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Citation: Braakman, Rogier et al. "Metabolic Evolution and the Self-Organization of Ecosystems." Proceedings of the National Academy of Sciences 114, 15 (March 2017): E3091–E3100 © 2017 National Academy of Sciences

As Published: <http://dx.doi.org/10.1073/PNAS.1619573114>

Publisher: National Academy of Sciences (U.S.)

Persistent URL: <http://hdl.handle.net/1721.1/112214>

Version: Final published version: final published article, as it appeared in a journal, conference proceedings, or other formally published context

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Metabolic evolution and the self-organization of ecosystems

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Metabolism mediates the flow of matter and energy through the biosphere. We examined how metabolic evolution shapes ecosystems by reconstructing it in the globally abundant oceanic phytoplankter *Prochlorococcus*. To understand what drove observed evolutionary patterns, we interpreted them in the context of its population dynamics, growth rate, and light adaptation, and the size and macromolecular and elemental composition of cells. This multilevel view suggests that, over the course of evolution, there was a steady increase in *Prochlorococcus*' metabolic rate and excretion of organic carbon. We derived a mathematical framework that suggests these adaptations lower the minimal subsistence nutrient concentration of cells, which results in a drawdown of nutrients in oceanic surface waters. This, in turn, increases total ecosystem biomass and promotes the coevolution of all cells in the ecosystem. Additional reconstructions suggest that *Prochlorococcus* and the dominant cooccurring heterotrophic bacterium SAR11 form a coevolved mutualism that maximizes their collective metabolic rate by recycling organic carbon through complementary excretion and uptake pathways. Moreover, the metabolic codependencies of *Prochlorococcus* and SAR11 are highly similar to those of chloroplasts and mitochondria within plant cells. These observations lead us to propose a general theory relating metabolic evolution to the self-amplification and self-organization of the biosphere. We discuss the implications of this framework for the evolution of Earth's biogeochemical cycles and the rise of atmospheric oxygen.

metabolic evolution | *Prochlorococcus* | microbial oceanography | mutualism | Earth history

Metabolism sustains the nonequilibrium chemical order of the biosphere by continually supplying the energy and building blocks of all cells on Earth (1–5). Here we ask: How does cellular metabolic evolution shape the mass and energy flows of ecosystems? The oceanic phytoplankter *Prochlorococcus* (6), the most abundant photosynthetic cell on Earth (7, 8), provides an ideal model system for addressing this question. *Prochlorococcus* and its deeper-branching sister lineage marine *Synechococcus* make up the marine picocyanobacteria and have a characteristic biogeography (9). *Prochlorococcus* “ecotypes” have geographically (10, 11) and seasonally (12) dynamic populations that in warm, stable stratified water columns always return to the same general structure: Recently diverging high-light-adapted (HL) ecotypes are most abundant toward the surface, whereas deeper branching low-light-adapted (LL) ecotypes are most abundant at depth (10–14) (Fig. 1).

What selective forces drove this niche partitioning in *Prochlorococcus*, and what were the consequences for the ocean ecosystem in general? To address these questions, we reconstructed (15, 16) (Fig. 2) the evolution of core metabolism in strains representing the major clades of *Prochlorococcus*. To interpret the observed patterns, we developed an evolutionary framework that illuminates the driving forces that produced many of the features of *Prochlorococcus*. Using this framework, we argue that the evolutionary patterns manifested in the *Prochlorococcus* collective reflect fundamental processes shaping the coevolution of the chemistry of the oligotrophic oceans and the microbial ecosystems they harbor.

Results and Discussion

Metabolism Provides Clues About Large-Scale Evolutionary Driving Forces. To examine the driving forces shaping the evolution of *Prochlorococcus*, we reconstructed (Fig. 2) the evolution of its metabolic core (SI Appendix, Fig. S1). Because all biosynthetic pathways originate there, its evolution is highly constrained, and any innovations likely reflect major driving forces (5). Previous studies identified some unique presence/absence patterns for core metabolic genes in both *Prochlorococcus* and *Synechococcus* and relative to other cyanobacteria. For example, both have replaced the cyanobacterial RuBisCO and related proteins of the CO₂-concentrating mechanism with proteobacterial variants (17–19), lost key genes in the cyanobacterial TCA cycle (20) and glycolysis (21), and acquired a menaquinone-based malate dehydrogenase (22). Photorespiration proteins are, in turn, universally preserved in *Synechococcus*, but unevenly distributed across *Prochlorococcus* ecotypes (9, 23).

Expanding on these studies, we surveyed cyanobacteria for the presence/absence of core metabolic genes (SI Appendix, Table S1 and Fig. S1) and mapped this distribution onto their phylogeny to reconstruct a phylometabolic tree (Fig. 2) that resolves the evolution of *Synechococcus* and *Prochlorococcus* (Fig. 1). All of the variations are part of a sequence of innovations that remodeled the metabolic core as *Prochlorococcus* diverged from the rest of cyanobacteria (Fig. 3). Key innovations occur in ancestral freshwater lineages (Fig. 3 and SI Appendix, Table S1), indicating that underlying selection pressures preceded the emergence of the marine lineages.

Significance

Understanding what drives self-organization in complex systems and how it arises is a major challenge. We addressed this challenge using dominant oceanic photosynthetic and heterotrophic microbes as a model system. Reconstructing the metabolic evolution of this system suggests that its self-organization and self-amplification were coupled and driven by an increasing cellular energy flux. Specifically, the evolution of cells steadily increased their metabolic rate and excretion of organic carbon. We describe how this increases cellular nutrient uptake and thereby ecosystem biomass. The release of organic carbon, in turn, promotes positive feedbacks among species that reinforce this evolutionary drive at the ecosystem level. We propose the evolutionary self-organization of oceanic microbial ecosystems contributed to the oxygenation of Earth.

Author contributions: R.B. designed research; R.B. performed research; R.B. contributed new reagents/analytic tools; R.B. developed the model; M.J.F. contributed to model development; R.B. and S.W.C. analyzed data; and R.B., M.J.F., and S.W.C. wrote the paper. Reviewers: R.M., Weizmann Institute; and J.A.R., University of Dundee, U.K.

The authors declare no conflict of interest.

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This article contains supporting information online at www.pnas.org/lookup/suppl/doi:10.1073/pnas.1619573114/-DCSupplemental.

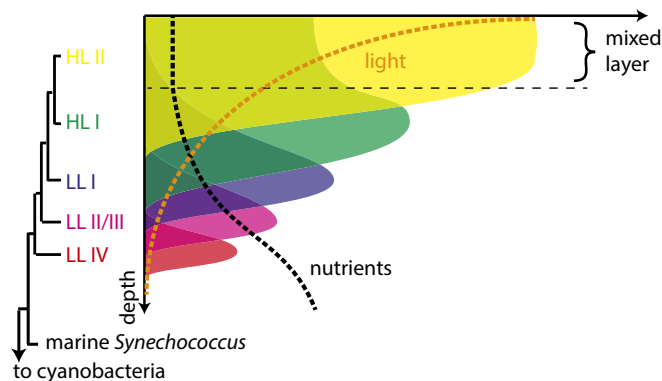


Fig. 1. Typical relative abundance distributions of *Prochlorococcus* ecotypes as a function of depth and accompanying light intensity and nutrient concentration profiles in stratified oceanic waters. Ecotype populations are geographically and temporally dynamic, but in warm, stable water columns return to this same depth-differentiated state (12). The deepest branching ecotypes are most abundant at the bottom of the euphotic zone, where nutrient concentrations are high and light energy low. The most recently diverging ecotypes are most abundant near the surface, where the reverse is true (10–14). HL, high-light-adapted; LL, low-light-adapted.

The remodeling of *Prochlorococcus*' metabolic core includes the disruption of photorespiration and the TCA cycle (Fig. 3 and *SI Appendix, Fig. S1*), raising the possibility that intermediates of truncated pathways are excreted from the cell. Phytoplankton commonly excrete organic carbon as an outlet of excess reducing power under nutrient limitation or intense light (24). Analogously, when facing a large energy supply from organic carbon, some heterotrophs will effectively drain reducing power into the environment by excreting partially oxidized organic carbon rather than fully oxidizing it to CO_2 (25–28), whereas some photoheterotrophs use CO_2 -fixation as a sink for excess reducing power (29). For photosynthetic cells in the oligotrophic surface oceans, where the solar energy supply may commonly outpace the nutrient supply (Fig. 1), the combination of increased CO_2 fixation and increased excretion of organic carbon could thus well be under strong selection. This is consistent with observations that *Prochlorococcus* has the most efficient carbon-concentrating mechanism (30) and highest known rate of CO_2 fixation per photosynthetic pigment (31) of any phytoplankton, even though its small size and slow growth (6, 32) suggest a relatively small carbon flux requirement. Lastly, selection to rid the system of excess reducing power is consistent with the acquisition of the Plastocyanin Terminal Oxidase (PTOX) in the LLI and HL clades of *Prochlorococcus* (Fig. 3 and *SI Appendix, Table S1*), which adds an additional outlet for excess reducing power in their photosynthetic electron transfer chain (*SI Appendix, Fig. S1*) (33–36).

That the evolution of *Prochlorococcus* permanently increased the excretion of organic carbon in its late-branching strains is also consistent with limited laboratory observations. At moderate light levels in nutrient replete medium, strains from HL clades that dominate surface waters (Fig. 1) excrete up to 20% of fixed carbon, and recently diverging strains excrete the most (37). Cells excrete significant amounts of glycolate (37), one of the dead ends in the metabolic network (*SI Appendix, Fig. S1*). P-starved cells excrete slightly less carbon (37), but this could be due to the coincident cessation of growth. Nevertheless, a higher fraction of the carbon excreted by P-starved cells consists of glycolate and other small carboxylic acids (37). Many phytoplankton excrete glycolate under intense light or nutrient limitation (24, 38), but retain the capacity to recycle it by using the three-subunit iron-sulfur protein glycolate oxidase (GlcDEF, rxn 13 in *SI Appendix, Fig. S1*), shown to be the enzymatic workhorse for this function in cyanobacteria (23). However, this gene is absent in all but the deepest-branching LLIV clade of *Prochlorococcus* (*SI Appendix, Table S1*), suggesting a permanent opening of this pathway early in its evolution (Fig. 3). Finally, the bulk of organic carbon excreted by *Synechococcus* consists of polysaccharides (39), which are commonly excreted by nutrient-limited microbes (26, 40), suggesting that these compounds could similarly act as a redox safety valve in *Prochlorococcus*, which dominates in one of the most nutrient-poor environments on Earth (7–9) (Fig. 1).

We further examined the possibility that the evolution of *Prochlorococcus* increased its excretion of organic carbon as an outlet of excess reducing power (Fig. 3) through additional genomic analyses. Functionally related and coexpressed genes are commonly located near each other, so we searched the genomic neighborhoods of core metabolic genes for transporters across clades (*SI Appendix, Fig. S2*). We identified chromosome rearrangements repositioning a series of transporters, including three export and one import transporters, near key metabolic genes, consistent with selection acting to fortify the control of transport pathways (*SI Appendix, Fig. S2*). Chromosome rearrangements are seen in freshwater picocyanobacteria (*SI Appendix, Fig. S2*), again suggesting that these pathways came under selection before the emergence of the marine lineages and the loss of photorespiration in the LLII/III clade of *Prochlorococcus* (Fig. 3 and *SI Appendix, Fig. S1*). This analysis suggests that, in addition to glycolate, pyruvate and citrate (or isocitrate) are exported, whereas malate is imported (*SI Appendix, Figs. S1 and S2*).

Furthermore, because environmentally driven variations in the expression of genes can give insight into their function [i.e., “reverse ecology” (41)], and the metabolism of *Prochlorococcus* is highly choreographed to the diel light:dark cycle (42, 43), we examined the gene expression of a HL-adapted strain grown under a diel cycle (42). All putative export pathway genes have mRNA expression maxima at sunrise (*SI Appendix, Fig. S2*), consistent with their acting as redox safety valves, activated when the supply of reducing power increases at dawn. Similarly,

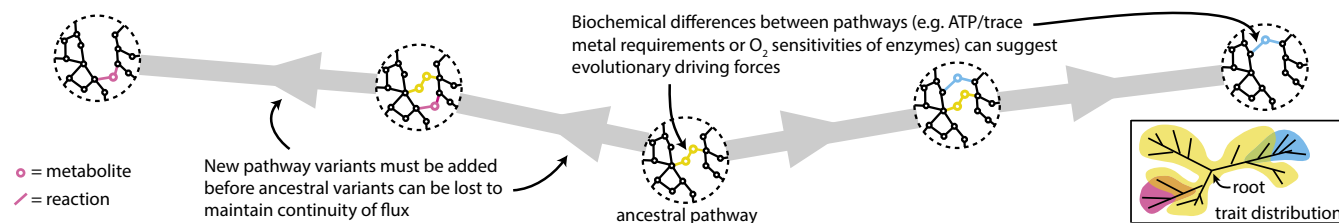


Fig. 2. Illustration of approach to metabolic reconstructions. Phylometabolic trees reflect the evolution of metabolic network phenotypes because they integrate constraints from phylogenetics and metabolism (15, 16). All sequenced genomes within a given clade are searched for the presence/absence of enzymes catalyzing the reactions of different pathways. Mapping pathway variability patterns onto phylogenies of the clade suggests the order of metabolic innovations. In this example, three alternative pathways (pink, yellow, and blue) connect essential and universal pathways (black). Genes for the yellow pathway are nearly universally distributed (*Inset*), suggesting that it is the ancestral pathway, with the pink and blue pathways deriving from it. Maintaining continuity of flux results in trees of functional phenotypes. Biochemical differences between alternative pathways (e.g., ATP/trace metal requirements or oxygen sensitivities of their enzymes) suggest evolutionary driving forces (15, 16).

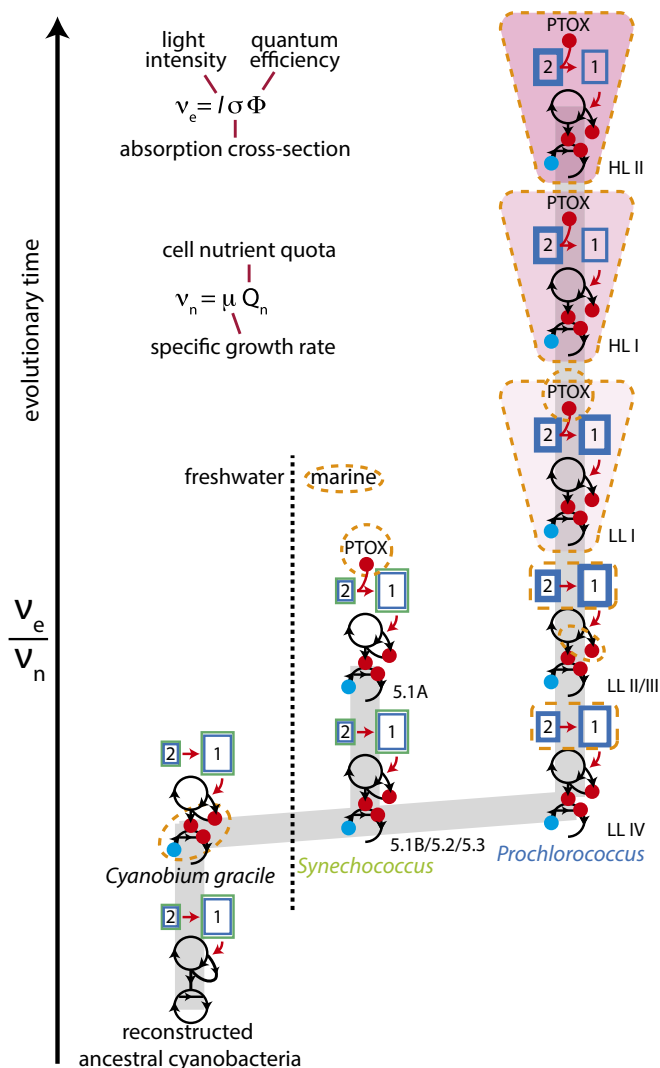


Fig. 3. Metabolic evolution of *Prochlorococcus*. The metabolic variants are represented in simplified form along a gradient of the ratio of electron flux to nutrient flux ν_e/ν_n . Innovations (*SI Appendix*, Fig. S1 and Table S1) are highlighted with dashed orange lines. Cellular electron drains are represented as red dots and uptake pathways as blue dots. Photosystems are represented as boxed 2's and 1's, and their colors reflect the major absorption wavelengths, their line thickness reflects their absorption cross-section, and their relative size reflects the PSII:PSI ratio (*SI Appendix*, Fig. S1). Key innovations in the HLI and HLII clades are related to protection/repair of direct photodamage (52, 53), including UV damage, as indicated by a darkening protective pink shade.

the expression of PTOX, the most immediate outlet for excess reducing power in the electron transfer chain (*SI Appendix*, Fig. S1), is maximal at midday (35, 42). Collective evidence thus suggests that the evolution of *Prochlorococcus* steadily added redox outlets, leading to increased excretion of organic carbon. Lastly, expression of the putative malate uptake pathway peaks at dusk (*SI Appendix*, Fig. S2), suggesting that the rest of the ecosystem may return some organic carbon to *Prochlorococcus* at night.

What drove *Prochlorococcus*' increased excretion of organics over the course of its evolution? If maintaining redox balance was the driving force, why not enhance the ability to lower photosynthetic electron flux (44, 45), which many photosynthesizers use to prevent redox imbalance (46, 47)? Because *Prochlorococcus* carries out a sizable fraction of the primary production in these ecosystems, perhaps metabolic interdependencies with cooccur-

ring heterotrophs are in play. Invoking cross-feeding interactions to explain excretion of organic carbon by *Prochlorococcus* is problematic, however, because in unstructured populations typical of oceanic microbes, the apparent cost of fixing and excreting carbon leads to a "public goods" dilemma: Nonproducer cells that avoid the cost of production would outcompete producer cells and suppress cross-feeding (48). Direct benefits for the excretion of organic carbon by individual *Prochlorococcus* cells are therefore needed to explain cross-feeding interactions. Benefits of excreting organic carbon hypothesized for aerobic heterotrophs (26, 49–51) provide relevant clues, but do not fully translate to *Prochlorococcus* evolution (see *SI Appendix*, *SI Text* for further discussion).

Integrating Evidence Across Levels of Organization Suggests a Unifying View of Evolution in *Prochlorococcus*. To gain a better understanding of the evolution of *Prochlorococcus*, we mapped observations of its metabolism (Fig. 3 and *SI Appendix*, Fig. S1) onto its population structure (Fig. 1), growth and light physiology, and the macromolecular and elemental composition of cells. This suggests two evolutionary trends. First, in stable water columns, more recently diverged (HL) ecotypes experience a higher photosynthetic electron flux density [$\mu\text{mol electrons (g dry weight)}^{-1} \text{ time}^{-1}$], which can be parameterized as:

$$\nu_e = I\sigma_{PSII}\Phi_{PSII}, \quad [1]$$

where I is light intensity ($\mu\text{mol photons m}^{-2} \text{ time}^{-1}$), σ_{PSII} is the mass-normalized whole-cell photosystem II absorption cross-section [$\text{m}^2 \text{ of PSII (g dry weight)}^{-1}$], and Φ_{PSII} is the quantum efficiency of PSII (electrons photon $^{-1}$). An increase in the cellular electron flux density along the *Prochlorococcus* phylogeny is suggested most obviously by the layering of the relative abundance of its ecotypes that emerges when the water column stratifies after a mixing event in the open ocean (Fig. 1), with the HL cells always dominating the surface waters (10–14). However, σ_{PSII} also increases along the phylogeny, suggesting that more recently diverging strains also experience an increased electron flux at a given light level I . This increase in σ_{PSII} is manifested in changes in the pigments (6, 9, 54, 55) and stoichiometry (56, 57) of the photosystems (Fig. 3 and *SI Appendix*, Fig. S1). These changes largely relinquish absorption of green and yellow light, but maintain or increase absorption of blue light—typifying the open ocean—in *Prochlorococcus* relative to *Synechococcus* (32, 58). At the same time *Prochlorococcus* is significantly smaller than *Synechococcus* (6), and cell mass decreases along its phylogeny (59), thus suggesting an increase in the mass-normalized absorption cross-section σ_{PSII} . Fewer comparative studies exist for the quantum efficiency of PSII (Φ_{PSII}) but it appears to be roughly similar along the *Prochlorococcus* phylogeny (57). The increasing excretion of organic carbon by *Prochlorococcus* thus appears to reflect a more general trend that increases the total cellular throughput of electrons (Fig. 3). Various genes related to the protection and repair of light damage have been added in HL clades of *Prochlorococcus* (52, 53), suggesting that later divergences were primarily related to fortifying cellular machinery to allow cells to generate an increased electron flux at the highest photon fluxes near the surface (Fig. 3).

In addition to an increasing electron flux density (ν_e), integrating evidence across levels suggests the evolution of *Prochlorococcus* also decreased its limiting nutrient flux density, ν_n [moles of n (g dry weight) $^{-1} \text{ time}^{-1}$]. For steady-state growth under constant conditions, ν_n can be parameterized as:

$$\nu_n = \mu Q_n, \quad [2]$$

where μ is the specific growth rate (time^{-1}) and Q_n is the mass-normalized cellular quota of limiting nutrients [moles of n (g dry weight) $^{-1}$]. A decreasing nutrient flux density during the evolution of *Prochlorococcus* is suggested by a decrease in its maximal intrinsic growth rate μ_{max} (time^{-1}) relative to *Synechococcus*

(6, 32) and an increased efficiency in its use of limiting nutrients. The latter is inferred from many features of *Prochlorococcus*, including less investment of N and P in the genome by reducing its size and guanine–cytosine content (60), decreased use of amino acids with N-rich side chains (61), a switch from P- to sulfolipid membranes (62, 63), and less Fe use in metabolism and photosynthetic machinery (33, 34) (*SI Appendix, Fig. S1*). These changes suggest a decreasing Q_n for N, Fe, and P, the three main limiting nutrients in the oligotrophic oceans (64). An increasing ν_e coinciding with a decreasing ν_n over the course of *Prochlorococcus* evolution can be expressed as a single variable: the electron-to-nutrient flux ratio, ν_e/ν_n (Fig. 3).

Electron Flux and Nutrient Acquisition. What are the selective pressures maximizing ν_e/ν_n in *Prochlorococcus*? Photosynthetic electron flux transfers solar energy into metabolism via cofactors such as ATP and NAD(P)H (65). An increased electron flux [$\mu\text{mol electrons (g dry weight)}^{-1} \text{ time}^{-1}$] thus suggests an increased metabolic rate [$\text{kJ of absorbed solar energy (g dry weight)}^{-1} \text{ time}^{-1}$]. This principle is exemplified by the increasing metabolic activity and ATP/ADP ratios of plant chloroplasts and cyanobacteria when they shift from darkness into light (66, 67). Furthermore, the highest ν_e/ν_n phenotypes are the most recently diverging clades (Fig. 3), dominating near the surface where the solar energy supply is greatest and nutrient levels are lowest (Fig. 1), suggesting an advantage to higher metabolic rates at lower nutrient concentrations.

How metabolic rate affects nutrient uptake can be understood from Michaelis–Menten kinetics, which under strong nutrient limitation ($[n] \ll K_{M,n}$) simplifies to (68) (*SI Appendix, SI Text*):

$$\nu_n = [n]a_n^0 = V_{max}[n]/K_{M,n}, \quad [3]$$

where $[n]$ is the nutrient concentration (moles of $n \text{ L}^{-1}$) and a_n^0 is the specific nutrient affinity [$\text{L (g dry weight)}^{-1} \text{ time}^{-1}$]. Specific affinity indicates how strongly cells absorb limiting nutrients (analogous to the pumping speed of a vacuum pump) and is equivalent to the saturated maximal nutrient uptake rate V_{max} [moles of $n \text{ (g dry weight)}^{-1} \text{ time}^{-1}$] over the Michaelis constant $K_{M,n}$ (moles of $n \text{ L}^{-1}$) (68) (*SI Appendix, Fig. S3*). V_{max} reflects the maximum handling rate for absorbed nutrients and is proportional to metabolic rate (68, 69). The free energy cost $\Delta_r G$ (kJ mol^{-1}) of nutrient transport scales with natural log of the ratio of internal to external nutrient concentrations (65) (*SI Appendix, SI Text*) and can become very large in the extremely nutrient-poor oligotrophic oceans (69). For example, the free energy cost of phosphate uptake in the oligotrophic oceans is far greater than the free energy gain of ATP hydrolysis unless the ATP/(ADP \times P_i) ratio is increased drastically from commonly assumed metabolite concentrations of 1 mM for ATP, ADP, and P_i (*SI Appendix, SI Text*). An analogous, but slightly different, situation occurs for ammonia (NH_4^+) uptake, which has the potential for a major futile cycle in the oligotrophic ocean, because its conjugate base NH_3 passively diffuses out of the cell (70) (*SI Appendix, SI Text*). It is therefore thought that cells poise their internal $[\text{NH}_4^+]$ at the minimal viable value (71), which in turn similarly requires a significant increase in the ATP/(ADP \times P_i) ratio to drive forward glutamine synthesis (glutamate + NH_3 + ATP \rightleftharpoons glutamine + ADP + P_i), the central highway for nitrogen into metabolism (65). (See *SI Appendix, SI Text* for detailed calculations and discussion.) Finally, while kinetic bottleneck reactions can be driven forward by increasing the ATP/(ADP \times P_i) ratio, this simultaneously increases the free energy cost of ATP synthesis, therefore requiring a greater proton motive force, and thus ultimately a greater photosynthetic electron flux (69).

The principles just outlined suggests that innovations increasing ν_e may lower the minimal subsistence concentration of limiting nutrients $[n]^*$ (the lowest value of $[n]$ at which net positive growth is possible). When growth and loss processes are rela-

tively rapid and tightly coupled, microbial strains with the lowest $[n]^*$ dominate (72), suggesting that selection should favor such innovations. To understand how selection for lower $[n]^*$ shapes cells, we can substitute an expression for V_{max} that assumes reversible kinetics (*SI Appendix, SI Text*) into Eq. 3:

$$[n] = \frac{K_{M,n}}{V_{max}} \nu_n = \frac{K_{M,n}}{[E]k^+} \frac{\nu_n}{1 - e^{\Delta_r G/RT}}, \quad [4]$$

where $[E]$ is the enzyme concentration [moles of enzyme (g dry weight) $^{-1}$]. k^+ is the rate constant [moles of n (moles of enzyme) $^{-1} \text{ time}^{-1}$], R is the gas constant ($\text{J K}^{-1} \text{ mol}^{-1}$), and T is temperature (K). Eq. 4 suggests that there are two strategies for lowering $[n]^*$ (*SI Appendix, Fig. S3*). First, cells can modify the kinetics (decreasing $K_{M,n}$ and/or increasing k^+ , $[E]$) and thermodynamics (decreasing $\Delta_r G$) of their metabolism, the latter by increasing ν_e as we have just argued. Second, cells can lower their $[n]^*$ by lowering their required flux of limiting nutrients ν_n (Eq. 4), which can be achieved by lowering their minimal Q_n [i.e., streamlining (73)] or their μ (Eq. 2), both of which are apparent in the evolution of *Prochlorococcus* as discussed above. Optimizing kinetics/thermodynamics and decreasing ν_n can work synergistically, and both are helped by a decrease in cell mass, as is observed along the *Prochlorococcus* phylogeny (6, 59). That is, selection to lower ν_n allows a decrease in total cell mass by minimizing the mass dedicated to nonessential components (73). If the amount of metabolic enzyme and photosynthetic machinery is kept fixed, this decrease in total cell mass would increase both $[E]$ (68) and σ_{PSII} (and thus ν_e , which makes $\Delta_r G$ more negative). For nitrogen, increasing ν_e in turn provides an additional avenue for lowering Q_N (and thus ν_N), because pathways with a more negative $\Delta_r G$ require less protein biomass for a given flux (74). Thus, maximizing ν_e/ν_n lowers the $[n]^*$ of cells by making the $\Delta_r G$ more negative (driving kinetic bottleneck reactions forward), thereby increasing the nutrient affinity (68, 69). As a result, the evolution of *Prochlorococcus* has driven nutrients to vanishingly low levels ($<0.1 \text{ nM}$) in the oligotrophic oceans (64).

Benefits of Excreting Organic Carbon. If selection to lower $[n]^*$ drives the maximization of ν_e/ν_n , why should it lead to an increased excretion of organic carbon? We argue this ultimately emerges from mass and energy conservation. That is, in the presence of kinetic bottlenecks, cells can drive up their ATP/ADP ratio by increasing their ATP supply rate, but to maintain steady state, they must also increase ATP consumption rates (49, 69). [The same argument applies to the NAD(P)H/NAD(P) ratio.] This is consistent with observations that some nutrient-limited aerobic chemoheterotrophs have increased glucose uptake rates, increased levels of respiration, increased excretion of polysaccharides (whose synthesis from glucose requires ATP consumption), and a high flux through various ATP-consuming futile cycles (25, 26, 49). For photosynthetic cells CO_2 -fixation is a major sink for ATP and NAD(P)H, but cells are limited in the carbon flux density, ν_C , they can accommodate. This limit is proportional to growth rate (Eq. 2), and, because selection to lower $[n]^*$ favors relatively slow-growing cells (Eq. 4), it is lower in the oligotrophic oceans. Thus, we argue that excreted organic carbon represents a kind of “carbon exhaust” that allows cells to maximize their nutrient affinity by increasing their metabolic rate above limits arising from carbon saturation.

This general mechanism is further illustrated by expanding Eq. 4. We assume that the electron flux must support carbon fixation sufficient to build biomass at a rate dictated by ν_n , that biomass has an elemental stoichiometry Q_n/Q_C , that the efficiency of fixation is $\#C/\#e$ (carbon atoms electron $^{-1}$), and that a fraction ($0 < \beta < 1$) of carbon is excreted/respired. $\#C/\#e$ depends on the oxidation states of the carbon source, biomass, and excreted carbon. For example, if the carbon source is CO_2 and biomass and excreted carbon have the oxidation state CH_2O , then $\#C/\#e$ is 1/4 (i.e., $\text{CO}_2 + 4e^- + 4\text{H}^+ \rightleftharpoons \text{CH}_2\text{O} +$

H₂O). Photoprotective mechanisms like the water–water cycle of PTOX (*SI Appendix, Fig. S1*) act to lower #C/#e. Together, these assumptions lead to the expression (see also *SI Appendix, SI Text*):

$$[n] = \frac{K_{M,n} \nu_n}{[E]k^+ 1 - e^{\Delta_r G/RT}} = \frac{K_{M,n} Q_n \nu_e (\#C/\#e) (1 - \beta)}{[E]k^+ Q_C 1 - e^{\Delta_r G/RT}} \quad [5]$$

Thus, while increasing metabolic rate (i.e., lowering $\Delta_r G$) lowers $[n]^*$, it also acts to increase $[n]^*$ by increasing the carbon flux density, unless excess carbon is excreted [$\beta > 0$ and increasing faster than $\nu_e (\#C/\#e)$] (Eq. 5). Selection to lower $[n]^*$ would thus favor a decrease in ν_n and simultaneous increases in ν_e and β (carbon excretion), exactly as we observed for the evolution of *Prochlorococcus* (Fig. 3).

Metabolic Rate, Ecosystem Biomass, and Coevolutionary Dynamics.

What are the evolutionary and ecological consequences of Eq. 5? Focusing on the evolution of *Prochlorococcus*, it suggests that the layered population structure observed in stable water columns (Fig. 1) (10–14) reflects the sequential evolution of new ecotypes near the surface, each with increasing metabolic rates, drawing down limiting nutrients and restricting ancestral ecotypes to ever deeper waters (Fig. 4). This type of evolutionary dynamic in which a key innovation causes a population to expand along a shifted adaptive landscape produces adaptive radiations (75, 76) and has been argued to be the dominant contributor to global microbial diversity (77). This is consistent with observations that the broadly defined ecotypes of *Prochlorococcus* consist of hundreds of stable subpopulations that diverged millions of years ago (78). These are “niche constructing” adaptive radiations, since the shift in the adaptive landscape arises because new ecotypes chemically modify their own environment (Fig. 4) (79). Subpopulations form clusters defined by the acquisition of small cassettes of genes involved in electron transfer, redox stress, and/or synthesis of membrane polysaccharides (78). Increased production of the latter is a recognized outlet of excess reducing power (25, 26, 40). In addition to providing an electron sink, membrane polysaccharides also mediate biological interactions (e.g., with phage, grazers, or other bacteria), suggesting that an increase in electron flux could have produced ecological feedbacks that drove further differentiation within *Prochlorococcus* ecotypes (78) and in those with which it interacts (80). Lastly,

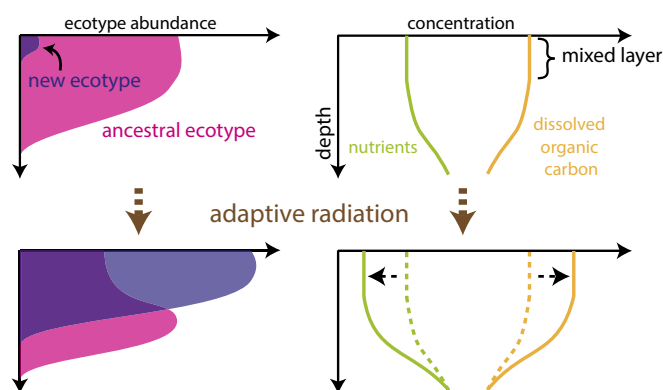


Fig. 4. The emergence of *Prochlorococcus* ecotypes. Eq. 5 suggests that innovations increasing ν_e/ν_n trigger the emergence of new ecotypes (purple) that draw down limiting nutrients (n , green depth profile) in oceanic surface waters as they go through adaptive radiation. Nutrient drawdown near the surface increases ecosystem biomass and restricts ancestral ecotypes (pink) with a higher minimum subsistence nutrient concentration $[n]$ to greater depths in the water column where nutrient levels remain higher. Finally, Eq. 5 suggests that by excreting increasing amounts of fixed carbon, while maximizing ν_e/ν_n , the evolution of *Prochlorococcus* has increased the long-term steady-state concentrations of DOC (orange depth profile).

this framework suggests that by promoting the fixation and excretion of increasing amounts of organic carbon as a by-product of increasing the nutrient affinity, the evolution of *Prochlorococcus* increased the long-term steady-state concentrations of dissolved organic carbon (DOC) in the oligotrophic oceans (Fig. 4), which are much higher (by a factor of up to ~ 2) than elsewhere in the ocean (81). Because several compounds that *Prochlorococcus* may be excreting are known iron-binding ligands, including polysaccharides (82) and carboxylic acids like citrate (83), it may, in turn, also play a key role in buffering trace metal bioavailability in these environments.

When extended to the ecosystem level, Eq. 5 provides a mechanistic view of how the evolution of all microbial cells in the open ocean is interconnected via the chemically coevolving environment (79, 84): Any innovation in any lineage that increases the energy flux and lowers nutrient concentrations (Fig. 4) pushes all other lineages to follow suit and adopt similar innovations. Unlike in the classic “Red Queen hypothesis,” in which evolution is a zero-sum game (85), here, the evolutionary dynamic increases resource capture, and thereby biomass, of the ecosystem. [Similarly, a paleontological survey of the distribution and body size of marine animals also led to the conclusion that the energy flux and biomass of ecosystems increases over evolutionary time (86).] The imprint of this collective dynamic can be seen in the broadly convergent features of all oceanic microbes. Slow growth, small cell size, streamlined genomes and proteomes, and use of nonphospholipid membranes—signature features of *Prochlorococcus*—are observed across both autotrophic and heterotrophic microbes in the oligotrophic oceans (61, 73, 87–89). Some of the photosynthetic machinery modifications that increased the electron flux of *Prochlorococcus* (Fig. 3) are also seen in diatoms and picoeukaryotes (33, 90, 91). Many oceanic heterotrophs in turn supplement their energy supply from organics by capturing sunlight with proteorodhopsins (92) and possess electron drains in their respiratory electron transfer chains (93). In general, heterotrophic growth in the oligotrophic oceans favors metabolic rate over efficiency—the fraction of carbon taken up from the environment that is converted into biomass is one of the lowest of any aquatic environment on Earth (94). These observations are all consistent with the notion that the maximization of ν_e/ν_n , and thereby a lowering of steady-state nutrient concentrations (Eq. 5 and Fig. 4), has been a general evolutionary driving force in ocean ecosystems.

Emergence of Mutualism in Oceanic Microbial Ecosystems. Finally, if lowering $[n]^*$ increases the excretion of organic carbon into the environment, it could produce new opportunities for cooccurring heterotrophs, like the ubiquitous and abundant SAR11 (73, 95). SAR11 requires pyruvate and either glycolate, glycine, serine, or glycine betaine (96); the latter four all feed into the same pathway that starts from glycolate (96). Furthermore, coastal SAR11 strains can replace pyruvate with glucose metabolized via glycolysis, whereas open ocean strains lack glycolysis and have an obligate requirement for pyruvate (97). Similar adaptations are not observed in freshwater strains of SAR11 (98). Thus, oceanic SAR11 populations have evolved a dependency on exactly the compounds (pyruvate and glycolate) that our metabolic reconstructions suggest emerged as excretion pathways in *Prochlorococcus* (Fig. 3 and *SI Appendix, Fig. S1*).

We mapped the distribution of metabolic genes across clades (*SI Appendix, Table S2*) onto the phylogeny of marine SAR11 clades to further reconstruct the evolution of their metabolic core (96–98) (Fig. 5 and *SI Appendix, Fig. S4*). We aimed to resolve the innovations of open ocean lineages and look for evidence of selection on pathway controls by looking for transporters in the vicinity of metabolic genes. As in *Prochlorococcus* (Fig. 3 and *SI Appendix, Table S1*), the metabolic core of SAR11 (*SI Appendix, Fig. S4*) evolved through a sequence of innovations (Fig. 5 and *SI Appendix, Table S2*), including the step-wise completion of the glyoxylate shunt and the well-documented switch

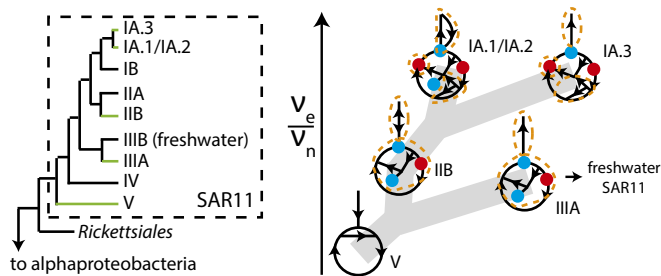


Fig. 5. Metabolic evolution of SAR11. Variants are shown in simplified form along a gradient of ν_e/ν_n . Dashed orange lines highlight innovations (SI Appendix, Fig. S4 and Table S2), which include disruptions or replacements of glycolysis (all branches), the step-wise completion of the glyoxylate shunt (from group V to IIB to all groups IA) and the gain and loss of excretion/uptake pathways. See main text for details. Excretion pathways are represented as red dots and uptake pathways as blue dots. This reconstruction is less certain than for *Prochlorococcus* because SAR11 has greater diversity (99, 100), and as highlighted in the phylogeny (included clades in green), many clades have no or few sequenced representatives (SI Appendix, Table S2). Several branches were therefore drawn as a polytomy.

from Emden–Meyerhoff–Parnass (EMP) to Entner–Doudoroff (ED) glycolysis (97). Glycolysis is disrupted in the 1.3A clade (97) (Fig. 5 and SI Appendix, Table S2), which is most abundant in surface waters where the *Prochlorococcus* HLII clade dominates (Fig. 1) (99, 100). Glycolate uptake is also lost in this SAR11 clade (Fig. 5 and SI Appendix, Table S2), which suggests that pyruvate produced by the *Prochlorococcus* HLII ecotype may have become its central source of carbon and energy (97). Finally, similar to *Prochlorococcus* (SI Appendix, Fig. S2), chromosome rearrangements positioned a series of transporter proteins near key metabolic genes in SAR11 (SI Appendix, Fig. S5), consistent with selection on the control of transport pathways. We identified putative import transporters for pyruvate and glycolate, and putative export transporters for citrate and malate, the latter exclusive to the IA clade (Fig. 5 and SI Appendix, Fig. S5). The emergence of malate export in SAR11 would provide a potential source for the emergent malate uptake pathway of *Prochlorococcus* that is putatively activated at night (Fig. 3 and SI Appendix, Figs. S1 and S2).

Furthermore, the glyoxylate shunt is a TCA cycle bypass activated under redox stress in some microbes (101–103), while the evolutionary switch from EMP to the higher-rate ED variant of glycolysis has, in other microbes, been attributed to an increased energy supply (74). These observations are consistent with selection acting to maximize ν_e/ν_n in both systems and thereby producing pathways that transfer pyruvate and glycolate from *Prochlorococcus* to SAR11 and malate from SAR11 to *Prochlorococcus*. Oligotrophic waters have nanomolar concentrations of pyruvate and glycolate, with midday maxima for the latter, consistent with biological cross-feeding synchronized with the input of sunlight, although abiotic photochemistry may also contribute (104, 105).

Additional evidence for metabolic mutualism in these systems comes from observations regarding hydrogen peroxide (HOOH), a by-product of biological electron transport, photochemistry, and other abiotic processes (106). *Prochlorococcus* and some later-diverging clades of SAR11 have lost HOOH-detoxifying catalase (SI Appendix, Tables S1 and S2), and *Prochlorococcus* grows better in the presence of bacteria retaining catalase (107, 108). This led to the “Black Queen” hypothesis, which argues that subpopulations of ecosystems can save essential nutrients by giving up inevitably shared functions (such as detoxifying the freely diffusible HOOH), so long as they are preserved by others in the ecosystem (109). Our reconstructions suggest that the loss of catalase in *Prochlorococcus* and SAR11 coincides with increased excretion of organic carbon, which provides carbon and energy for catalase-containing bacteria (109).

These observations suggest that metabolic mutualisms are self-amplifying feedback loops (110) that maximize the collective ν_e/ν_n , and thus total productivity, of ecosystems. Specifically, recycling otherwise wasted electrons through complementary excretion/uptake pathways increases the average ν_e of participating cells. Similarly, the loss of functions that are shared and require limiting nutrients (e.g., iron in catalase) in some members of the community decreases the average ν_n . Mutualisms thus raise the ν_e/ν_n of ecosystems—and lower the subsistence nutrient requirements of their cells (Eq. 5 and Fig. 4)—beyond what is possible for individual lineages in isolation. Because excreting organic carbon lowers the minimal subsistence nutrient requirements of individual cells, mutualisms of this kind are emergent properties of ecosystems and avoid public goods dilemmas (48). An upper bound may exist on the maximization of ν_e/ν_n due to the minimal requirements of being an autonomous cell. As the smallest photosynthetic cell (6), *Prochlorococcus* may be closest to this limit, reinforcing the notion that it has a central role in shaping the features of ecosystems in the surface oceans.

Are Plant Cells Microscopic Analogs of Oceanic Microbial Ecosystems?

It occurred to us that the metabolic organization of oceanic microbial ecosystems and green plant cells are similar (Fig. 6): Intermediates of lower glycolysis and photorespiration are central conduits of electron transfer from *Prochlorococcus* to SAR11 and from chloroplasts to mitochondria in plant cells, while TCA cycle intermediates facilitate electron transfer in the opposite direction in both systems (111–113). Similar patterns emerge at other levels of organization. Microbes other than *Prochlorococcus*/SAR11 in ocean communities, and organelles other than chloroplasts/mitochondria in plant cells, appear central to peroxide detoxification (108, 113). The heterotrophic bacteria SAR86 and SAR116 may be important for this function in the oligotrophic oceans, because they both possess catalase (114–116) and are abundant in warm stratified waters (117). *Prochlorococcus* and chloroplasts both have PTOX as an electron drain in their photosynthetic electron transfer chain, whereas SAR11 and mitochondria both have alternative oxidase (AOX) as an electron drain in their respiratory electron transfer chain (93, 113). Furthermore, like chloroplasts, *Prochlorococcus* uses chlorophyll *b* in addition to chlorophyll *a*, which has not been observed in cyanobacteria other than *Prochloron* and *Prochlorotrix* (6). Finally, organelles of plant cells have also undergone reductive genome evolution, and, for mitochondria, it has been argued that this increased the cellular power density of eukaryotes (118)—similar to what we argued for oceanic microbial ecosystems. The extensive convergence of plant cells and oceanic microbial ecosystems highlights the constraints that metabolism imposes on the large-scale structure of evolution (5) and suggests that the metabolic dependencies of eukaryotic organelles can evolve without the physical intimacy of endosymbiosis.

Biospheric Self-Amplification and the Rise of Atmospheric Oxygen.

We have proposed that maximizing the metabolic rate of cells lowers their minimal subsistence nutrient requirements and that this is achieved by maximizing the cellular electron-to-nutrient flux ratio (ν_e/ν_n), while increasing the excretion of organic carbon (Eq. 5). This leads to an evolutionary dynamic that increases total ecosystem biomass (Fig. 4) and paves the way for self-amplifying feedback loops that recycle organic carbon (Fig. 6) and reinforce the maximization of cellular metabolic rate at the ecosystem level. It has been argued that the hierarchical organization of pathways within metabolism reflects the outgrowth of self-amplifying feedbacks that increased the free energy consumption of the emerging biosphere (5, 119). Our framework extends those arguments into the world of phenotypically differentiated cells and microbial ecosystems, and is consistent with the theorem that the flow of energy through the biosphere promotes its self-organization into chemical cycles (2).

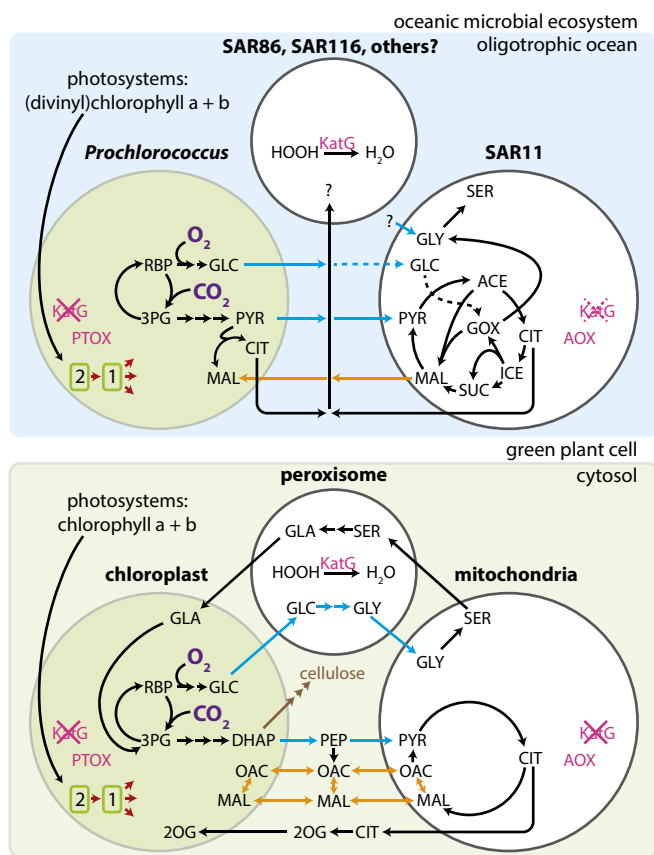


Fig. 6. The metabolic organization of *Prochlorococcus* and SAR11 and that of green plant cells. For simplicity, many reaction sequences are shown as single arrows without accurate stoichiometry. Blue/orange arrows indicate organic carbon flux through similar pathways between *Prochlorococcus* and SAR11 and between chloroplasts and mitochondria. The loss of glycolate uptake in the open ocean 1A.3 clade (Fig. 5) is indicated by a dashed arrow. The loss of catalase in *Prochlorococcus* is indicated by a crossed-out KatG gene, while this cross-out is dashed for SAR11 to indicate the loss of catalase in some of its later branching clades. In plant cells, catalase is similarly located in the peroxisome and not chloroplasts/mitochondria (111–113) as indicated by the crossed-out KatG genes. PTOX and AOX provide electron drains in the electron transport chains of both systems, and chloroplasts and *Prochlorococcus* both use chlorophyll *b* as well as *a* (6). 2OG, 2-oxoglutarate; 3PG, 3-phosphoglycerate; ACE, acetyl-CoA; KatG, catalase; CIT, citrate; DHAP, dihydroxyacetone phosphate; GLA, glycylate; GLC, glycolate; GLY, glycine; GOX, glyoxylate; ICE, isocitrate; MAL, malate; OAC, oxaloacetate; PEP, phosphoenolpyruvate; PYR, pyruvate; RBP, ribulose biphosphate; SER, serine; SUC, succinate.

Our framework has implications for Earth history. If biospheric self-amplification driven by the sun enhanced the burial of organics and carbonates (120, 121) simply by increasing their production, it would help explain the drawdown of atmospheric CO_2 and the rise of atmospheric O_2 across several stages of Earth history (122, 123). Perhaps not coincidentally, marine picocyanobacteria are estimated to have emerged near the transition from the Neoproterozoic (1,000–541 Ma) to the Phanerozoic (541 Ma to present) (124, 125), when sediments indicate the occurrence of global glaciations (126, 127), global carbon cycle perturbations (128–130), enhanced organic carbon burial (120, 121), and a second major rise in atmospheric O_2 toward modern levels (123, 131–133).

The proposed evolutionary dynamic (Fig. 4) may provide insights into the Neoproterozoic–Phanerozoic rise in atmospheric oxygen. It has been argued that after the Great Oxidation Event (GOE) the deep Proterozoic oceans remained largely anoxic and euxinic (rich in H_2S) (134). Recent studies

support anoxia, but suggest that the Proterozoic oceans were instead largely ferruginous (rich in Fe^{2+}), with euxinia restricted to productive continental margins (135–137). It was argued this general redox structure persisted because of N/Mo limitation (138, 139), with Mo scavenged by H_2S even with only limited global euxinia (137), and P limitation due to its scavenging by abundant Fe^{2+} (140). We suggest an additional negative feedback involving iron, the most widely used metal cofactor for biological electron transfer (141). The free O_2 produced during the rise of oxygenic photosynthesis transformed iron from its soluble Fe^{2+} form into its insoluble Fe^{3+} form (142), effectively causing early oxygenic photosynthesizers to self-limit their expansion in the global oceans by locally extinguishing available iron. Extant oceanic microbes surmount this negative feedback on photosynthetic electron transfer through reduced cellular Fe demands (33, 90, 91) (SI Appendix, Fig. S1), and through Fe-ligation by DOC (143), including polysaccharides (82), citrate (83), and other carboxylic acids, all of which may be excreted by marine picocyanobacteria (37, 39) (SI Appendix, Fig. S1). Polysaccharides and small carboxylic acids also enhance the dissolution of minerals (144), and minerals in wind-blown dust are a major source of Fe (145) and P (146) to the surface oceans. Thus, we hypothesize that the evolution of marine picocyanobacteria (Fig. 4) increased both the bioavailability and the overall supply of iron under aerobic conditions and helped transform the oceans from an anoxic state rich in free iron (135, 136) to an oxygenated state (131, 133) with DOC-bound iron. This positive iron–DOC feedback, strengthened by an increased metabolic rate, was critical in pushing the marine biosphere past a major evolutionary bottleneck and paved the way for an expansion of oceanic oxygenic photosynthesis and a rise in atmospheric O_2 (123).

Sedimentary and genomic records suggest several additional positive feedbacks that could have pushed forth the Phanerozoic oxygenation of the ocean. Increased Fe bioavailability under aerobic conditions coinciding with the drawdown of nitrogen (Fig. 4) would have created opportunities for N_2 -fixers, while ocean oxygenation would have lifted their Mo-limitation by suppressing euxinia (137, 138), together increasing the supply of nitrogen to the oceans (124, 138, 139). This is consistent with the suggested overlap in the rise of marine picocyanobacteria and planktonic N_2 -fixers (124, 125). Furthermore, sediments suggest an increase in oceanic P levels after the Neoproterozoic, and it was argued that this was because a drop in Fe^{2+} concentrations lessened the scavenging of P (140). We add that if enhanced DOC-dissolution of minerals from dust enhanced the oceanic iron supply under aerobic conditions, as we argued above, it could have also enhanced the P supply (146), thus contributing to the reconstructed rise in P levels (140).

These scenarios are similar to those of how the rise of land plants impacted the Earth system. Nutrient-limited plants leach Fe, P, and other nutrients from rocks by excreting small carboxylic acids from their roots (which suggests an increased metabolic rate—Eq. 5) (147). It has therefore been argued that plant colonization of the continents during the Phanerozoic increased the weathering of rocks, and, in turn, the precipitation of carbonates, which, together with the increased burial of plant-derived organics, resulted in a drawdown of atmospheric CO_2 and a rise in atmospheric O_2 (148–150).

The convergent metabolic evolution of oceanic microbial ecosystems and land plants (Fig. 6), which, we have argued, may have impacted the Earth system in similar ways, suggests the temporal profile of the Earth's oxygenation (122, 123) may be constrained by two biological stages. The first consisted of the expansion of self-damping cyanobacterial O_2 -photosynthesis in shallow aquatic environments and is associated with the GOE (122, 123). The second consisted of the global expansion of eukaryotic or “eukaryote-like” O_2 -photosynthesis—both onto the continents and into the deep open ocean (Fig. 6)—with a higher metabolic rate and correspondingly greater ability to mobilize Fe and P under oxidizing conditions, and is associated

with the Neoproterozoic Oxidation Event (122, 123, 132, 142). Genomic studies estimate that chloroplast endosymbiosis leading to the rise of all photosynthetic eukaryotes occurred between the late Paleoproterozoic (2,500–1,600 Ma) and the early Neoproterozoic (151–154). Paleontological studies in turn find a significantly increased fossil diversity of eukaryotes (including photosynthetic eukaryotes) in the Neoproterozoic (155, 156), whose increasing body sizes and fecal pellets could have moreover strengthened the export and burial of organic carbon from the oceans (157, 158).

As a final note, one could argue that the emergence of modern human societies is a variant of the general framework we propose. As our populations expanded and extracted ever more electrons from fossil fuels, we have increased global CO₂, while drawing down natural resources and global O₂ (albeit a small amount relative to the contemporary inventory for the latter) (159, 160). In the process, we have become increasingly socially, technologically, and economically interconnected, analogous to what we have observed in the evolution of oceanic microbial ecosystems and plant cells. As in those systems, this has increased our collective ability to harvest more difficult-to-access natural resources. Managing the biogeochemical perturbation that our

global emergence is imposing on the Earth system is one of humanity's greatest challenges.

Materials and Methods

We analyzed 56 genomes of cyanobacteria representative of the diversity of this clade (161) obtained from the UniProt website, 56 genomes of *Prochlorococcus* and marine *Synechococcus* (162), and 16 SAR11 genomes obtained from the National Center for Biotechnology Information website. We searched these genomes for the presence and absence of a set of reference enzymes (*SI Appendix, Tables S1 and S2*) using the BLASTp and tBLASTn algorithms (163). We reconstructed phylogenetic trees (Fig. 2) by mapping the distributions of metabolic genes onto the phylogenies of *Prochlorococcus* (Figs. 1 and 3) and SAR11 (Fig. 5).

ACKNOWLEDGMENTS. We thank the reviewers for critical and insightful reviews, and we thank Jamie Becker, Shane Hogle, and John Casey for comments on an earlier version of our manuscript. R.B. is a Simons Foundation Fellow of the Life Sciences Research Foundation. This work was supported by Simons Foundation Grant SCOPE 329108 (to S.W.C. and M.J.F.); Gordon and Betty Moore Foundation Grants 3778 (to M.J.F.) and 495.01 (to S.W.C.); and by the NSF Center for Microbial Oceanography Research and Education (S.W.C.). R.B. dedicates this work to the memory of Harold Morowitz, a pioneer in the study of biological energy flow, whose theories and approach to studying life were central in shaping the core ideas of this work.

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