMW DOM constituents, including alkylphosphonates, making it an appropriate model system for DOM degradation studies.

Growth of *P. stutzeri* HI00D01 on HMW DOM polysaccharides

In seawater based media, *P. stutzeri* HI00D01 observed increased cell yields concomitantly with HMW DOM polysaccharides concentrations, although the increment was not linear (Fig. 4.2). Cultures reached maximum cell densities after 2 days of incubation. The final cell yields of cultures amended with <1 mM DOC were not significantly different from control treatments (Student's T-Test; $H_0: \bar{X}_1 = \bar{X}_2$; P-value < 0.05). The observed cell yields ($1.5\text{-}2 \times 10^5$ cells ml⁻¹) were similar to yields of cultures grown in the presence of bulk HMW DOM (Fig. 3.5 in Chapter 3). The low cell yields observed compared to the total DOC available indicate that only a limited portion of the HMW DOM F1 polysaccharides is bioavailable for *P. stutzeri* HI00D01 growth, at least under the conditions and timescale of these experiments.

Growth of *P. stutzeri* strain HI00D01 on alkylphosphonates as phosphorus source

Strain HI00D01 tested positive for growth on methyl phosphonate, ethyl phosphonate, 2-aminoethyl phosphonate, 2-phosphonopropionate, 3-phosphonopropionate, and 3-aminopropyl

phosphonate as the only available phosphorus source in agarose media containing glucose as sole carbon source (Fig. 4.3). The compounds aminomethyl phosphonate and 2-phosphonobutyric acid did not support growth (Fig. 4.3). The lack of growth on aminomethyl phosphonate is consistent with the absence of *phnO* in the *P. stutzeri* operon, which in *Escherichia coli* is required for catabolism and detoxification of 1-aminoalkylphosphonic acids (Errey & Blanchard, 2006; Hove-Jensen, McSorley, & Zechel, 2012; Martínez et al., 2013). On the other hand, 2-phosphonobutyric acid may fall outside the substrate range of the C-P lyase or phosphonate transporters and thus cannot support growth.

Identification of *P. stutzeri* HI00D01 transposon mutants deficient in phosphonate utilization

A transposon mutant library of the phosphonate-degrading *P. stutzeri* strain HI00D01 (clone RA1) was prepared. The vast majority of the mutants were viable. Only 3 of 1344 mutants obtained failed to grow on marine agar. A total of 28 viable mutants were identified that failed to grow or grew poorly on the phosphonate screening media. Of these 28 viable mutants, only 12 showed clear growth on the glucose medium with inorganic phosphate indicating that the rest seemed to carry mutations that rendered them auxotrophic for other nutrients. Further testing identified 3 strains (14-E11, 15-B9, and 16-A10) that could clearly grow on glucose and inorganic phosphate but failed to metabolize 2-aminoethyl phosphonate. Assembled Illumina reads and BLASTn analysis revealed that strains 14-E11 and 15-B9 had transposon insertions in *phnK* and strain 16-A10 had an insertion in *phnM*, both genes required for catalysis of the C-P cleavage reaction (Metcalf and Wanner, 1993b). Mutations in the other phosphonate genes encoding catalytic function were not identified most likely because of the relatively small size of the library screened (1344 mutants compared to the draft genome's 4676 predicted CDS; see

Appendix I Table A4 and A5). Similarly to the wildtype, these mutants were able to form large colonies when grown in agarose media containing inorganic phosphate and glucose. Further testing confirmed that *P. stutzeri* HI00D01 mutant strain 14-E11 failed to utilize either of the phosphonate compounds for which the wildtype strain tested positive for growth (Fig. 4.3).

Screening for production of hydrocarbon gases from bacterial phosphonate degradation

The first screening experiment revealed the ability of *P. stutzeri* HI00D01 of producing hydrocarbons when grown in the presence of several phosphonates and HMW DOM polysaccharides (Table 4.1). However, hydrocarbon gases were not measurable in batch cultures amended with HMW DOM polysaccharides unless glucose was available, most likely because the labile carbon stimulated bacterial growth until trace inorganic phosphorus became limiting (Table 4.1). As expected, methane was not detected in *P. stutzeri* mutant strain (phn^-) batch cultures amended with methyl phosphonates or HMW DOM polysaccharides. The wildtype culture's methane yields were lower than the methyl phosphonate available. The culture amended with methyl phosphonate (104 μM) only evolved 94.8 nM (0.09%) methane (Table 4.1). While the culture amended with HMW DOM F1 polysaccharides (1 mM DOC) and glucose only evolved 4.6 nM (Table 4.1). HMW DOM contains 1 mol of methyl phosphonate per 4000 mols of carbon (C:Methyl phosphonate ratio = 4000:1; Daniel Repeta, personal communication). Therefore, the methane yield in the HMW DOM-amended cultures was approximately 1.84% of the estimated methyl phosphonate available in HMW DOM polysaccharides (~250 nM). In addition to methane, the wildtype cultures supplemented with HMW DOM polysaccharides and glucose produced comparable amounts of ethylene (Table 4.1) above the background levels of the sterile media. Ethylene gas is most likely produced from the degradation of 2-hydroxyethyl phosphonates, estimated to have comparable abundance as methyl phosphonates in HMW DOM

polysaccharides (Repeta et al. unpublished). As expected, methane was not produced from aminomethyl phosphonate since this phosphonate could not be used as a phosphorus source by the wildtype strain in agarose plates. Cultures did not produce methane, ethylene, or propylene from 2-aminoethyl phosphonate or 3-aminopropyl phosphonate, compounds that the wildtype strain can utilize as phosphorus source (Table 4.1). It may be necessary to use a medium without nitrogen in order to observe the release of methane and ethylene from to stimulate the utilization of these nitrogen-containing phosphonates as a nitrogen source. Alternatively, the C-P bond cleavage of these phosphonates may produce ethylamine and propylamine that were consumed as carbon and nitrogen sources. Small amounts of methane were detected in the control culture grown with glucose and no phosphorus source; although without replicates this measurement cannot be verified as significantly different from the control sterile media. The high concentration of glucose likely made cultures anoxic as suggested by the observed drawdown of nitrous oxide compared to the initial concentration in the sterile medium. Overall, this production of methane and ethylene from HMW DOM polysaccharides by *P. stutzeri* HI00D01 is consistent with previous observations from microcosm experiments prepared with Hawaii oligotrophic seawater (Repeta et al., unpublished).

Inhibition of phosphonate degradation in *P. stutzeri* HI00D01 with inorganic phosphate

In *Escherichia coli*, the C-P lyase pathway is under the control of the phosphorus starvation Pho regulon (Metcalf & Wanner, 1991, 1993b; Wanner & Metcalf, 1992). On the other hand, the *Pseudomonas* sp. 4ASW was observed to utilize a combination of orthophosphate and phenylphosphonate when both phosphorus sources were present in the medium (Kononova & Nesmeyanova, 2002). In this experiment, inorganic phosphate was added to phosphonate-amended *P. stutzeri* HI00D01 cultures to determine if gas production from

phosphonate degradation was inhibited. Under phosphate-deficient conditions, approximately 50% of the methyl phosphonate supplied to cultures (1 μM) was accumulated as methane gas over the course of 1.75 days (Table 4.2). Addition of inorganic phosphate (1 μM) suppressed methyl phosphonate utilization and methane production to about 5% of the total methyl phosphonate available most likely because it is a more labile than phosphonates. Similarly, ethane yields were suppressed by an order of magnitude in cultures amended with ethyl phosphonate in the presence of inorganic phosphate (1 μM), though ethane gas concentrations could not be quantified in this experiment for lack of an appropriate standard (Table 4.2).

Consistent with the first screening experiment, cultures amended with HMW DOM polysaccharides produced methane and ethylene gas. Methane concentrations [10.3 nM without inorganic phosphate and 6 nM with the addition of inorganic phosphate (Table 4.2)] were similar to the concentrations observed in the first experiment (4 nM methane in cultures supplemented with glucose and HMW DOM polysaccharides). These concentrations represent only a small fraction (1.3% and 0.75%, respectively) of the methyl phosphonate added as HMW DOM polysaccharides (800 nM). Ethylene concentrations (17.3 nM) in cultures amended with HMW DOM polysaccharides and glucose were slightly higher than methane concentrations (10.3 nM) but again represented a small fraction of the 2-hydroxyethyl phosphonate present in HMW DOM (a similar amount to the methyl phosphonate in HMW DOM). Even though only a small fraction of the HMW DOM phosphonates available were converted to hydrocarbons, the phosphate obtained from these phosphonates may have been sufficient to meet the phosphorus demand of bacteria. The final cell density reached in batch cultures amended with HMW DOM polysaccharides (and glucose) was $\sim 1 \times 10^6$ cells ml^{-1} higher than the cell density in unamended control batch cultures ($4.61 \times 10^6 - 3.67 \times 10^6$ cells ml^{-1} ; Table 4.4). Assuming a cellular carbon

content of 20 fg cell⁻¹ (Lee and Fuhrman, 1987) and a C:P cellular ratio for marine heterotrophic bacteria of ~50 (Goldman, Caron, & Dennett, 1987), the phosphorus required to account for this growth is approximately 33.4 nM. The evolution of methane (10.3 nM) and ethylene (17.3 nM) totaled 27.6 nM which is equivalent to the concentration of phosphorus released during phosphonate degradation and represents ~82.6% of the bacterial phosphorus demand. It is possible that once cells meet their phosphorus demand HMW DOM phosphonate degradation stops. This may explain the low methane and ethylene concentrations evolved in the presence of HMW DOM polysaccharides. The higher concentrations of methane (~448 nM) evolved in cultures amended with free methyl phosphonate (1 μM) respectively, may be a surplus produced to the continued activity of the C-P lyase. Measurements of inorganic phosphate in the spent media should be included in future experiments to determine if it accumulates once cultures meet their phosphorus quota.

The estimates of bacterial phosphorus demand based on the cellular carbon content of 20 fg cell⁻¹ (or 1.67 fmols C cell⁻¹) may be too low considering that batch cultures were grown in high concentrations of glucose (100-200 μM carbon). Using oxygen consumption measurements (Table 4.3) the carbon consumed by bacteria can be calculated from a theoretical respiratory quotient (CO₂ evolved per mol of O₂ consumed) for carbohydrates consumption of ~1 and a growth yield constant, (biomass produced per mol of substrate consumed) for glucose, $Y_{glucose}$ of 0.5 (White, 2007). During the incubation period, 67.7 μM O₂ was consumed in cultures amended with HMW DOM polysaccharides (237.2 μM O₂ at the beginning of the experiment and 169.5 μM O₂ when cultures were killed; Table 4.3). Assuming 1 mol of CO₂ is produced per mol of O₂ consumed, 67.7 μM CO₂ were produced. Given a $Y_{glucose}$ of 0.5, the total glucose carbon incorporated into bacterial biomass was also 67.7 μM. Dividing 67.7 μM by the bacterial

growth in these batch cultures ($\sim 3 \times 10^6$ cells ml^{-1} or the difference between the initial and final cell concentrations; Table 4.4) results in a bacterial carbon content of 272 fg C cell^{-1} (22.6 fmols C cell^{-1}). This value is 13.6 times greater than the 20 fg C cell^{-1} (1.67 fmols C cell^{-1}). Based on this result, and assuming a C:P ratio of heterotrophic bacteria of 50, the bacterial phosphorus demand is of 0.452 fmols P cell^{-1} . Using this phosphorus quota, a culture with 1×10^6 cells ml^{-1} requires 452 nM phosphorus. This value is in closer agreement with the methane concentrations evolved in batch cultures amended with free methyl phosphonate (~ 448 nM). If these assumptions are correct, the data indicates that only a small fraction of HMW DOM phosphonates were utilized, most likely because bacteria met their phosphorus demand with additional sources of phosphorus in HMW DOM. It may also be that the majority of HMW DOM phosphonates are simply inaccessible to *P. stutzeri*.

Inorganic phosphate additions suppressed methane and ethane production ten-fold (relative to cultures with no phosphate added) in cultures amended with methyl phosphonate and ethyl phosphonate (Table 4.2). In contrast, methane and ethylene concentrations were similar in HMW DOM polysaccharide-amended cultures with and without inorganic phosphate additions (Table 4.2). These results indicate that inorganic phosphate does not suppress HMW DOM phosphonate degradation suggesting perhaps a regulation mechanism based on substrate concentration. It is possible though that degradation of HMW DOM phosphonates was suppressed because other forms of organic phosphorus were more readily available in HMW DOM. Phosphate esters and pyrophosphate comprise approximately 80% of the organic phosphorus in HMW DOM F1 polysaccharides while phosphonates are only 20% (Daniel Repeta, personal communication). If only a small fraction of these compounds are more readily available than the phosphonates the bacteria may have met their phosphorus quota. Alternatively,

longer incubation periods may be required to degrade all the available phosphonates in HMW DOM. In the other words, the hydrolysis of phosphonate moieties from HMW DOM polysaccharides is the limiting step in the degradation of phosphonates by the C-P lyase. The incubation period in this experiment was limited to 1.75 days in order to prevent cultures from becoming anoxic. In addition, a large fraction of the HMW DOM phosphonates may be inaccessible to *P. stutzeri* HI00D01 due to molecular size constraints of HMW polysaccharides or the specific linkage of phosphonates and polysaccharides. Another explanation to the low hydrocarbon concentrations observed is that *P. stutzeri* HI00D01 utilized these as a carbon source. Incubations with ¹⁴C radio-labeled phosphonates can be performed to determine if phosphonate carbon is incorporated into bacterial biomass.

Batch cultures of *P. stutzeri* phn⁻ strain 14-E11 amended with phosphonates and HMW DOM polysaccharides did not accumulate hydrocarbon gases as predicted (Table 4.2). Oxygen drawdown measurements (Table 4.3) and total cell counts (Table 4.4) indicated that mutant cultures were as active as the wildtype strain during the incubation period. The mutant strain growth yields were lower than the wildtype strain when phosphonates were the sole source of phosphorus available. This was expected as the mutant strain could not metabolize phosphonates. In addition, the mutant strain seemed to have higher respiration rates than the wild type strain when no phosphorus was available (Table 4.3). Both of these observations suggest that the mutant strain had overall lower growth efficiencies than the wildtype strain most likely due the presence of antibiotic in the media. Growth yields were similar in both strains when inorganic phosphate was available in the media. These results suggest that the pre-culture conditions were suitable to test utilization of phosphonates and HMW DOM polysaccharides as the source of phosphorus.

The modifications to the previous experiment, namely adding a lower concentration of glucose (100 μ M carbon) and replenishing oxygen in the media at the beginning of the incubation, demonstrated that isolate *P. stutzeri* HI00D01, a phosphonate degrader, is capable of producing methane and ethylene from HMW DOM polysaccharides under aerobic conditions. As expected, methane was produced under aerobic conditions as indicated by the oxygen concentrations present at the end of the incubation (Table 4.3). The disruption of the phosphonate degradation gene *phnK* in mutant strain 14-E11 clearly disabled the production of trace gases from HMW DOM and model phosphonate compounds.

Methane and ethylene accumulation in *P. stutzeri* HI00D01 cultures amended with methyl phosphonate, 2-hydroxyethyl phosphonate, and HMW DOM polysaccharides

A third experiment was prepared with triplicate batch cultures of *P. stutzeri* HI00D01 to further verify the production of hydrocarbon gases from HMW DOM phosphonate degradation. Control treatments consisting of phosphorus-deficient cultures of the wildtype strain accumulated a minimal, though significant, amount of methane (0.5 nM) over the initial gas concentration (Table 4.5; T-Test, P-value<0.001). Consistent with previous observations, gases were not detected in phosphonate-amended cultures of the *phn*- strain 14-E11 (Table 4.5). The methane concentrations detected in mutant cultures were not significantly different from the gas concentrations at the beginning of the incubation (T-Test, P-value>0.32). Ethylene was only detected in wildtype cultures supplemented with HMW DOM polysaccharides or 2-hydroxyethyl phosphonate. Approximately 76% of the methyl phosphonate (1 μ M) available to cultures accumulated as methane gas over the course of the 3.8 day incubation period. In the previous experiment, the incubation time was limited to 1.75 days to prevent anoxia and the fraction of methyl phosphonate added that was converted to methane was 50%. It is possible that the

extended incubation period in this experiment allowed further degradation of methyl phosphonate. In contrast, the fraction of 2-hydroxyethylene converted to ethylene was only 6.8% of the total in wildtype cultures amended with 2-hydroxyethyl phosphonate (Table 4.5). The lower yields of ethylene relative to methane suggest that a fraction of the ethylene was consumed, assuming both phosphonates are used at equal rates. Consistent with this observation, the yields of ethylene from HMW DOM (2.2 nM) were slightly lower than for methane (7.3 nM) (Table 4.5) even though both phosphonates are found at similar concentrations in HMW DOM (Daniel Repeta, personal communication). Alternatively, ethylene can be detrimental to cells or inhibit growth. No ethane or other hydrocarbons were detected in phosphonate-amended cultures. The overall poor yields of methane and ethylene evolved from HWM DOM polysaccharides are consistent with the low yields observed in previous experiments (Table 4.1 and 4.2) and may be an indication of the diagenetic state of HMW DOM. These results confirm the observation that only a small fraction of phosphonates associated with HMW DOM are available to the phosphonate-degrading *P. stutzeri* isolate HI00D01.

Degradation of HMW DOM phosphonates by *Sulfitobacter* strain HI0054

Batch cultures experiments of *Sulfitobacter* strain HI0054 confirmed that this bacterium is also able to degrade phosphonates associated with HMW DOM. Both methane and ethylene accumulated in the medium amended with HMW DOM polysaccharides (Table 4.6) and yields were similar to those observed in *P. stutzeri* HI00D01 batch cultures (Table 4.2 and 4.5), albeit ethylene yields were ten times higher. In addition, a small amount of propylene was detected in batch cultures above the background concentration at the beginning of the incubation period. *Sulfitobacter* HI0054 cultures converted approximately 10% of methyl phosphonate to methane and 8–35% of 2-hydroxyethyl phosphonate to ethylene (Table 4.6). Additionally, ethane was

released from cultures amended with 2-phosphonopropionate though yields were only 1.1% of the total. The higher yield of ethylene (20.6 nM) from HMW DOM relative to methane (7.2 nM) may indicate that this isolate cannot consume ethylene. The higher yields of ethylene relative to methane may be due to the use of a different sample of HMW DOM F1 polysaccharides for this experiment which were collected from Station ALOHA rather than in the National Energy Laboratory Hawaii Authority, Kona, Hawaii.

Distribution of the C-P lyase pathway in the global ocean

The C-P lyase genes identified in the draft genomes of Hawaii open-ocean isolates (Appendix I Table A8) and in phosphonate-amended microcosms experiments (Martinez et al., 2013; Appendix I Table A7) were used to interrogate the OM-RGC, a reference gene database constructed with the metagenomic information of the *Tara* Oceans Expedition as well as the GOS database and publicly available marine genomes (Sunagawa et al., 2015). The analysis identified *Tara* Oceans metagenome sampling stations with OM-RGC hits to at least 7 out of 11 C-P lyase genes (*phnCDEFGHIJKLM*) predicted in the genomes of Hawaii bacterial isolates or in the microcosm metagenomes (Fig. 4.4). OM-RGC hits with a normalized coverage $\geq 10^{-6}$ (that is, with respect to the total coverage of the reference genes from each metagenomic sample) were used to determine presence or absence and thus represents a conservative view of the occurrence of the C-P lyase pathway throughout the ocean. By this metric, the C-P lyase genes were present in 69 out of 243 metagenomes of the *Tara* Oceans Expedition. Of these 69 metagenomes, 15 occurred in the Mediterranean Sea (or 79% of Mediterranean Sea metagenomes), 16 in the South Atlantic Ocean (or 30% of South Atlantic Ocean metagenomes), and 11 in the North Atlantic Ocean (or 52% North Atlantic Ocean metagenomes) (Fig. 4.4). Of the 243 metagenomes interrogated, a quarter of them (60 total) were obtained from coastal environments, as defined by

Longhurst (2007). These coastal stations spanned the borders of the Indian Ocean, the North and South Atlantic Ocean, the North and South Pacific Ocean, and the Red Sea. Of these 60 coastal metagenomes, the C-P lyase genes were present in 13 (22%), five of which were located in the Red Sea, three in the Caribbean, two off the Eastern Arica coast, one off the coast of Chile and Peru, and one off the Guianas coast.

The remaining 56 metagenomes where the C-P lyase genes were present were obtained from open sea regions in the Mediterranean Sea (15 metagenomes), the South and North Atlantic Ocean Gyres (11 and 7 metagenomes, respectively), the South Pacific Ocean Gyre (15 metagenomes), and the Indian Ocean (7 metagenomes). In some of these regions planktonic communities are known to be co-limited by phosphorus and another macronutrient, typically nitrogen (Moore et al., 2013). In the Mediterranean Sea for example, bacterioplankton communities are limited by phosphorus especially in the spring and summer months when the water column is more stratified (Pinhassi et al., 2006). Among the bacterioplankton, Pinhassi et al. (2006) found that the *Roseobacter* and *Flavobacteria* phylotypes were particularly sensitive to inorganic nitrogen and phosphorus additions. Similarly, in Mediterranean Sea waters Sebastián & Gasol, (2013) found that the *Roseobacter* clade was co-limited by nitrogen and phosphorus and the *Gammaproteobacteria* were the most phosphorus limited, whereas SAR11 bacteria were not sensitive to nutrient levels. In the North Atlantic Ocean, studies in the region of the Bermuda Atlantic Time Series have found that plankton can have a strong deficiency in phosphorus as revealed by measurements of phosphorus availability, phosphorus assimilation rates, and the high N:P ratio of both inorganic and organic nutrient pools (Ammerman, Hood, Case, & Cotner, 2003 and references therein). There is less information regarding the mechanism of nutrient limitation in the South Atlantic Ocean and Indian Ocean (Moore et al., 2013).

Additionally, the analysis did not detect the C-P lyase genes in North Pacific Ocean stations representative of the subtropical gyre, a region in which plankton may be under phosphorus control, especially in times of net nitrogen-fixation (Karl et al., 2001).

Based on the depth of the metagenomic samples, the presence of the C-P lyase genes was relatively similar (25-27% of samples of the same depth category) across surface waters (5 m), deep chlorophyll maximum waters (71 m; s.d. 41 m), and mixed layer waters (120-150 m). Interestingly, half of surface water metagenomes with C-P lyase genes were located in the Mediterranean Sea and North Atlantic Ocean. However, the C-P lyase genes were also present in 42% of mesopelagic samples (660m; s.d. 220 m) suggesting that phosphonate cycling may have special importance even in areas where inorganic phosphorus is not a limiting nutrient.

The analysis presented here is by no means a complete assessment of the distribution of C-P lyase genes across marine regions. It simply corroborates that C-P lyase genes are widespread and begins to explore the hypothesis that phosphonates may be an important source of phosphorus to bacterioplankton especially when inorganic phosphate levels are low such as in oligotrophic stratified systems. The higher percentage of metagenomes with C-P lyase genes presence in open sea regions compared to coastal regions may be an indication of the importance of DOP cycling for the microbial community in waters with low phosphate concentrations. Phosphorus enters the ocean primarily through the riverine flux carrying dissolved and particulate phosphorus derived originally from continental weathering. Most of this phosphorus cycles within the continental shelves and thus other phosphorus inputs like atmospheric deposition (aerosols, volcanic ash, and mineral dust) are more important for open sea regions (Paytan & Mclaughlin, 2007). As shown in Figure 4.3, the mean annual phosphate concentration in coastal regions ($>2 \mu\text{g L}^{-1}$) is often an order of magnitude higher than in the oligotrophic open

sea regions ($<0.2 \mu\text{g L}^{-1}$), though not all of the riverine phosphorus is bio-available (Paytan & Mclaughlin, 2007). Differences in phosphonate gene representation are also evident between two distinct oligotrophic regions. Martinez et al. (2010) explored the prevalence of phosphonate degradation genes, including representative genes of the C-P lyase pathway, in metagenomes from the Sargasso Sea and from Station ALOHA. Based on the abundance metrics used in their bioinformatics analysis, the C-P lyase gene homologues (*phnI* and *phnJ*) were present in 1% or less of the Station ALOHA microbial community in surface waters (upper 100 meters), while in the Sargasso Sea, 5-20% of the community contained C-P lyase genes. The C-P lyase genes in the Sargasso Sea samples are dominated by the *Alphaproteobacteria* SAR11 clade and *Rhodobacterales*. The prevalence of C-P lyase genes in the Sargasso Sea may be associated with phosphorus availability. In the Sargasso Sea, phosphonates may be more important as an alternative source of phosphorus since the Western North Atlantic is considered phosphorus-limited whereas Station ALOHA waters have higher dissolved inorganic phosphate (DIP) concentrations. However, C-P lyase genes abundance seems to increase at depths below 500 meters where DIP is not limiting (Martinez et al., 2010). As Martinez et al. (2010) discuss, this observation may be related to the use of phosphonates not only as a phosphorus source but as a carbon and nitrogen substrate. In fact, phosphonate utilization is not always controlled by levels of inorganic phosphorus; it can be regulated by substrate concentrations. This seems to be consistent with the observation that phosphate additions did not suppress HMW DOM phosphonate degradation in batch cultures. The C-P lyase genes at depth may thus enable microorganisms to utilize phosphonates produced at those depths or phosphonates exported via sinking particles. Further investigation is needed to determine the sources and sinks driving phosphonate cycling in the ocean.

Conclusion

Batch cultures of *P. stutzeri* HI00D01, *P. stutzeri* mutant strain 14-E11, and *Sulfitobacter* HI0054 were prepared to test for production of hydrocarbon gases from the degradation of free phosphonates and HMW DOM-bound phosphonates. The HI00D01 and HI0054 wildtype strains were predicted to release methane and other hydrocarbons as a result of degrading HMW DOM alkylphosphonates while the *P. stutzeri* *phn*⁻ mutant would have this function inhibited. This potential was also suggested by the presence of C-P lyase genes in the genome of these two wildtype strains. As expected, methane was detectable only in cultures of *P. stutzeri* HI00D01 and *Sulfitobacter* HI0054 wildtype strains when grown with a labile source of organic carbon and methyl phosphonate and HMW DOM polysaccharides as a source of phosphorus (Table 4.1, 4.2, 4.5, and 4.6). Similar tests with a genetic mutant strain of *P. stutzeri* with a disabled phosphonate degradation pathway demonstrated the participation of the C-P lyase pathway in the production of hydrocarbons from HMW DOM phosphonates.

The exact molecular details of the linkage of HMW DOM phosphonates are not clear. NMR spectra of HMW DOM indicates that methyl phosphonate and 2-hydroxyethyl phosphonate account for >85% of the phosphonates on HMW DOM polysaccharides (Daniel Repeta, personal communication). The enzymes necessary to cleave HMW DOM-bound phosphonates, however, have not been identified. *P. stutzeri* strain HI00D01 and other HMW DOM phosphonate-degrading microorganisms may secrete cell-surface enzymes similar to alkaline phosphatase or diphosphoesterases that can cleave phosphonate moieties from polysaccharides to make them available to the C-P lyase pathway. These experiments however, did not aim to identify such enzymes. They solely served to test the potential of HMW DOM phosphonates to serve as a source of phosphorus and to produce environmentally-relevant

hydrocarbons such as methane. Future work must focus on studying the activity of this pathway under normal environmental conditions (perhaps *in situ*) rather than under carbon and nutrient-replete conditions (nitrogen and trace elements were supplemented in defined media). Elucidating the genes and enzymes required to cleave HMW DOM-bound phosphonates to make them available for cell uptake will be key in understanding phosphonate and HMW DOM cycling.

One strategy that may prove useful may be to induce the expression of phosphorus-acquiring genes of *P. stutzeri* HI00D01 and *Sulfitobacter* HI0054 cultures grown under phosphorus-limited conditions and in the presence of HMW DOM. However, it may be that HMW DOM phosphonate degradation occurs independently of phosphate starvation as suggested by the relatively low inhibition effect of hydrocarbon production with the addition of inorganic phosphate (Table 4.2) and the alternative regulators (blue-light activated proteins and TetR) clustered with the C-P lyase genes in the isolates studied (Chapter 3). An alternative approach is to target genes in the C-P lyase clusters that are not required for phosphonate catabolism but encode auxiliary functions such as *phnN*, *phnO*, and *phnP* (Metcalf & Wanner, 1993b). For example, *phnP* encodes a phosphodiesterase with high cyclic nucleotide activity (Podzelinska et al., 2009). It is possible that PhnP also exhibits phosphodiesterase activity towards HMW DOM phosphonates since these are linked to sugar residues by ester bonds (Repeta et al., unpublished).

Event though methyl phosphonate and 2-hydroxyethyl phosphonate account for >85% of the phosphonates present in HMW DOM from open ocean surface waters (Daniel Repeta, personal communication), the data gathered in this work suggests that only a small fraction of these HMW DOM-bound phosphonates (<1% to 3%) are bio-available to heterotrophic bacteria.

It is possible though that the results of batch culture experiments can be interpreted differently assuming that the phosphorus demand of these bacteria is lower than estimated. Representative members of two abundant and widespread lineages of heterotrophic bacteria, the *Roseobacter* clade and *Pseudomonas*, were shown to utilize these phosphonates. What limits the degradation of HMW DOM phosphonates, and how prevalent this metabolic property is among heterotrophic microorganisms, remain open questions. The widespread occurrence of the C-P lyase genes in marine bacterial genomes (Fig. 4.4; see also Martinez, Tyson, & Delong, 2010; Martinez et al., 2013; Villarreal-Chiu, 2012) suggests that many bacterial lineages, at least within the *Protoeobacteria*, may participate in phosphonate cycling. These observations and the accumulation of HMW DOM in the water column support the hypothesis that methyl phosphonate degradation is an important contributor of the observed sub-surface methane maxima across the global ocean and of natural methane emissions (Karl et al., 2008). The methyl phosphonate associated with HMW DOM (~10% of HMW DOP; Daniel Repeta, personal communications) is sufficient to account for the methane observed in the upper 300 meters of Station ALOHA and in the sub-surface methane maxima (Fig. 4.1). The sub-surface methane maxima concentrations at Station ALOHA ranged from 2.9-3.5 nM between 100-300 meters (Fig. 4.1; Holmes et al., 2000). Given a HMW DOC concentration of 24 μ M (30% of total DOC) in open ocean surface waters, and assuming a HMW DOC:Methyl phosphonate ratio of 4000:1 and that all methyl phosphonate is bio-available, HMW DOM would produce 6 nM methane, which is enough to account for the sub-surface methane maxima concentrations. Methane evolution in HMW DOM-amended batch cultures was approximately <1% to 3% of the methyl phosphonate available. Thus, not all phosphonates may be bio-available. Based on the methane gas measurements obtained from *P. stutzeri* HI00D01 batch culture experiments, the HMW DOC

to methane molar yield ratio is ca. 3.1×10^5 to 1, that is for every ~300,00 mols of HMW DOC one mol of methane is produced (Table 4.7). Assuming this represents the bio-available fraction of methyl phosphonate, the 24 μM HMW DOC normally present in seawater would only produce 0.08 nM methane (Table 4.7), two orders of magnitude lower than the concentrations measured in surface waters (Fig. 4.1). In turn, if the ethylene yields in batch cultures of *Sulfitobacter* HI0054 (1 mol of ethylene per $\sim 1.26 \times 10^5$ mol HMW DOC) are representative of the 2-hydroxyethyl phosphonate fraction available, HMW DOM would account for 0.19 nM concentrations of ethylene, an order of magnitude higher than the 0.01-0.07 nM concentrations detected in open ocean regions (Fig. 4.1). Under these assumptions, HMW DOM (ca. 24 μM DOC) can only account for the ethylene levels in the upper water column but not of the methane levels, suggesting that additional methane sources exist. It remains to be tested if methane and ethylene yields are higher when HMW DOM is incubated in the presence of several groups of bacteria. It may be that each bacterial group has the ability to access different phosphonates. Furthermore, the step-wise degradation of HMW DOM by multiple bacterial groups may expose and make bio-available additional phosphonates. Additional information regarding the biological source and bio-availability of of HMW DOM phosphonates may further our understanding of their implication in the marine DOP cycle and the greenhouse climate effect.

Despite the progress made in understanding the role of bacteria in DOP turnover it remains unclear why phosphonates are most prevalent in HMW DOM compared to POM and plankton (Kolowitz et al., 2001; Paytan et al., 2003). Diverse organisms have been shown to synthesize or contain phosphonates including *Nitrosopumilus maritimus*, a representative of Marine Group I Archaea (Metcalf et al., 2012) the cyanobacteria *Trichodesmium* (Dyhrman, Benitez-Nelson, Orchard, Haley, & Pellechia, 2009) and various dinoflagellates and

coccolithophorids (Kittredge, Horiguchi, & Williams, 1969). In addition, phosphonate biosynthesis genes identified by Metcalf and colleagues have been detected in genomic and metagenomic sequences belonging to two of the most abundant bacteria in marine surface waters, the *Alphaproteobacteria* SAR11 clade (Metcalf et al., 2012) and the photosynthetic cyanobacteria *Prochlorococcus* (Yu et al., 2013). In open-ocean waters below depths of 100 m, the Marine Group I Archaea (aka, *Thaumarchaeota*) rival bacteria in abundance (DeLong, 2003) and may also contribute significantly to the phosphonate content of HMW DOM. Interestingly, the chemical nature of the phosphonates detected in *N. maritimus* biomass – methyl phosphonate likely associated to exopolysaccharides through an ester bond (Metcalf et al., 2012) – is similar to the phosphonates detected in HMW DOM polysaccharides (Repeta et al., unpublished, personal communication). Perhaps the production of methyl phosphonates by marine Archaea like *N. maritimus* and their high abundance in mesopelagic waters may partly explain the presence of the C-P lyase genes in many of the mesopelagic metagenomes of the *Tara* Oceans Expedition (Fig. 4.4). However, a mechanistic explanation of why phosphonates accumulate in HMW DOM is needed. It may be that in the environment, secreted exopolysaccharides decorated with phosphonates [as suggested by Metcalf et al. (2012)] dissociate from prokaryotic cells as they are grazed or lysed by viruses thus becoming part of the DOM pool. It may also be that some microorganisms secrete these phosphonated polysaccharides directly into the water column as a way to store phosphorus in a semi-degradable molecule that can be accessed under phosphorus stress, though this strategy seems detrimental to cells as it involves the loss of cellular phosphorus that can be otherwise used for growth. These hypotheses and considering the abundance of these bacterial and archaeal populations in the ocean may explain in part why phosphonates preferentially accumulate in the DOM pool relative to POM. Shedding light on the

cellular localization and function of phosphonates as well as on the environmental parameters that regulate phosphonate degradation may further our understanding of the implication of these molecules in the carbon and phosphorus cycles of the ocean.

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Table 4.1. Screening of *P. stutzeri* HI00D01 phosphonate-amended cultures for accumulation of hydrocarbon gases. Treatments consisted of single replicates. Initial gas concentrations were subtracted from the end-point measured concentrations.

	Treatment	Methane (nM)	Ethylene (nM)	ΔN_2O (counts*)	Cell counts (ml ⁻¹)
1	Methyl phn + glucose	94.8	-	392.5	4.08E+06
2	Methyl phn + glucose (phn ⁻)	-0.2	-	376.5	4.04E+06
3	2-Aminoethyl phn + glucose	0.3	-	293.2	4.16E+06
4	3-Aminopropyl phn +glucose	0.3	-	350.3	4.15E+06
5	Aminomethyl phn +glucose	0.9	-	339.1	3.98E+06
6	Phosphate + glucose	0.0	-	329.6	4.01E+06
7	Glucose (No phosphorus)	1.7	-	374.6	4.08E+06
8	HMW DOM polysaccharides	-0.1	-	10.2	1.43E+06
9	HMW DOM polysaccharides + glucose	4.6	4.8	299.4	4.10E+06

Abbreviations: N₂O, nitrous oxide; phn, phosphonate, phn⁻, mutant strain 14-E11

*N₂O concentrations were not quantified. The table shows the difference in N₂O area counts observed at the start and end of incubations.

Table 4.2. Production of hydrocarbon gases by *P. stutzeri* HI00D01 wildtype and mutant cultures amended with methyl phosphonate, ethyl phosphonate, and HMW DOM polysaccharides. Average gas concentrations of duplicate samples incubated over 1.75 days are indicated.

Treatment	Methane (nM)		Ethylene (nM)		Ethane*	
	WT	phn-	WT	phn-	WT	phn-
1 Methyl phosphonate	447.8	0.0	0.1	0.0	nd	nd
2 Methyl phosphonate + iP	56.4	0.1	0.1	0.1	nd	nd
3 Ethyl phosphonate	0.1	0.0	0.4	0.1	10150	nd
4 Ethyl phosphonate + iP	0.0	0.2	0.2	0.1	1006	nd
5 HMW DOM polysaccharides	10.3	0.0	17.3	0.1	nd	nd
6 HMW DOM polysaccharides + iP	6.5	0.1	14.9	0.1	nd	nd
7 No additions	0.5	0.1	0.1	0.1	nd	nd

Abbreviations: WT, wild type strain; phn-, mutant strain 14-E11; iP, inorganic phosphate;
nd, not determined
*Ethane area counts

Table 4.3. Oxygen consumption in phosphonate-amended cultures of *P. stutzeri* HI00D01 wildtype and phn⁻ 14-E11 mutant strains. Measurements correspond to the second set of batch culture experiments presented in Table 4.2. Dissolved oxygen concentrations are calculated based on the change in oxygen/argon ratios (O_2/Ar) and the initial concentration of oxygen. Respiration (Resp) estimates were determined from the total oxygen drawdown ($T_0 O_2 - T_f O_2$) and the incubation period of 1.75 days.

Treatment	Wild Type			Mutant		
	O_2/Ar	O_2 (μM)	Resp (μM d-1)	O_2/Ar	O_2 (μM)	Resp (μM d-1)
T ₀ Initial conditions	19.8	237.3		20.0	238.4	
1 Methyl phosphonate	14.5	173.5	36.5	13.7	164.0	42.5
2 Methyl phosphonate + iP	14.8	177.7	34.1	14.5	173.4	37.1
3 Ethyl phosphonate	14.7	175.8	35.2	13.6	162.0	43.6
4 Ethyl phosphonate + iP	14.8	176.7	34.6	14.7	175.6	35.9
5 HMW DOM polysaccharides	14.2	169.5	38.7	13.4	160.6	44.4
6 HMW DOM polysaccharides + iP	14.4	172.2	37.2	14.5	173.5	37.1
7 No additions	14.1	168.6	39.3	13.6	162.5	43.3

Table 4.4. Growth of phosphonate-amended cultures of *P. stutzeri* HI00D01 wildtype and *phn*⁻ mutant strains. The growth measurements correspond to the second set of batch culture experiments presented in Table 4.2. Cultures were grown for a period of 1.75 days.

Treatment		Cell density cells ml ⁻¹		Growth cells ml ⁻¹ d ⁻¹	
		WT	<i>phn</i> ⁻	WT	<i>phn</i> ⁻
T ₀	Initial conditions	1.19E+06	1.88E+06		
1	Methyl phosphonate	4.71E+06	3.16E+06	2.01E+06	1.12E+06
2	Methyl phosphonate + iP	5.06E+06	4.31E+06	2.21E+06	1.78E+06
3	Ethyl phosphonate	4.16E+06	3.01E+06	1.69E+06	1.04E+06
4	Ethyl phosphonate + iP	4.56E+06	4.43E+06	1.93E+06	1.85E+06
5	HMW DOM polysaccharides	4.12E+06	3.80E+06	1.67E+06	1.49E+06
6	HMW DOM polysaccharides + iP	4.61E+06	4.82E+06	1.95E+06	2.07E+06
7	No additions	3.67E+06	3.19E+06	1.41E+06	1.14E+06

Abbreviations: WT, wild type strain; *phn*⁻, mutant strain 14-E11; iP, inorganic phosphate;

Table 4.5. *P. stutzeri* HI00D01 methane and ethylene yields (nM) from phosphonate degradation. Cultures were amended with 1 μM methyl phosphonate and 2-hydroxyethyl phosphonate. Hydrocarbon gas concentrations (nM) represent the mean and SEM for triplicate samples.

Treatment	Methane		Ethylene	
	WT	phn-	WT	phn-
1 No inorganic phosphate	0.5±0.1	-0.2±0.3	-	-
2 HMW DOM polysaccharides	7.3±0.1	0.1±0.1	2.2±0.6	-
3 Methyl phosphonate	762.7±12.8	0.1±0.0	-	-
4 2-Hydroxyethyl phosphonate	0.6±0.0	-0.1±0.3	68.3±10.1	-

Abbreviations: WT, wildtype; phn-, mutant strain 14-E11

Table 4.6. Degradation of phosphonates by *Sulfitobacter* strain HI0054. Accumulated hydrocarbon gas concentrations (nM) were measured in duplicate phosphonate-amended cultures after 2 days of growth. The media consisted of Hawaii seawater supplemented with 100 μ M carbon (glycerol).

Treatments	Methane	Ethylene	Ethane
1 Inorganic phosphate	0.3	0.1	0.0
2 HMW DOM polysaccharides	7.2	20.6	0.0
3 2-Phosphonopropionate	0.1	0.1	11.3
4 Methyl phosphonate	100.7	0.1	0.0
5 2-Hydroxyethyl phosphonate	0.3	216.2	0.0

Table 4.7. Bio-availability of HMW DOM methyl phosphonate and methane source strength in the upper ocean. Estimates were calculated based on methane concentrations evolved in batch culture experiments.

Isolate & Experiment	HMW DOC F1 (mM) ^a	HMW DOM MePhn (nM) ^b	Methane (nM)	% Methane yield	HMW DOC:MePhn available ^c	Methane source strength (nM) ^d
<i>P. stutzeri</i> Table 4.1	1	250	4.6	1.8	217391	0.11
<i>P. stutzeri</i> Table 4.2	3.2	800	10.3	1.3	310680	0.08
<i>P. stutzeri</i> Table 4.5	9.8	2450	7.3	0.3	1342466	0.02
<i>Sulfitobacter</i> Table 4.6	2.6	650	7.2	1.1	361111	0.07

^aBased on percent carbon of HMW DOM (~36%)

^bMePhn, methyl phosphonate. Concentration of HMW DOM methyl phosphonate amendment based on a ratio of HMW DOC:MePhn ratio of 4000:1 ($=[\text{HMW DOC}]/4000$; Daniel Repeta, personal communication).

^cRatio of HMW DOC:Methyl phosphonate available to cultures based on concentrations of methane evolved ($=[\text{HMW DOC}]/[\text{Methane}]$)

^dEstimated methane production from 24 μM HMW DOC in surface waters based on HMW DOC:MePhn available

Figures

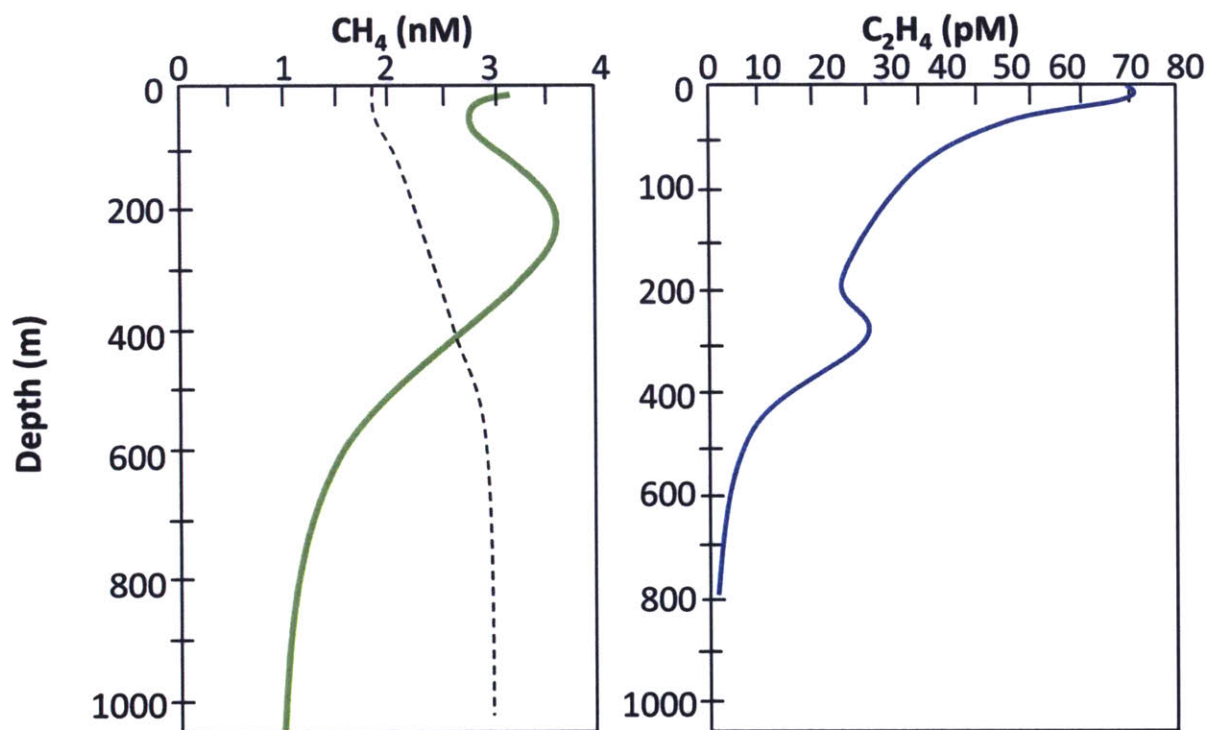


Figure 4.1. Depth profiles of methane (CH₄) and ethylene (C₂H₄) gas concentrations characteristic of the oligotrophic ocean. The dotted line in the methane profile indicates the expected atmospheric equilibrium concentrations. In the North Pacific, methane concentrations (green) in the upper 300 meters are consistently 37-67% (up to 90% in the North Atlantic) higher than the expected atmospheric equilibrium concentrations over the course of the year (Holmes, Sansone, Rust, & Popp, 2000). Ethylene concentrations (blue) are two orders of magnitude lower than methane. Methane and ethylene profiles are based on measurements in Holmes, Sansone, Rust, & Popp (2000) and Tsurushima, Watanabe, & Tsunogai (1999), respectively.

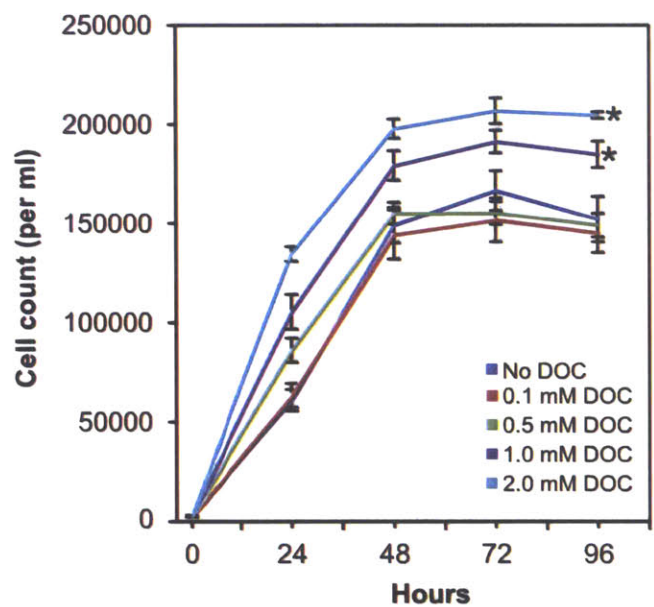


Figure 4.2. Growth of *P. stutzeri* strain HI00D01 on HMW DOM F1 polysaccharides. The growth media consisted of MOPS minimal media amended with 50 μ M phosphate. DOC amendments were estimated from the carbon content of HMW DOM polysaccharides. Treatment end-point cell yields marked with (*) were significantly different from the No DOC controls (Student's T-Test; $H_0: \bar{X}_1 = \bar{X}_2$; P-value < 0.05). Error bars indicate one standard deviation (n=3).

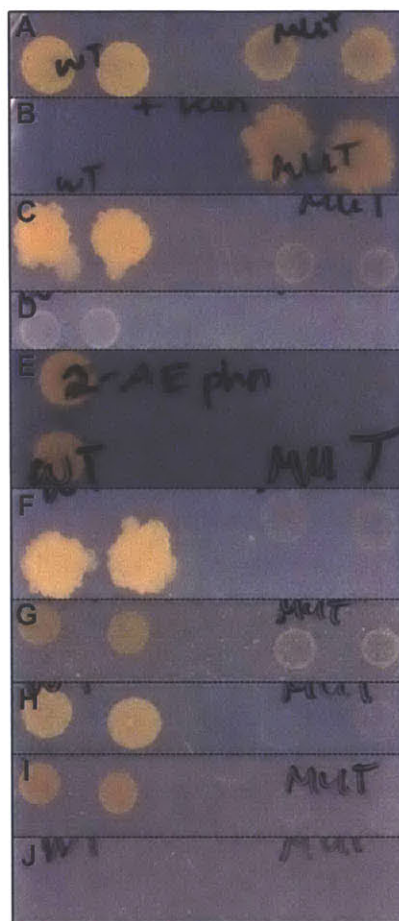
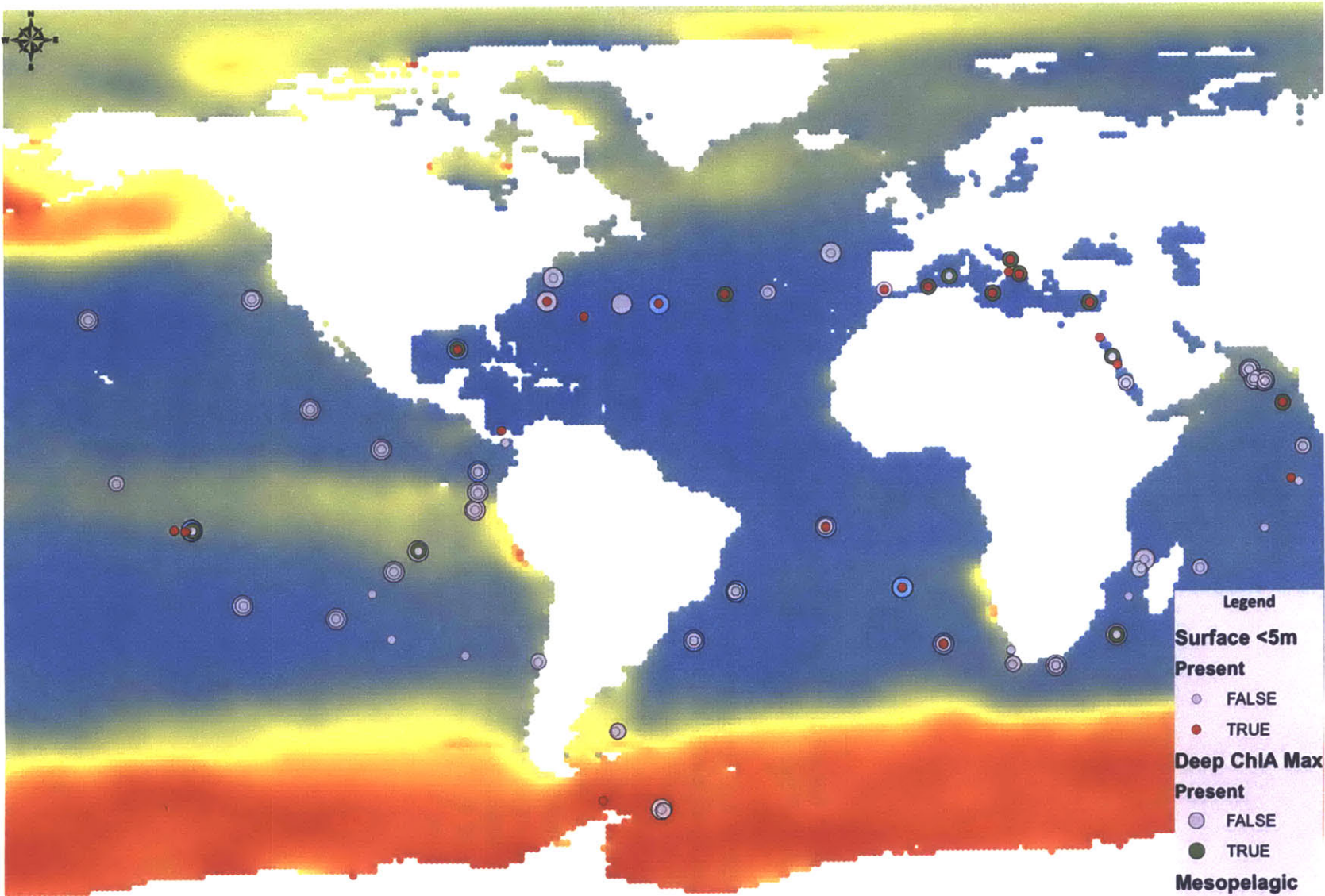


Figure 4.3. Growth of a *P. stutzeri* HI00D01 wildtype strain (WT, left) and *phn*⁻ mutant strain 14-E11 (MUT, right) on phosphonate-amended agarose media. Duplicate samples of each strain were spotted onto the agarose surface and grown in darkness at 26°C for 5-7 days. The media consisted of 1X MOPS buffer in high-purity agarose supplemented with 0.2% glucose and 100 mM phosphorus. The phosphorus sources tested included: A) inorganic phosphate (K_2HPO_4) and B) inorganic phosphate (K_2HPO_4) with kanamycin ($50 \mu g ml^{-1}$) control treatments, C) methyl phosphonate, D) aminomethyl phosphonate, E) 2-aminoethyl phosphonate, F) 3-aminopropyl phosphonate, G) ethyl phosphonate, H) 2-phosphonopropionate, I) 3-phosphonopropionate, and J) 2-phosphonobutyrate.

(Opposite page)

Figure 4.4. Global ocean distribution of bacterial C-P lyase genes and dissolved inorganic phosphate. The C-P lyase genes identified in open-ocean bacterial isolates (Chapter 3) and metagenomes (Martínez et al., 2013) were searched in the Ocean Microbiome Reference Gene Catalogue (OM-RGC; Sunagawa et al., 2015) by BLAST analysis with a bitscore ≥ 60 and percent identity ≥ 60 based on their amino acid sequence. Stations highlighted in color (red, green, or blue) indicate the presence of ≥ 7 of 11 C-P lyase gene homologs (*phnCDEFGHIJKLM*) involved in phosphonate degradation. In the map, the circles represent the geographic location of metagenomic samples included in the OM-RGC and the color indicates the sampling depth category: red inner circle, surface waters (5 m); green middle circle, deep chlorophyll maximum (DCM; 71 m; s.d. 41 m); blue outer circle, mesopelagic waters (660 m; s.d. 220 m). The background color range of global ocean indicates the mean annual dissolved inorganic phosphate concentration in surface waters (depth < 5 meters; World Ocean Atlas 2013), from high ($> 2 \mu\text{g L}^{-1}$, red) to low ($< 0.2 \mu\text{g L}^{-1}$, blue) values. Yellow and green represent intermediate values.



References

- Ammerman, J. W., Hood, R. R., Case, D. a., & Cotner, J. B. (2003). Phosphorus deficiency in the Atlantic: An emerging paradigm in oceanography. *Eos, Transactions American Geophysical Union*, 84(18), 165.
- Björkman, K., & Karl, D. (2003). Bioavailability of dissolved organic phosphorus in the euphotic zone at Station ALOHA, North Pacific Subtropical Gyre. *Limnology and Oceanography*, 48(3), 1049–1057.
- Buchan, A., González, J. M., & Moran, M. A. (2005). Overview of the marine Roseobacter lineage. *Applied and Environmental Microbiology*, 71(10), 5665–5677.
- Clark, L. L., & Ingall, E. D. (1998). Marine phosphorus is selectively remineralized. *Nature*, 393, 426.
- Clark, L. L., Ingall, E. D., & Benner, R. (1999). Marine organic phosphorus cycling: Novel insights from nuclear magnetic resonance. *American Journal of Science*, 29999, 724–737.
- DeLong, E. F. (2003). Oceans of Archaea. *ASM News*, 69(10), 503–511.
- Dyrhman, S. T., Benitez-Nelson, C. R., Orchard, E. D., Haley, S. T., & Pellechia, P. J. (2009). A microbial source of phosphonates in oligotrophic marine systems. *Nature Geoscience*, 2(10), 696–699.
- Ferrón, S., Wilson, S. T., Martínez-García, S., Quay, P. D., & Karl, D. M. (2015). Metabolic balance in the mixed layer of the oligotrophic North Pacific Ocean from diel changes in O₂/Ar saturation ratios. *Geophysical Research Letters*, 42, 3421–3430.
- Goldman, J. C., Caron, D. a., & Dennett, M. R. (1987). Regulation of gross growth efficiency and ammonium regeneration in bacteria by substrate C:N ratio. *Limnology and Oceanography*, 32(6), 1239–1252.
- Holmes, M., Sansone, F., Rust, T., & Popp, B. N. (2000). Methane production, consumption, and air-sea exchange in the open ocean: An evaluation based on carbon isotopic ratios. *Global Biogeochemical Cycles*, 14(1), 1–10.
- Ivanova, E. P., Gorshkova, N. M., Sawabe, T., Zhukova, N. V., Hayashi, K., Kurinlenko, V., ... Christen, R. (2004). *Sulfitobacter delicatus* sp. nov. and *Sulfitobacter dubius* sp. nov., respectively from a starfish (*Stellaster equestris*) and sea grass (*Zostera marina*). *International Journal of Systematic and Evolutionary Microbiology*, 54(2004), 475–480.
- Kana, T. M., Darkangelo, C., Hunt, M. D., Oldham, J. B., Bennett, G. E., & Cornwell, J. C. (1994). Membrane inlet mass spectrometer for rapid high-precision determination of N₂, O₂, and Ar in environmental water samples. *Analytical Chemistry*, 66(23), 4166–4170.

- Karl, D. M. (2014). Microbially mediated transformations of phosphorus in the sea: New views of an old cycle. *Annual Review of Marine Science*, 6, 279–337.
- Karl, D. M., Beversdorf, L., Björkman, K. M., Church, M. J., Martinez, A., & Delong, E. F. (2008). Aerobic production of methane in the sea. *Nature Geoscience*, 1, 473–478.
- Karl, D. M., Björkman, K. M., Dore, J. E., Fujieki, L., Hebel, D. V., Houlihan, T., ... Tupas, L. M. (2001). Ecological nitrogen-to-phosphorus stoichiometry at station ALOHA. *Deep-Sea Research Part II*, 48, 1529–1566.
- Kittredge, J. S., Horiguchi, M., & Williams, P. M. (1969). Aminophosphonic acids: Biosynthesis by marine phytoplankton. *Comparative Biochemistry and Physiology*, 29, 859–863.
- Kolowith, L. C., Ingall, E. D., & Benner, R. (2001). Composition and cycling of marine organic phosphorus. *Limnology and Oceanography*, 46(2), 309–320.
- Kononova, S. V, & Nesmeyanova, M. a. (2002). Phosphonates and their degradation by microorganisms. *Biochemistry. Biokhimiia*, 67(2), 184–195.
- Lalucat, J., Bennasar, A., Bosch, R., García-Valdés, E., & Palleroni, N. J. (2006). Biology of *Pseudomonas stutzeri*. *Microbiology and Molecular Biology Reviews : MMBR*, 70(2), 510–547.
- Longhurst, A. R. (2007). *Ecological Geography of the Sea*. (pp. 327–441). Academic Press, San Diego.
- Luo, H., & Moran, M. A. (2014). Evolutionary ecology of the marine Roseobacter clade. *Microbiology and Molecular Biology Reviews*, 78(4), 573–587.
- Martinez, A., Tyson, G. W., & Delong, E. F. (2010). Widespread known and novel phosphonate utilization pathways in marine bacteria revealed by functional screening and metagenomic analyses. *Environmental Microbiology*, 12, 222–238.
- Martínez, A., Ventouras, L. A., Wilson, S. T., Karl, D. M., & DeLong, E. F. (2013). Metatranscriptomic and functional metagenomic analysis of methylphosphonate utilization by marine bacteria. *Frontiers in Microbiology*, 4, 340.
- Metcalf, W. W., Griffin, B. M., Cicchillo, R. M., Gao, J., Janga, S. C., Cooke, H. A., ... Donk, W. a Van Der. (2012). Synthesis of methylphosphonic acid by marine microbes: A source for methane in the aerobic ocea. *Science*, 337(6098), 1104–1107.
- Metcalf, W. W., & Wanner, B. L. (1991). Involvement of the *Escherichia coli* phn (psiD) gene cluster in assimilation of phosphorus in the form of phosphonates, phosphite, P(i) esters, and P(i). *Journal of Bacteriology*, 173(2), 587–600.

- Metcalf, W. W., & Wanner, B. L. (1993a). Evidence for a fourteen-gene, *phnC* to *phnP* locus for phosphonate metabolism in *Escherichia coli*. *Gene*, *129*, 27–32.
- Metcalf, W. W., & Wanner, B. L. (1993b). Mutational analysis of an *Escherichia coli* fourteen-gene operon for phosphonate degradation, using *TnphoA'* elements. *Journal of Bacteriology*, *175*(11), 3430–3442.
- Moore, C. M., Mills, M. M., Arrigo, K. R., Berman-Frank, I., Bopp, L., Boyd, P. W., ... Ulloa, O. (2013). Processes and patterns of oceanic nutrient limitation. *Nature Geoscience*, *6*(9), 701–710.
- Paytan, A., Cade-Menun, B. J., McLaughlin, K., & Faul, K. L. (2003). Selective phosphorus regeneration of sinking marine particles: evidence from ^{31}P -NMR. *Marine Chemistry*, *82*, 55–70.
- Paytan, A., & McLaughlin, K. (2007). The Oceanic Phosphorus Cycle. *Chemical Reviews*, *107*(2), 563–576.
- Pinhassi, J., Gómez-Consarnau, L., Alonso-Sáez, L., Sala, M., Vidal, M., Pedrós-Alió, C., & Gasol, J. (2006). Seasonal changes in bacterioplankton nutrient limitation and their effects on bacterial community composition in the NW Mediterranean Sea. *Aquatic Microbial Ecology*, *44*, 241–252.
- Podzelinska, K., He, S.-M., Wathier, M., Yakunin, a., Proudfoot, M., Hove-Jensen, B., ... Jia, Z. (2009). Structure of *PhnP*, a phosphodiesterase of the carbon-phosphorus lyase pathway for phosphonate degradation. *Journal of Biological Chemistry*, *284*(25), 17216–17226.
- Repeta, D. J., Ferron, S., Sosa, O. A., DeLong, E. F., & Karl, D. M. (2016). Title of paper. Unpublished manuscript.
- Sebastián, M., & Gasol, J. M. (2013). Heterogeneity in the nutrient limitation of different bacterioplankton groups in the Eastern Mediterranean Sea. *The ISME Journal*, *7*, 1665–1668.
- Sunagawa, S., Coelho, L. P., Chaffron, S., Kultima, J. R., Labadie, K., Salazar, G., ... Lepoivre, C. (2015). Structure and function of the global ocean microbiome. *Science*, *348*(6237), 1–10.
- Tilbrook, B. D., & Karl, D. M. (1995). Methane sources, distributions and sinks from California coastal waters to the oligotrophic North Pacific gyre. *Marine Chemistry*, *49*, 51–64.
- Villarreal-Chiu, J. F. (2012). The genes and enzymes of phosphonate metabolism by bacteria, and their distribution in the marine environment. *Frontiers in Microbiology*, *3*, 1–13.
- Wagner-Döbler, I., Rheims, H., Felske, A., El-Ghezal, A., Flade-Schröder, D., Laatsch, H., ... Tindall, B. J. (2004). *Oceanibulbus indolifex* gen. nov., sp. nov., a North Sea

- alphaproteobacterium that produces bioactive metabolites. *International Journal of Systematic and Evolutionary Microbiology*, 54(2004), 1177–1184.
- Wanner, B. L., & Metcalf, W. W. (1992). Molecular genetic studies of a 10.9-kb operon in *Escherichia coli* for phosphonate uptake and biodegradation. *FEMS Microbiology Letters*, 79, 133–9.
- White, D. (2007). *The physiology and biochemistry of prokaryotes* (3rd ed., pp. 70–71). New York: Oxford University Press.
- White, A. K., & Metcalf, W. W. (2007). Microbial metabolism of reduced phosphorus compounds. *Annual Review of Microbiology*, 61, 379–400.
- Yu, X., Doroghazi, J. R., Janga, S. C., Zhang, J. K., Circello, B., Griffin, B. M., ... Metcalf, W. W. (2013). Diversity and abundance of phosphonate biosynthetic genes in nature. *Proceedings of the National Academy of Sciences of the USA*, 110(51), 20759–20764.
- Zhang, Q., & van der Donk, W. (2012). Answers to the carbon-phosphorus lyase conundrum. *Chembiochem*, 13(5), 627–629.

CHAPTER FIVE

Summary and future directions

Modeling studies of the marine carbon cycle often exclude or sparingly mention the role of microorganisms, specifically of heterotrophic bacteria which live through the oxidation of organic substrates, because the net carbon fluxes they mediate are negligible on the timescales relevant to the ocean-atmosphere CO₂ equilibrium, i.e., thousands of years (Sarmiento & Gruber, 2006). The biological pump – the biological and physical processes that transport inorganic carbon fixed into organic matter to the interior of the deep blue ocean – for example, is left out of the equation because it is assumed to operate in steady-state with the upward transport of dissolved inorganic carbon and because phytoplankton, a major component of the sinking export flux, are primarily controlled by nutrients and light, not by CO₂ levels, and therefore does not contribute to anthropogenic CO₂ sequestration (Sarmiento & Gruber, 2006). These assumptions however are prone to fail with the onset of feedback mechanisms that result from a changing climate influencing both food web structure and dynamics and the capacity of the ocean to take up CO₂. For example, the domain shift hypothesis – the alteration of the native phytoplankton community composition from eukaryotic and bacterial to bacterial-dominated assemblages (primarily *Prochlorococcus*) as a result of ecosystem changes driven by climate change – predicts consequent changes in the efficiency of the biological pump – the fraction of photosynthetically-fixed carbon exported from the surface layer (Karl, Bidigare, & Letelier, 2001). Moreover, the possible impact of ocean warming on the marine oxygen cycle can have future implications on the metabolic state of the ocean – the balance between gross primary production and community respiration – a topic that has remained controversial over a decade (Ducklow & Doney, 2011). These examples warrant the need not only to include biology in

current models (Reid, 2011) but to expand our understanding of the functions performed by microorganisms in the carbon cycle.

The role of heterotrophic bacteria in regulating the biological pump as well as the contribution of dissolved organic carbon (DOC) to the export flux are now well recognized (Ducklow, Steinberg, & Buesseler, 2001). New insights from cultivation-independent techniques in oxygenated marine waters have gained insight into the aerobic microbial loop (McCarren et al., 2010a; Ottesen et al., 2013, 2014; Rinta-Kanto, Sun, Sharma, Kiene, & Moran, 2012; Sharma et al., 2014; Teeling et al., 2012), wherein heterotrophic bacterioplankton play a central role in the turnover of dissolved organic matter (DOM) and serve as a link to higher trophic levels (Azam et al., 1983). These studies are extremely valuable because they provide information about the microbial taxa, biochemical pathways, and temporal and spatial scales that are relevant to DOM turnover, albeit they lack details illustrating the exact molecular transformations underlying the bacterial degradation of DOM substrates. Even though thousands of microbial isolates (including marine bacteria) have been tested to determine the types of organic substrates they can degrade or use for growth, only a few model microorganisms exist today that have enabled detailed studies of the exact transformations of organic substrates in the marine environment. An important example is that of cultured members of the *Roseobacter* lineage which have been used to describe the degradation of organic sulfur compounds such as the phytoplankton-derived osmolyte dimethyl sulfoniopropionate into the greenhouse gas dimethyl sulfide and other byproducts (Gonzalez, Kiene, & Moran, 1999). Another case is that of cultured strains of the ubiquitous SAR11 clade which have been used to investigate the degradation of one-carbon compounds (Sun et al., 2011) including methyl phosphonate which releases methane, another greenhouse gas (Carini, White, Campbell, & Giovannoni, 2014). However, detailed

information regarding the microorganisms relevant to the slower decomposition cycles of less labile pools of DOM are lacking.

A primary objective of this work was to identify appropriate model systems, e.i., marine heterotrophic microorganisms, with which to study DOM decomposition under laboratory conditions. To attain this goal, an ecologically relevant organic carbon substrate, namely the high-molecular-weight fraction (HMW) of marine DOM obtained by tangential flow ultrafiltration (Aluwihare et al., 1997; Mccarthy et al., 1996) was selected to grow and isolate suitable model microorganisms. HMW DOM is of special interest to microbial oceanographers and biogeochemists because its proton NMR-based molecular signature and predominant carbohydrate constitution is highly conserved throughout the global ocean (Aluwihare et al., 1997; Mccarthy et al., 1996) indicating that similar food web processes, if not the same organisms, produce this material. HMW DOM is also thought to be representative of the semi-labile reservoir of DOM due its periodic accumulation in surface waters and its particular nutrient-deplete elemental stoichiometry. Thus, understanding the molecular details of HMW DOM degradation has the potential of shedding light on the cycling of one of the largest organic carbon reservoirs of the planet.

In Chapter 2, I tested the hypothesis that HMW DOM isolated from marine surface waters contains bio-available carbon substrates that can stimulate the growth of bacteria. The enrichment of dilution-to-extinction cultures prepared with coastal seawater with HMW DOM resulted in the isolation of several clades of *Beta*- and *Gammaproteobacteria* that could support higher cell yields with increasing concentrations of HMW DOM. The dominant bacterial group isolated was a member of the methylotrophic *Betaproteobacteria* OM43 clade (Giovannoni et al., 2008) with a metabolism geared towards the oxidation of C1 compounds like methanol and

formaldehyde. This result was unexpected given that the major bioreactive constituents of HMW DOM are thought to be carbohydrates (Benner et al., 1992; Mccarthy et al., 1996). However, previous open-ocean microcosm experiments amended with HMW DOM also resulted in the enrichment of methylotrophic *Gammaproteobacteria* (McCarren et al., 2010a) suggesting the presence of a common degradation pathway supporting C1 metabolism in both microbial communities. Thus a logical next step was to reproduce this experimental framework in an open-ocean environment.

In Chapter 3, a series of similar cultivation experiments conducted at Station ALOHA near Hawaii resulted in the enrichment of several *Alpha*- and *Gammaproteobacteria* strains with putative HMW DOM-degrading capabilities, though none of the cultures identified matched known methylotrophic strains. Even though HMW DOM clearly stimulated the recovery of open-ocean dilution-to-extinction cultures, further experiments indicated that the majority of isolates did not exhibit significant increases in cell yields in response to HMW DOM additions, a result that contrasted with the more robust growth response of the coastal methylotrophs described in Chapter 2. The Station ALOHA isolate *Pseudomonas stutzeri* strain HI00D01, however, was studied further and was shown to increase its growth concomitantly with HMW DOM additions (Fig. 4.2 in Chapter 4) though maximum cell yields were an order of magnitude lower than the Nahant Bay OM43 clade isolate NB0046 (described in Chapter 2) suggesting that a limited fraction of HMW DOM is available to this bacterium.

A factor that may have contributed to the differences in growth yields observed between the coastal methylotrophs identified in Chapter 2 and open-ocean isolates explored in Chapter 3 is their specific bacterial carbon demand (BCD). BCD is equal to the gross bacterial production expressed in carbon units and includes the carbon incorporated into biomass and the carbon used

for respiration (Ducklow, 2000). Even though estimates of BCD were not obtained from cultures grown on HMW DOM, the cell size of these isolates may provide an indirect assessment of their carbon quota. NB0046 is similar to other cultured members of the oligotrophic OM43 clade in having a small genome (<1400 genes) and relatively small cell size (0.1-0.3 μm wide and 0.6-1.8 μm long) (Huggett et al., 2012). These small cells were in fact often difficult to enumerate by flow cytometry due their low scattering signal and small genome. Thus, one can speculate that NB0046 was likely able to achieve the highest cell yields with HMW DOM additions due to its relatively low BCD. In turn, other Hawaii isolates like *P. stutzeri* strain HI00D01 (5 Mbps genome) and *Sulfitobacter* strain HI0054 (4.2 Mbps genome) are more representative of the so-called “high-DNA” bacteria (Gasol, Giorgio, & del Giorgio, 2000) and of copiotrophic bacteria. Pedler, Aluwihare, & Azam (2014) large bacterial cells ($\sim 0.3355 \mu\text{m}^3$ and $>40 \text{ fg carbon cell}^{-1}$) from the *Alteromonas* clade that can rapidly consume most of the labile DOC pool indicating that these bacteria may require high levels of carbon to sustain their populations. Consistent with this observation, several bacterial production studies indicate that high-DNA cells are responsible for incorporating most of the radio-labeled thymidine and leucine (Gasol et al., 2000). To test this hypothesis the carbon content of OM43 clade and *Pseudomonas stutzeri* bacteria were estimated based on their theoretical cell volumes. Based on the cell dimensions measured by Huggett et al. (2012), the OM43 clade bacterial cells have a theoretical maximum volume of $0.162 \mu\text{m}^3$. Applying the allometric relationship between cell volume and bacterial cellular carbon of Loferer-Kröbächer, Klima, & Psenner (1998), the carbon content of the OM43 clade bacteria is of $45.5 \text{ fg C cell}^{-1}$. Using this carbon conversion factor and assuming a bacterial growth efficiency of 20%, cells may require up to $227.5 \text{ fg C cell}^{-1}$ to grow. Considering the ca. $5 \times 10^6 \text{ cells ml}^{-1}$ cell yields observed in 10X and 50X DOC treatments (700-

3500 μM DOC; Fig. 2.5A), the maximum carbon required to account for this bacterial biomass is estimated to be 0.095 $\mu\text{mol C ml}^{-1}$, or 2.7 to 13.6% of the carbon in 50X and 10X DOC treatments, respectively. A more conservative estimate, based on the mean cellular carbon content of 20 fg per cell⁻¹ for a natural bacterial assemblage (Lee and Fuhrman, 1987; Loferer-Kröbächer, Klima, & Psenner, 1998) and a growth efficiency of 50%, results in carbon utilization fractions of 0.04 to 2.4%, in 50X and 10X DOC treatments, respectively. With respect to *Pseudomonas stutzeri*, in the DOM growth response experiment (Fig. 4.2) cultures reached cell concentrations of 2×10^5 cells ml^{-1} in 2 mM DOC (ca. 25X DOC) treatments, 25 times less than the OM43 clade bacteria grown under similar DOC concentrations. Assuming *P. stutzeri* cells are in the large part of size spectrum of natural bacterial samples ($\sim 0.3355 \mu\text{m}^3$; Pedler, Aluwihare, & Azam, 2014), the carbon to volume conversion results in a cell carbon content of 85 fg C cell⁻¹. Applying again a growth efficiency of 20%, *P. stutzeri* cells would require approximately 425 fg C cell⁻¹ to grow. Therefore, the maximum carbon utilized by these cultures is estimated to be 0.007 $\mu\text{mol C ml}^{-1}$, or 0.35% of the 2 mM DOC in the media. Despite having similar BCD, OM43 clade and *P. stutzeri* bacteria cell yields (per DOC) differed by an order of magnitude (1.42×10^6 cells μmol^{-1} and 1×10^5 cells μmol^{-1} , respectively), indicating that less DOC was available to the *P. stutzeri* cells. Incorporating measurements of respiration and carbon consumption in future studies will help tease apart these growth differences and improve the detection of HMW DOM-degrading bacteria.

The non-linear growth response observed in DOM dose experiments remains unclear. In the case of the OM43 clade NB0046 strain, the 50X DOC treatment did not yield a corresponding five-fold increase in cell yields with respect to the 10X DOC treatment (Fig. 2.5). There are several explanations to this behavior. For example, nutrients such as vitamins may be limiting

growth. This is possible considering that the cultivation media was prepared with UV-oxidized seawater to lower the total DOC content. Another explanation for the DOM dose response observed is that the methylotrophs' growth is inhibited at high DOC concentrations. This could occur because at high DOC concentrations the probability of intermolecular interactions is increased and molecules tend to form aggregates (Zsolnay, 2003) that reduce the available substrate concentration. Another possibility is that the non-linear growth observed is related to the abiotic release of methanol; if a fraction of abiotically-produced methanol escapes from the medium a smaller amount of methanol would be available for bacterial consumption. The abiotic production of methanol from HMW DOM most likely did not influence the isolation of methylotrophs from dilution-to-extinction cultures nor their growth response in DOM dose experiments considering that neither of the cultivation experiments performed with samples from Hawaii resulted in the isolation of methylotrophic strains like the ones observed before in microcosm experiments (McCarren et al., 2010). Transcriptome analysis of methylotrophs and other bacterial isolates grown in HMW DOM-amended media may shed light on the biochemical pathways and HMW substrates these bacteria degrade and metabolize.

In contrast to other dilution-to-extinction cultivation studies, the experiments conducted in this work did not identify cultures related to the abundant *Alphaproteobacteria* SAR11 clade (Giovannoni et al., 2005). Members of the SAR11 clade are among the most common, though difficult to culture, bacterial isolates in high throughput dilution-to-extinction experiments because of their dominance in bacterioplankton communities (Connon & Giovannoni, 2002; Rappé et al., 2002; Song et al., 2009). However, neither the coastal or open-ocean cultivation experiments performed in this work were able to isolate members of this clade. This result is not unexpected given the sensitivity and recalcitrance of oligotrophic clades to grow under

laboratory conditions. Additionally, SAR11 clade strains seem to be streamlined to consume low-molecular-weight organic compounds to support their metabolism (Grote et al., 2012; Sun et al., 2011a) rather than HMW DOM. Interestingly, members of the OM43 clade share similar characteristics with the SAR11 clade; yet NB0046 cultures were able to reach high cell yields in the presence of HMW DOM. It may well be that only a few genes are required to confer HMW DOM-degradation properties to these oligotrophic isolates. Future genomic studies and gene expression experiments will be necessary to identify the enzymes encoded by NB0046 that degrade HMW DOM substrates.

Physiological parameters and genome content are both implicated in the ability of bacteria to degrade HMW DOM. In these studies, despite supplementing cultures with high concentrations of DOC and sufficient nutrients (nitrogen and phosphorus) to induce carbon consumption, cell yields were usually low and the fraction of DOC consumed was estimated to be >1% (up to 13% in non-conservative estimates) of the total organic carbon available. These results indicate that in its native form HMW DOM offers a limited supply of labile carbon to monocultures. Whether inherent physicochemical properties of HMW DOM (e.g., formation of nanogels, intermolecular interactions, or steric hindrance; see Verdugo et al., 2008) or the genomic potential of isolates are the main restriction to HMW DOM degradation remains to be determined. In addition to the apparent difference in the fraction of HMW DOM substrates available to the bacterial groups tested, the freeze-drying process used to preserve HMW DOM for analysis and culturing experiments may have affected its physicochemical properties and influenced the degree to which it is labile to bacterial degradation. Even though freeze-drying is the best method available to preserve organic matter, as it minimizes physicochemical changes and microbial degradation, the possibility that the procedure results in, for example, the

dehydration of HMW DOM moieties, cannot be neglected. Such reactions may or may not be reversible and therefore ultimately affect the way natural bacterial assemblages interact and breakdown HMW DOM. However, freeze-drying is unlikely to exert physicochemical alterations on HMW DOM given that the powder extract readily dissolves in water whereas air-drying or heat-drying of other forms of organic matter are much harder to dissolve (Malcolm, 1968).

The genomic analysis performed in Chapter 3 suggests that the bacterioplankton clades identified may provide complementary functions in the aerobic decomposition of HMW DOM involving the degradation of different moieties from HMW polysaccharides like acetyl, amino, and phosphonate residues, and methyl or methoxy groups in the case of methylotrophs (Fig. 5.1). It is hypothesized that OM43 clade bacteria encode enzymes that cleave single carbon atom moieties from HMW DOM such as methyl and methoxy groups, as these are enriched in the sugars making up HMW polysaccharides (Panagiotopoulos, Repeta, & Johnson, 2007; Quan & Repeta, 2007). The role of OM43 clade methylotrophs in HMW DOM cycling is intriguing. Methylotrophic bacteria are widespread in the marine environment and are thought to be an important sink of atmospheric methanol as it enters the ocean (Dixon, Beale, & Nightingale, 2011; Dixon, Sargeant, Nightingale, & Colin Murrell, 2013). The OM43 clade is characteristic of bacterioplankton communities in coastal waters (Giovannoni et al., 2008; Rappe, Kemp, & Giovannoni, 1997) and has been observed to increase significantly in abundance during phytoplankton blooms (Morris, Longnecker, & Giovannoni, 2006; Rich, Pham, Eppley, Shi, & DeLong, 2011) presumably due to the production of one-carbon compounds like methanol, though the exact nature of this association is not well constrained. Even though both the atmosphere and phytoplankton supply methanol to methylotrophic bacteria in the ocean, the OM43 clade isolates described in Chapter 2 may gain a competitive advantage over other

methylootrophs by obtaining one-carbon compounds from the degradation of HMW DOM, especially during non-bloom conditions. Furthermore, by contributing to the aerobic decomposition chain of HMW DOM (Fig 5.1), methylootrophs like the OM43 clade bacteria may gain access to additional carbon sources as well as to organic forms of nitrogen and phosphorus. Future work focused on co-culture experiments and artificial (chemical) alteration of HMW DOM polymers may shed some light as to why HMW DOM seems inaccessible to monocultures.

In Chapter 4, two microbial isolates were explored further: the alpha-proteobacterium *Sulfitobacter* HI0054 and the gamma-proteobacterium *P. stutzeri* HI00D01. These HMW DOM isolates were selected as candidate model systems to study HMW DOM metabolism due to their ease of growth under laboratory conditions, including robust growth on defined agar-based, nutrient-rich media, and the potential to modify them genetically. Strain HI00D01, a member of the genus *Pseudomonas* sensu stricto, was particularly amenable to genetic manipulation. In Chapter 4, I used genomic information gathered in Chapter 3 and previous chemical information derived from HMW DOM NMR spectra to test the hypothesis that bacteria carrying a known phosphonate degradation pathway were capable of utilizing alkylphosphonates associated to HMW DOM polysaccharides. Both isolates tested were capable of using phosphonates in HMW DOM as shown by the accumulation of methane and other hydrocarbons expected to be released as byproducts in batch cultures grown under phosphorus-limited conditions. The methane concentrations evolved during HMW DOM phosphonate degradation were considerably lower than expected considering the methyl phosphonate content of HMW DOM and than the total phosphorus required to account for the cell growth observed. However, it is possible that the bacterial phosphorus demand was met and bacteria down-regulated the degradation of additional

phosphonates. It is also possible that alternative sources of organic phosphorus in HMW DOM, such as phosphate esters, were more readily accessible to these bacteria and suppressed utilization of phosphonates. HMW DOM contains four times more phosphate esters and pyrophosphate than phosphonates (Daniel Repeta, personal communication). Given that HMW DOM methyl phosphonate additions were estimated to be 0.25-2.5 μM , an additional 1-10 μM organic phosphorus was available to bacteria and could have been sufficient to suppress phosphonate degradation (only 1 μM inorganic phosphate was required to considerably reduce methane evolution; Table 4.2). Alternatively, only a limited fraction of phosphonates in HMW DOM polysaccharides may be accessible to bacteria, a pattern that is consistent with the apparent low availability of HMW DOM carbon substrates to monocultures as indicated by their poor DOM dose responses (Figures 3.4, 3.5, and 3.6).

Additionally, using a genetic mutant of strain HI00D01 deficient in phosphonate catabolism, I demonstrated the involvement of a well described phosphonate-degrading enzyme complex, the C-P lyase pathway (Metcalf & Wanner, 1993), in the production of methane gas from the degradation of HMW DOM phosphonates. The C-P lyase pathway is present in the genomes of diverse *Proteobacteria* lineages (Martínez et al., 2013) and may be a primary mechanism controlling the sub-surface methane maxima observed across ocean basins (Holmes et al., 2000). Like other potent greenhouse gases, methane absorbs terrestrial radiation in wavelengths outside the spectrum already absorbed by CO_2 meaning that small increases in the concentration of this trace gases in the atmosphere can have profound effects on climate change. Therefore, confirming the involvement of the bacterial C-P lyase pathway and HMW DOM phosphobates in aerobic methanogenesis represents an important step in understanding the potential of the ocean's DOM cycle as a natural source of methane. The experiments conducted

as part of this study however, focused on assessing the potential of bacterial methane production from HMW DOM phosphonates under conditions not characteristic of the environment (e.g., glucose- and nutrient-replete cultures) and thus do not provide a thorough assessment of its ecological and climatic implications. Future work must focus on identifying additional sources and sinks of marine phosphonates and methane, as well as characterizing this methanogenic pathway under relevant environmental conditions so that such processes can be appropriately parameterized and evaluated in climate and ecosystem models.

Investigating the role of HMW DOM in the environment remains difficult in great part because its exact source is not well constrained. However, the presence of phosphonates in HMW DOM may be evidence of its microbial origin and provide clues regarding its function. The fact that at least a small fraction of HMW DOM phosphonates were readily used by heterotrophic bacteria like *P. stutzeri* HI00D01 and *Sulfitobacter* HI0054 suggest that production of phosphonates may be a mechanism by which microorganisms lock away phosphorus for later use. This strategy seems to be contradicted by the observation that POM does not usually contain phosphonates as it would be expected if these served as a cellular phosphorus bank. In a more complex scenario, however, HMW DOM phosphonates may be produced or exported extracellularly as semi-labile forms (e.g., phosphono-sugars) that cannot be easily degraded. This strategy would benefit the microbial community by alleviating phosphorus limitation for example, under periods of net nitrogen fixation and by maintaining phosphorus levels in the upper ocean that would otherwise be exported via sinking particular organic matter. Hypothetically, during these stress periods, heterotrophic microorganisms capable of degrading phosphonates like *P. stutzeri* and other bacteria encoding the C-P lyase pathway, regenerate phosphorus from this HMW DOM phosphonate reservoir for their direct benefit and indirectly

for the original producer of phosphonates. Expanding this scenario, HMW DOM nitrogen may serve a similar function as phosphorus. HMW DOM polysaccharides contain semi-labile amino sugars that may serve as a source of nitrogen to microbial communities when inorganic forms are scarce. Using measurements of stable nitrogen isotope in HMW DOM, Benner, Biddanda, Black, & McCarthy (1997) estimated turnover times of HMW dissolved organic nitrogen (DON) of 4-6 nM d^{-1} , a rate that can support a large fraction of the phytoplankton nitrogen demand if is representative of the total DON pool, suggesting that DON is an important part of the upper ocean nitrogen cycle. If this is the case, heterotrophic bacteria capable of remineralizing phosphorus and nitrogen from HMW DOM may be vital to maintain productivity in marine surface waters.

In this thesis, the model system-based approach proved useful in identifying a microbial HMW DOM decomposition pathway that can have significant implications on climate. Additionally, several candidate model microorganisms were identified such as the methylotrophic betaproteobacterium strain NB0046, the oligotrophic SAR92 clade strain NB0015, several *Oceanospirillacea* hydrocarbon degraders, and the aforementioned phosphonate-degrading bacteria which can be used to explore further the role of bacteria in the aerobic decomposition of semi-labile DOM. Continuation of the development of microbial model systems identified in HMW DOM enrichment experiments may prove useful in identifying novel pathways and carbon transformations that may have important implications on the ecology of the sea.

Figures

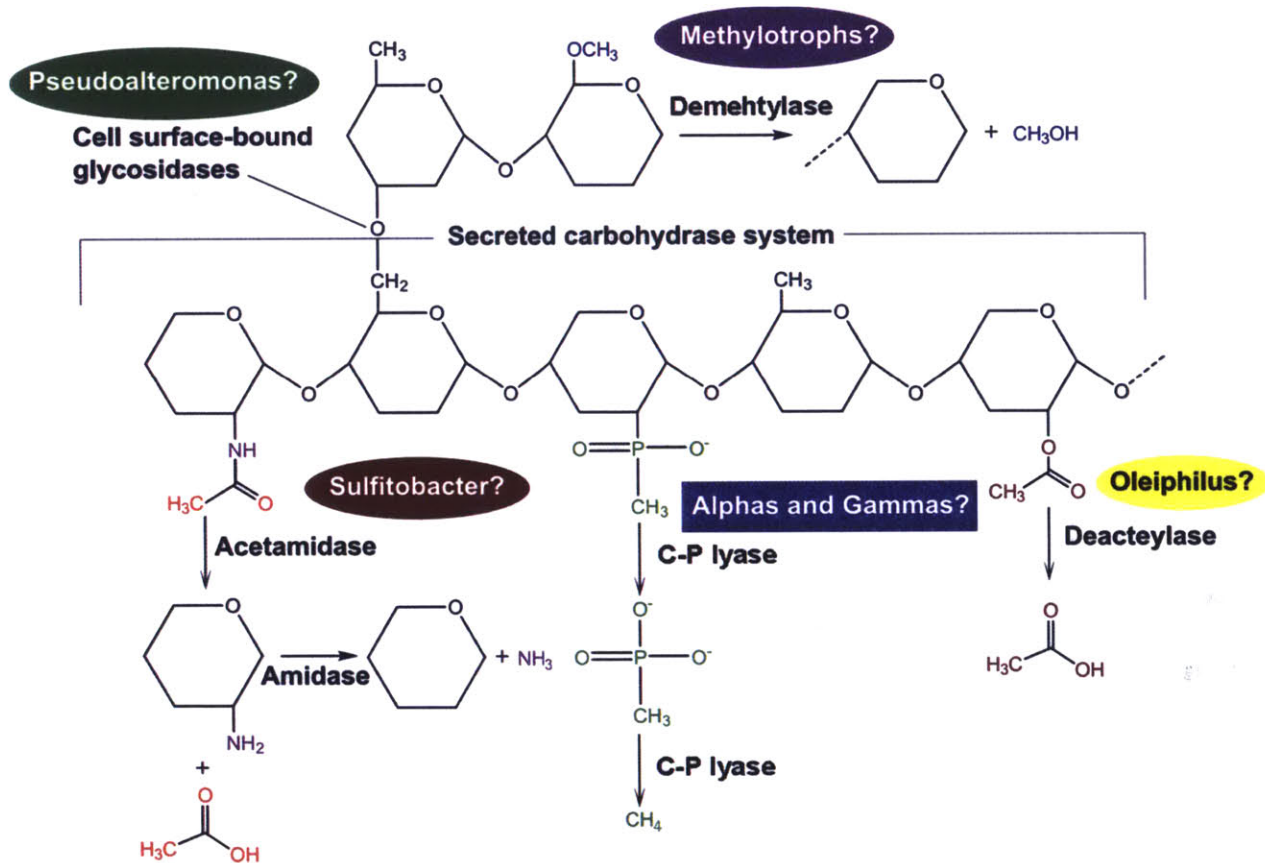


Figure 5.1. Hypothetical model of the microbial decomposition of high-molecular-weight (HMW) marine dissolved organic matter (DOM). HMW DOM is enriched in acyl-polysaccharides (Aluwihare et al., 1997) and amino (Aluwihare et al., 2005), deoxy, methyl, and phosphono sugars (Panagiotopoulos, Repeta, & Johnson, 2007; Quan & Repeta, 2007, Repeta et al., unpublished, personal communication) with moderate branching. A suite of diverse bacterial groups is hypothesized to participate in the partial degradation of HMW DOM. The cleavage of the organic moieties decorating HMW polysaccharides serves as carbon and nutrient sources to heterotrophic bacteria (e.g., methanol, acetate, and the phosphate group from phosphonates) and enables larger enzyme complexes such as secreted carbohydrases to bind and breakdown the backbone polysaccharides. In addition, the degradation of alkylphosphonates results in the release of methane and other hydrocarbons, that may or may not be consumed by bacteria.

References

- Aluwihare, L. I., Repeta, D. J., & Chen, R. F. (1997). A major biopolymeric component to dissolved organic carbon in surface sea water. *Nature* 387, 166–169.
- Aluwihare, L. I., Repeta, D. J., Pantoja, S., & Johnson, C. G. (2005). Two chemically distinct pools of organic nitrogen accumulate in the ocean. *Science*, 308(May), 1007–1010.
- Azam, F., Fenchel, T., Field, J., Gray, J., Meyer-Reil, L., & Thingstad, F. (1983). The ecological role of water-column microbes in the sea. *Marine Ecology Progress Series*, 10, 257–263.
- Benner, R., Pakulski, J. D., McCarthy, M., Hedges, J. I., & Hatcher, P. G. (1992). Bulk chemical characteristics of dissolved organic matter in the ocean. *Science*, 255, 1561–1564.
- Carini, P., White, A. E., Campbell, E. O., & Giovannoni, S. J. (2014). Methane production by phosphate-starved SAR11 chemoheterotrophic marine bacteria. *Nature Communications*, 5, 1–8.
- Connon, S. A., & Giovannoni, S. J. (2002). High-throughput methods for culturing microorganisms in very-low-nutrient media yield diverse new marine isolates. *Applied and Environmental Microbiology*, 68(8), 3878–3885.
- Dixon, J. L., Beale, R., & Nightingale, P. D. (2011). Microbial methanol uptake in northeast Atlantic waters. *The ISME Journal*, 5(4), 704–16.
- Dixon, J. L., Sargeant, S., Nightingale, P. D., & Colin Murrell, J. (2013). Gradients in microbial methanol uptake: productive coastal upwelling waters to oligotrophic gyres in the Atlantic Ocean. *The ISME Journal*, 7(3), 568–80.
- Ducklow, H. (2000). Bacterial production and biomass in the oceans. *Microbial Ecology of the Oceans*, 85–120.
- Ducklow, H. W., & Doney, S. C. (2013). What Is the metabolic state of the oligotrophic ocean? A debate. *Annual Review of Marine Science*, 5, 525–33.
- Ducklow, H. W., Steinberg, D. K., & Buesseler, K. O. (2001). Upper ocean carbon export and the biological pump. *Oceanography*, 14(4), 50–58.
- Gasol, J. M., Giorgio, P. a D. E. L., & del Giorgio, P. a. (2000). Using flow cytometry for counting natural planktonic bacteria and understanding the structure of planktonic bacterial communities. *Scientia Marina*, 64(2), 197–224.
- Giovannoni, S. J., Hayakawa, D. H., Tripp, H. J., Stingl, U., Givan, S. a., Cho, J. C., ... Rappé, M. S. (2008). The small genome of an abundant coastal ocean methylotroph. *Environmental Microbiology*, 10, 1771–1782.

- Giovannoni, S. J., Tripp, H. J., Givan, S., Podar, M., Vergin, K. L., Baptista, D., ... Mathur, E. J. (2005). Genome streamlining in a cosmopolitan oceanic bacterium. *Science*, *309*, 1242–1245.
- Gonzalez, J. M., Kiene, R. P., & Moran, M. A. (1999). Transformation of sulfur compounds by an abundant lineage of marine bacteria in the α -subclass of the class *Proteobacteria*. *Applied and Environmental Microbiology*, *65*(9), 3810–3819.
- Grote, J., Cameron Thrash, J., Huggett, M. J., Landry, Z. C., Carini, P., Giovannoni, S. J., & Rappé, M. S. (2012). Streamlining and core genome conservation among highly divergent members of the SAR11 clade. *mBio*, *3*, e00252–12.
- Holmes, M., Sansone, F., Rust, T., & Popp, B. N. (2000). Methane production, consumption, and air-sea exchange in the open ocean: An evaluation based on carbon isotopic ratios. *Global Biogeochemical Cycles*, *14*(1), 1–10.
- Huggett, M. J., Hayakawa, D. H., & Rappé, M. S. (2012). Genome sequence of strain HIMB624, a cultured representative from the OM43 clade of marine *Betaproteobacteria*. *Standards in Genomic Sciences*, *6*, 11–20.
- Karl, D. M., Bidigare, R. R., & Letelier, R. M. (2001). Long-term changes in plankton community structure and productivity in the North Pacific Subtropical Gyre: The domain shift hypothesis. *Deep Sea Research Part II: Topical Studies in Oceanography*, *48*, 1449–1470.
- Loferer-Kröbber, M., Klima, J., & Psenner, R. (1998). Determination of bacterial cell dry mass by transmission electron microscopy and densitometric image analysis, *64*(2), 688–694.
- Malcolm, R. L. (1968). Freeze-drying of organic matter, clays, and other earth materials. *U. S. Geological Survey Research - Professional Paper 600-C*, C211–C216.
- Martínez, A., Ventouras, L. A., Wilson, S. T., Karl, D. M., & DeLong, E. F. (2013). Metatranscriptomic and functional metagenomic analysis of methylphosphonate utilization by marine bacteria. *Frontiers in Microbiology*, *4*, 1–18.
- McCarren, J., Becker, J. W., Repeta, D. J., Shi, Y., Young, C. R., Malmstrom, R. R., ... DeLong, E. F. (2010). Microbial community transcriptomes reveal microbes and metabolic pathways associated with dissolved organic matter turnover in the sea. *Proceedings of the National Academy of Sciences of the USA*, *107*, 16420–16427.
- Mccarthy, M., Hedges, J., & Benner, R. (1996). Major biochemical composition of dissolved high molecular weight organic matter in seawater. *Marine Chemistry*, *55*, 281–297.

- Metcalf, W. W., & Wanner, B. L. (1993). Mutational analysis of an *Escherichia coli* fourteen-gene operon for phosphonate degradation, using Tnp_{phoA}' elements. *Journal of Bacteriology*, 175(11), 3430–3442.
- Morris, R. M., Longnecker, K., & Giovannoni, S. J. (2006). *Pirellula* and OM43 are among the dominant lineages identified in an Oregon coast diatom bloom. *Environmental Microbiology*, 8, 1361–1370.
- Ottesen, E. A., Young, C. R., Eppley, J. M., Ryan, J. P., Chavez, F. P., Scholin, C.A., & DeLong, E. F. (2013). Pattern and synchrony of gene expression among sympatric marine microbial populations. *Proceedings of the National Academy of Sciences of the USA*, 110, E488–97.
- Ottesen, E. a, Young, C. R., Gifford, S. M., Eppley, J. M., Marin, R., Schuster, S. C., ... DeLong, E. F. (2014). Multispecies diel transcriptional oscillations in open ocean heterotrophic bacterial assemblages. *Science*, 345, 207–12.
- Panagiotopoulos, C., Repeta, D. J., & Johnson, C. G. (2007). Characterization of methyl sugars, 3-deoxysugars and methyl deoxysugars in marine high molecular weight dissolved organic matter. *Organic Geochemistry*, 38, 884–896.
- Pedler, B. E., Aluwihare, L. I., & Azam, F. (2014). Single bacterial strain capable of significant contribution to carbon cycling in the surface ocean. *Proceedings of the National Academy of Sciences of the USA*, 111(20), 7202–7207.
- Quan, T. M., & Repeta, D. J. (2007). Periodate oxidation of marine high molecular weight dissolved organic matter: Evidence for a major contribution from 6-deoxy- and methyl sugars. *Marine Chemistry*, 105, 183–193.
- Rappé, M. S., Connon, S. a, Vergin, K. L., & Giovannoni, S. J. (2002). Cultivation of the ubiquitous SAR11 marine bacterioplankton clade. *Nature*, 418, 630–633.
- Reid, A. (2011). Incorporating microbial processes into climate models. *A Report from the American Academy of Microbiology* (p. 16). Dallas, TX. Retrieved from http://academy.asm.org/images/stories/documents/Incorporating_Microbial_Processes_In_to_Climate_Models.pdf
- Repeta, D. J., Ferron, S., Sosa, O. A., DeLong, E. F., & Karl, D. M. (2016). Title of paper. Unpublished manuscript.
- Rich, V. I., Pham, V. D., Eppley, J., Shi, Y., & DeLong, E. F. (2011). Time-series analyses of Monterey Bay coastal microbial picoplankton using a “genome proxy” microarray. *Environmental Microbiology*, 13, 116–134.

- Rinta-Kanto, J. M., Sun, S., Sharma, S., Kiene, R. P., & Moran, M. A. (2012). Bacterial community transcription patterns during a marine phytoplankton bloom. *Environmental Microbiology*, *14*, 228–239.
- Sarmiento, J. L., & Gruber, N. (2006). Oceanic carbon cycle, atmospheric CO₂, and climate. *Ocean Biogeochemical Dynamics*. Chapter 10 (pp. 1–99). Princeton University Press, Princeton, NJ.
- Sharma, A. K., Becker, J. W., Ottesen, E. a, Bryant, J. a, Duhamel, S., Karl, D. M., ... DeLong, E. F. (2014). Distinct dissolved organic matter sources induce rapid transcriptional responses in coexisting populations of *Prochlorococcus*, *Pelagibacter* and the OM60 clade. *Environmental Microbiology*, *16*, 2815–30.
- Song, J., Oh, H. M., & Cho, J. C. (2009). Improved culturability of SAR11 strains in dilution-to-extinction culturing from the East Sea, West Pacific Ocean. *FEMS Microbiology Letters*, *295*, 141–147.
- Sun, J., Steindler, L., Thrash, J. C., Halsey, K. H., Smith, D. P., Carter, A. E., ... Giovannoni, S. J. (2011). One carbon metabolism in SAR11 pelagic marine bacteria. *PLoS ONE*, *6*(8), e23973.
- Teeling, H., Fuchs, B. M., Becher, D., Klockow, C., Gardebrecht, a., Bennke, C. M., ... Amann, R. (2012). Substrate-controlled succession of marine bacterioplankton populations induced by a phytoplankton bloom. *Science*, *336*, 608–611.
- Verdugo, P., Orellana, M. V, Chin, W.-C., Petersen, T. W., van den Eng, G., Benner, R., & Hedges, J. I. (2008). Marine biopolymer self-assembly: implications for carbon cycling in the ocean. *Faraday Discussions*, *139*, 393–398.

APPENDIX I

This section presents supplementary data for Chapter 3 which includes the complete flow cytometry growth screen results for open-ocean dilution-to-extinction cultivation experiments I and II (Tables A1 and A2, respectively). The tables indicate the total number of wells screened for each DOC treatment in DCM and Mesopelagic samples. The phylogeny of SSU rRNA gene sequences of open-ocean isolates is presented in Figures A1 and A2.

Table A1. Experiment I flow cytometry growth screen and recovery of cultures from control and HMW DOM-amended samples. Seawater samples for cultivation were collected during cruise HOT 250 on March of 2013.

<i>Treatments</i>	Control media			DCM (110 m)			Mesopelagic (200 m)		
	<i>No DOC</i>	<i>2.5X DOC</i>	<i>7X DOC</i>	<i>No DOC</i>	<i>2.5X DOC</i>	<i>7X DOC</i>	<i>No DOC</i>	<i>2.5X DOC</i>	<i>7X DOC</i>
Cell count (per ml)									
<5x10 ⁴	207	209	208	625	1254	1061	54	380	7
5x10 ⁴ to 10 ⁵	1	0	0	5	5	18	158	84	70
10 ⁵ to 5x10 ⁵	2	0	1	0	1	13	263	15	317
5x10 ⁵ to 10 ⁶	0	1	1	0	0	0	28	17	66
> 10 ⁶	0	0	0	0	0	0	1	8	44
Total wells	210	210	210	630	1260	1092	504	504	504
Percent Recovery									
>5x10 ⁴	1.4	0.5	1.0	0.8	0.5	2.8	89.3	24.6	98.6

Table A2. Experiment II flow cytometry growth screen and recovery of cultures from control and HMW DOM-amended samples. Seawater samples for cultivation were collected during C-MORE HOE-BOE I on March of 2014.

<i>Treatment</i>	Control media				DCM (95 m)				Mesopelagic (250 m)			
	<i>No DOC</i>	<i>4X DOC</i>	<i>10X DOC</i>	<i>4X DOC F1</i>	<i>No DOC</i>	<i>4X DOC</i>	<i>10X DOC</i>	<i>4X DOC F1</i>	<i>No DOC</i>	<i>4X DOC</i>	<i>10X DOC</i>	<i>4X DOC F1</i>
Cell count (per ml)												
<10 ³	141	143	139	144	694	710	705	704	704	683	667	662
10 ³ to 5x10 ³	3	1	5	0	19	4	4	6	11	23	15	12
5x10 ³ to 10 ⁴	0	0	0	0	1	0	2	0	0	2	12	6
10 ⁴ to 5x10 ⁴	0	0	0	0	2	4	4	4	3	6	6	13
5x10 ⁴ to 10 ⁵	0	0	0	0	1	0	2	1	2	2	7	9
10 ⁵ to 5x10 ⁵	0	0	0	0	2	1	2	4	0	3	10	12
5x10 ⁵ to 10 ⁶	0	0	0	0	0	0	0	1	0	0	2	2
10 ⁶ to 5x10 ⁶	0	0	0	0	1	1	1	0	0	1	1	4
>5x10 ⁶	0	0	0	0	0	0	0	0	0	0	0	0
Total wells	144	144	144	144	720	720	720	720	720	720	720	720
Percent Recovery												
>10 ⁴	0	0	0	0	0.83	0.83	1.25	1.39	0.69	1.67	3.61	5.56

Table A3. Vitamin and trace metal supplements for Hawaii seawater cultivation media.

Vitamins	Final concentration (nM)
Pyrodoxine-HCl	243
Clacium pantothenate	420
PABA-paraaminobenzoic acid	36
Folic acid	2
Vitamin B12	0
Biotin	2
Thiamine	297

Tracel metals	Final concentration (nM)
Manganese sulfate (monohydrate)	179
Zinc chloride	20
Cobalt chloride (hexahydrate)	2
Sodium molybdate (dihydrate)	20
Nickel chloride (hexahydrate)	2
Sodium selenite (pentahydrate)	1

(Opposite page)

Table A4. Genome draft assembly summary of 59 representative open-ocean isolates of cultivation Experiments I and II, as well as of the coastal SAR92 clade isolate NB0015. Assembled genomes were deposited in NCBI as Bioproject PRJNA305749.

Isolate ID	Phylogenetic Id ^a	Reads assembled	%GC	Coverage ^b	Total contigs	Largest contig	N50	Total consensus ^c	Quality ^d
NB0015	SAR92 clade	222378	50	15	470	49503	11234	2630413	83
HI00S01-18 ^e	Halioglobus	8224731	57	330	834	460777	342731	4643869	72
HI00D01-58 ^e	Pseudomonas	7574112	63	257	699	564926	231407	5056098	74
HI00D59-64 ^e	Erythrobacter	3326750	63	172	151	623674	306780	3352205	75
HI00D65-71 ^e	Vibrio	3168151	44	95	748	361131	195360	5339401	73
HI0003	Alcanivorax	552606	59	27	510	89908	25632	3656522	81
HI0007	Alcanivorax	716839	59	35	528	124870	46304	3656241	81
HI0009	Oleiphilus	723138	44	30	2693	241593	16347	6322614	55
HI0011	Alcanivorax	627875	59	26	2529	156362	21670	5638842	58
HI0013	Alcanivorax	850375	59	36	3241	163490	18494	6614072	60
HI0019	Erythrobacter	320380	64	18	984	28738	8598	3360862	81
HI0020	Erythrobacter	586146	63	32	461	146840	41339	3217465	81
HI0021	Oceanibulbus	631044	61	24	928	92580	18120	4494867	81
HI0023	Oceanibulbus	596480	63	26	857	67329	18387	4335530	81
HI0027	Oceanibulbus	1124339	61	45	732	134133	46668	4482328	81
HI0028	Erythrobacter	324735	64	19	668	40850	12736	3317849	44
HI0033	Alcanivorax	488095	59	25	708	94967	20152	3690776	81
HI0035	Alcanivorax	822176	59	35	3156	205575	17383	6694373	59
HI0037	Erythrobacter	774140	63	37	3032	143343	21899	5774452	62
HI0038	Erythrobacter	475039	63	22	2043	87689	14945	4662640	58
HI0040	Oceanibulbus	646904	63	27	932	90981	20339	4230235	81
HI0043	Oleiphilus	707561	44	26	2044	143440	25101	5481930	54
HI0044	Alcanivorax	488420	58	19	2059	48075	10187	5164445	60
HI0049	Roseovarius	567627	64	16	3524	57962	4960	7104809	66
HI0050	Oleiphilus	344906	44	14	2072	28608	4972	4825708	71
HI0053	Pseudoalteromonas	695383	42	19	3130	40367	5334	6742281	70
HI0054	Sulfitobacter	731891	62	31	777	91112	23268	4207901	81
HI0061	Oleiphilus	434002	44	18	1925	62339	8507	5100920	60
HI0063	Erythrobacter	630550	63	33	586	161626	41202	3293092	81
HI0065	Oleiphilus	818810	48	35	2711	336635	28355	6401110	56
HI0066	Oleiphilus	560893	44	25	1926	150802	22410	5050227	54
HI0067	Oleiphilus	503329	44	22	1789	119737	20427	4845662	55
HI0068	Oleiphilus	365219	44	14	2039	21553	5157	4848124	71
HI0069	Oleiphilus	590352	44	22	2264	96705	14127	5905670	53
HI0071	Oleiphilus	601234	48	26	2311	143853	24676	5701391	53
HI0072	Oleiphilus	337845	44	14	2241	25254	4687	4928926	71
HI0073	Oleiphilus	784137	48	34	2660	418796	56040	6107507	55
HI0074	Erythrobacter	379698	64	22	747	48613	12515	3174080	81
HI0075	Oceanobacter	535677	51	22	2622	98856	12664	6125644	56
HI0076	Oceanibulbus	574934	61	23	1070	47234	13536	4400770	81
HI0077	Erythrobacter	365940	64	21	667	74328	13723	3173227	81
HI0078	Oleiphilus	303674	44	12	2589	16558	2900	4440923	78
HI0079	Oleiphilus	426827	47	19	1492	48425	13007	4413786	58
HI0080	Oleiphilus	377861	47	16	1308	33760	8068	3946656	72
HI0081	Oleiphilus	298111	44	12	2517	15225	3026	4411065	78
HI0082	Oceanibulbus	616668	61	21	2364	49576	9483	5450880	62
HI0083	Alcanivorax	345236	58	14	1829	29024	5595	4452193	73
HI0085	Oleiphilus	253579	44	11	3293	15089	2070	4492359	78
HI0086	Oleiphilus	506633	44	20	1604	52346	10771	4764100	63
HI0117	Oleiphilus	546950	44	19	1632	35441	7173	4636056	75
HI0118	Oleiphilus	1204101	48	49	2842	443600	15926	6562827	55
HI0120	Oceanospirillaceae	766145	49	27	2130	225416	35705	5682561	55
HI0122	Oleiphilus	743601	48	32	2536	358202	37188	6045884	55
HI0123	Oleiphilus	545524	44	17	2100	21032	4967	4493975	79
HI0125	Oleiphilus	605193	44	25	1616	109922	19459	4565113	55
HI0128	Oleiphilus	568083	44	20	1725	98189	11624	4898028	59
HI0129	Oceanibulbus	395966	63	17	1228	81523	8497	4223361	81
HI0130	Oleiphilus	433776	48	18	2114	62064	10256	5203078	57
HI0132	Oleiphilus	501301	44	20	1963	51478	9469	5179518	61
HI0133	Oleiphilus	869569	48	36	2428	329080	47727	5733204	55

^aGenus of most closely related sequence based on LAST analysis against SILVA 119

^bAverage total coverage assessment calculated from contigs >= 5000 with uniform coverage

^cContig size >=500

^dAverage consensus quality

^eAssemblies for which the raw reads of several dilution-to-extinction cultures that appeared clonal based on SSU rRNA identity were pooled

(Opposite page)

Table A5. PROKKA annotation summary of the draft genomes of 59 representative open-ocean isolates and the SAR92 clade isolate NB0015. The number of PROKKA-predicted CDSs matching known and hypothetical proteins in the RefSeq and KEGG databases are also indicated. Draft annotations of genomes are available under NCBI BioProject PRJNA305749.

Isolate ID	Phylogenetic ID ^a	Conitgs ^b	Bases ^b	rRNAs	SSU	LSU	tRNAs	CDS	RefSeq		KEGG	
									Annot. ^c	Hyp. ^d	Annot. ^c	Hyp. ^d
NB0015	SAR92 clade	467	2629892	3	1	1	29	2401	1466	672	1478	636
HI00S01-18e	Halioglobus	415	4588515	6	1	1	57	4293	4236	1266	2006	1248
HI00D01-58e	Pseudomonas	607	5044166	10	1	1	77	4676	2749	1851	2700	1871
HI00D59-64e	Erythrobacter	121	3347358	4	1	1	48	3217	3169	1218	1641	1245
HI00D65-71e	Vibrio	614	5319010	21	2	1	187	4707	4520	2146	2133	2320
HI0003	Alcanivorax	378	3635468	38	1	1	38	3338	1709	1389	1717	1364
HI0007	Alcanivorax	286	3620498	5	1	1	45	3303	1699	1369	1703	1350
HI0009	Oleiphilus	2650	6316839	9	2	1	90	5721	2716	1682	2696	1636
HI0011	Alcanivorax	2482	5632014	5	1	0	65	5387	2797	2143	2812	2102
HI0013	Alcanivorax	3199	6607668	7	2	1	72	6449	3444	2511	3471	2456
HI0019	Erythrobacter	880	3344147	3	1	1	41	3290	1672	1249	1658	1240
HI0020	Erythrobacter	275	3189298	4	1	1	45	3091	1580	1186	1565	1187
HI0021	Oceanibulbus	717	4462052	6	1	1	46	4403	2481	1542	2405	1575
HI0023	Oceanibulbus	652	4303630	3	1	1	46	4219	2377	1534	2308	1560
HI0027	Oceanibulbus	458	4441089	6	1	1	48	4343	2428	1542	2351	1575
HI0028	Erythrobacter	610	3308918	3	1	1	44	3239	1637	1234	1620	1230
HI0033	Alcanivorax	509	3659598	5	1	1	34	3355	1729	1379	1735	1358
HI0035	Alcanivorax	3128	6690527	7	2	2	68	6566	3471	2558	3486	2513
HI0037	Erythrobacter	2993	5769061	7	2	2	77	5795	3060	2036	3030	2039
HI0038	Erythrobacter	2004	4656539	5	1	2	70	4641	2405	1688	2383	1686
HI0040	Oceanibulbus	611	4181428	2	1	0	42	4079	2310	1475	2240	1505
HI0043	Oleiphilus	1975	5472755	10	2	2	49	4693	2396	1515	2378	1500
HI0044	Alcanivorax	1990	5153888	7	2	1	7	5021	2572	2072	2599	2032
HI0049	Roseovarius	3470	7096493	4	2	0	45	7101	3883	2403	3822	2358
HI0050	Oleiphilus	2036	4819895	9	1	2	38	4043	2022	1286	2020	1269
HI0053	Pseudoalteromonas	2964	6718721	6	2	3	117	5484	2939	2055	2860	2039
HI0054	Sulfitobacter	555	4174729	6	1	1	47	4099	2304	1427	2217	1465
HI0061	Oleiphilus	1885	5095272	8	1	2	57	4352	2186	1439	2167	1433
HI0063	Erythrobacter	356	3258146	4	1	1	48	3059	1559	1126	1546	1132
HI0065	Oleiphilus	2693	6398614	10	2	2	67	5497	2868	1800	2880	1719
HI0066	Oleiphilus	1887	5044867	10	2	2	52	4447	2341	1428	2350	1370
HI0067	Oleiphilus	1752	4840075	9	2	2	53	4291	2245	1369	2246	1323
HI0068	Oleiphilus	2001	4842151	9	2	2	46	4173	2083	1338	2047	1348
HI0069	Oleiphilus	2238	5901752	10	2	2	57	5080	2540	1634	2517	1623
HI0071	Oleiphilus	2289	5698317	10	2	2	63	4940	2570	1620	2569	1568
HI0072	Oleiphilus	2192	4921543	7	2	1	38	4140	2087	1321	2069	1303
HI0073	Oleiphilus	2628	6103244	12	2	2	73	5230	2706	1736	2702	1676
HI0074	Erythrobacter	583	3148101	4	1	1	42	3112	1585	1151	1573	1146
HI0075	Oceanobacter	2573	6118559	9	2	2	57	5706	2772	2367	2803	2306
HI0076	Oceanibulbus	875	4370296	5	1	1	46	4312	2423	1507	2358	1528
HI0077	Erythrobacter	592	3161113	5	1	1	45	3091	1568	1152	1556	1146
HI0078	Oleiphilus	2530	4431832	9	1	1	22	3679	1839	1163	1807	1166
HI0079	Oleiphilus	1448	4407017	9	2	2	43	3867	2011	1257	2009	1214
HI0080	Oleiphilus	1249	3938001	7	1	1	43	3439	1803	1143	1810	1101
HI0081	Oleiphilus	2463	4402732	9	2	1	30	3655	1839	1159	1837	1131
HI0082	Oceanibulbus	2310	5443144	8	2	1	68	5472	3165	1788	3065	1825
HI0083	Alcanivorax	1795	4446909	7	2	2	50	4404	2293	1811	2312	1782
HI0085	Oleiphilus	3230	4482542	8	1	1	40	3669	1831	1115	1824	1093
HI0086	Oleiphilus	1534	4753747	10	2	2	41	4080	2063	1351	2035	1349
HI0117	Oleiphilus	1521	4620570	9	2	2	37	3984	1959	1287	1947	1277
HI0118	Oleiphilus	2826	6560891	11	2	2	77	5687	2946	1819	2945	1754
HI0120	Thalassolituus	2071	5674715	11	2	2	73	4960	2430	1876	2422	1843
HI0122	Oleiphilus	2490	6039240	11	2	2	61	5133	2681	1677	2673	1623
HI0123	Oleiphilus	1909	4468126	8	2	1	41	3752	1846	1203	1829	1198
HI0125	Oleiphilus	1531	4553828	8	2	2	41	4033	2148	1286	2147	1247
HI0128	Oleiphilus	1636	4885910	7	1	1	46	4166	2114	1354	2128	1312
HI0129	Oceanibulbus	1145	4210493	3	1	1	36	4170	2358	1488	2285	1522
HI0130	Oleiphilus	2071	5196716	9	2	2	65	4544	2313	1455	2278	1433
HI0132	Oleiphilus	1901	5170237	8	2	1	48	4436	2215	1441	2195	1436
HI0133	Oleiphilus	2393	5728633	10	2	2	68	4947	2587	1604	2569	1556

^aGenus of most closely related sequence based on LAST analysis against SILVA 119

^bThe contigs annotated by PROKKA and the total nucleotides comprising those contigs

^cCDSs matching a known function

^dCDSs matching a hypothetical function

^eAssemblies for which the raw reads of several dilution-to-extinction cultures that appeared clonal based on SSU rRNA identity were pooled

Table A6. One-way ANOVA comparisons of the mean number of CAZy-predicted genes for each CAZy class (AA, CBM, CE, GH, GT, and PL). Hawaii isolate phylotypes Oleiphilus 1, Oleiphilus 2, Alcanivorax, Erythrobacter 1, Erythrobacter 2, and Sulfitobacter/Oceanibulbus were compared.

	DF	Sum Square	Mean Square	F-value	P-value
AA	5	1561.8	312.36	16.49	4.54E-09
Residuals	43	814.6	18.94		
CBM	5	132.66	26.531	18.11	1.23E-09
Residuals	43	62.98	1.465		
CE	5	1200.7	240.14	11.69	3.62E-07
Residuals	43	883.3	20.54		
GH	5	761.9	152.39	17.66	1.76E-09
Residuals	43	371.1	8.63		
GT	5	8769	1754	23.69	2.33E-11
Residuals	43	3184	74		
PL	5	2.646	0.5293	1.217	0.318
Residuals	43	18.701	0.4349		

Abbreviations: DF, degrees of freedom; AA, auxiliary activities; CBM, carbohydrate binding modules; CE, carbohydrate esterases; GH, glycoside hydrolases; GT, glycosyl transferases; PL, polysaccharide lyases

Table A7. The three C-P lyase gene clusters identified by Martinez et al., 2013 in microcosm experiments amended with methyl phosphonate. The amino acid sequences of these genes were used as queries to find homologs in Hawaii isolate genomes.

C-P lyase protein	Accession	C-P lyase protein	Accession	C-P lyase protein	Accession
B3TF_MPn1 (<i>gamma</i> IMCC1989)		B3TF_MPn2 (<i>Rhodobacterales</i>)		B3TF_MPn8 (<i>Vibrio nigripulchritudo</i> ATCC27043)	
PT	AHG53029.1	PhnC	AHG53005.1	PhnC	AHG52968.1
PhnN	AHG53030.1	PhnD	AHG53006.1	PhnD	AHG52969.1
PhnM	AHG53031.1	PhnE1	AHG53007.1	PhnE	AHG52970.1
PhnL	AHG53032.1	PhnE2	AHG53008.1	PhnF	AHG52971.1
PhnK	AHG53033.1	PT	AHG53009.1	PhnG	AHG52972.1
PhnJ	AHG53034.1	PhnM1	AHG53010.1	PhnH	AHG52973.1
PhnI	AHG53035.1	PhnF	AHG53011.1	PhnI	AHG52974.1
PhnH	AHG53036.1	PhnG	AHG53012.1	PhnJ	AHG52975.1
PhnG	AHG53037.1	PhnH	AHG53013.1	PhnK	AHG52976.1
PhnE	AHG53038.1	PhnI	AHG53014.1	PhnL	AHG52977.1
PhnD	AHG53039.1	PhnJ	AHG53016.1	PhnM	AHG52978.1
PhnC	AHG53040.1	PhnK	AHG53017.1	PhnN	AHG52979.1
PAT	AHG53041.1	PhnL	AHG53018.1	PhnP	AHG53043.1
PhnF	AHG53042.1	PhnN	AHG53019.1		
		RcsF	AHG53020.1		
		PhnM2	AHG53021.1		

Abbreviations: GI, GenBank identifier; PT, putative transferase; PAT, putative acyl-CoA N-acetyltransferase

(Opposite page)

Table A8. The predicted C-P lyase genes in Hawaii isolate genomes and their PROKKA CDS identification number. The phylogeny and isolate ID are indicated for each gene cluster. PROKKA CDSs are available under NCBI BioProject PRJNA305749.

Isolate ID	PROKKA CDS	C-P lyase gene	Isolate ID	PROKKA CDS	C-P lyase gene	Isolate ID	PROKKA CDS	C-P lyase gene
<i>Pseudomonas stutzeri</i>			<i>Rhodobacteraceae</i>			<i>Oceanospirillaceae</i>		
HI00D01	874	<i>phnC</i>	HI0021	574	<i>phnC</i>	HI0009	4933	<i>phnC</i>
	875	<i>phnD</i>		573	<i>phnD</i>		4932	<i>phnD</i>
	876	<i>phnE</i>		571	<i>phnE</i>		4934	<i>phnE</i>
	877	<i>phnF</i>		572	<i>phnE</i>		4931	<i>phnF</i>
	878	<i>phnG</i>		570	<i>phnF</i>		4935	<i>phnG</i>
	879	<i>phnH</i>		569	<i>phnG</i>		4936	<i>phnH</i>
	880	<i>phnI</i>		568	<i>phnH</i>		4937	<i>phnI</i>
	881	<i>phnJ</i>		567	<i>phnI</i>		4938	<i>phnJ</i>
	882	<i>phnK</i>		566	<i>phnJ</i>		4939	<i>phnK</i>
	883	<i>phnL</i>		565	<i>phnK</i>		4940	<i>phnL</i>
	884	<i>phnM</i>		564	<i>phnL</i>		4941	<i>phnM</i>
				561	<i>phnM</i>			
<i>Rhodobacteraceae</i>			HI0027	2282	<i>phnC</i>	HI0066	2175	<i>phnC</i>
HI0049	5135	<i>phnC</i>		2283	<i>phnD</i>		2176	<i>phnD</i>
	5136	<i>phnD</i>		2284	<i>phnE</i>		2174	<i>phnE</i>
	5137	<i>phnE</i>		2285	<i>phnE</i>		2177	<i>phnF</i>
	5138	<i>phnE</i>		2286	<i>phnF</i>		2173	<i>phnG</i>
	3370	<i>phnF</i>		2287	<i>phnG</i>		2172	<i>phnH</i>
	3371	<i>phnG</i>		2288	<i>phnH</i>		2171	<i>phnI</i>
	3372	<i>phnH</i>		2289	<i>phnI</i>		2170	<i>phnJ</i>
	3373	<i>phnI</i>		2290	<i>phnJ</i>		2169	<i>phnK</i>
	3375	<i>phnJ</i>		2291	<i>phnK</i>		2168	<i>phnL</i>
	3376	<i>phnK</i>		2292	<i>phnL</i>		2167	<i>phnM</i>
	4894	<i>phnL</i>		2295	<i>phnM</i>			
	5140	<i>phnM</i>	HI0076	3373	<i>phnC</i>	HI0067	2616	<i>phnC</i>
	4891	<i>phnM</i>		3374	<i>phnD</i>		2617	<i>phnE</i>
HI0054	442	<i>phnC</i>		3375	<i>phnE</i>		2618	<i>phnG</i>
	441	<i>PhnD</i>		3376	<i>phnE</i>		2619	<i>phnH</i>
	439	<i>PhnE</i>		3377	<i>phnF</i>		2620	<i>phnI</i>
	440	<i>phnE</i>		3378	<i>phnG</i>		2621	<i>phnJ</i>
	437	<i>phnF</i>		3379	<i>phnH</i>		2622	<i>phnK</i>
	436	<i>phnG</i>		3380	<i>phnI</i>		2623	<i>phnL</i>
	435	<i>phnH</i>		3381	<i>phnJ</i>		2624	<i>phnM</i>
	434	<i>phnI</i>		3382	<i>phnK</i>			
	433	<i>phnJ</i>		3383	<i>phnL</i>	HI0125	1618	<i>phnC</i>
	432	<i>phnK</i>		3386	<i>phnM</i>		1619	<i>phnD</i>
	431	<i>phnL</i>	HI0082	428	<i>phnC</i>		1617	<i>phnE</i>
	428	<i>phnM</i>		429	<i>phnD</i>		1620	<i>phnF</i>
				430	<i>phnE</i>		616	<i>phnG</i>
				431	<i>phnE</i>		1616	<i>phnG</i>
				432	<i>phnF</i>		617	<i>phnH</i>
				433	<i>phnG</i>		618	<i>phnI</i>
				434	<i>phnH</i>		619	<i>phnJ</i>
				435	<i>phnI</i>		3860	<i>phnK</i>
				436	<i>phnJ</i>		3861	<i>phnL</i>
				437	<i>phnK</i>		3862	<i>phnM</i>
				438	<i>phnL</i>			
				441	<i>phnM</i>			

(Opposite page)

Table A9. SSU rRNA genes of 59 Hawaii isolates and coastal isolate NB0015 identified in PROKKA draft genome assemblies. Sequences are available in the draft genome assemblies deposited in NCBI as BioProject PRJNA305749. The draft assemblies contained one or two copies of the SSU rRNA gene and belonged to the same phylogenetic group.

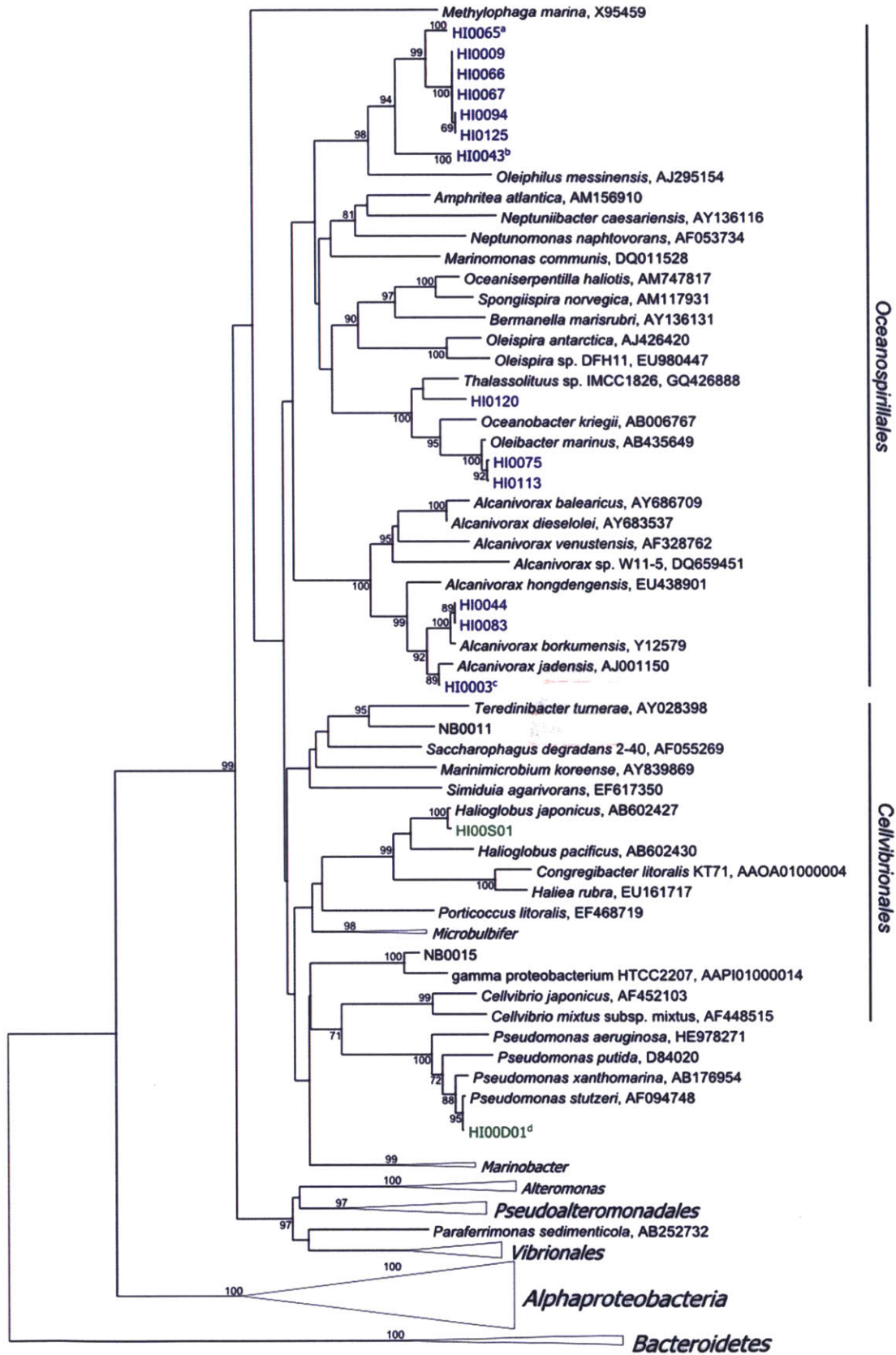
Isolate ID	PROKKA ID	Length (bps)	Isolate ID	PROKKA ID	Length (bps)
NB0015 ^a	763	1506	HI0068	4113	1379
HI00S01-18 ^b	4326	1492	HI0069	5063	1530
HI00D01-58 ^b	4617	1533	HI0069	5123	1530
HI00D59-64 ^b	3259	1481	HI0071	4986	1530
HI00D65-71 ^b	4734	1545	HI0071	4994	1515
HI00D65-71 ^b	4752	1551	HI0072	4043	1530
HI0003	3376	1537	HI0072	4172	1239
HI0007	3348	1537	HI0073	5290	1530
HI0009	5772	1530	HI0073	5291	1530
HI0009	5788	1530	HI0074	3138	1486
HI0011	5370	1537	HI0075	5700	1529
HI0013	6467	1537	HI0075	5707	1340
HI0013	6518	1405	HI0076	4259	1461
HI0019	3309	1486	HI0077	3117	1486
HI0020	3121	1484	HI0078	3567	1530
HI0021	4354	1461	HI0079	3841	1530
HI0023	2280	1457	HI0079	3850	1530
HI0027	4296	1461	HI0080	3472	1530
HI0028	3264	1486	HI0081	3547	1530
HI0033	3383	1537	HI0081	3558	1024
HI0035	6594	1537	HI0082	5253	1461
HI0035	6629	1537	HI0082	5469	1461
HI0037	5825	1484	HI0085	3515	1530
HI0037	5839	1484	HI0086	4018	1530
HI0038	4638	1484	HI0086	4124	1530
HI0040	1485	1457	HI0117	3928	1530
HI0043	4662	1530	HI0117	4019	1495
HI0043	4737	1530	HI0118	5741	1530
HI0044	4931	1537	HI0118	5750	1530
HI0044	5064	1537	HI0120	5003	1535
HI0049	1640	1329	HI0120	5004	1535
HI0049	6443	1379	HI0122	5178	1530
HI0050	3958	1530	HI0122	5188	1530
HI0053	5419	1538	HI0123	3677	1530
HI0053	5431	1538	HI0123	3794	1463
HI0054	4070	1459	HI0125	4036	1530
HI0061	4312	1530	HI0125	4043	1530
HI0063	972	1063	HI0128	4137	1530
HI0065	5547	1530	HI0129	3432	1295
HI0065	5555	1530	HI0130	4507	1530
HI0066	4459	1530	HI0130	4542	1530
HI0066	4460	1467	HI0132	4394	1530
HI0067	4307	1530	HI0132	4476	1125
HI0067	4350	1530	HI0133	4995	1530
HI0068	4099	1530	HI0133	5001	1530

^aSAR92 clade gammaproteobacteria NB0015 (Chapter 2)

^bExperiment I Hawaii isolates. Assemblies of related isolates were pooled.

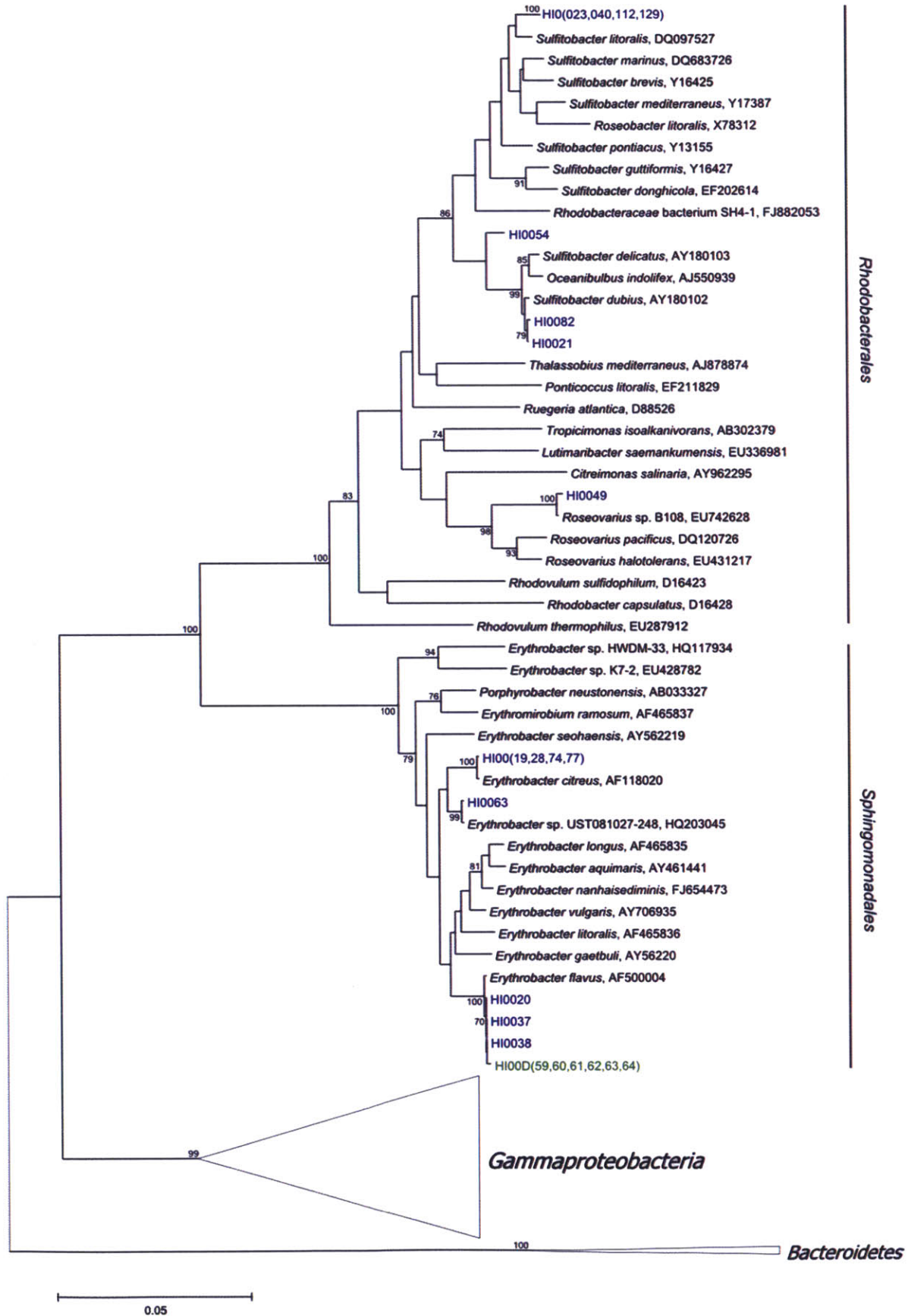
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Figure A1. SSU rRNA gene phylogeny of Hawaii open-ocean isolates belonging to the *Gammaproteobacteria* orders *Oceanospirillales* and *Cellvibrionales*. Nodes with >70% agreement (1000 replicates) are shown. The scale bar indicates substitutions per site. Marked in green are sequences identified from Experiment I samples and in green, sequences from Experiment II. ^aGrouping of HI0065 and 14 closely related sequences. ^bGrouping of HI0043 and 17 closely related sequences. ^cGrouping of HI0003 and 5 closely related sequences from the DCM. ^dGrouping of 41 *P. stutzeri* sequences.



(Opposite page)

Figure A2. The SSU rRNA gene phylogeny of Hawaii open-ocean isolates belonging to the *Alphaproteobacteria* orders *Rhodobacterales* and *Sphingomonadales*. Nodes with >70% agreement (1000 replicates) are shown. The scale bar indicates substitutions per site. Only sequences >1400 bps long were included. Marked in green are sequences identified from Experiment I samples and in green, sequences from Experiment II.



APPENDIX II

This section describes the defined media formulation for growth of *Pseudomonas stutzeri* HI00D01 (Table A1).

Table A1. MOPS minimal formulations (<http://www.teknova.com/category-s/519.htm>) used for the growth of *P. stutzeri* strain H00D01.

M2101 MOPS Modified Rich Buffer	10X Concentration	M2105/M2106 1X Concentration
MOPS (MW 209.3)	400mM	40mM
Tricine (MW 179.2)	40mM	4.0mM
Iron Sulfate Stock	0.1mM	0.01mM
Ammonium Chloride	95mM	9.5mM
Potassium Sulfate	2.76mM	0.276mM
Calcium Chloride	0.005mM	0.0005mM
Magnesium Chloride	5.25mM	0.525mM
Sodium Chloride	500mM	50mM
Ammonium Molybdate	3×10^{-8} M	3×10^{-9} M
Boric Acid	4×10^{-6} M	4×10^{-7} M
Cobalt Chloride	3×10^{-7} M	3×10^{-8} M
Cupric Sulfate	10^{-7} M	10^{-8} M
Manganese Chloride	8×10^{-7} M	8×10^{-8} M
Zinc Sulfate	10^{-7} M	10^{-8} M