

MIT Open Access Articles

Ground Motion in Kuwait from Regional and Local Earthquakes: Potential Effects on Tall Buildings

The MIT Faculty has made this article openly available. **Please share** how this access benefits you. Your story matters.

As Published: <https://doi.org/10.1007/s00024-018-1943-5>

Publisher: Springer International Publishing

Persistent URL: <https://hdl.handle.net/1721.1/131414>

Version: Author's final manuscript: final author's manuscript post peer review, without publisher's formatting or copy editing

Terms of use: Creative Commons Attribution-Noncommercial-Share Alike



1 **Ground Motion in Kuwait from Regional and Local Earthquakes:**
2 **Potential Effects on Tall Buildings**

3 Chen Gu¹, Germán A Prieto⁴, Abdullah Al-Enezi³, Farah Al-Jeri³, Jamal Al-
4 Qazweeni³, Hasan Kamal³, Sadi Kuleli¹, Aurélien Mordret¹, Oral Büyüköztürk²,
5 M. Nafi Toksöz¹

6 ¹Earth Resources Laboratory, Department of Earth, Atmospheric, and Planetary
7 Sciences, Massachusetts Institute of Technology, Cambridge, MA, USA.

8 ²Laboratory for Infrastructure Science and Sustainability, Department of Civil
9 and Environmental Engineering, Massachusetts Institute of Technology,
10 Cambridge, MA, USA.

11 ³Kuwait Institute for Scientific Research, Kuwait City, Kuwait.

12 ⁴Departamento de Geociencias, Facultad de Ciencias, Universidad Nacional de
13 Colombia, Bogotá, Colombia

14 **Abstract** In recent years, the construction of tall buildings has been increasing in
15 many countries, including Kuwait and other Gulf states. These tall buildings are
16 especially sensitive to ground shaking due to long period seismic surface waves.
17 Although Kuwait is relatively aseismic, it has been affected by large ($M_w > 6$)
18 regional earthquakes in the Zagros Fold-Thrust Belt (ZFTB). Accurate ground
19 motion prediction for large earthquakes is important to assess the seismic hazard to
20 tall buildings. In this study, we first analyze the observed ground motions due to
21 two earthquakes widely felt in Kuwait: the 08/18/2014 M_w 6.2 earthquake, 360 km
22 NNE of Kuwait City, and the 11/12/2017 M_w 7.3 earthquake, 642 km NNE of
23 Kuwait City. The peak spectral displacement periods of the ground motion from
24 the 08/18/2014 M_w 6.2 earthquake matched well with the ambient vibration
25 spectra of the tallest building - the Al-Hamra Tower. We calculate the ground
26 motions from potential regional and local earthquakes. We use a velocity model
27 obtained by matching the observed seismograms of the 2014 and 2017
28 earthquakes. We calculate ground motions in Kuwait due to a regional $M_w = 7.5$
29 earthquake, and a local $M_w = 5.0$ earthquakes. Our study shows that a significant
30 source of seismic hazard to tall buildings in Kuwait comes from the regional
31 tectonic earthquakes. However, local earthquakes have the potential to generate
32 high peak ground accelerations ($\sim 98 \text{ cm/sec}^2$) close to their epicenters.

33

34

1. Introduction

35 There has been a rapid increase in the total number and construction of tall
36 buildings worldwide, including in Kuwait. Tall buildings are greatly affected by
37 long trains of low-frequency seismic surface waves from regional earthquakes.
38 Ground motions from regional large earthquakes have frequency contents that may
39 be in the same range as the natural frequencies of some of these tall buildings; and
40 this may have damaging consequences (Shakal et al., 1996; Çelebi and Liu, 1998;
41 Çelebi et al., 2014). The building response to seismic shaking may be obtained by
42 seismic interferometry using ambient vibrations or recorded earthquakes from the
43 building monitoring system (Kohler et al., 2005; Snieder and Şafak, 2006; Prieto et
44 al., 2010; Sun et al., 2017).

45 Previous studies used different methods to calculate the ground motion in a
46 region (Abrahamson and Shedlock, 1997; Olsen et al., 1995; Pitarka et al., 1998;
47 Olsen, 2000; Olsen et al., 2006, 2009; Prieto and Beroza, 2008; Denolle et al.,
48 2013, 2014). The ground motion prediction equation (GMPE) is one method to
49 estimate ground motion intensity based on the observed seismic data and
50 attenuation laws (Abrahamson and Shedlock, 1997; Cauzzi et al., 2014; Danciu et
51 al., 2017, 2018; Şeşetyan et al., 2018). However, in many areas with limited
52 information this approach is primarily directed to acceleration and does not include
53 long period surface waves and the earthquake source rupture process, especially

54 important for large earthquakes; it also ignores the effects of regional variations in
55 crustal structures along the wave propagation path. To make the prediction more
56 physical and realistic, many studies use numerical methods to predict ground
57 motions from a simulated large earthquake (Olsen et al., 1995; Pitarka et al., 1998;
58 Olsen, 2000; Olsen et al., 2006, 2009). The reliability of these ground motion
59 simulations depends heavily on the accuracy of the earthquake source rupture
60 model and the velocity and attenuation models used to propagate seismic waves. In
61 the last decade, studies have started using the ambient seismic field to predict the
62 ground motions from sources around seismic stations (Prieto and Beroza, 2008;
63 Denolle et al., 2013, 2014). This method, although appealing, is limited because it
64 only predicts the ground motions near available seismic stations and thus requires a
65 high-density network. In this report, we use numerically simulated ground motions
66 for both regional and local earthquakes in and near Kuwait.

67 Ground motion predictions in Kuwait are possible thanks to well-documented
68 studies on regional tectonics, earthquake source mechanisms and seismic velocity
69 structures in and around Kuwait. These data provide the starting models for the
70 ground motion calculations (Jackson and McKenzie, 1984; Ni and Barazangi,
71 1986; Bou-Rabee, 2000; Gök et al., 2000; Talebian and Jackson, 2004; Pasyanos et
72 al., 2007; Sadek, 2008; Abbas and Al-Sabri, 2015; Kottmeier et al., 2016; Çaktı et
73 al., 2016). In addition, some previous research has been conducted on long-period

74 ground motions in the Arabian Gulf (Pitarka et al., 2012, 2015). El-Hussain et al.
75 (2012, 2013, 2015) carried out extensive studies of earthquake catalogs for the
76 Zagros Fold-Thrust Belt (ZFTB) and Gulf region as part of their study of seismic
77 hazard in Oman. In the State of Kuwait, hazard studies have been confined mostly
78 to local regions and earthquakes (Bou-Rabee, 2000).

79 In this paper, we calculate ground motions from large regional earthquakes
80 around Kuwait, mostly in the ZFTB and from the largest local earthquakes in
81 Kuwait. To minimize uncertainties associated with crustal and sedimentary basin
82 structures, we determine a regional velocity model using data from two regional
83 earthquakes, 08/18/2014 Mw 6.2 and 11/12/2017 Mw 7.3, recorded by seismic
84 stations in Kuwait. To deterministically address seismic hazard from regional
85 earthquakes in Kuwait City, we simulate ground motions from a $M_w = 7.5$ regional
86 earthquake, located 300 km east of Kuwait City. The response spectra (spectral
87 displacement/pseudo-acceleration spectra) are also calculated and compared with
88 the mode vibrations obtained from ambient vibrations spectra of the Al-Hamra
89 Tower to assess the seismic hazard to tall buildings in Kuwait City. In addition, the
90 ground motions due to recent local earthquakes are calculated.

2. *Ground Motion Data and Model*

91
92
93
94
95
96
97
98
99
100
101
102
103
104
105
106
107
108

The Kuwait National Seismic Network (KNSN), started in 1997, has been monitoring earthquakes since then (Figure 2b). The KNSN initially consisted of seven three-component short-period stations (AB, MI, QR, RD, RS, SA, UM) and one three-component broadband station (KB). In 2013, the KNSN upgraded their seismic monitoring network by replacing five short-period stations (MI, QR, RD, RS, UM) with broadband stations (RefTek's 151-120 broadband seismometers with a natural period of 120 sec). In this study, we use only the broadband seismic data, since the instrument response files for short-period stations AB and SA are not available. We deconvolve the instrument response from the raw digital data to convert them to ground velocity seismograms. A quarter-cycle cosine taper, with unit response between 0.0125 Hz and 8.0 Hz and zero below 0.01 Hz and above 10.0 Hz, was applied during the deconvolution to dampen the response at very low and high frequencies. The displacement and acceleration seismograms are obtained by integrating and differentiating the velocity seismograms, respectively. The spectral displacement, pseudo-velocity spectra and pseudo-acceleration spectra are calculated to estimate the potential effects of earthquakes on civil structures (Gupta, 1992).

109
110

To calculate the ground motion, we used the discrete-wavenumber method (Bouchon, 1981; 2003). This method, which converts the continuous wavenumber

111 integration to a summation of discrete wavenumbers, is an accurate and
112 computationally efficient method for generating synthetic seismograms for a 1-D
113 layered velocity model. We could have used a finite difference method for 2-D or
114 3-D velocity models if such models were available (Olsen, 2000; Olsen et al.,
115 2006; Maeda and Furumura, 2013; Almuheidib and Toksöz, 2014).

116 The 1-D layered velocity model we used is based on the KUW1 model of
117 Pasyanos et al. (2007). The KUW1 model was modified to include the
118 sedimentary layers and fit the observed seismograms from the two large
119 earthquakes in 2014 and 2017. The new model, designated KUW-P or preferred
120 model, is listed in Table 1 and shown in Figure 3. More detailed velocity model
121 improvement procedure is included in the supplemental materials. The large
122 regional earthquakes occur mostly in the Zagros Fold-Thrust Belt (ZFTB) and
123 surrounding areas (Pasyanos et al., 2007; Sadek, 2008). Figure 1 shows significant
124 regional earthquakes around Kuwait in the period 1997-2017. On average, about
125 one earthquake of magnitude 6 or greater occurs in the region per year. In Figure 1,
126 we also show sample seismograms (vertical components) recorded at station MI.
127 Unfortunately, the length of the recordings is limited to 200 seconds and
128 seismograms are truncated. For closer events (e.g. at about 300 km distance), we
129 note the long duration of surface waves.

130 The observed long-period, long-duration ground motion in Kuwait shows that
131 regional earthquakes are one source of hazards to tall buildings in Kuwait. To
132 demonstrate the impact of regional earthquakes on tall buildings in Kuwait, we
133 analyze seismograms from two large Zagros belt earthquakes that occurred in 2014
134 and 2017.

135 The 08/18/2014 Mw 6.2 earthquake, which happened at the Iran-Iraq border,
136 360 km NNE of Kuwait City, was widely felt in Kuwait. We show the three-
137 component seismograms of this earthquake from four stations in Kuwait and
138 RAYN in Saudi Arabia in Figure 2. Since this earthquake occurred to the north of
139 Kuwait and from station RYAN in Saudi Arabia, the Z and N component
140 seismograms are alike and are with a 90-degree phase shift, mostly Rayleigh
141 waves. The seismograms at the station RD and QR in Kuwait are truncated
142 because the recordings were limited to 200 secs. The E component is primarily
143 Love waves.

144 To calculate synthetic seismograms, we use a 1-D layered velocity model
145 (KUW-P). The comparison between the observed and calculated seismograms for
146 the Mw 6.2 are shown in Figure 4. The figure also includes the spectral
147 displacements. The peak displacement response peaks at periods 3.4 s, 6.9 s, and
148 16.0 s are due to shear, Rayleigh and Love waves, respectively.

149 We do the similar comparison of synthetic and observed seismograms for the
150 recent 11/12/2017 Mw 7.3 earthquake, 642 km NNE of Kuwait City, at the Iran-
151 Iraq border. The comparison of the observed and modeled ground displacements
152 and spectral displacements of the Mw 7.3 earthquake at stations KB, QR and RS is
153 shown in Figure 5. The horizontal components of broad band recordings were
154 slightly clipped due to the large magnitude. In spite of this, the match between the
155 observed displacement seismograms and spectral displacements is good, especially
156 for the vertical components least impacted by the clipping. The peak spectral
157 displacements are above 10 s period. The peak ground displacement is about 4 cm
158 and the duration of the ground motions is longer than 300 sec.

159 *3. Ground Motions in Kuwait Due to Regional and local Earthquakes*

160 The determination of ground motion requires the maximum credible magnitude,
161 source properties and the most likely locations. For Kuwait, the primary site for a
162 regional earthquake is the segment of the ZFTB, 300 km east of the Kuwait City.
163 The maximum credible magnitude is $M_w = 7.5$, based on the regional earthquake
164 catalog from the KNSN and the work of El-Hussain et al. (2012, 2013), who have
165 conducted extensive studies of regional earthquakes in the ZFTB for seismic
166 hazard zonation in the Sultanate of Oman. A $M_w = 7.5$ earthquake corresponds to
167 a moment of $M_0 = 2.24 \times 10^{27}$ dyne \square cm. In our calculation, source (fault rupture

168 area) was assumed to be $20 \text{ km} \times 40 \text{ km}$ with the top of the rupture 5 km below the
169 surface. The modeled earthquake is a thrust event with fault plane: strike = 300° ,
170 dip = 20° , and rake = 90° . Using the KUW-P velocity model (Table 1), we
171 calculate synthetic seismograms with the epicenter located 300 km east of Kuwait
172 City.

173 The calculated ground accelerations and displacements in Kuwait are shown in
174 Figure 6. Our calculations show that in Kuwait City, the maximum acceleration is
175 46 cm/sec^2 and maximum displacement is 23 cm. The peaks of the displacement
176 response spectra are above 10 s period. The duration of the ground motion is
177 longer than 300 seconds.

178 To demonstrate the effects of ground motions in Kuwait due to regional
179 earthquakes on tall buildings in Kuwait City, we look at the Al-Hamra Tower. The
180 Al-Hamra Tower is a 414 m tall sculpted concrete steel structure. The resonant
181 frequencies of the tower were obtained by finite element modeling and by ambient
182 vibrations measurements. For the ambient vibrations monitoring, Kuwait Institute
183 for Scientific Research (KISR) deployed two accelerometers at the top terrace of
184 the Al-Hamra Tower in Kuwait City (two Kinometrics systems, with 3-component
185 Episensor and Q330 dataloggers). We used ambient vibrations to obtain vibration
186 spectra of the Al-Hamra Tower on day 329 of 2013, shown in Figure 7.

187 The peak periods obtained from the two instruments are quite consistent. The
188 periods of the first and second strongest building spectral peaks (7.1 sec and 5.7
189 sec) are close to the calculated peak spectral displacement shown in Figure 4. This
190 overlap between the dominant frequency of ground motions from regional
191 earthquakes, similar to the Mw 6.2 earthquake, and the natural frequencies of the
192 Al-Hamra Tower suggests that regional earthquakes will cause strong resonance
193 vibrations in the Al-Hamra Tower and other tall buildings of similar height. The
194 observed ground motions due to the Mw 7.3 earthquake, 642 km from Kuwait City,
195 have large (about 4 cm) amplitudes at periods less the 10 sec, but the peak
196 displacement response periods are larger than 10 sec; however, the large peak
197 ground displacements, ~4 cm due to the Mw 7.3 earthquake could still cause large
198 motions at the higher floors of the building. The observed building motions due to
199 the Mw 7.3 earthquake from the first GPS measurements of the Al-Hamra tower
200 (Herring et al., 2018, submitted) indicate a building displacement of ~20 cm at the
201 top of the 414 m tall building. This amount of motion did not cause any damage to
202 the Al-Hamra tower.

203 It is important to mention that the calculated ground motions shown in the
204 figures are for a competent rock layer with shear velocity of 1.4 km/sec. “Soft”
205 sediments with lower shear velocity would further amplify the ground motions.
206 More detailed geotechnical information about the soil properties and sedimentary

207 structures at the site are needed to better quantify and predict the ground motions
208 (Bowden and Tsai, 2017). However, this is beyond the scope of this paper.

209 Local earthquakes in Kuwait occur frequently but tend to have small
210 magnitudes (Pasyanos, 2007). Since 1997, the KNSN has recorded more than 1000
211 earthquakes ($M_w < 5$). Most of the local earthquakes are distributed close to the
212 oil/gas fields in the northeast and southwest of Kuwait (Figure 8) and are probably
213 induced by oil production activities (Gu et al., 2017). Two of the largest local
214 earthquakes - M_w 4.5 on 03/21/2015 and M_w 4.1 on 08/18/2015 - are located near
215 the southern oil fields and near the northern oil fields, respectively. Such seismicity
216 related to oil/gas fields has also been observed in North America, Europe and the
217 Middle East (Sarkar, 2008; Li et al., 2011a, b; Zoback and Gorelick, 2012;
218 Ellsworth, 2013; Shapiro, 2015; Rubinstein, 2015; Bommer et al., 2017).

219 The seismograms of the M_w 4.5 earthquake from three selected stations and the
220 corresponding pseudo acceleration response spectra are shown in Figure 9. The
221 peak acceleration response periods are in the range from 0.1 s to 1 s (10 Hz to 1
222 Hz).

223 So far, no earthquakes of magnitude greater than M_w 4.5 have been recorded in
224 Kuwait. To determine the ground motions due to the largest local earthquake in
225 Kuwait, we assumed a maximum credible magnitude M_w 5.0. We generated 3-
226 component synthetics for a local M_w 5.0 earthquake at depth of 3 km (see Figure

227 10), with a reverse fault and source mechanism (strike = 85° , dip = 20° , and rake =
228 10°) that are consistent with regional stress (Carman, 1996; El-Hussain et al., 2012,
229 2015; Abbes and Al-Sabri, 2015; Gu et al., 2017). Figure 10 shows the peak
230 ground acceleration (PGA), representing the root mean square (RMS) of the east
231 and north components. Unlike the large regional earthquakes, the local
232 earthquakes, which excite short period ground motions, would not affect the tall
233 buildings. These earthquakes could generate accelerations ($\sim 98 \text{ cm/sec}^2$) high
234 enough to affect small structures, built without seismic design codes. Because of
235 the shallow depths, the maximum shaking generates high ground motions but is
236 confined to a small area around these local earthquakes with a steep fall-off. Note
237 that the maximum frequencies for the synthetics are 2Hz, this frequency limitation
238 may under-estimate the PGA.

239 *4. Conclusions*

240 For tall buildings and large structures in Kuwait, the primary hazard is due to
241 regional earthquakes (e.g., from Zagros Fold-Thrust Belt), especially from surface
242 waves. Our ground motion calculation shows a maximum acceleration of 46
243 cm/sec^2 and a maximum displacement of 23 cm in Kuwait City from a virtual Mw
244 7.5 earthquake in the Zagros region. The presence of near surface “soft” layers
245 increases duration and amplitudes of the surface waves (site amplification) and,

246 hence, increases the hazard. Knowing the detailed shallow structure at the building
247 site would help to better estimate the seismic site specific ground motions.

248 The local earthquakes, most likely induced by oil/gas field activities, can also
249 represent seismic hazard (Gu et al., 2017). They can generate high peak ground
250 accelerations ($\sim 0.1g$) at periods less than 1 second, close to their epicenters. These
251 local earthquakes may potentially affect local structures.

252 *Acknowledgements*

253 This project was sponsored by the Kuwait Foundation for the Advancement of
254 Sciences. The project was conducted as part of the Kuwait-MIT signature project
255 on sustainability of Kuwait's built environment under the direction of Oral
256 Büyüköztürk.

257 REFERENCES

258 Abbas, M. and N. A. Al-Sabri (2015). Focal mechanism and stress tensor analysis
259 in the south Red Sea, *9th Gulf Seismic Forum*.

260 Abrahamson, N., and K. Shedlock (1997). Overview of ground motion attenuation
261 relationships. *Seismol. Res. Lett.*, 68(1).

262 Anderson, S. and R. E. Grimm (1998). Rift processes at the Valles Marineris:
263 Constraints from gravity on necking and rate-dependent strength evolution. *J.*
264 *Geophys. Res.*, 103, 11113-11124.

265 Almuheidib, A. M., and M.N. Toksöz (2014). Numerical modeling of elastic-wave
266 scattering by near-surface heterogeneities. *Geophysics*, 79(4), T199–T217.

267 Bommer, J. J., P. J. Stafford, B. Edwards, B. Dost, E. van Dedem, A. Rodriguez-
268 Marek, P. Kruiver, J. van Elk, D. Doornhof, and M. Ntinalexis (2017). Framework
269 for a Ground-Motion Model for Induced Seismic Hazard and Risk Analysis in the
270 Groningen Gas Field, The Netherlands. *Earthquake Spectra*, Vol. 33, No. 2, pp.
271 481-498.

272 Bou-Rabee, F. (1999). Site selection for the field stations of the Kuwait National
273 Seismic Network, *Seismol. Res. Lett.*, 70(6), 712–717.

274 Bou-Rabee, F. (2000). Seismotectonics and earthquake activity of Kuwait. *J.*
275 *Seismol.*, 4(2), 133–141.

276 Bouchon, M. (1981). A simple method to calculate green's functions for elastic
277 layered media. *Bull. Seismol. Soc. Am.*, 71(4), 959–971.

278 Bouchon, M. (2003). A review of the discrete wavenumber method. *Pure. and*
279 *Appl. Geophy.*, 160(3-4), 445–465.

280 Bowden, D. C., and V. C. Tsai (2017). Earthquake ground motion amplification for
281 surface waves. *Geophys. Res. Lett.*, 44, 121-127, doi:10.1002/2016GL071885.

282 Çaktı, E., I. El-Hussain, K. Şeşetyan, A. Deif, U. Hancılar, G. Al-Rawas, Y. Kamer
283 & Al-Jabri, K. (2016). Development of ground-shaking maps for the Sultanate of
284 Oman. *Nat. Hazards*, 82(2), 1357-1373.

285 Carman, G. J. (1996). Structural elements of onshore Kuwait. *GeoArabia*, 1, no. 2,
286 239–266.

287 Cauzzi, C., Faccioli, E., Vanini, M., & Bianchini, A. (2015). Updated predictive
288 equations for broadband (0.01-10 s) horizontal response spectra and peak ground
289 motions, based on a global dataset of digital acceleration records. *Bull. Earthquake*
290 *Eng.*, 13(6), 1587.

291 Çelebi, M., and H.-P. Liu (1998). Before and after retrofit—response of a building
292 during ambient and strong motions. *J. Wind Eng. Ind. Aerodyn.*, 204 77, 259–268.

293 Çelebi, M., M. N. Toksöz, and O. Büyüköztürk (2014). Rocking behavior of an
294 instrumented unique building on the MIT campus identified from ambient shaking
295 data. *Earthq. Spectra*, 30(2), 705–720.

296 Danciu, L., Kale, Ö., & Akkar, S. (2016), The 2014 Earthquake Model of the
297 Middle East: ground motion model and uncertainties. *Bull. Earthquake Eng.*, 1-37.

298 Giardini, D., Danciu, L., Erdik, M., Şeşetyan, K., Tümsa, M. B. D., Akkar, S., ... &
299 Zare, M. (2018), Seismic hazard map of the Middle East. *Bull. Earthquake Eng.*, 1-
300 4.

301 Denolle, M., E. Dunham, G. Prieto, and G. Beroza (2013). Ground motion
302 prediction of realistic earthquake sources using the ambient seismic field. *J.*
303 *Geophys. Res.: Solid Earth*, 118(5), 2102–2118.

304 Denolle, M., E. Dunham, G. Prieto, and G. Beroza (2014). Strong ground motion
305 prediction using virtual earthquakes. *Science*, 343(6169), 399–403.

306 Dziewonski, A. M., T.-A. Chou and J. H. Woodhouse (1981). Determination of
307 earthquake source parameters from waveform data for studies of global and
308 regional seismicity. *J. Geophys. Res.*, 86, 2825-2852.

309 El-Hussain, I., A. Deif, K. Al-Jabri, M. N. Toksoz, S. El-Hady, S. Al-Hashmi, K.
310 Al-Toubi, Y. Al-Shijbi, M. Al-Saifi, S. Kuleli (2012). Probabilistic seismic hazard
311 maps for the sultanate of Oman. *Nat. hazards*, 64(1), 173-210.

312 El-Hussain, A. Deif, K. Al-Jabri, A. M. E. Mohamed, G. Al-Rawas, M. N. Toksöz,
313 N. Sundararajan, S. El-Hady, S. Al-Hashmi, K. Al-Toubi, M. Al-Saifi, Z. Al-
314 Habsi, (2013). Seismic microzonation for Muscat region. Sultanate of Oman. *Nat.*
315 *hazards*, 69(3), 1919-1950.

316 El-Hussain, I., A. Deif, A. M. E. Mohamed, K. Al-Jabri, B. Gareth, H. Yaqoup
317 (2015). Deterministic seismic hazard assessment close to a gas field in northern
318 Oman. *Arab J Geosci*, 8(7), 4299-4316.

319 Ekström, G., M. Nettles, and A. M. Dziewonski. The global CMT project 2004-
320 2010: Centroid-moment tensors for 13,017 earthquakes. *Phys. Earth Planet. Inter.*,
321 200-201, 1-9, 2012.

322 Gök, R., N. Türkelli, E. Sandvol, D. Seber, and M. Barazangi (2000). Regional
323 wave propagation in Turkey and surrounding regions. *Geophys. Res. Lett.*, 27(3),
324 429-432.

325 Gu, C., F. Al-Jeri, A. Al-Enezi, Oral Büyüköztürk, M.N. Toksöz (2017). Source
326 mechanism study of local earthquakes in Kuwait. *Seismol. Res. Lett.*, 88(6), 1465-
327 1471.

328 Gupta, A. K. (1992). Response spectrum method in seismic analysis and design of
329 structures (Vol. 4). CRC press.

330 Herring, T. A., C. Gu, M. N. Toköz, O. Büyüköztürk, J. Parol, A. Enezi, F. Al-Jeri,
331 J. Al-Qazweeni, H. Kamal (2018). GPS Measurements of Large Oscillations of a
332 Tall Building Due to a Magnitude 7.3 Earthquake. submitted.

333 Jackson, J., and D. McKenzie (1984). Active tectonics of the Alpine—Himalayan
334 belt between western Turkey and Pakistan, *Geophys. J. Int.*, 77(1), 185–264.

335 Kohler, M. D., P. M. Davis, and E. Safak (2005). Earthquake and ambient
336 vibration monitoring of the steel-frame UCLA Factor Building, *Earthquake*
337 *Spectra*, 21(3), 715–736.

338 Kottmeier, C., A. Agnon, D. Al-Halbouni, P. Alpert, U. Corsmeier, T. Dahm, ... &
339 N. Kalthoff (2016). New perspectives on interdisciplinary earth science at the Dead
340 Sea: The DESERVE project. *Science of the Total Environment*, 544, 1045-1058.

341 Maeda T., and Furumura T. (2013). FDM simulation of seismic waves, ocean
342 acoustic waves, and tsunamis based on tsunami-coupled equations of motion[J].
343 *Pure Appl. Geophys.*, 170(1-2): 109-127.

344 Li, J., H. S. Kuleli, H. Zhang, and M. N. Toksöz (2011a). Focal mechanism
345 determination of induced microearthquakes in an oil field using full waveforms
346 from shallow and deep seismic networks, *Geophysics*, 76, no. 6, WC87–WC101.

347 Li, J., H. Zhang, H. S. Kuleli, and M. N. Toksöz (2011b). Focal mechanism
348 determination using high-frequency waveform matching and its application to
349 small magnitude induced earthquakes, *Geophys. J. Int.*, 184, no. 3, 1261–1274.

350 Ni, J., and Barazangi, M. (1986). Seismotectonics of the Zagros continental
351 collision zone and a comparison with the Himalayas. *J. Geophys. Res.: Solid Earth*,
352 91(B8), 8205-8218.

353 Olsen, K. (2000). Site amplification in the Los Angeles basin from three-
354 dimensional modeling of ground motion. *Bull. Seismol. Soc. Am.*, 90(6B), S77–
355 S94.

356 Olsen, K., S. Day, J. Minster, Y. Cui, A. Chourasia, M. Faerman, R. Moore, P.
357 Maechling, and T. Jordan (2006). Strong shaking in Los Angeles expected from
358 southern San Andreas earthquake. *Geophy. Res. Lett.*, 33(7).

359 Olsen, K., S. Day, L. Dalguer, J. Mayhew, Y. Cui, J. Zhu, V. M. Cruz-Atienza, D.
360 Roten, P. Maechling, T. H. Jordan, D. Okaya, A. Chourasia (2009). Shakeout-d:
361 Ground motion estimates using an ensemble of large earthquakes on the southern
362 San Andreas fault with spontaneous rupture propagation. *Geophy. Res. Lett.*, 36(4).

363 Olsen, K. B., R. J. Archuleta, and J. R. Matarese (1995). Three-dimensional
364 simulation of a magnitude 7.75 earthquake on the San Andreas fault. *Science*,
365 270(5242), 1628.

366 Pasyanos, M. E., H. Tkalčić, R. Gök, A. Al-Enezi, and A. J. Rodgers (2007).
367 Seismic structure of Kuwait. *Geophys. J. Int.*, 170(1), 299–312.

368 Pitarka, A., K. Irikura, T. Iwata, and H. Sekiguchi (1998), Three-dimensional
369 simulation of the near-fault ground motion for the 1995 Hyogo-Ken Nanbu
370 (Kobe), Japan, earthquake, *Bull. Seismol. Soc. Am.*, 88(2), 428–440.

371 Pitarka, A., M. Pasyanos, R. Mellors, and A. Rodgers (2012). Observation and
372 simulation of long-period ground motions in the Persian/Arabian Gulf from

373 earthquakes in Zagros thrust belt. Tech. rep., Lawrence Livermore National
374 Laboratory (LLNL), Livermore, CA.

375 Pitarka, A., A. Al-Amri, M. E. Pasyanos, A. J. Rodgers, and R. J. Mellors (2015).
376 Long period ground motion in the Arabian Gulf from earthquakes in the Zagros
377 mountains thrust belt. *Pure. and Appl. Geophy.*, 172(10), 2517–2532.

378 Prieto, G. A., and G. C. Beroza (2008). Earthquake ground motion prediction using
379 the ambient seismic field, *Geophy. Res. Lett.*, 35(14).

380 Prieto, G. A., J. F. Lawrence, A. I. Chung, and M. D. Kohler (2010). Impulse
381 response of civil structures from ambient noise analysis. *Bull. Seismol. Soc. Am.*,
382 100(5A), 2322–2328.

383 Rubinstein, J. L., and A. B. Mahani (2015). Myths and facts on wastewater
384 injection, hydraulic fracturing, enhanced oil recovery, and induced seismicity,
385 *Seismol. Res. Lett.*, 86, no. 4, 1060–1067.

386 Sadek, A. (2004). Seismic map for the State of Kuwait. *Emirates J. Eng. Res.*, 9(2),
387 53–58.

388 Sarkar, S. (2008). Reservoir monitoring using induced seismicity at a petroleum
389 field in Oman, Ph.D. thesis, Massachusetts Institute of Technology.

390 Shapiro, S. A. (2015). *Fluid-Induced Seismicity*, Cambridge University Press.

391 Shakal, A. F., M. J. Huang, and R. B. Darragh (1996). Interpretation of significant
392 ground-response and structure strong motions recorded during the 1994 Northridge
393 earthquake. *Bull. Seismol. Soc. Am.*, 86(1B), S231-S246.

394 Şeşetyan, K., Danciu, L., Demircioğlu Tümsa, M.B. et al. (2018). The
395 2014 Earthquake Model of the Middle East: overview and results. *Bull Earthquake*
396 *Eng.*, 1-34

397 Snieder, R., and E. Sáfak (2006). Extracting the building response using seismic
398 interferometry: Theory and application to the Millikan Library in Pasadena,
399 California. *Bull. Seismol. Soc. Am.*, 96(2), 586–598.

400 Sun, H., A. Mordret, G. A. Prieto, M. N. Toksöz, and O. Büyüköztürk (2017).
401 Bayesian characterization of buildings using seismic interferometry on ambient
402 vibrations. *Mechanical Systems and Signal Processing*, 85, 468-486.

403 Talebian, M., & Jackson, J. (2004). A reappraisal of earthquake focal mechanisms
404 and active shortening in the Zagros mountains of Iran. *Geophy. J. Int.*, 156(3), 506-
405 526.

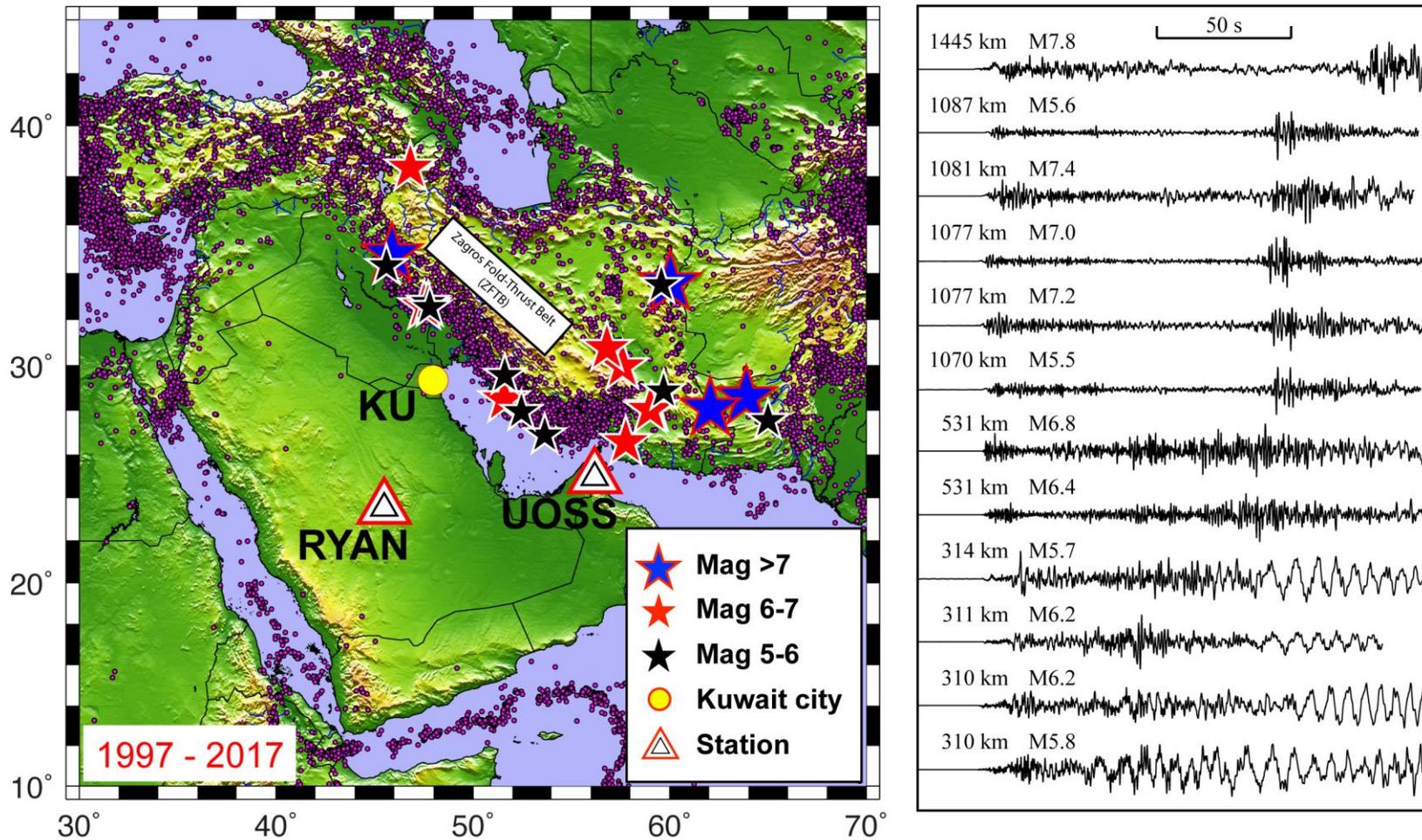
406 Zoback, M. D., and S. M. Gorelick (2012). Earthquake triggering and large-scale
407 geologic storage of carbon dioxide, *Proc. Natl. Acad. Sci.*, 109, no. 26, 10164–
408 10168.

409 **Table 1. Preferred 1-D elastic model (KUW-P) for ground motion calculation^a**

Layer number	Thickness (km)	V _P (km/s)	V _S (km/s)	Density (g/cm ²)	Q _P (s ⁻¹)	Q _S (s ⁻¹)
1	2.0	2.25	1.40	2.50	40	20
2	2.0	4.76	2.75	2.60	100	50
3	17.0	5.89	3.40	2.70	300	150
4	9.0	6.41	3.70	2.70	300	150
5	11.0	6.95	3.90	2.70	300	150
6	∞	7.80	4.40	2.70	300	150

410

411 ^aTo better fit the observed seismograms, we varied the thickness, P-wave velocity
 412 (V_P), and S-wave velocity (V_S) of the first layer of the KUW1 model of Pasyanos
 413 et al. (2007).

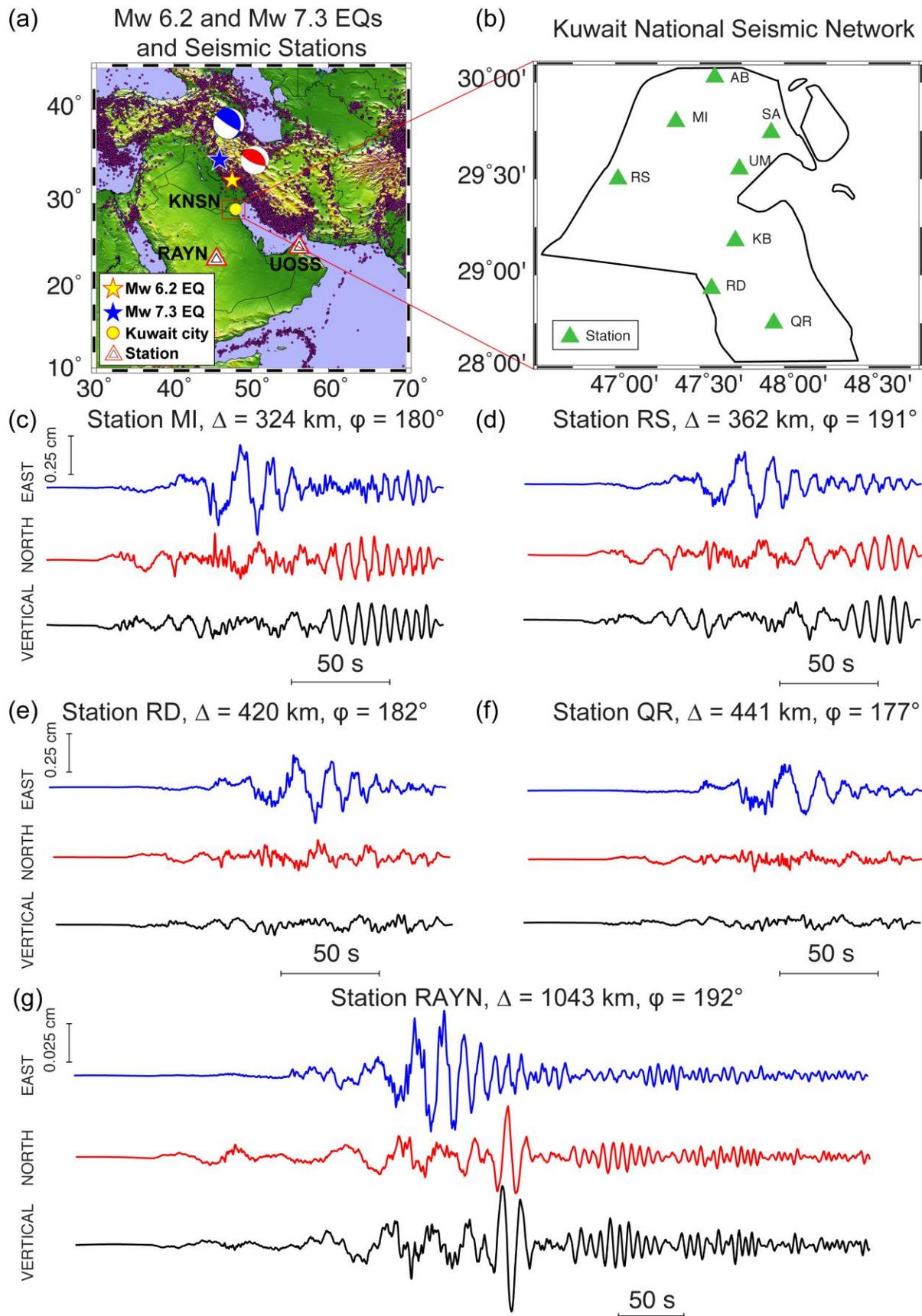


414

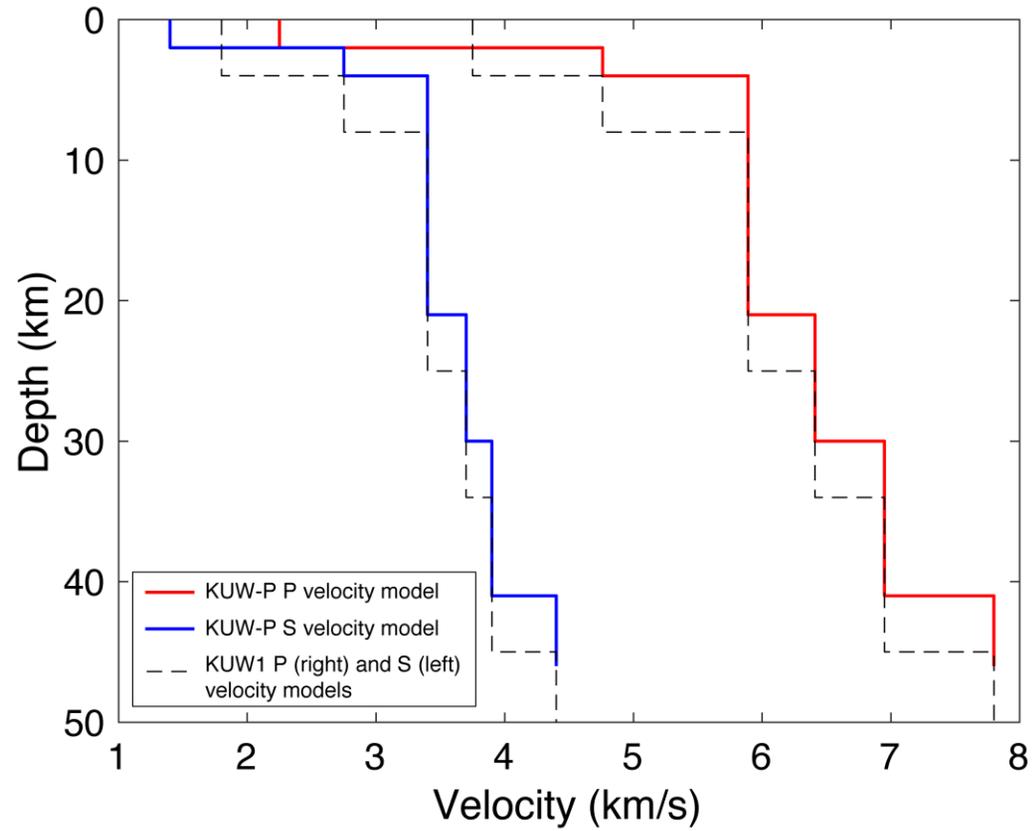
415 **Figure 1.** Significant regional earthquakes since 1997. Seismograms are recorded by the broadband station MI in

416 Kuwait. Duration of recording is limited to 200 seconds for each event. For events farther than 300 km, surface

417 waves are not recorded because the recordings are limited to 200 seconds and seismograms are truncated.

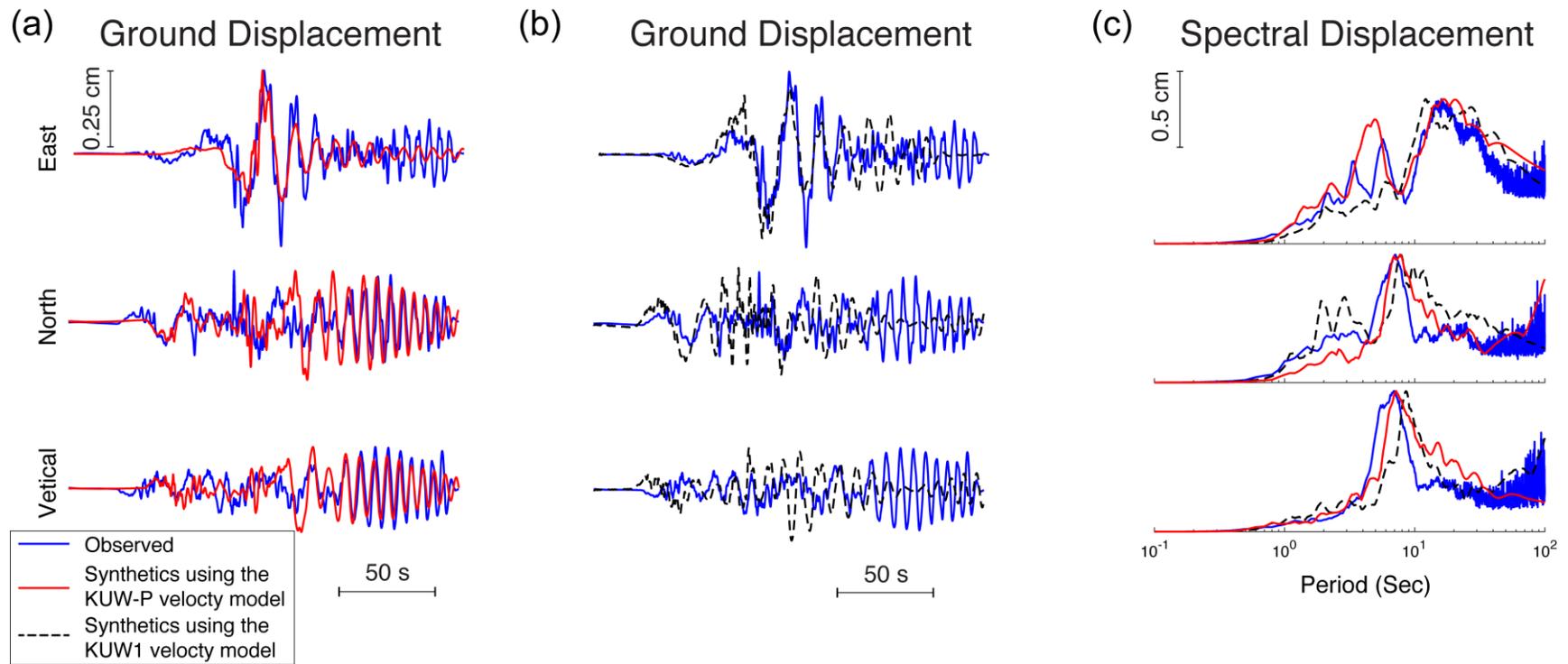


419 **Figure 2.** (a) The 2014/08/18 Mw 6.2 and 2017/11/12 Mw 7.3 earthquakes and
420 seismic stations in a geographic map. The yellow and blue stars show the
421 hypocenters of the Mw 6.2 and Mw 7.3 earthquakes, the yellow circle shows
422 Kuwait City, the red rectangle shows the region covered by Kuwait National
423 Seismic Network (KNSN), the white triangles show station RAYN in Saudi Arabia
424 and station UOSS in UAE and the red and blue beach balls show the source
425 mechanism of the Mw 6.2 and Mw 7.3 earthquake. This earthquake was widely
426 felt in Kuwait. (b) The eight stations of KNSN throughout Kuwait. The five
427 broadband Stations MI, QR, RD, RS and UM were operational when the Mw 6.2
428 earthquake occurred. (c) The observed seismograms of Mw 6.2 earthquake at MI
429 in Kuwait. Since the epicenter is north of the station, the Z and N component
430 seismograms are alike, mostly Rayleigh waves. The E component is primarily
431 Love waves. (d)-(f) The three-component seismograms of Mw 6.2 earthquake at
432 station RS, RD and QR. Because all the stations of KNSN are to the south of the
433 hypocenter of Mw 6.2 earthquake, all the waveforms are very similar showing very
434 consistent Love and Rayleigh waves. The Rayleigh waves of RD and QR are
435 truncated because of the 200 sec recording limits (g) The three-component
436 seismograms of stations RAYN. Because RAYN is almost to the south of the
437 hypocenter of Mw 6.2 earthquake, the seismograms show similar Love and
438 Rayleigh waves as stations of KNSN.



439

440 **Figure 3.** The comparison of the K UW1 and the preferred velocity (K UW-P) models. The red and blue lines show
 441 the P and S velocities of the K UW-P model. The black dashed line denotes the K UW1 P (right) and S (left)
 442 velocity models.

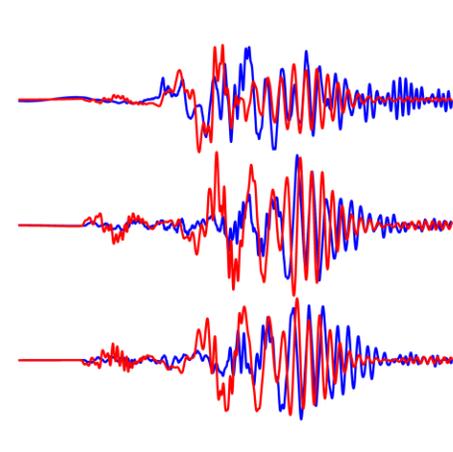
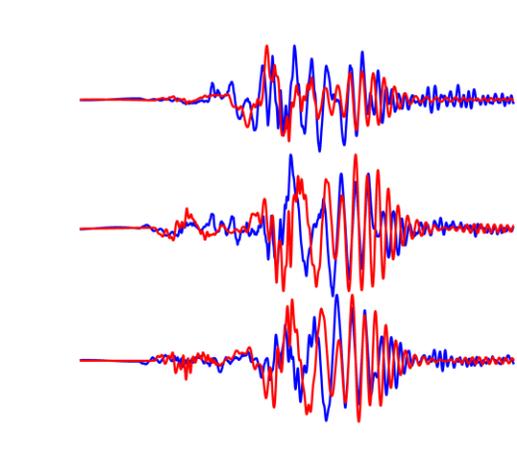
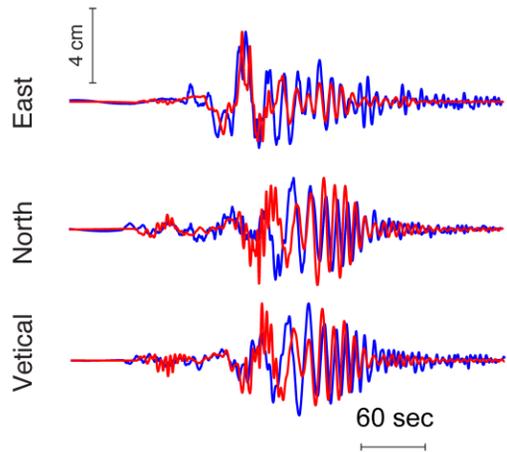


443

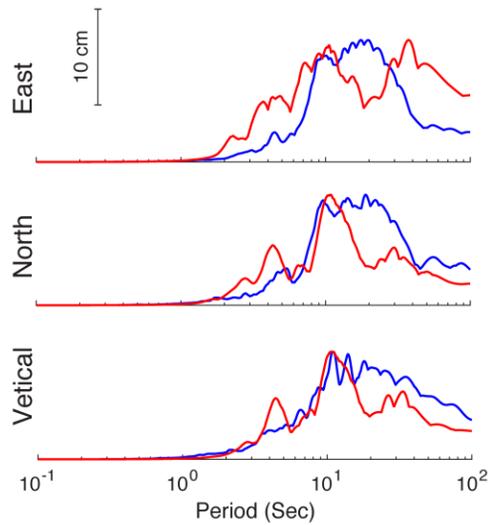
444 **Figure 4.** The observed and modeled ground displacement and spectral displacement of the 08/18/2014 Mw 6.2
 445 earthquake at MI. (a) The blue lines show the observed seismograms and the red lines show the synthetic
 446 seismograms using the K UW-P velocity model, (b) The blue lines show the observed seismograms and the dashed
 447 black lines show the synthetic seismograms using the K UW1 velocity model. (c) The blue lines show the observed

448 spectral displacement, the red lines show synthetic spectral displacement using the preferred velocity model, and
449 the dashed black lines show the synthetic seismograms using the KUW1 velocity model.

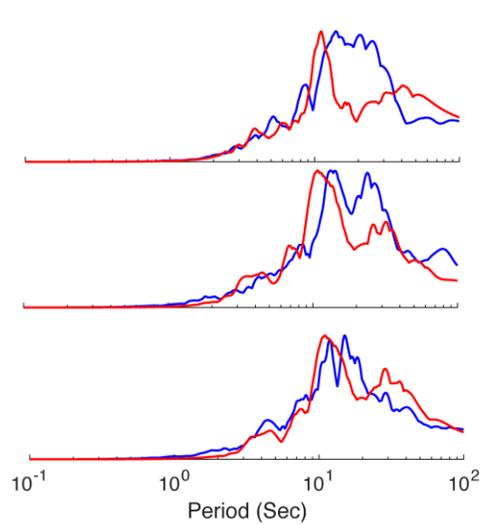
(a) Station RS, $\Delta = 597$ km, $\phi = 170^\circ$ (b) Station KB, $\Delta = 647$ km, $\phi = 164^\circ$ (c) Station QR, $\Delta = 694$ km, $\phi = 165^\circ$



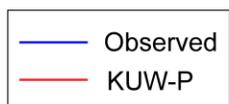
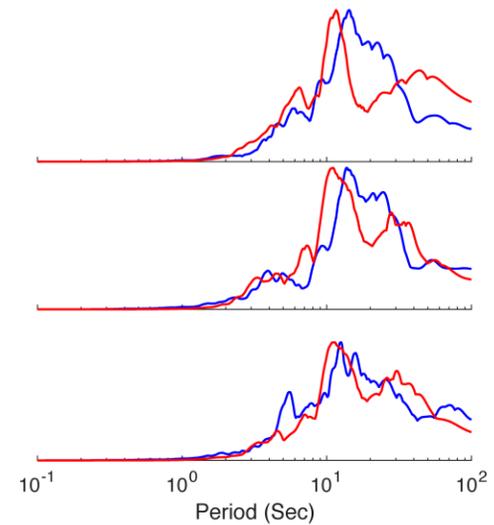
(d) Spectral Displacement



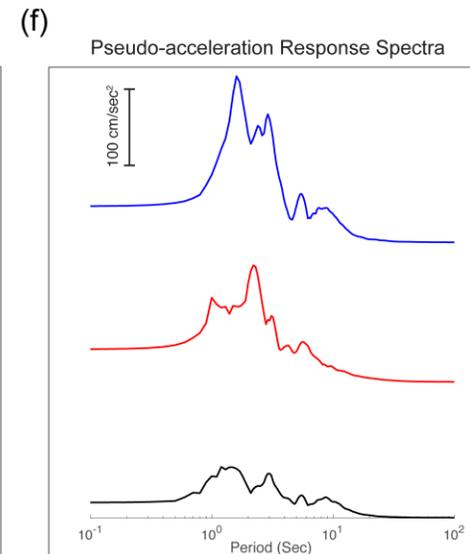
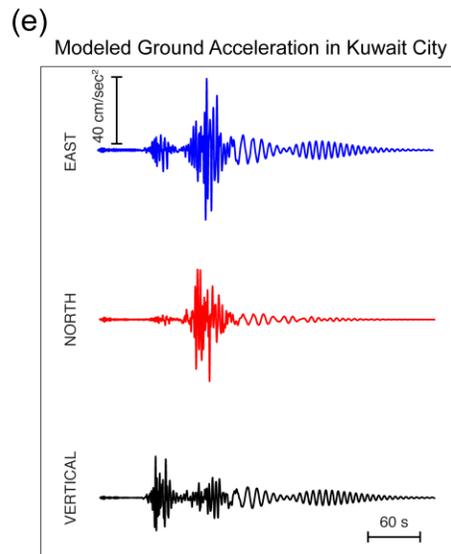
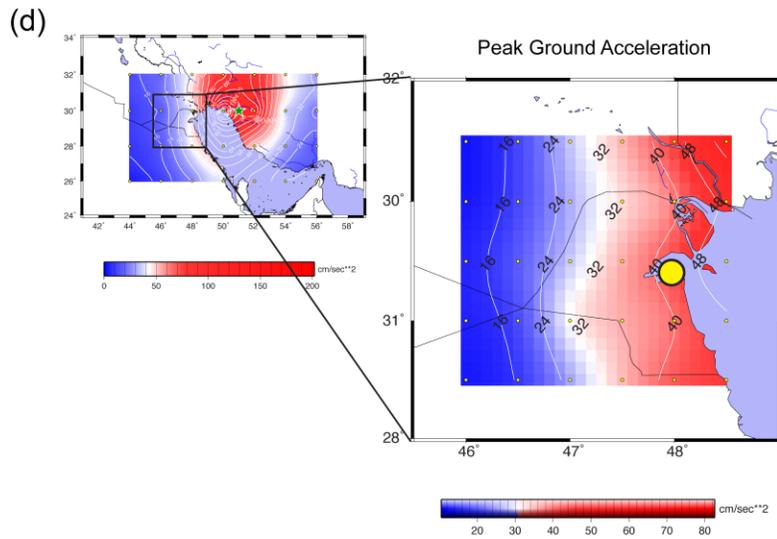
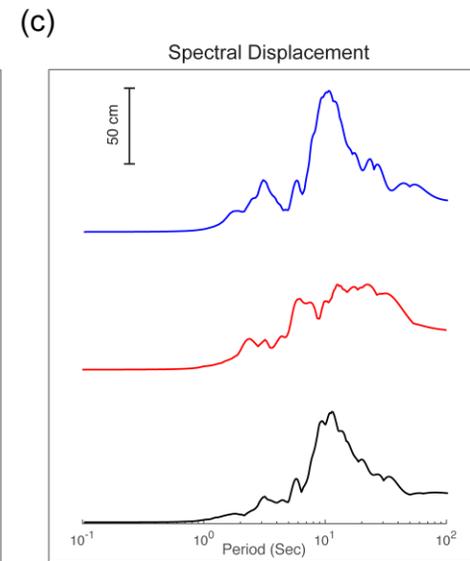
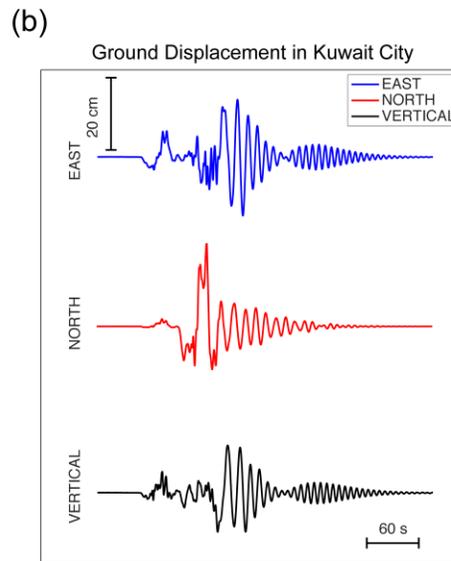
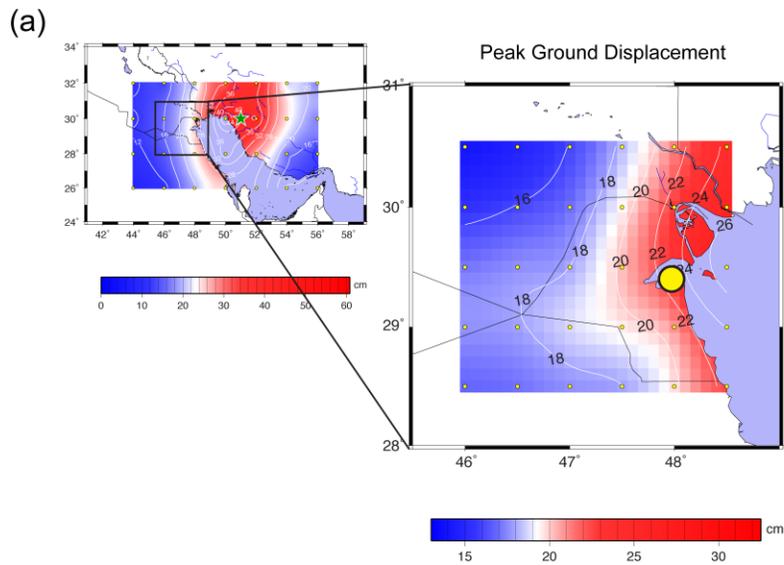
(e) Spectral Displacement



(f) Spectral Displacement



451 **Figure 5.** The comparison of the observed (blue) and modeled (red) ground displacement and spectral
452 displacement of the 11/12/2017 Mw 7.3 earthquake at stations KB, QR and RS using the KUW-P. (a)-(c) 3-
453 Component displacement seismograms. (d)-(f) 3-Component spectral displacements.

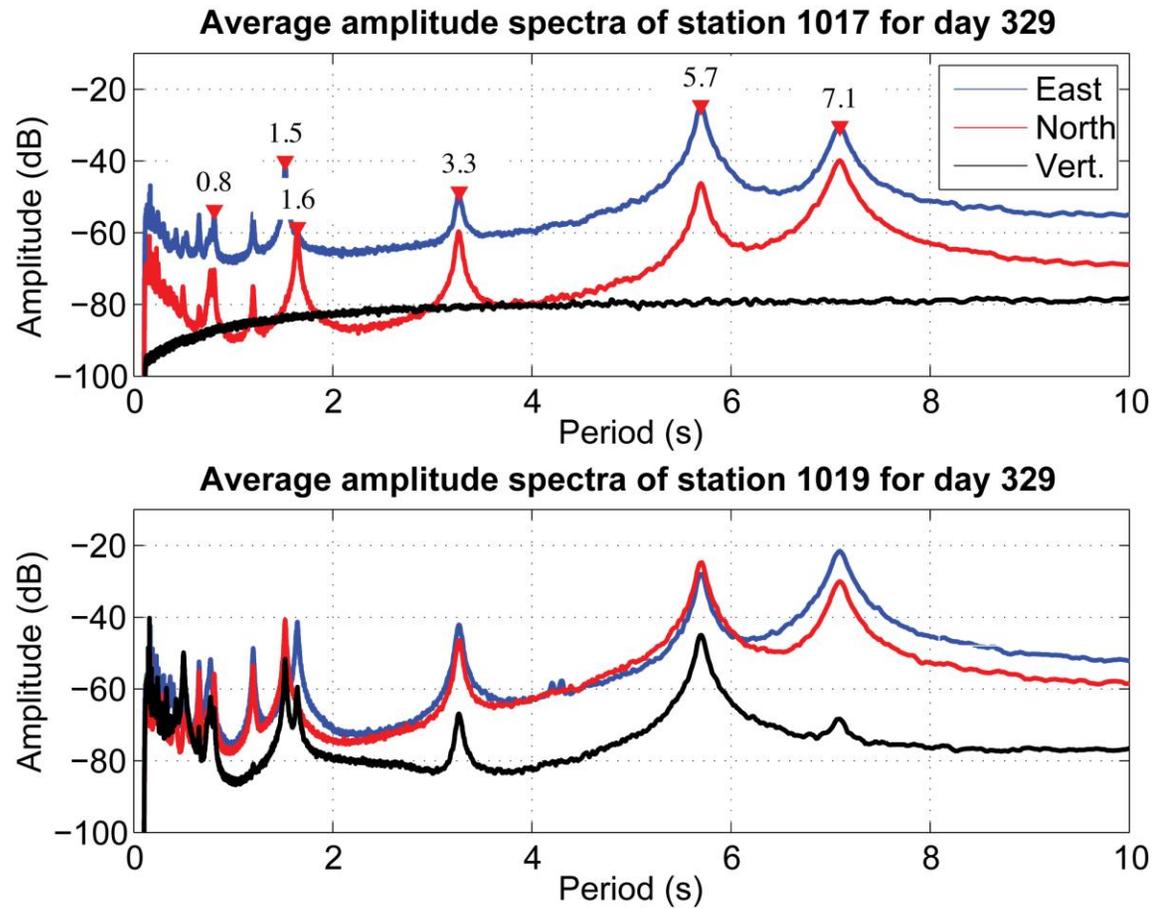


455 **Figure 6.** Modeled ground motions in Kuwait City (yellow circle) due to magnitude $M_w=7.5$ thrust earthquake in
456 Zagros with strike = 300° , Dip= 20° , and Rake= 90° . (a) The contour of peak ground displacement (PGD; cm) in
457 the north-south direction around the epicenter (green star); (b) the 3-component displacement (cm) at Kuwait City;
458 (c) the spectral displacement (cm); (d) the contour of peak ground accelerations (PGA; cm/s^2) in the north-south
459 direction around the epicenter (green star); (e) the 3-component acceleration (cm/s^2) at Kuwait City; (f) the
460 pseudo-acceleration response spectra (cm/s^2).

Al-Hamra Tower



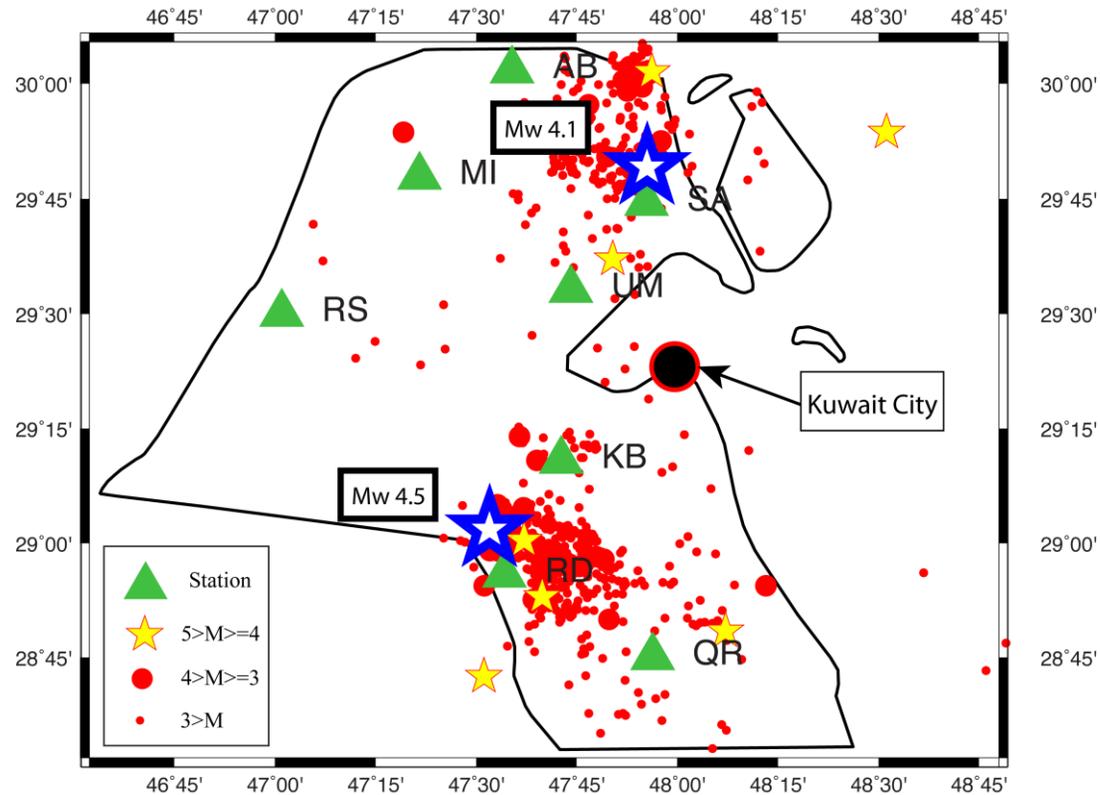
Height = 414 m



461

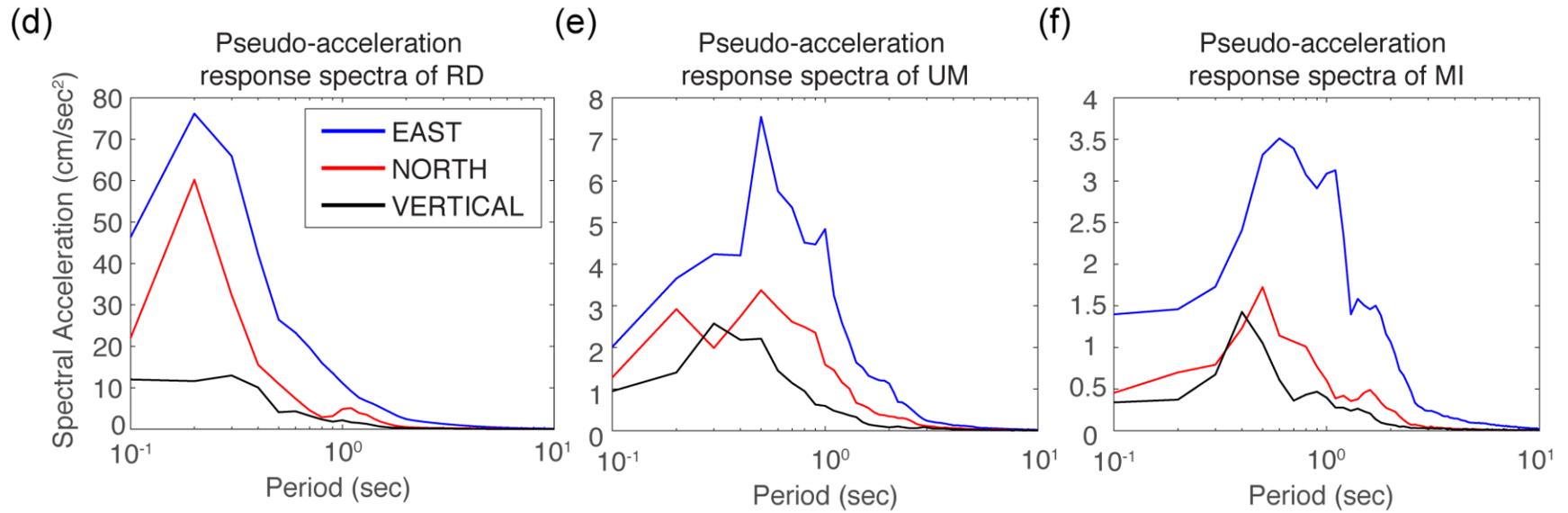
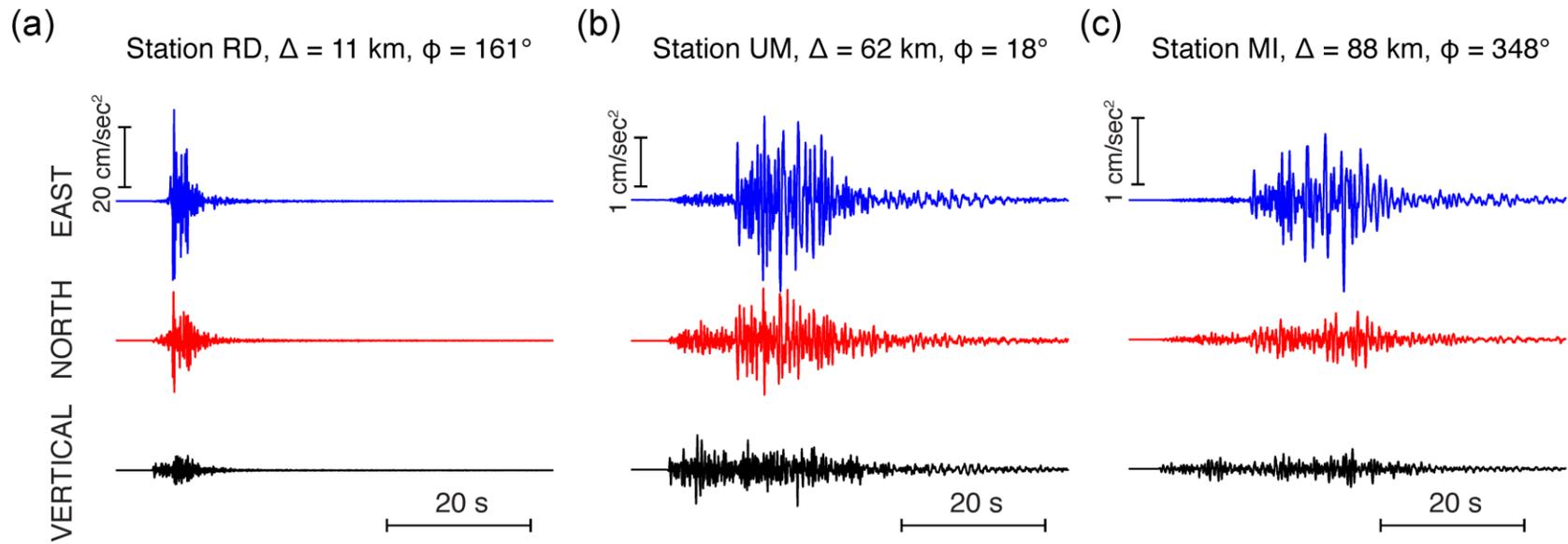
462 **Figure 7.** Amplitude spectra, calculated using ambient vibrations data recorded by two Kinometrics EpiSensor
463 instruments, at the top of the Al-Hamra Tower. Average spectra for the day 329 were computed with a 600 s
464 moving window. The vertical component of station 1017 was not working.

Kuwait Seismicity 1997-2015

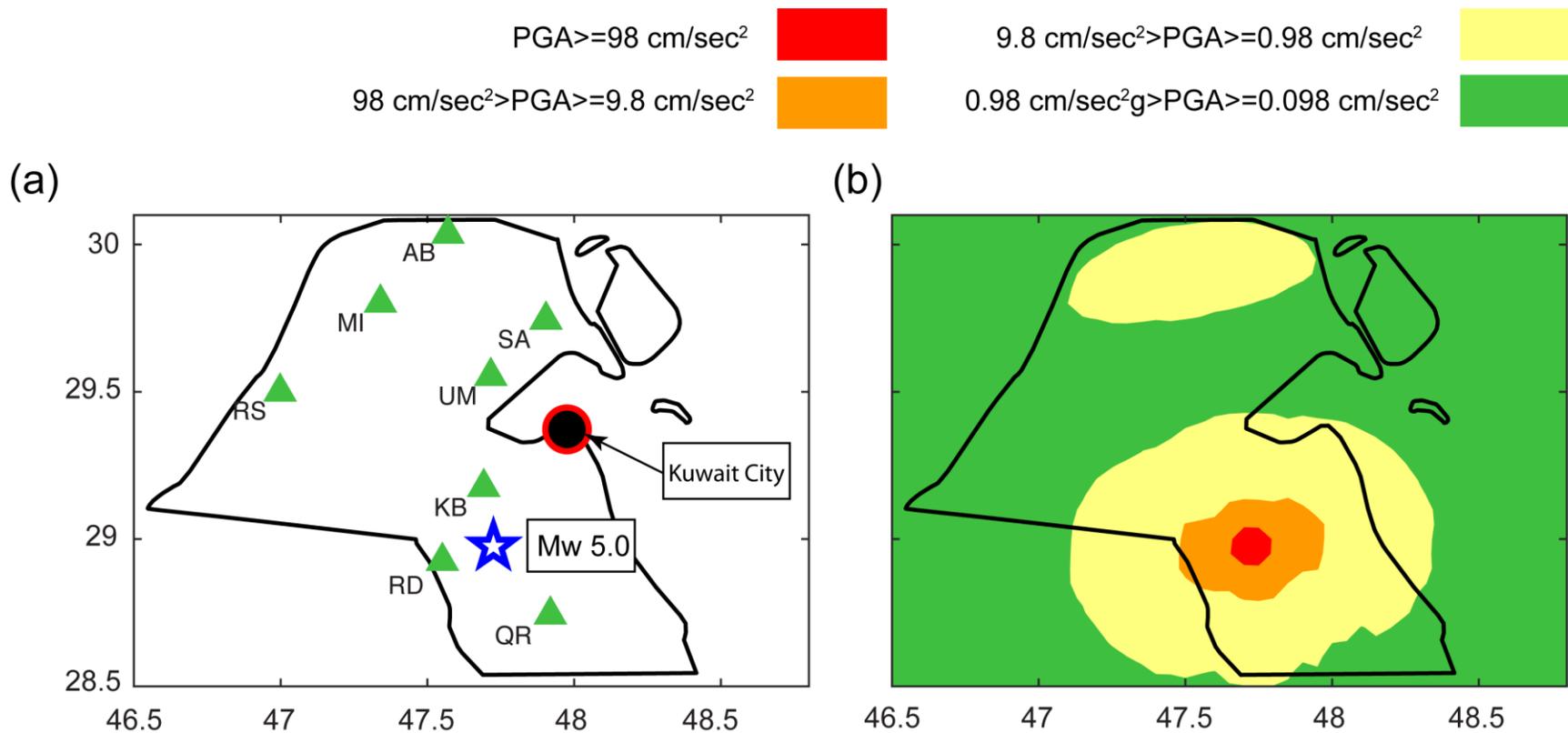


465

466 **Figure 8.** Local earthquakes (left) during 1997 to 2015 in Kuwait. We show earthquakes with $M \geq 4$ as yellow
467 stars, $4 > M \geq 3$ as large red circles, and $M < 3$ in small red circles. The 03/21/2015 Mw 4.5 and 08/18/2015 Mw
468 4.1 earthquakes are marked as blue stars.



470 **Figure 9.** (a)-(c) 3-Component acceleration seismograms for the Mw 4.5 local event (from Kuwait National
471 Seismic Network stations. (d)-(f) Pseudo-acceleration response spectra of 3-Component seismograms for the Mw
472 4.5 local event (from Kuwait National Seismic Network stations).



473

474 **Figure 10.** (a) The location of the modeled Mw 5.0 local earthquake (blue star). (b) The contour of peak ground
 475 acceleration (PGA; cm/sec²) around the epicenter (blue star). We generate the synthetic acceleration seismograms
 476 in the map region for this Mw 5.0 earthquake, a thrust event with fault plane: strike = 85°, dip = 20°, and rake =
 477 10°. The PGAs are calculated from the acceleration seismograms.