

MIT Open Access Articles

Influence of the NO/NO2 Ratio on Oxidation Product Distributions under High-NO Conditions

The MIT Faculty has made this article openly available. *Please share* how this access benefits you. Your story matters.

Citation: Nihill, Kevin J. et al. "Influence of the NO/NO2 Ratio on Oxidation Product Distributions under High-NO Conditions." Forthcoming in Environmental Science and Technology. © 2021 American Chemical Society

As Published: https://doi.org/10.1021/acs.est.0c07621

Publisher: ACS Publications

Persistent URL: https://hdl.handle.net/1721.1/130544

Version: Author's final manuscript: final author's manuscript post peer review, without publisher's formatting or copy editing

Terms of use: Creative Commons Attribution-Noncommercial-Share Alike



This document is confidential and is proprietary to the American Chemical Society and its authors. Do not copy or disclose without written permission. If you have received this item in error, notify the sender and delete all copies.

Influence of the NO/NO₂ Ratio on Oxidation Product Distributions under High-NO Conditions

Journal:	Environmental Science & Technology
Manuscript ID	es-2020-07621y.R2
Manuscript Type:	Article
Date Submitted by the Author:	n/a
Complete List of Authors:	Nihill, Kevin; Massachusetts Institute of Technology, Civil and Environmental Engineering Ye, Qing; Massachusetts Institute of Technology, Civil and Environmental Engineering Majluf, Francesca; Aerodyne Research Inc, Krechmer, Jordan; Aerodyne Research, Inc., Center for Aerosol and Cloud Chemistry Canagaratna, Manjula; Aerodyne Research Inc, Kroll, Jesse; Massachusetts Institute of Technology, Civil and Environmental Engineering



1	"Influence of the NO/NO2 Ratio on Oxidation Product Distributions under High-NO Conditions"
2	
3	Kevin J. Nihill, ¹ * Qing Ye, ¹ Francesca Majluf, ² Jordan E. Krechmer, ² Manjula R. Canagaratna, ²
4	Jesse H. Kroll ¹
5	
6	1. Department of Civil and Environmental Engineering, Massachusetts Institute of Technology,
7	77 Massachusetts Avenue, Cambridge, MA, USA
8	2. Aerodyne Research, Inc., 45 Manning Rd., Billerica, MA, USA
9	
10	*Email address for corresponding author: kevin.j.nihill@gmail.com
11	
12	Synopsis: Simple, well-constrained organic oxidation systems are probed to measure the
13	influence of the NO/NO ₂ ratio on product distributions.
14	Keywords: oxidation; NO _x ; peroxy radicals; alkyl nitrites; secondary organic aerosol
15	TOC Art:



17 Abstract

18 Organic oxidation reactions in the atmosphere can be challenging to parse due to the large number 19 of branching points within each molecule's reaction mechanism. This complexity can complicate 20 the attribution of observed effects to a particular chemical pathway. In this study, we simplify the 21 chemistry of atmospherically relevant systems, and particularly the role of NO_x, by generating 22 individual alkoxy radicals via alkyl nitrite photolysis (in order to limit the number of accessible 23 reaction pathways) and measuring their product distributions under different NO/NO₂ ratios. 24 Known concentrations of NO in the classically "high NO" range are maintained in the chamber, 25 thereby constraining first-generation RO₂ (peroxy radicals) to react nearly exclusively with NO. 26 Products are measured in both the gas phase (with a Proton-Transfer Reaction Mass Spectrometer) 27 and the particle phase (with an Aerosol Mass Spectrometer). We observe substantial differences 28 in measured products under varying NO/NO₂ ratios (from ~ 0.1 to >1); along with modeling 29 simulations using the Master Chemical Mechanism (MCM), these results suggest indirect effects 30 of NO_x chemistry beyond the commonly-cited RO_2 + NO reaction. Specifically, lower NO/NO₂ 31 ratios foster higher concentrations of secondary OH, higher concentrations of peroxyacyl nitrates 32 (PAN, an atmospheric reservoir species), and a more highly oxidized product distribution that 33 results in more secondary organic aerosol (SOA). The impact of NO_x concentration beyond simple 34 RO₂ branching must be considered when planning laboratory oxidation experiments and applying 35 their results to atmospheric conditions.

36 Introduction

Atmospheric organic oxidation mechanisms are highly complex, involving numerous reaction branching points and multiple generations of oxidation for an individual compound.^{1,2} The large number of products formed from a given compound, which are a strong function of the compound's structure and of reaction conditions, poses substantial challenges for the elucidation of detailed mechanisms and the prediction of major secondary species such as ozone and secondary organic aerosol (SOA).^{3–5}

43 A key branch point in atmospheric oxidation mechanisms involves organic peroxy (RO₂) 44 radicals, which can react bimolecularly with NO_x, HO₂, or other RO₂, or undergo unimolecular 45 reactions.^{5,6} The role of NO_x in RO₂ fate is of particular interest as NO_x is present across a wide range of concentrations in the atmosphere, varying from ppt levels in remote regions⁷ to tens or 46 47 even hundreds of ppb in urban settings and in biomass burning plumes.⁸⁻¹¹ Under high NO 48 concentrations (i.e., NO mixing ratios in the ppb level or higher), the dominant reaction pathway for peroxy radicals is $RO_2 + NO \rightarrow RO + NO_2$, $^{4,5,12-15}$ with a minor contribution from the reaction 49 $RO_2 + NO \rightarrow RONO_2$.¹⁶⁻¹⁸ Recent work on NO_x has gone beyond absolute NO_x levels in order to 50 51 focus on the role of the NO/NO₂ ratio in reaction mixtures. While some studies have explored the 52 role of this ratio in terms of important subsets of atmospheric mechanisms (e.g., SOA, highly oxidized molecules),^{19–21} the NO/NO₂ ratio has not been investigated in terms of its effects on the 53 54 overall product distribution. This limits our ability to accurately predict how reaction systems 55 respond to changes in NO_x levels, and risks leading to inaccurate recreations of "polluted 56 conditions" in laboratory studies.

57 Here, we seek to better understand the detailed role of NO_x , and specifically the NO/NO_2 58 ratio, in influencing product distributions; this requires a reaction scheme in which the initiating

59 chemistry is independent of NO_x and the product distribution has a manageable complexity. We 60 accomplish this via the photolysis of alkyl nitrite (RONO) compounds^{22,23} to directly generate 61 alkoxy radicals (key intermediates in organic oxidation) in the presence of known concentrations 62 of NO. For larger RO radicals, such as the *n*-butoxy radical shown in Figure 1, the dominant 63 channel is isomerization to form an RO₂ radical, which can subsequently undergo a number of 64 reactions. This method involves no direct introduction of gas-phase oxidants, and the generation 65 of a single initial organic radical (as opposed to a mixture of radicals arising from multiple potential 66 OH-reaction sites, which is typical for oxidant-initiated chemistry), greatly simplifying the product distribution compared to traditional laboratory oxidation studies.^{22–24} Moreover, it enables control 67 68 over NO_x levels in a manner that does not affect the initial reaction rate, thus facilitating the role 69 of NO_x to be studied directly.

70 These experiments are run under two NO concentrations, both within the classical "high 71 NO" limit ([NO] >> 1 ppb), but representing NO/NO₂ ratios that differ by over an order of 72 magnitude. Such high concentrations of NO ensure that the initially-formed RO₂ reacts almost 73 exclusively with NO, thus making it possible to probe these simple RO₂ systems as a function of changing NO/NO2 ratio. Such systems can provide insight into the mechanisms underlying the 74 75 NO_x -dependence of VOC oxidation chemistry, specifically elucidating the role of the NO/NO_2 76 ratio in environmental chamber studies; this in turn may help to foster more realistic NO_x 77 conditions in chamber studies simulating the formation of SOA and other products under high-78 NO_x reaction conditions.

79



80

Figure 1. Major first-generation products of the photolysis of *n*-butyl nitrite. The resulting alkoxy radical will primarily isomerize, yielding an RO₂ radical that can undergo a number of different reactions. Red arrows indicate dominant reaction pathways under the high-NO conditions employed herein.

86 Materials & Methods

87 <u>Chamber Conditions</u>

Reactions were carried out in a 150 L PFA chamber (described in detail elsewhere²⁵)
surrounded by an array of twelve 340 nm UV lights (Q-lab). The spectral distribution of these
lamps overlaps well with the absorption spectrum of precursor RONO compounds,²⁶ ensuring
rapid photolysis. UV irradiation in this wavelength range (290-400 nm) involves relatively lowenergy photons, limiting the extent of vibrationally/electronically excited products.^{22,26}
Experiments were run at room temperature (~25 °C) and pressure (~1 atm) in semi-batch

94 mode, with sampling flows balanced by an equal input of pure, low-RH (< 1%) air, resulting in a 95 chamber residence time of approximately 15 minutes. Prior to each experiment, the chamber was 96 flushed with pure, dry air for at least one hour. Additionally, the internal walls of the chamber 97 were cleaned between groups of experiments by flooding with O_3 and H_2O while irradiating 98 overnight.

99 Experiments were run under one of two NO concentrations to ensure the dominance of the 100 $RO_2 + NO$ reaction. In "higher-NO" experiments (NO/NO₂ > 1), the chamber was maintained at a 101 constant concentration of ~1 ppm NO by continual addition of NO prior to and throughout the run. 102 In "lower-NO" experiments (NO/NO₂ \approx 0.1, roughly representative of NO/NO₂ ratios in ambient conditions²⁷), the only source of NO was from the photolysis of the RONO precursor, resulting in 103 104 a steady state concentration of ~40 ppb with the lights on. (Full NO/NO₂ ratios throughout a typical 105 experiment are provided in the SL) At these classically "high NO" conditions, reactions with HO₂ 106 and isomerization reactions cannot compete with the $RO_2 + NO$ pathway.^{28–30} In addition to 107 restricting accessible RO₂ reaction pathways, these high NO concentrations further limit reaction 108 complexity by shortening the lifetime of secondary oxidants O₃ and NO₃, which could otherwise 109 contribute to oxidation and SOA formation.³¹ However, as described below, there is still sufficient 110 secondary OH formation in the reaction mixture to affect the product distributions.

Prior to injection of RONO, the chamber was filled with dry ammonium sulfate seed particles to provide surface area to promote condensation of low-volatility products, and to allow for correction for particle losses due to dilution and wall loss. Polydisperse $(NH_4)_2SO_4$ seed was added to the system by atomizing 1 g/L aqueous solution with a constant output atomizer (TSI) and passing the output through a desiccant prior to entering the chamber. Following this, ~400 ppb of the RONO compound (described below) was injected into a septum and carried by a stream of air into the chamber where it was allowed to mix in the dark for two minutes. Finally, the lights 118 were turned on to initiate the reaction and remained on for the duration of the experiment 119 (approximately one hour).

120

121 Instrumentation

Product distributions were measured by two real-time mass spectrometric instruments. Particle mass and composition were measured by an Aerodyne high-resolution Aerosol Mass Spectrometer (AMS),³² run in "V mode" (mass resolving power of ~3000). Known ion fragmentation of various ions detected by the AMS enabled extraction of the elemental ratios H/C and O/C,³³ thereby allowing the ensemble oxidation state of the SOA to be measured throughout the course of the reaction.³⁴ AMS organic signal was normalized to sulfate concentration in order to account for chamber dilution, wall loss, and changes in the AMS collection efficiency.

129 Products in the gas phase were measured by a Vocus Proton-Transfer Reaction High Resolution Time-of-Flight Mass Spectrometer (PTR-MS),³⁵ which is capable of providing 130 131 speciated measurements of individual molecules and is exceptionally sensitive to volatile 132 compounds with relatively low carbon oxidation states.³⁶ In order to maximize its sensitivity to 133 low-volatility compounds, the Vocus inlet is heated to 100 °C to reduce wall losses due to gas-wall 134 partitioning. (The loss of gas-phase species to chamber walls and instrument inlets is expected to 135 be minor in these experiments, as described in the SI.) The pure RONO precursor is itself only 136 weakly detected by the Vocus as a protonated molecule ([M+H]⁺); it is instead primarily detected as a combination of an aldehyde (via loss of -NO, [M - NO]⁺) and an alkene (via loss of -ONO, 137 [M - ONO]⁺), as observed in previous work.³⁷ One challenge is that the aldehyde species is also a 138 139 product formed from the oxidation of the alkoxy radical. In order to deconvolute the contributions 140 to this ion from the RONO precursor and aldehyde product, the aldehyde time series was fit with a function that included a decay factor for the precursor and a growth factor for the product, asshown in Figure S4.

143 The Vocus was calibrated by equating the total precursor signal (counts per second) prior 144 to photolysis to the known amount of precursor injected into the chamber (~400 ppb); this ratio 145 was then directly applied to all product compounds as an approximate calibration factor. Given the 146 relatively limited range of oxidized functionalities and the tendency of PTR calibration factors to 147 vary only up to a factor of ~ 2 in either direction,³⁸ the use of a single calibration factor for all 148 species is assumed to be a reasonable approximation. While this approach introduces some error 149 into the quantification of individual product species, differences in measured levels of a given 150 compound in both the higher- and lower-NO experiments are independent of calibration, thus 151 allowing for a direct comparison between experiments run under different NO/NO₂ ratios.

152 In addition to the mass spectrometric measurements of the organic species, concentrations 153 of NO and NO₂ were measured with one of two NO_x monitors (Thermo Fisher Scientific, Model 154 42*i* for measuring NO and NO_x, or 2B Technologies Model 405 nm for measuring NO and NO₂; 155 see SI for more details). The presence of NO_v in the chamber interfered with precise NO_2 156 measurements (details regarding the deconvolution of interfering RONO signal from the pure NO_x 157 signal can be found in the SI); however, an order-of-magnitude difference in NO/NO₂ ratios 158 between the two sets of experiments was still clearly observed. All gas-phase data collected by the 159 NO_x monitor and Vocus-PTR were corrected for dilution (with the exception of NO in the higher-160 NO experiments, in which it is part of the dilution flow) using an experiment-specific dilution rate 161 based on chamber volume and input flow-rates.

162

163 Alkyl Nitrite Precursors

Experiments were carried out with four straight-chain alkyl nitrites (*n*-butyl, *n*-pentyl, *n*hexyl, and *n*-decyl nitrite). This study focuses on *n*-butyl nitrite as a simple model for gas-phase systems; *n*-pentyl nitrite was employed in order to examine trends across another gas-phase system, whereas the larger nitrites (*n*-hexyl and *n*-decyl nitrite) were studied to examine SOA formation.^{23,39,40}

169 N-butyl nitrite and n-pentyl nitrite were purchased directly (Sigma-Aldrich) and used 170 without further purification; *n*-hexyl nitrite and *n*-decyl nitrite were not commercially available, 171 and so were synthesized in the laboratory. Synthesis of alkyl nitrites was carried out by Onitrosation of the parent alcohol species (Sigma-Aldrich), as described elsewhere.^{23,41,42} 172 173 Confirmation of the conversion of alcohol to alkyl nitrite was made by UV-Vis spectroscopy of the RONO mixture, with spectra similar to those reported by Heicklen.²⁶ After synthesis, RONO 174 175 species were wrapped in foil to limit exposure to ambient light and stored in the refrigerator until 176 they were used in an experiment, which typically occurred within 3 hours of synthesis to maintain 177 integrity.

178

179 Master Chemical Mechanism Simulations

Simulations using the Master Chemical Mechanism (MCM v3.2)^{43,44} run using the F0AM
 package⁴⁵ in MATLAB were employed in order to map out the reaction mechanisms for individual
 NO regimes and precursors. These simulations were exploited to further probe differences between
 lower- and higher-NO conditions, and for estimating species that are not detectable by our
 instruments (e.g., OH).

Because the precursor RONO species used in these experiments are not included in the MCM, the experimentally-determined photolysis rate of the RONO (as measured by the Vocus) 187 was used to introduce RO and NO into the simulation at a controlled rate; in higher-NO 188 experiments, the concentration of NO in the simulation was fixed at 1 ppm. Simulations included 189 a dilution factor in order to recreate chamber conditions.

190

191 Results & Discussion

Average NO/NO₂ ratios are determined by comparisons of NO_x monitor measurements and MCM simulations. For lower-NO experiments, the NO/NO₂ ratio quickly reaches a steady-state value of ~0.1 with the lights turned on; for higher-NO experiments, the constant flow of 1 ppm NO into the chamber results in NO/NO₂ > 1 throughout the experiment (Table S1 and Figures S1-2).

197 Figure 2 shows the major gas-phase products (weighted by ppb carbon) from the photolysis 198 of *n*-butyl nitrite under lower-NO (panel a) and higher-NO (panel b) conditions. These major ions 199 account for ~60% (lower-NO) and ~75% (higher-NO) of measured secondary carbon; stacked 200 plots of all detected product traces are provided in Figure S5. In both experiments, the precursor reacts away at a roughly equivalent rate (average decay constant $\sim 2 \times 10^{-3} \text{ s}^{-1}$), which is also 201 202 the case for all other precursors in these experiments; this decay is consistent with precursor loss 203 by photolysis that exhibits no dependence on NO. While the precursor RONO is capable of 204 reacting directly with OH generated in the reaction mixture, the concentrations of OH (predicted 205 by MCM simulations and shown in Figure S6) and small rate constant ($k_{\text{RONO+OH}}$) $< 3 \times 10^{-12}$ cm³ molec⁻¹ s⁻¹)⁴⁶ suggest that this pathway is minor, accounting for only 5-10% of 206 207 RONO loss.



Figure 2. Carbon-weighted concentrations of the precursor and the three major product ions from the photolysis of *n*-butyl nitrite under (a) lower-NO and (b) higher-NO conditions, as measured by Vocus-PTR. Traces are labeled with the ion detected by the Vocus and the corresponding chemical identity; see SI for a detailed discussion of molecular assignments for Vocus ions.

214 The RO radicals formed from RONO photolysis are expected to undergo the same reactions 215 in both the lower-NO and higher-NO experiments (Figure 1). A fraction (~20%) of the RO radicals 216 are expected to react directly with O₂ to form butanal, but because the detected ion is the same as 217 one of the ions from the precursor (as discussed above), the exact contribution of this minor 218 channel is not well-constrained in these experiments. The majority of the RO radicals will 219 isomerize, forming a hydroxy-substituted RO₂ radical. The high concentration of NO in both cases 220 ensures that this RO₂ will react with NO, predominantly forming 4-hydroxybutanal ($C_4H_8O_2$, 221 primarily detected as the dehydrated $C_4H_7O^+$ ion by the Vocus⁴⁷). A fraction of the RO₂ will react 222 with NO to form the 4-hydroxynitrate product ($C_4H_9NO_4$), but the yield is expected to be very small $(\sim 1\%)$,⁵ and such oxygenated nitrates are poorly detected by the Vocus.³⁶ The other RO₂ 223 224 channels are not expected to be competitive: under both lower- and higher-NO conditions, the RO₂ 225 + HO₂, RO₂ + RO₂, and RO₂ isomerization channels are expected to contribute negligibly (<<1%)

to the reaction, and the peroxynitrate formed from $RO_2 + NO_2$ is too short-lived to contribute to the reaction mixture.^{5,20}

228 Despite the identical chemistry of the initially-formed RO and RO₂ radicals under the two 229 NO regimes, there are substantial differences in their product distributions (Figure 2). (These 230 differences are much larger than expected run-to-run variability, as duplicate runs show very little 231 variation, as shown in Figure S7.) Most notably, while the major product in both cases is 4-232 hydroxybutanal ($C_4H_7O^+$), it is present in much greater concentrations under higher-NO 233 conditions. This disparity arises from differences not in formation yield but in loss rates; as shown 234 in Figure S8, the initial formation rate of this compound is the same in the two cases, as expected 235 from the RO₂ chemistry (Figure 1). Because the main chemical sink of hydroxybutanal is oxidation 236 by OH (photolysis is only a very minor channel, estimated to be $\sim 2\%$ by MCM), the more rapid 237 loss of this species under lower-NO conditions implies that lower-NO experiments involve higher 238 concentrations of OH.

239 While the reaction system used here did not involve the initial generation of OH, secondary 240 OH can be formed from the reaction of HO₂ (formed after the isomerization of the hydroxyalkoxy 241 radical, Figure 1) with NO. MCM simulations predict that, under higher-NO conditions, OH is 242 produced at a greater rate (by a factor of ~ 2.5), but that the OH reactivity is higher still (largely 243 due to increased NO_x levels), leading to lower levels of OH overall (Figure S6). The prediction of 244 higher OH levels under lower-NO conditions is further confirmed by the higher mean carbon oxidation state $(\overline{OS_C})^{34}$ of the measured product distribution under lower-NO conditions; this trend 245 246 of increased oxidation under lower NO/NO₂ ratios is also observed for the photolysis of *n*-pentyl 247 nitrite (Figure S9).

248 The photolysis of *n*-butyl nitrite produces no observable SOA, consistent with the small 249 carbon skeleton and consequently high volatility of the products formed. Larger nitrites ($n_c > 5$), 250 however, can form products with sufficiently low volatilities to contribute to the formation of SOA.^{23,40} This is evident from Figure 3, which shows SOA formation from the photolysis of n-251 252 pentyl nitrite, *n*-hexyl nitrite, and *n*-decyl nitrite under lower- and higher-NO conditions. As in the 253 gas-phase, the particle-phase measurements exhibit differences under the two NO regimes. All 254 three precursors exhibit higher SOA production under lower-NO conditions as measured in the 255 plateau region (i.e., after 10 minutes). Most notably, n-pentyl nitrite produces no measurable SOA 256 under higher-NO conditions but measurable levels under lower-NO conditions. Additionally, n-257 hexyl nitrite and *n*-decyl nitrite produce approximately 64% and 78% more SOA under lower-NO 258 conditions, respectively. (As shown in the SI, the observed differences are greater than the 259 uncertainty in the measurements.) As with results in the gas phase, this can be attributed to higher 260 levels of secondary OH under lower-NO conditions, leading to more highly oxidized products that 261 partition into the particle phase. This observation is also in agreement with previous studies that 262 see SOA yields for most systems as being inversely correlated with NO concentrations.^{12,19} 263 Further, the mean oxidation state of measured SOA formed from *n*-decyl nitrite (the only precursor for which SOA formation is large enough for a precise measurement of $\overline{OS_C}$ is greater under 264 265 lower-NO conditions (-1.34) than under higher-NO conditions (-1.45); this is consistent with the 266 observed mean oxidation states in the gas phase product distributions.

267



Figure 3. AMS total organic time series under lower-NO (dashed lines) and higher-NO (solid lines) concentrations. Traces are normalized to sulfate concentration to account for wall losses, collection efficiency, and dilution. *N*-butyl nitrite produces no organic aerosol, similar to *n*-pentyl nitrite under higher-NO conditions, and so it is not shown.

273

274 The difference in OH in the lower-NO and higher-NO experiments thus leads to differences 275 in formation of later-generation products, the extent of oxidation of the product distribution, and 276 the formation of SOA. While this effect is mostly important in laboratory studies (this chemistry 277 does not have a controlling influence on OH in the atmosphere), the NO/NO₂ ratio is likely to 278 affect product distributions via other effects as well. To further investigate the influence of NO_x 279 on VOC oxidation product distributions, Figure 4a compares the average (carbon-weighted) 280 concentrations of different gas-phase products from *n*-butyl nitrite photolysis under the two 281 different NO/NO₂ ratios. The higher-NO regime is characterized by a dominant concentration of 282 4-hydroxybutanal ($C_4H_7O^+$) due to lower levels of OH and therefore a longer lifetime, as discussed 283 above (Figure 2 and S8). Conversely, greater OH concentrations under lower-NO condition result 284 in a wider variety of products and greater concentrations of products with higher oxidation states

(Figures 4 and S10), providing the basis for the larger OS_C (Figure S9). This includes multigenerational oxidation products such as propanal ($C_3H_7O^+$, formed from the reaction of butanal + OH), which has a greater concentration under lower-NO conditions (Figures 2, 4).



Figure 4. Correlation plots for fractions of gas-phase products under lower-NO vs. higher-NO regimes, averaged over the entire experiment. Panel (a): Vocus-PTR observations. Panel (b): MCM predictions. Markers are shaped according to class of molecule and are colored by mean carbon oxidation state ($\overline{OS_C}$). Labeled compounds in panel (a) are the same as those in Figure 2; all other compounds are discussed in detail in the SI. MCM results (panel b) include labels for PAN species (solid triangles); potential PAN fragments are represented by filled markers in panel (a). Dashed line represents the 1:1 ratio.

297

Figure 4b shows the same comparison of product distributions in the two NO regimes, but based on MCM predictions rather than experimental data. While detailed comparisons between Vocus and MCM distributions are beyond the scope of this work, the fact that the MCM predicts fewer major products than are measured by the Vocus is likely attributable to individual molecules being detected as multiple fragment ions by the Vocus, and to the generally simplified chemistry of the MCM. Overall consistencies between the MCM simulations and Vocus data include a predominance of 4-hydroxybutanal ($C_4H_8O_2$) occurring under both NO regimes, with a greater concentration of this species under higher-NO conditions. These general results are similar to those from the *n*-pentyl nitrite system (Figures S11-12).

307 The most pronounced differences between the two NO_x regimes in Figure 4b are the PAN 308 compounds (e.g., C₄H₇NO₆), which are considerably more prevalent under lower-NO conditions 309 and, as discussed below, are second-generation oxidation products. PANs are not detected directly by PTR-MS, but can be detected by known fragmentation patterns.^{48–50} One example is $C_4H_7O_2^+$, 310 311 a predicted tracer ion for $C_4H_7NO_6$ (analogous to $C_2H_3O^+$ serving as a tracer ion for peroxyacetyl 312 nitrate, C₂H₃NO₅);⁴⁸ its identity as a PAN is further suggested by the induction period observed in 313 its time series (Figure 2), which is indicative of later-generation products. Potential PAN fragments 314 are represented by shaded squares in Figure 4a. (A more detailed discussion of these other PAN-315 related ions detected by the Vocus can be found in the SI.) Although these compounds are more 316 prevalent under lower-NO conditions, the measured differences are not as dramatic as predicted 317 by MCM simulations. This may be because these ions are not unique to PAN fragments, as they 318 may be formed from the fragmentation of other product ions (e.g., acyl compounds), potentially 319 resulting in a shift towards the 1:1 line. For example, while $C_4H_7O_2^+$ is a tracer for $C_4H_7NO_6$, it 320 could be a tracer for 4-hydroxybutanoic acid and similar species as well.

PAN formation, from the reaction of acylperoxy radicals with NO₂, is not shown in Figure
1, since acylperoxy radicals are not formed as first-generation radicals from alkoxy radical

isomerization. Instead, acylperoxy radicals will be formed from the oxidation of first-generationaldehyde species, such as 4-hydroxybutanal:

325
$$C_4H_8O_2$$
 (hydroxybutanal) + OH + $O_2 \rightarrow C_4H_7O_4$ (acylperoxy radical) (1)

326
$$C_4H_7O_4$$
 (acylperoxy radical) + $NO_2 \rightleftharpoons C_4H_7NO_6$ (PAN) (2)

327 Rates of PAN formation are observed (Figure 4a) and predicted (Figure 4b) to be substantially 328 greater under lower-NO conditions. This is a result of two factors: The difference in OH levels (as 329 discussed above), which controls the formation of acylperoxy radicals and the subsequent 330 chemistry of the acylperoxy radical. When NO_x is present, acylperoxy radicals are limited to two 331 reactions: reaction with NO₂ to form PAN, and reaction with NO to form acyloxy radicals. The 332 concentrations of PAN species are thus a strong function of the NO/NO₂ ratio, as discussed elsewhere.^{4,14,19,20} Under higher-NO conditions, this ratio is sufficiently high (NO/NO₂ > 1) that 333 334 the acylperoxy + NO pathway is dominant, limiting the formation of PAN. Under lower-NO 335 conditions (NO/NO₂ \approx 0.1), there is considerable competition from the acylperoxy + NO₂ pathway, 336 resulting in the accumulation of PAN species. Thus, PAN formation is more favored under lower-337 NO than under higher-NO conditions because it fosters higher OH levels and a greater rate of the 338 acylperoxy + NO₂ reaction. This additional PAN formation, which is observed in the measurements and predicted by MCM simulations (Figure 4), sequesters RO2 radicals from the 339 340 reaction mixture, terminating the oxidation chain and decreasing the extent to which subsequent 341 chemistry occurs over the timescales of the experiments.^{4,14,20}

342

343 Implications

344 These results are broadly consistent with previous studies that demonstrate the importance 345 of NO_x in controlling product distributions, beyond the prevailing RO_2 + NO reaction in "high

NO" systems. Hoffmann et al.²¹ suggested the role of NO_x in controlling the relative concentrations 346 347 of secondary oxidants, and the subsequent influence on SOA formation. A number of recent 348 studies have examined the specific role that PAN formation and the NO/NO₂ ratio may play in laboratory studies of hydrocarbon oxidation. Chan et al.¹⁹ demonstrated the importance of the 349 350 NO/NO₂ ratio in governing SOA production from isoprene oxidation, in which PAN is an 351 intermediate in SOA formation. Specifically, they found that even under high NO conditions, low 352 NO/NO₂ ratios foster increased production of PAN, which subsequently contributes to SOA 353 generation. Rissanen et al.²⁰ expanded on that work by characterizing the individual contributions 354 of NO and NO₂ in the formation of highly oxidized multifunctional compounds (HOMs), finding 355 that the NO/NO₂ ratio effectively controls the identities of HOMs in oxidation systems. Similarly, 356 the modeling study of Peng et al.⁴ highlighted the importance of accurately representing NO_x 357 chemistry (including PAN formation) and oxidant levels in oxidation flow reactors (typically used 358 for measurements of SOA formation) in order to reflect atmospheric conditions.

This work builds onto these previous studies by showing that the entire product distribution (not only the formation of HOMs and SOA) can be impacted by NO_x effects that go beyond the standard RO_2 branching ($RO_2 + NO$ vs. $RO_2 + HO_2$ vs. RO_2 isomerization). We find that, even in a high NO regime, a lower NO/NO₂ ratio fosters higher concentrations of secondary OH, higher PAN concentrations, and a more highly oxidized product distribution. Together, these results can affect the entire product distribution in chamber experiments, leading to a potential disconnect between chamber results and product distributions expected in the atmosphere.

PAN plays an important role in atmospheric systems by sequestering HO_x and NO_x , thereby influencing the kinetics of organic carbon evolution. Here, PAN formation is observed to be highly sensitive to the NO/NO₂ ratio. Under lower-NO conditions, PAN is observed to form 369 preferentially, limiting the extent of subsequent chemistry. As such, PAN formation can affect 370 SOA formation, even when it does not serve as a direct intermediate in SOA formation (as is the 371 case in isoprene oxidation¹⁹). This has implications for recreating high-NO atmospheric conditions 372 in chamber experiments. Specifically, when trying to achieve "polluted conditions" it is not 373 sufficient to flood the reactor with NO; while this ensures that $RO_2 + HO_2$ and RO_2 isomerization 374 reactions cannot compete with $RO_2 + NO$, it risks leading to PAN concentrations that may not be 375 representative of atmospheric conditions. Rather, the atmospheric NO/NO₂ ratio has an important 376 influence on the relevance of chamber results to atmospheric conditions; a lower NO/NO₂ ratio results in increased levels of SOA (via the increasingly competitive $RO_2 + NO_2$ reaction channel), 377 378 while a higher NO/NO₂ ratio results in fewer products and lower mean oxidation states. This work 379 therefore highlights the need for experimental studies of product distributions and SOA formation 380 to be carried out under atmospherically relevant NO/NO2 ratios. This has been suggested previously for the accurate simulation of SOA and HOM formation;^{19,20} here we show that the use 381 382 of atmospherically relevant NO/NO₂ ratios is important in virtually all oxidation systems in order 383 to better simulate the complex, multiphase product distributions generated during atmospheric 384 oxidation processes. It is thus important that laboratory product studies be carried out under 385 conditions in which both the NO/NO₂ ratio and RO₂ chemistry are accurately representative of the 386 atmosphere. Both the absolute NO_x level and the NO/NO₂ ratios may be important in controlling 387 product distributions, and future study of these effects should focus on how product distributions 388 depend on both.

389

390 Supporting Information

391 Further information regarding the molecular assignments of PTR data; concentrations of NO and 392 NO₂ for each of the experiments (Table S1; Figures S1-2); discussion of wall and tubing losses 393 (Figure S3); deconvolution of precursor and product contributions to $C_4H_9O^+$ (Figure S4); PTR 394 stacked plots for *n*-butyl nitrite photolysis (Figure S5); concentrations, sources, and sinks of OH 395 in *n*-butyl nitrite photolysis as predicted by MCM simulations (Figure S6); experimental uncertainty and reproducibility (Figure S7); C₄H₇O⁺ product time series under different NO/NO₂ 396 397 ratios (Figure S8); oxidation states of gas-phase product distributions (Figure S9); detailed 398 correlation plots for the photolysis of *n*-butyl nitrite (Figure S10); product distributions from the 399 photolysis *n*-pentyl nitrite (Figures S11-12). 400

401 Acknowledgments.

This work was supported by NSF grant CHE-1709993. The authors also wish to acknowledge Josh
Moss for his insight on MCM simulations and PTR fragmentation patterns; Andrew Lambe and
Anthony Carrasquillo for their assistance in precursor synthesis; and Megan Claflin for sharing
her knowledge on detection of alkyl nitrites in the PTR.

406

407 **References**

408 (1) Lannuque, V.; Camredon, M.; Couvidat, F.; Hodzic, A.; Valorso, R.; Madronich, S.;

409 Bessagnet, B.; Aumont, B. Exploration of the Influence of Environmental Conditions on

- 410 Secondary Organic Aerosol Formation and Organic Species Properties Using Explicit
- 411 Simulations: Development of the VBS-GECKO Parameterization. *Atmos. Chem. Phys.*
- **2018**, *18* (18), 13411–13428.
- 413 (2) Goldstein, A. H.; Galbally, I. E. Known and Unexplored Organic Constituents in the

414		Earth's Atmosphere. Environ. Sci. Technol. 2007, 41 (5), 1514–1521.
415	(3)	Hallquist, M.; Wenger, J. C.; Baltensperger, U.; Rudich, Y.; Simpson, D.; Claeys, M.;
416		Dommen, J.; Donahue, N. M.; George, C.; Goldstein, A. H.; Hamilton, J. H.; Herrmann,
417		H.; Hoffmann, T.; Iinuma, Y.; Jang, M.; Jenkin, M. E.; Jimenez, J. L.; Kiendler-Scharr,
418		A.; Maenhaut, W.; McFiggans, G.; Mentel, Th. F.; Monod, A.; Prévôt, A. S. H.; Seinfeld,
419		J. H.; Surratt, J. D.; Szmigielski, R.; Wildt, J. The Formation, Properties and Impact of
420		Secondary Organic Aerosol: Current and Emerging Issues. Atmos. Chem. Phys. 2009, 9
421		(June), 5155–5236.
422	(4)	Peng, Z.; Lee-Taylor, J.; Orlando, J. J.; Tyndall, G. S.; Jimenez, J. L. Organic Peroxy
423		Radical Chemistry in Oxidation Flow Reactors and Environmental Chambers and Their
424		Atmospheric Relevance. Atmos. Chem. Phys. 2019, 19 (2), 813-834.
425	(5)	Orlando, J. J.; Tyndall, G. S. Laboratory Studies of Organic Peroxy Radical Chemistry:
426		An Overview with Emphasis on Recent Issues of Atmospheric Significance. Chem. Soc.
427		<i>Rev.</i> 2012 , <i>41</i> (19), 6294–6317.
428	(6)	Crounse, J. D.; Nielsen, L. B.; Jørgensen, S.; Kjaergaard, H. G.; Wennberg, P. O.
429		Autoxidation of Organic Compounds in the Atmosphere. J. Phys. Chem. Lett. 2013, 4
430		(20), 3513–3520.
431	(7)	Lelieveld, J.; Butler, T. M.; Crowley, J. N.; Dillon, T. J.; Fischer, H.; Ganzeveld, L.;
432		Harder, H.; Lawrence, M. G.; Martinez, M.; Taraborrelli, D.; Williams, J. Atmospheric
433		Oxidation Capacity Sustained by a Tropical Forest. Nature 2008, 452 (7188), 737–740.
434	(8)	Ren, X.; Harder, H.; Martinez, M.; Lesher, R. L.; Oliger, A.; Shirley, T.; Adams, J.;
435		Simpas, J. B.; Brune, W. H. HO _x Concentrations and OH Reactivity Observations in New
436		York City during PMTACS-NY2001. Atmos. Environ. 2003, 37 (26), 3627-3637.

- 437 (9) Martinsson, J.; Eriksson, A. C.; Nielsen, I. E.; Malmborg, V. B.; Ahlberg, E.; Andersen,
- 438 C.; Lindgren, R.; Nyström, R.; Nordin, E. Z.; Brune, W. H.; Svenningsson, B.; Swietlicki,
- 439 E.; Boman, C.; Pagels, J. H. Impacts of Combustion Conditions and Photochemical
- 440 Processing on the Light Absorption of Biomass Combustion Aerosol. *Environ. Sci.*
- 441 *Technol.* **2015**, *49* (24), 14663–14671.
- 442 (10) Tkacik, D. S.; Lambe, A. T.; Jathar, S.; Li, X.; Presto, A. A.; Zhao, Y.; Blake, D. R.;
- 443 Meinardi, S.; Jayne, J. T.; Croteau, P. L.; Robinson, A. L. Secondary Organic Aerosol
- 444 Formation from In-Use Motor Vehicle Emissions Using a Potential Aerosol Mass Reactor.
- 445 Environ. Sci. Technol. **2014**, 48 (19), 11235–11242.
- 446 (11) Ortega, A. M.; Day, D. A.; Cubison, M. J.; Brune, W. H.; Bon, D.; De Gouw, J. A.;
- 447 Jimenez, J. L. Secondary Organic Aerosol Formation and Primary Organic Aerosol
- 448 Oxidation from Biomass-Burning Smoke in a Flow Reactor during FLAME-3. *Atmos.*
- 449 *Chem. Phys.* **2013**, *13* (22), 11551–11571.
- 450 (12) Kroll, J. H.; Ng, N. L.; Murphy, S. M.; Flagan, R. C.; Seinfeld, J. H. Secondary Organic
- 451 Aerosol Formation from Isoprene Photooxidation. *Environ. Sci. Technol.* 2006, 40 (6),
 452 1869–1877.
- 453 (13) Crounse, J. D.; Paulot, F.; Kjaergaard, H. G.; Wennberg, P. O. Peroxy Radical
- 454 Isomerization in the Oxidation of Isoprene. *Phys. Chem. Chem. Phys.* **2011**, *13* (30),
- 455 13607–13613.
- 456 (14) Peng, Z.; Jimenez, J. L. Modeling of the Chemistry in Oxidation Flow Reactors with High
 457 Initial NO. *Atmos. Chem. Phys.* 2017, *17* (19), 11991–12010.
- 458 (15) Kroll, J. H.; Seinfeld, J. H. Chemistry of Secondary Organic Aerosol: Formation and
- 459 Evolution of Low-Volatility Organics in the Atmosphere. *Atmos. Environ.* 2008, 42 (16),

- 460 3593–3624.
- 461 (16) Ziemann, P. J.; Atkinson, R. Kinetics, Products, and Mechanisms of Secondary Organic
 462 Aerosol Formation. *Chem. Soc. Rev.* 2012, *41* (19), 6582–6605.
- 463 (17) Arey, J.; Aschmann, S. M.; Kwok, E. S. C.; Atkinson, R. Alkyl Nitrate, Hydroxyalkyl
- 464 Nitrate, and Hydroxycarbonyl Formation from the NO_x -Air Photooxidations of C_5 - $C_8 n$ -
- 465 Alkanes. J. Phys. Chem. A **2001**, 105 (6), 1020–1027.
- 466 (18) Zhang, X.; Cappa, C. D.; Jathar, S. H.; McVay, R. C.; Ensberg, J. J.; Kleeman, M. J.;
- 467 Seinfeld, J. H. Influence of Vapor Wall Loss in Laboratory Chambers on Yields of
- 468 Secondary Organic Aerosol. *Proc. Natl. Acad. Sci.* **2014**, *111* (16), 5802–5807.
- 469 (19) Chan, A. W. H.; Chan, M. N.; Surratt, J. D.; Chhabra, P. S.; Loza, C. L.; Crounse, J. D.;
- Yee, L. D.; Flagan, R. C. Role of Aldehyde Chemistry and NO_x Concentrations in
 Secondary Organic Aerosol Formation. *Atmos. Chem. Phys.* 2010, *10*, 7169–7188.
- 472 (20) Rissanen, M. P. NO2 Suppression of Autoxidation Inhibition of Gas-Phase Highly
- 473 Oxidized Dimer Product Formation. ACS Earth Sp. Chem. 2018, 2 (11), 1211–1219.
- 474 (21) Hoffmann, T.; Odum, J. A. Y. R.; Bowman, F.; Collins, D.; Klockow, D.; Flagan, R. C.;
- 475 Seinfeld, J. H. Formation of Organic Aerosols from the Oxidation of Biogenic
- 476 Hydrocarbons. J. Atmos. Chem. **1997**, 1, 189–222.
- 477 (22) Kessler, S. H.; Nah, T.; Carrasquillo, A. J.; Jayne, J. T.; Worsnop, D. R.; Wilson, K. R.;
- 478 Kroll, J. H. Formation of Secondary Organic Aerosol from the Direct Photolytic
- 479 Generation of Organic Radicals. J. Phys. Chem. Lett. 2011, 2 (11), 1295–1300.
- 480 (23) Carrasquillo, A. J.; Hunter, J. F.; Daumit, K. E.; Kroll, J. H. Secondary Organic Aerosol
- 481 Formation via the Isolation of Individual Reactive Intermediates: Role of Alkoxy Radical
- 482 Structure. J. Phys. Chem. A **2014**, 118 (38), 8807–8816.

483	(24)	Carrasquillo, A. J.; Daumit, K. E.; Kroll, J. H. Radical Reactivity in the Condensed Phase:
484		Intermolecular versus Intramolecular Reactions of Alkoxy Radicals. J. Phys. Chem. Lett.
485		2015 , <i>6</i> (12), 2388–2392.
486	(25)	Lim, C. Y.; Hagan, D. H.; Coggon, M. M.; Koss, A. R.; Sekimoto, K.; De Gouw, J. A.;
487		Warneke, C.; Cappa, C. D.; Kroll, J. H. Secondary Organic Aerosol Formation from the
488		Laboratory Oxidation of Biomass Burning Emissions. Atmos. Chem. Phys. 2019, 19 (19),
489		12797–12809.
490	(26)	Heicklen, J. The Decomposition of Alkyl Nitrites and the Reaction of Alkoxy Radicals. In
491		Advances in Photochemistry; 1988; Vol. 14, pp 177–272.
492	(27)	Silvern, R. F.; Jacob, D. J.; Travis, K. R.; Sherwen, T.; Evans, M. J.; Cohen, R. C.;
493		Laughner, J. L.; Hall, S. R.; Ullmann, K.; Crounse, J. D.; Wennberg, P. O.; Peischl, J.;
494		Pollack, I. B. Observed NO/NO ₂ Ratios in the Upper Troposphere Imply Errors in NO-
495		NO ₂ -O ₃ Cycling Kinetics or an Unaccounted NO _x Reservoir. <i>Geophys. Res. Lett.</i> 2018 , 45
496		(9), 4466–4474.
497	(28)	Denisova, T. G.; Denisov, E. T. Kinetic Parameters of Alkyl, Alkoxy, and Peroxy Radical
498		Isomerization. Kinet. Catal. 2001, 42 (5), 620-630.
499	(29)	Praske, E.; Otkjær, R. V.; Crounse, J. D.; Hethcox, J. C.; Stoltz, B. M.; Kjaergaard, H. G.;
500		Wennberg, P. O. Atmospheric Autoxidation Is Increasingly Important in Urban and
501		Suburban North America. Proc. Natl. Acad. Sci. 2018, 115 (1), 64-69.
502	(30)	Otkjær, R. V.; Jakobsen, H. H.; Tram, C. M.; Kjaergaard, H. G. Calculated Hydrogen
503		Shift Rate Constants in Substituted Alkyl Peroxy Radicals. J. Phys. Chem. A 2018, 122,
504		8665–8673.
505	(31)	Griffin, R. J.; Cocker, D. R.; Flagan, R. C.; Seinfeld, J. H. Organic Aerosol Formation

	from the Oxidation of Biogenic Hydrocarbons. J. Geophys. Res. 1999, 104 (D3), 3555-
	3567.
(32)	DeCarlo, P. F.; Kimmel, J. R.; Trimborn, A.; Northway, M. J.; Jayne, J. T.; Aiken, A. C.;
	Gonin, M.; Fuhrer, K.; Horvath, T.; Docherty, K. S.; Worsnop, D. R.; Jimenez, J. L.;
	Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer. Anal.
	Chem. 2006, 78 (24), 8281–8289.
(33)	Canagaratna, M. R.; Jimenez, J. L.; Kroll, J. H.; Chen, Q.; Kessler, S. H.; Massoli, P.;
	Hildebrandt Ruiz, L.; Fortner, E.; Williams, L. R.; Wilson, K. R.; Surratt, J. D.; Donahue,
	N. M.; Jayne, J. T.; Worsnop, D. R. Elemental Ratio Measurements of Organic
	Compounds Using Aerosol Mass Spectrometry: Characterization, Improved Calibration,
	and Implications. Atmos. Chem. Phys. 2015, 15 (1), 253-272.
(34)	Kroll, J. H.; Donahue, N. M.; Jimenez, J. L.; Kessler, S. H.; Canagaratna, M. R.; Wilson,
	K. R.; Altieri, K. E.; Mazzoleni, L. R.; Wozniak, A. S.; Bluhm, H.; Mysak, E. R.; Smith,
	J. D.; Kolb, C. E.; Worsnop, D. R. Carbon Oxidation State as a Metric for Describing the
	Chemistry of Atmospheric Organic Aerosol. Nat. Chem. 2011, 3 (2), 133-139.
(35)	Krechmer, J. E.; Lopez-Hilfiker, F. D.; Koss, A. R.; Hutterli, M.; Stoermer, C.; Deming,
	B.; Kimmel, J. R.; Warneke, C.; Holzinger, R.; Jayne, J.; Worsnop, D.; Fuhrer, K.; Gonin,
	M.; de Gouw, J. Evaluation of a New Reagent-Ion Source and Focusing Ion-Molecule
	Reactor for Use in Proton-Transfer-Reaction Mass Spectrometry. Anal. Chem. 2018, 90,
	12011–12018.
(36)	Isaacman-VanWertz, G.; Massoli, P.; O'Brien, R. E.; Nowak, J. B.; Canagaratna, M. R.;
	Jayne, J. T.; Worsnop, D. R.; Su, L.; Knopf, D. A.; Misztal, P.; Arata, C.; Goldstein, A.
	H.; Kroll, J. H. Using Advanced Mass Spectrometry Techniques to Fully Characterize
	(32) (33) (34) (35)

- 529 Atmospheric Organic Carbon: Current Capabilities and Remaining Gaps. *Faraday*530 *Discuss.* 2017, 200, 579–598.
- 531 (37) Aoki, N.; Inomata, S.; Tanimoto, H. Detection of C₁-C₅ Alkyl Nitrates by Proton Transfer
- 532 Reaction Time-of-Flight Mass Spectrometry. Int. J. Mass Spectrom. 2007, 263 (1), 12–21.
- 533 (38) Yuan, B.; Koss, A. R.; Warneke, C.; Coggon, M. M.; Sekimoto, K.; De Gouw, J. A.
- 534 Proton-Transfer-Reaction Mass Spectrometry: Applications in Atmospheric Sciences.
- 535 *Chem. Rev.* **2017**, *117* (21), 13187–13229.
- 536 (39) Lim, Y. Bin; Ziemann, P. J. Effects of Molecular Structure on Aerosol Yields from OH
- 537 Radical-Initiated Reactions of Linear, Branched, and Cyclic Alkanes in the Presence of
- 538 NO_x. Environ. Sci. Technol. **2009**, *43*, 2328–2334.
- 539 (40) Lim, Y. Bin; Ziemann, P. J. Products and Mechanism of Secondary Organic Aerosol
 540 Formation from Reactions of *n*-Alkanes with OH Radicals in the Presence of NO_x.
- 541 Environ. Sci. Technol. 2005, 39 (23), 9229–9236.
- 542 (41) Hunter, J. F.; Carrasquillo, A. J.; Daumit, K. E.; Kroll, J. H. Secondary Organic Aerosol
- Formation from Acyclic, Monocyclic, and Polycyclic Alkanes. *Environ. Sci. Technol.*2014, 48 (17), 10227–10234.
- 545 (42) Noyes, W. A. Explanation of the Formation of Alkyl Nitrites in Dilute Solutions; Butyl
 546 and Amyl Nitrites. *J. Am. Chem. Soc.* 1933, 55 (9), 3888–3889.
- 547 (43) Bloss, C.; Wagner, V.; Jenkin, M. E.; Volkamer, R.; Bloss, W. J.; Lee, J. D.; Heard, D. E.;
- 548 Wirtz, K.; Rea, G.; Wenger, J. C.; Pilling, M. J. Development of a Detailed Chemical
- 549 Mechanism (MCMv3.1) for the Atmospheric Oxidation of Aromatic Hydrocarbons.
- 550 Atmos. Chem. Phys. 2005, 3 (2002), 641–664.
- 551 (44) Saunders, S. M.; Jenkin, M. E.; Derwent, R. G.; Pilling, M. J. Protocol for the

552		Development of the Master Chemical Mechanism, MCM v3 (Part A): Tropospheric
553		Degradation of Non-Aromatic Volatile Organic Compounds. Atmos. Chem. Phys. 2003, 3,
554		161–180.
555	(45)	Wolfe, G. M.; Marvin, M. R.; Roberts, S. J.; Travis, K. R.; Liao, J. The Framework for 0-
556		D Atmospheric Modeling (F0AM) v3.1. Geosci. Model Dev. 2016, 9 (9), 3309-3319.
557	(46)	Calvert, J. G.; Orlando, J. J.; Stockwell, W. R.; Wallington, T. J. The Mechanisms of
558		Reactions Influencing Atmospheric Ozone; Oxford University Press, 2015.
559	(47)	Demarcke, M.; Amelynck, C.; Schoon, N.; Dhooghe, F.; Rimetz-Planchon, J.; Van
560		Langenhove, H.; Dewulf, J. Laboratory Studies in Support of the Detection of Biogenic
561		Unsaturated Alcohols by Proton Transfer Reaction-Mass Spectrometry. Int. J. Mass
562		Spectrom. 2010, 290 (1), 14–21.
563	(48)	Hastie, D. R.; Gray, J.; Langford, V. S.; Maclagan, R. G. A. R.; Milligan, D. B.; Mcewan,
564		M. J. Real-Time Measurement of Peroxyacetyl Nitrate Using Selected Ion Flow Tube
565		Mass Spectrometry. Rapid Commun. Mass Spectrom. 2010, 24, 343-348.
566	(49)	Hansel, A.; Wisthaler, A. A Method for Real-Time Detection of PAN, PPN and MPAN in
567		Ambient Air. Geophys. Res. Lett. 2000, 27 (6), 895-898.
568	(50)	Blake, R. S.; Monks, P. S.; Ellis, A. M. Proton-Transfer Reaction Mass Spectrometry.
569		Chem. Rev. 2009, 109 (0), 861–896.