

Phosphate-Mineral Interactions and Potential Consequences for Nutrient Cycling

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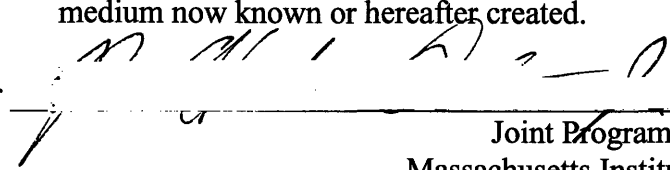
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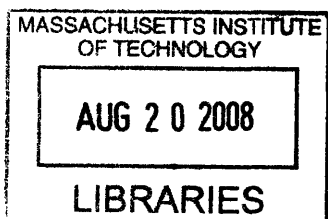
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PHOSPHATE-MINERAL INTERACTIONS AND POTENTIAL CONSEQUENCES FOR NUTRIENT CYCLING

By

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ABSTRACT:

Biogeochemical cycling of phosphate is a key component in the overall production rate of coastal ecosystems. Mineral phases in the near-shore sediments play a significant role in the return of phosphate remineralized in the upper sediments to the water column. Sequential Extraction (SEDEX) of the solid-phase associated PO_4^{3-} yielded reservoir profiles of phosphate at three sites off of the Massachusetts coast. These extractions found Fe-associated PO_4 to be the dominant phase associated with rapid porewater-solid P exchange. Additionally, a seasonal enrichment/depletion pattern of phosphate fluxes relative to total carbon was observed from the sediments. These observations established the behavior of phosphate in coastal sediments as interconnected with the ongoing Fe-cycling in the sediments as well.

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Introduction:

Phosphorus Cycling:

Phosphorus is one of the essential elements required for life on this planet and is a primary component of numerous biologically-necessary materials. Thus, phosphorus (elemental symbol = P) and its availability are a principal constraint on productivity, both on local and global scales. The global phosphorus cycle consists of three primary parts: the uplift and subsequent weathering of continental P-bearing minerals, the transport of released P to the oceans via rivers, and the utilization and subsequent burial of P in marine sediments until tectonic processes move the P-bearing sediments back to the beginning of the cycle. (Ruttenberg 2003) On shorter timescales, regional cycling between bioavailable P and solid-bound-P may determine a region's level of potential productivity. Phosphorus occurs in nature in the (+5) oxidation state, generally in the form of the phosphate ion (PO_4^{-3}). This behaves as a triprotic acid, with a majority speciation of HPO_4^{-2} and a minority speciation of H_2PO_4^- at average ocean water pH (Morel and Hering 1993). The accepted Redfield Ratio of C:N:P of 106:16:1 in oceanic organic carbon (C_{org}) suggests that P should act as a limiting nutrient in many systems. Unlike N, which can be fixed from dissolved N_2 by microorganisms, the sources of P in a system are either terrestrial input or a return flux from diagenesis of C_{org} . Therefore, it is crucial to understand the local cycling of P in order to establish a complete picture of the coastal productivity regime.

In the open ocean, most remineralization occurs in the water column, with only ~1% of C_{org} reaching the sediments. (Druffel, Williams et al. 1992) This is not the case in near-shore shallow environments where a significant portion of the organic matter

produced in the water column reaches the sediments. (Schuffert, Kastner et al. 1998; Slomp, Malschaert et al. 1998) Here, organisms use a variety of methods to oxidize this C_{org} , remineralizing the nutrients in the process. These nutrients then are able to return to the overlying waters via diffusive fluxes. This process is not 100% efficient, and thus P is lost as increasing burial removes it further and further from the ecosystem above. (Ingall and Van Cappellen 1990) Additionally, as PO_4 passes through the pore waters, chemical reactions causing the PO_4 to become associated with the solid-phase portion of the sediments and removed from the porewaters. (Krom and Berner 1980; Lehtoranta and Pitkänen 2003)

A major focus of investigation has been the effect that Fe-bearing minerals and their cycling have on the return of PO_4 to the water column. In the sediment column, Fe^{3+} acts as an electron receptor after exhaustion of oxygen, nitrate and manganese. Respiration reduces the iron to Fe^{2+} , which is soluble in water. Fe^{2+} builds up in the porewaters, generating a concentration gradient and producing an upward flux of dissolved iron. When this flux encounters the oxic/anoxic boundary layer, the Fe^{2+} reacts with the dissolved O_2 and oxidizes back to Fe^{3+} . This form of iron is highly insoluble and precipitates out, generally taking the form of a ferric oxy-hydroxide ($Fe-OOH$) or other oxide species. The cycle is completed by subsequent overlaying and movement of the reduction zone upwards in the sediment column, or by downward mixing through bioturbation back down into regions where Fe reduction is underway. (Burdige 2006) These Fe-bearing minerals have a very high affinity for PO_4 , encouraging adsorption of porewater PO_4 and forming a pool of Fe-associated or “bound” P in the sediments. (Anschutz, Zhong et al. 1998; Rozan, Taillefert et al. 2002) This suggests a possible

coupling between the Fe and P cycles in the sediments. How do these Fe oxides and other solid phases in the sediments ultimately affect PO₄ return or burial? To answer this, we need a separation of the solid-phase PO₄ into its component phases.

A method to measure the component fractions of PO₄ was developed in the early 1990s by Kathleen Ruttenberg, who realized the difficulties involved with physically separating and identifying the minerals from sediment samples. This method employs techniques developed for soil analysis and lake sediment to assemble a series of chemical extractions that can provide the desired separation of the solid-phase PO₄. Five phases are identified in this method: a “loosely sorbed” phase, an Fe-bound phase, a Ca-bound phase, a “detrital/refractory” phase, and an organic carbon (C_{org})-bound phase (P_{org}). (Ruttenberg 1992) This procedure is referred to as the **Sediment Extraction Method (SEDEX)**. We chose this method as the best means of identifying the solid phases of PO₄ that are commonly regarded as potential PO₄-sinks: the Fe-bound, Ca-bound and P_{org} phases.

Sites:

Three sites in Buzzards Bay, Hingham Bay and Massachusetts Bay were selected for study, based on their different seasonal variations in C_{org} and Fe cycling. Additional data for each of the sites includes porewater profiles of tCO₂, PO₄ and Fe²⁺. H₂S profiles were also available for Hingham Bay. The cores were collected between 2002 and 2004.

Buzzards Bay: This site, collected at a water depth of ~10 m, possesses an intermediate rate of organic matter oxidation (400 umol C/cm² yr). It exhibits a wide seasonal variability in the both the Oxygen Penetration Depth (OPD) and in

bioturbation/sediment irrigation rates. (Morford et al., submitted) No measurable amount of sulfide was found in the porewaters.

Hingham Bay: This site is located off the east coast of Paddock's Island in greater Boston Harbor at a depth of ~5 m. This site possesses a very high rate of organic deposition and respiration (~850 $\mu\text{mol C}/\text{cm}^2 \text{ yr}$). The porewaters contain sulfide at a relatively shallow depth, around 6 - 8 cm. (Morford, Martin et al. 2007)

Massachusetts Bay (Mass Bay): This site is located approximately 10 km from shore at water depth of ~30 m. C_{org} deposition and respiration at this site is still rapid (~600 $\mu\text{mol C}/\text{cm}^2 \text{ yr}$) although slower than at the Hingham Bay site. Sulfide was not observed in the porewaters until ~30 cm depth, which is beyond the range of this study. (Kalnejais 2005; Kalnejais, Martin et al. 2007; Sayles 2007)

Through analysis of these sites at two separate times of the year, I have been able to construct a picture of P and its numerous reservoirs in the fine-grain sediments of coastal Massachusetts. I combine my results with supplemental data on porewater concentrations and additional solid phase data to produce a more detailed illustration of the solid-phase/porewater coupled system.

Methods:

Sample Gathering and Preparation

Samples were collected from each location using either a ship-based coring device or hand collection by diver. Core sectioning and porewater extraction were performed under a nitrogen atmosphere. Solid samples were freeze-dried and ground to homogeneity.

Extraction:

8 depth points from the top 10 to 20 cm in each core were selected for extraction and analysis. The samples were chosen to provide greater resolution for the upper portions of the cores and still produce data for the lower portions which were not expected to exhibit many of the processes of interest. 0.50 g of each sediment sample was placed into 50 ml polypropylene Falcon centrifuge tubes. The samples then underwent a modified version of the sequential extraction process developed and described in *Ruttenberg*. (Ruttenberg 1992) Each step of the extraction involved addition of reagent, agitation for a set time on a shaker table, centrifugation for 10 min to separate the aqueous fraction, and removal of the aqueous layer via a syringe and Luer-lock sipper. Each aqueous sample was filtered through a 0.45 micron Millipore syringe filter and, with one exception, acidified with concentrated HCl (432 ul per 50 ml of extractant) and stored refrigerated.

Step I of the sequence is designed to extract loosely-sorbed phosphate. This involves two extractions with 1 M $MgCl_2$ (pH adjusted to 8) and one extraction with distilled, deionized water (DW), each for two hours.

Step II is designed to extract phosphate associated with iron oxyhydroxides and other oxidized Fe-compounds. This step involves one extraction with 0.3 M sodium citrate & 1.0 M sodium bicarbonate solution (buffered to pH = 7.6). Sodium dithionite is added in a ratio of 1.125 g per 45 ml of solution, and each sample only uses 45 ml per tube. This extraction is eight hours in duration, with small pauses to release built-up sulfur dioxide. Afterwards, the extractants are not acidified. The step concludes with one extraction in the 1M MgCl₂ solution and one extraction in DW, each for two hours.

Step III is designed to release carbonate-related phosphate. This reservoir includes chloro-fluoro-apatite, biogenic hydroxyapatite, and phosphate associated with CaCO₃-bearing minerals. This step involves one extraction with 1.0 M sodium acetate buffered with acetic acid to pH = 4 for six hours. The step concludes with two extractions with the MgCl₂ solution and one extraction in DW, each for two hours.

Step IV is designed to extract “detrital” phosphate and other inorganic forms of PO₄. “Detrital” is operationally defined as the phosphate left over after the first 3 extraction steps. This step consists of a single wash in 1 N HCl for 16 hours.

Step V involved some additional alteration of the solid sediments. The samples were removed from the centrifuge tubes and into pre-weighed Pyrex beakers using a DW squeeze bottle. These beakers were then placed into a 70 degree C drying oven overnight to remove all moisture from them. The samples were cooled and weighed before being placed inside a muffle furnace. The samples were ashed at a temperature of 550 degrees C for one hour, cooled, then reweighed to account for weight loss by C_{org} combustion. Following combustion, one 16 hour extraction in 1N HCl extracts organic bound phosphate (P_{org}).

Analysis:

All extractants, with the exception of the citrate-dithionite-bicarbonate (CDB) step, are analyzed using the single solution colorimetric molybdate-blue method described in *Methods of Seawater Analysis*. (Strickland and Parsons 1960) Before analysis, samples were diluted per recommendations in Ruttenberg. The dilutions are as follows: 1:2 dilutions for MgCl₂, Na-Acetate and H₂O supernatants, 1:5 dilution for HCl supernatants, no dilution for the CDB supernatants. Further dilutions were taken beyond these initial ones as needed to bring analyzed supernatants into concentrations covered by the calibration curve generated from a range of phosphate standards.

For the analysis of the CDB supernatant, a modified version of the stannous chloride-butanol method described in *Watanabe and Olsen* and expanded upon in *Anschutz et al* was developed and utilized. (Watanabe and Olsen 1962; Anschutz, Zhong et al. 1998) 25 ml of the supernatant CDB solution is reacted with a 0.25ml of 1M ferric chloride solution and allowed to neutralize for 1 week in refrigeration. Before analysis, the samples are aerated for 15 minutes each with ambient air to neutralize any remaining dithionite. Properly neutralized solution will take on the color of golden to dark-golden amber.

2 ml of the supernatant is placed in a 15 ml polypropylene Falcon tube. Then 4 ml of a molybdate solution (0.0405 M ammonium molybdate in 4.0 N sulfuric acid) and 3 ml of iso-butanol are added. This is then mixed by hand agitation for 2 minutes, pausing to expel gas generated from the neutralizing acid. The solution is allowed to separate for 1 minute. A syringe coupled with a Luer-lock sipper attachment is used to remove the aqueous portion of the tube, leaving the iso-butanol layer. Next, the iso-

butanol is rinsed with an addition of 3 ml of 1 N sulfuric acid, shaken once and then allowed to separate for 1 minute. Again, the aqueous layer is removed, leaving the iso-butanol layer intact. Following this, 5 ml of a stannous chloride solution (7.56 g of $\text{SnCl}_2 \cdot 2\text{H}_2\text{O}$ in 25 ml of concentrated HCl < 48 hrs. old) diluted 0.5 ml into 100ml of 1 N H_2SO_4 is added and shaken for 1 minute. After separation (~1 minute), 1 ml of the blue iso-butanol layer is removed and the rest of the liquid is discarded. The 1 ml of Isobutanol is then diluted with 3 ml of pure (200 proof) ethyl alcohol and allowed to develop in darkness for 2 hours. The absorbance is read on a spectrophotometer at a path length of 1 cm and a wavelength of 725 nm.

Each analysis was compared against a standard series of concentrations that were adjusted for both pH and concentration of supernatants post-dilution.

Results:

Solid Phases:

In this section, I will present the results of the SEDEX analysis on the selected sediment cores. I will also introduce other solid-phase data that is relevant to phosphate cycling in the sediments. The results of the SEDEX analysis on the sediment cores are presented in Figures 1 – 3 and in Tables 1 and 2. Additional solid phase data for Hingham Bay and Massachusetts Bay is presented in Figures 4 – 5.

Reproducibility: Due to the time investment that the SEDEX method requires, only a few samples were analyzed multiple times to establish the reproducibility of the results obtained in the study. Two samples from the Buzzards Bay (March 2003) site were run under the SEDEX procedure in either quadruplicate or triplicate and the resulting values are compared in Table 3. Step I is the most variable. This is likely a consequence of the poorly defined nature of the “loosely sorbed” phase, which is strictly operationally defined as any PO_4 released by either dissolution or ion exchange with the Cl^- ion in the MgCl_2 wash solution. It is likely that PO_4 from other phases is removed during this extraction, particularly from the organic-P fraction. (Ruttenberg 1992) As this amount may differ from extraction to extraction, a higher variance in this step is to be expected. This phase typically constitutes a small portion of the total solid-phase PO_4 and thus it does not distort the PO_4 -content profiles of other phases. The variability in the Step II results may be caused by relatively poor reproducibility in the PO_4 analysis of the CDB fraction. Ruttenberg notes that the Watanabe and Olsen Method (including the modified version utilized during this study) generally exhibits variance between 10% and 20%, which is in line with the results from this study (Ruttenberg 1992). The Ca-bound

fraction measured in Step III exhibits less variance than the Fe-bound fraction, but is still prone to interferences from other sediment components, particularly clay Al-bearing minerals, per Ruttenberg. A primary component of the Ca-fraction is apatite, which is also the principal component of fish bones. Detrital material homogenized in the samples can have a sizeable effect on this value. The other fractions exhibit a fairly reliable reproducibility, consistent with values presented in Ruttenberg.

SEDEX vs. Total Phosphorus:

A check of the SEDEX Method is to compare the SEDEX-derived totals with a “Total Phosphorus” measurement. After collection, samples of cores from Hingham Bay and Massachusetts Bay were completely digested and were analyzed on an Inductively-Coupled Plasma Mass Spectrometer (ICP-MS) to generate a Total P profile. (data from M. Bothner, personal communication; Briggs 1999). In Mass Bay, there is a very good agreement between the Total P and SEDEX Total profiles (Figure 6). In contrast, the profiles for Hingham Bay exhibit a significant difference. In both the January and September cores, the Total P values are ~1.75 times larger than the SEDEX Totals at the surface and ~1.3 times larger downcore. Thus, it would appear that there is a considerable addition to the pool of measurable phosphorus during the Total P analysis which may be explained by differences in the analysis methods and the reactions from the different sites. The SEDEX method utilizes chemical extraction techniques to liberate phosphate ion (PO_4^{-3}) which is then measured with a colorimetric process. The ICP-MS method uses total sediment dissolution and atom counting, which includes P that is not measured by the five SEDEX extractions. Hingham Bay and Mass Bay sediments vary mineralogically in ways which could generate the differences in the measurements.

Almost twice as much iron is present in the sediments at Hingham Bay as there is Mass Bay. Different mineral phases that contain P are not thought to be measured by the SEDEX process, including reduced-metal minerals such as vivianite ($\text{Fe}_3(\text{PO}_4)_2 \cdot 8\text{H}_2\text{O}$) and struvite ($\text{MnNH}_4\text{PO}_4 \cdot 6\text{H}_2\text{O}$). As these minerals are based around an already reduced form of metal, it is possible that a significant fraction will survive the extraction process. The CDB/Dithionite extraction utilized in the SEDEX method are primarily targeted at “easily-reduced ferric minerals” such as goethite, lepidocrocite, and hematite. (Ruttenberg 1992) Additionally, the SEDEX method is not designed to quantify Al-bound PO_4 , which would also be included in the Total P value. Finally, Fe has also been suggested to interfere with the colorimetric method of phosphate analysis, so the presence of Fe in sample solutions may act to lower the SRP measured. This is due to Fe^{2+} oxidizing to Fe^{3+} in the presence of oxygen, producing ferric oxyhydroxides (Fe-OOH) which precipitate out and sorb SRP onto surface binding sites during the analytical process. This may have occurred during storage of the CDB extractions, as they were not acidified during storage, as per Ruttenburg. Instead, the reducing potential of the dithionite and the chelating properties of the citrate were relied upon to keep the Fe in the (2+) state and soluble. It is possible that during this time, as the dithionite began to neutralize, that some Fe-precipitation may have occurred, especially due to the addition of the ferric chloride (FeCl_3) to the solution. During the analysis performed on the solution, acidic reagents are added while the dithionite is neutralized, which should prevent ferric precipitation.

Phosphate Distribution In Solid Phases

The proportions of the SRP between the extracted phases exhibit similar characteristics at all three locations. (Table 1 and 2) The dominant phase at the sediment-water interface (SWI) is the Fe-bound phase and occasionally Ca-bound. Downcore, this phase decreases and is exceeded by others. In Buzzards Bay, the organic-P and Ca-bound PO₄ are the majority phases at depth (Figure 7), while in Mass Bay, the Ca-bound PO₄ and the refractory PO₄ become the majority phases (Figure 8). In Hingham Bay (Figure 9), Ca-bound PO₄, refractory PO₄ and organic-PO₄ all exceed the Fe-bound fraction downcore in January. However, in September, the Fe-bound fraction decreases but remains the dominant solid phase. Hingham Bay possesses a very high Fe content in the sediments, reaching up to 5% by mass in the downcore regions in question. This could give the Fe-bound phase a much higher “holding capacity” for the Fe-bound PO₄ at this site than at the other two.

There is a definite seasonal variation to the SEDEX Total PO₄. At two of the sites, solid-phase-associated PO₄ increases from the winter samples to the fall samples near the SWI. The largest change of this manner occurs at Buzzards Bay (Figure 1), with a smaller one in Hingham Bay (Figure 2). At Mass Bay (Figure 3), the total PO₄ decreases in the October core, but a broader feature of constant solid-phase PO₄ below the SWI develops with a decrease below 1 cm. This is in contrast to every other core analyzed, which exhibited either an immediate increase to a peak or a decrease below the surface. The phases of refractory-PO₄, Ca-bound PO₄ and organic PO₄ show little seasonal variance, with one exception, the Ca-bound PO₄ in Hingham Bay. There is a massive increase in Ca-bound PO₄ at around 2 cm depth in January that is not seen in the

September core or at any other site. This may potentially be a result of core to core variability, reflecting perhaps a switch in type of incoming sediments, as will be investigated further in the next section.

With a few exceptions, the Ca-bound, organic-PO₄ and refractory PO₄ do not exhibit significant changes in downcore concentrations. Loosely-sorbed PO₄ and Fe-bound PO₄ demonstrate consistent features across the selected samples. It follows that the shape of the profile of these two phases, primarily that of Fe-bound PO₄, determines the structure of the Total PO₄ curve. These phases therefore have the most potential for generating a dynamic cycling reservoir.

Features of Principle Phases (Loosely-sorbed and Fe-bound):

The Fe-bound PO₄ profile peaks shift in core depth depending on the time of the year. In Buzzards Bay (Figure 1) the Fe-bound profile maximum value occurs at or immediately below the SWI during the late winter/early spring. These peaks shift downwards by less than a centimeter in the late summer/fall cores. In Hingham Bay (Figure 2) and Mass Bay (Figure 3), this trend is reversed, with a sub-surface peak observed in January and a maximum observed at the SWI during September. The profile at Mass Bay is smeared down, which may be a consequence of mixing in the upper 2 cm or of the beginning of a seasonal shift. Due to the nature of the collection methods employed during recovery of the cores, there is a high certainty of the upper sediments near the SWI being relatively intact.

The shape of the loosely-sorbed fraction profile mimics that of the Fe-bound phase. This may indicate that the loosely-sorbed fraction reflects some exchange between the porewaters and the Fe-bound phase. As Fe-bound PO₄ is cycled, the newly

liberated PO_4 can become associated with binding sites on the mineral particles, becoming briefly re-associated with the solid phase. The time scale for this exchange between the porewaters and the loosely-sorbed phase is unknown. Hyacinthe has proposed that a portion of the Fe-bound phase is highly reactive. (Hyacinthe 2004) That study utilized an ascorbate extraction (sodium citrate (50g/l), sodium bicarbonate (50g/l) and L(+) Ascorbic Acid (10g/l), pH = 7.5) for determination of the Fe-bound PO_4 , a method which has been shown to selectively target poorly-crystallized oxy-hydroxides, while leaving well-crystallized oxides, such as hematite and lepidocrocite, relatively intact (extraction of 25% or less). (C. Hyacinthe 2004) This measurement technique also would include what Ruttenberg has defined as the “loosely sorbed” phase. Considering the complications with the SEDEX method of determining the exact nature of this phase, suggestions have been made to simply merge it with the Fe-bound phase as an “active PO_4 ” phase. (C. Hyacinthe 2004) This would yield a profile of a higher concentration ($\mu\text{mol PO}_4/\text{g sediment}$) but it would not significantly change the shape of our generated profiles or the model of linked Fe-P cycling that we develop below.

Secondary Solid Phases (Ca-Bound, Refractory and Organic PO_4):

The secondary solid phases show little downcore variation with the exception of P_{org} in Mass Bay. With a few exceptions, solid-phase concentrations decrease slightly downcore. An increase in Ca-bound PO_4 in Hingham Bay in January is not believed to be a diagenetic feature (Figure 2). This feature disappears when P is normalized to Al (Figure 10) indicating that it may reflect a change in type of deposited sediments for an unknown reason. This does not appear to be a mineral precipitation curve, as there is no

corresponding feature in the pore-water PO₄ profile to suggest the precipitation of apatite (Figure 11).

Calculating the phosphate inventory of the sediments (umol PO₄/cm² sediment) (Table 4) confirms that the Ca-bound, “refractory” and P_{org} phases are the primary reservoirs downcore contributing to phosphate burial and subsequent removal from the coastal ecosystem. (Figure 13) My results differ from those of Hyacinthe, which suggested that the Fe-bound phase acted as major sink in freshwater sections of an estuarine environment. (C. Hyacinthe 2004) This does not appear to be the case at our coastal saltwater sites.

Potential complications arise from the organic PO₄ measurement. It is known that some of the organic PO₄ dissolves in the initial extraction and becomes part of the “loosely sorbed” measurement, although the total size of the “loosely sorbed” pool is small relative to the P_{org} pool. (Ruttenberg 1992) Consequently, the organic PO₄ measured in Step V reflects a minimum amount depending on the type of organic matter being deposited at the site. Studies have shown that in coastal regions, the C:P ratio in C_{org} can vary widely due to the input of continental C_{org}. C:P ratios can vary from 800 to 2050, in contrast to the 106 taken as the general oceanic value per Redfield (Ingall and Van Capellen 1990). In the nearby Long Island Sound, Ingall found a C:P ratio of 371 in the upper 10 cm of the sediment column. Our sediments seem to reflect this variability as well. (Table 5) These values were calculated by comparing the measured amounts of P_{org} from the SEDEX analysis and P_{org} calculated by multiplying measured C_{org} by (C:P) Redfield values (106:1). Ingall explains high C:P values by a combination of contributions of terrigenous C_{org} into the coastal waters and preferential remineralization

of PO_4 relative to C. Although no major rivers discharge into Buzzards Bay, terrestrial inputs are likely to be significant and both Mass Bay and Hingham Bay have historically had significant anthropogenic input of C_{org} via sewage discharge and organic pollutants in the overlying seawater. (Morford, Martin et al. 2007) These results show that the oceanic Redfield ratio of 106:1 is an overestimation of the amount of organic- PO_4 contained in C_{org} in the sediment column, in agreement with Ingall's findings.

Table 6 presents solid-phase Fe weight percentages and Fe/Al ratios (HB, MB from Martin, W., unpublished results, BB from Morford et al. (submitted)). The core-tops are generally enriched with Fe relative to Al, with the exception of Hingham Bay, where there is no apparent shift in the average Fe:Al ratio down core. Also, the core-tops contain a higher percentage of Fe by weight, except for Hingham Bay. This would make sense as Fe, reduced and released from the lower sediment column, diffuses upwards and precipitates out at the oxic/anoxic boundary layer, very near the SWI. In Hingham Bay, there is little difference between the average SWI and downcore values of both %Fe and Fe/Al. An increase in the solid phase Fe coincides with porewater H_2S rising above zero and porewater Fe^{2+} dropping below detectable levels. (Figure 4, Figure 11) The sulfide can act as a check on the Fe cycling of the sediment column by reacting with the porewater Fe^{2+} , generated by C_{org} diagenesis, and forming insoluble Fe-S compounds. Such compounds would not release reduced Fe at depth, limiting the supply of Fe returning towards the SWI.

Solid Phase Fluxes:

Utilizing sediment mixing data generated from radioisotope analysis of the study sites allows us to calculate the mixing rates of sediments generated by bioturbation of the

upper sediment column. (personal communication, Martin, W,)(Morford et al, submitted) (Morford, Martin et al. 2007) A bioturbation mixing coefficient was calculated for each site and combined with our SEDEX data to determine the apparent flux of solid-bound PO_4 as a result of this biological activity (Table 7). (Parameters and equations for flux calculations can be found in the Appendix). With one exception at Mass Bay (October 2002), the Fe-bound phase accounts for a high percentage of solid-phase PO_4 being mixed down into the sediment column. A significant amount of PO_4 is being transported down into the sediment column as an Fe-bound solid-phase component.

Summary

Fe-bound PO_4 appears to be the most active of all the solid-phase phosphate in the sediment. Fe/Al ratios indicate a strong iron cycle in both Buzzards Bay and Mass Bay, but a muted cycle in Hingham Bay. The SEDEX profiles indicate that the most active region of solid-phase cycling is located around the oxic/anoxic boundary layer in the upper 1-2 cm of the sediment column. The fraction of Fe-bound PO_4 , relative to other solid-phases, is highest at the core tops and diminishes deeper into the sediments. The largest burial of PO_4 seems to be associated with the Ca-bound phase and P_{org} .

Results:

Porewaters

Porewater Profiles:

Porewater profiles of several chemical species relevant to phosphate cycling were compiled to investigate aqueous-phase processes. These species included total dissolved carbon (tCO₂), total dissolved Fe²⁺, dissolved PO₄, dissolved oxygen and dissolved H₂S (Figures 11 - 12, 14 - 17). The sites exhibited typical sediment profile structures for many of these species. For instance, every core exhibited a subsurface PO₄ peak. All dissolved component profiles indicate production in situ and subsequent flux out of the sediments into the overlying waters, except for oxygen, which is consumed in the upper centimeter. Dissolved oxygen was measured *in situ* via electrode probe inserted into the sediments. All other profiles were determined after extraction and chemical analysis of porewaters.

Hingham Bay:

In Hingham Bay, during January, there is a gradual increase in porewater PO₄ downcore from the SWI until reaching a peak of 325 umol/l at ~7 cm (Figure 11). From this depth to 16 cm, the concentration remains roughly constant. Below this section, there is a final increase of 200 umol/l to 23 cm, the deepest depth measured. This is in contrast to the September core (Figure 12), in which porewater PO₄ increases much more rapidly, with a decrease in slope between 2 to 5 cm, and a peak at ~450 umol/l around 7 cm. The tCO₂ profile in September is almost twice as steep as the profile from January. In January, the Fe concentration increases from 0 to ~300 umol/kg at approx. 1.5 cm. In

the September core, the Fe peak is half the size of the January peak, and is located above 1 cm.

Buzzards Bay:

Total CO₂ (tCO₂) in March (Figure 114) increases downcore in the upper 2 cm and holds a steady value below, while in August (Figure 15), the tCO₂ levels increase by almost 1500 umol/kg. Fe porewater values during March increase from 0 at the SWI to a peak of about ~150 umol/kg at around 2.5 cm. This value decreases back to ~80 uM around 10 cm. The Fe values in the August core increase from 0 at the surface to a peak of ~250 umol/kg at ~1 cm with a marked decline back to a near-zero level by ~10 cm. Oxygen penetrates to 1 cm during March. In August, the oxygen penetration depth (OPD) is about 0.3 cm. (Morford et al, submitted)

Massachusetts Bay:

In the February core (Figure 16), PO₄ increases downcore with a reduced slope between 2 and 4 cm. PO₄ increases to a peak of ~125 umol/l at around 9 cm depth followed by a decrease downcore through the sampling range. In October, the porewater PO₄ increases to about 60 umol/l around 5 cm depth (Figure 17). This profile also exhibits the decrease in slope between 2-4 cm. Unlike the February profile, there is no marked decrease in PO₄ below the peak values, but rather a noisy static value till the base of the analyzed depth range. Oxygen penetration at this site was 0.6 cm in February decreasing to 0.3 cm during October (Sayles and Goudreau 2007).

Porewater Ratios

A consequence of the hypothesized PO₄ cycling described in the Introduction would be the enrichment of PO₄ values in the porewaters beyond what would normally be

predicted by C_{org} decomposition alone, as shown by calculating the porewater ratios for dissolved species relative to the established ratios produced by C_{org} remineralization. I determined that it would be most useful to examine the $tCO_2:PO_4$ ratios for this procedure, as opposed to dissolved oxygen or nitrate. Significant dissolution of carbonate minerals is unlikely to occur in this near-shore environment, and is therefore unlikely to affect our $tCO_2:PO_4$ ratios.

An examination of Hingham Bay illustrates this well. In the September 2002 site, a plot of PO_4 vs. tCO_2 is compared to the plot predicted for the Redfield values. (Figure 14) The PO_4 plot shows deviation above the expected Redfield values. This is a strong indicator that the principal factor causing this enrichment is from a cycle in the sediments that is independent of other remineralization products, such as the hypothesized Fe-P cycle. Massachusetts Bay pore waters exhibited similar enrichment (Figure 18). Buzzards Bay porewaters are not plotted because removal of PO_4 and tCO_2 at depth from proposed mineral precipitation skew the plots into showing depletion at depth.

Porewater PO_4/Fe Interactions

Comparing the porewater profiles of PO_4 and dissolved Fe^{+2} reveals a common structure across the sediment profiles. An example of this is illustrated in the combined profiles in Hingham Bay, January (Figure 11). Between 2 and 4 cm depth, there is a decline in the PO_4 gradient. At the peak of the Fe^{+2} profile, which occurs at around ~2 cm, there is a noticeable increase in the PO_4 gradient. This trend continues upcore until the measurable Fe^{+2} goes to zero at around 0.75 cm. PO_4 then returns to its previous gradient. This supports a dynamic between the precipitation of Fe in the sediments and removal of PO_4 from the porewaters, as the removal of Fe from the porewaters via

precipitation correlates with increased rates of PO_4 removal. Also, changes in the porewater PO_4 gradients correspond with changes in the dissolved Fe profile, which indicates connected removal processes.

Another typical instance of this linking behavior between dissolved Fe and PO_4 is evident in the Buzzards Bay site (Figure 14 and 15). In the profiles, one notes that both the dissolved Fe and PO_4 both go to zero at the OPD; in March, this is around 1 cm while in August it is at the SWI. This pattern of generation and removal holds even though far more dissolved Fe and P are generated in the summer, with dissolved peaks more than double the March concentrations. This is strong evidence for a process that simultaneously removes both components from the porewaters, such as coupled Fe-precipitation and P-adsorption. A similar behavior is observed at Mass Bay.

Porewater $t\text{CO}_2$

An increase in C_{org} respiration during the summer months leads to a large increase in $t\text{CO}_2$ porewater values. As the region of C_{org} oxidation in the sediment profiles does not shift significantly between seasons, summer porewater profiles show a much higher $t\text{CO}_2$ gradient than winter profiles, but overall profile shape remains the same (compare Figures 11 and 12, 14 and 15, 16 and 17). The $t\text{CO}_2$ profiles also show a sizeable step difference between the porewater values immediately below the SWI and the bottom water concentrations. In contrast with porewater PO_4 , whose values at the SWI are almost identical to the bottom water values, the $t\text{CO}_2$ concentrations are elevated relative to the overlying water. This is potentially the result of the oxidation of a highly labile portion of the incoming C_{org} immediately at the SWI.

Porewater Fluxes

With the available porewater data, diffusive fluxes were calculated using Fick's First Law for diffusive flux through a porous medium. (Berner 1980) To correct for the effects of the sedimentary environment, standard diffusion coefficients were adjusted for sediment tortuosity using porosity and formation factor data (personal communication, Martin, W.)(Morford, Martin et al. 2007). Visual identification and a trial and error approach located the regions of highest slope for the flux determinations. Further discussion of flux calculations is located in the Appendix.

Tables 8 through 11 contain porewater flux data for both PO_4 and tCO_2 , with flux rates calculated both at maximum downcore gradient (8, 9) and across the SWI (10) and $\text{tCO}_2:\text{PO}_4$ ratios of these fluxes (10, 11). The maximum porewater gradient flux represents production of the dissolved species in the sediments, while the SWI flux measures actual flux of species up into the overlying water column. Both sets of PO_4 and tCO_2 fluxes increase from winter to summer. The PO_4 flux calculated from the maximum gradient is the flux upwards into the oxic/anoxic boundary layer, while the SWI flux is the amount of PO_4 that diffuses from this layer into the water column. The shifts in magnitude between winter and summer are expected as both higher production in the water column and warmer temperatures in the bottom waters during the summer contribute to much higher levels of remineralization within the sediments. The ratio of tCO_2 to PO_4 fluxes across the SWI greatly decreases during the summer/late fall. (Table 10) At Buzzards Bay and Hingham Bay, the shift is on the scale of two orders of magnitude. This reflects the increased release of PO_4 from Fe-reducing respiration and its subsequent escape from the sediments into the overlying waters. The tCO_2 fluxes

change only modestly, while the PO_4 fluxes increase dramatically. These increased fluxes also correspond with a decrease in the thickness of the oxic layer in the sediment column. Thinner oxic layers generate less effective removal of PO_4 upwardly diffusing through the sediments, by decreasing the region where fresh Fe-OOH is likely to be present for phosphate sorption.

Benthic Flux Chamber Data

Overview:

In order to obtain additional data about the sediment's role in coastal cycling processes, benthic flux chambers were deployed at the research sites at locations similar to those from which cores were taken. There are numerous complications involved with utilizing this data. These chambers generate a micro-environment over the SWI, which alters the conditions in the upper section of the sediment column, most dramatically with regards to the dissolved O₂ levels. Also altered is the effect that bottom water circulation has on the exchange rates between the bottom waters and the sediment column. In order to overcome these complications, some of the chambers were fitted with a gas-permeable tube that circulated bottom water through in order to prevent O₂ levels from dropping to zero. This caused the chamber to equilibrate with the rate of exchange across the tubing, resulting in a lower O₂ concentration but preventing an anoxic environment from developing. This kept chamber conditions closer to ambient, and prevented "enhanced" fluxes potentially generated by anoxic conditions, but changes in the OPD were still possible from the decreased overlying O₂ concentrations. Additionally, all chambers were equipped with a paddle stirrer device to approximate the diffusive sublayer above the SWI. Table 12 contains the fluxes calculated from each site from the chamber data.

The chambers deployed in Buzzards Bay both had oxygen additions from bottom water diffusion. For the chambers deployed in Mass Bay, the February chamber was not equipped with the O₂ tubing, while the October chamber was. For Hingham Bay, data is available from two chambers, one equipped with O₂ tubing and one without. The flux values for these chambers were averaged together. PO₄ released from reducing Fe⁺³

bearing minerals increased as anoxia overtakes the SWI in the non-ventilated chambers, another link between Fe and PO₄, but this did not affect the flux calculations.

Flux values were calculated from species concentration in the chamber and its change over time. In the time sequence plots from the chambers, the majority of PO₄ flux out of the sediments appears to occur immediately, slowing considerably beyond 20-30 hours.

The same enrichment of PO₄ relative to tCO₂ during the progression from winter to summer/fall that was observed in the porewater fluxes, both in the max gradient and SWI values, is also apparent in these fluxes. This trend is clearest in Mass Bay and Hingham Bay. Buzzards Bay does not seem to experience this shift, but one possibility is that this is a reflection of increased PO₄ sequestration in the upper sediments.

Discussion:

This study suggests a definite connection between the sedimentary iron and phosphorus cycles. Figure 19 gives a cartoon visualization of this intertwined Fe – P cycling.

Support for the correlation between Fe and P cycling is illustrated by graphing the porewater PO₄ and Fe-bound PO₄ profiles on the same depth scale. (Figures 20, 21, 22) An inverse pattern between the shapes of the Fe-bound and porewater PO₄ is noticeable across the sites. This indicates that there is simultaneous removal of porewater PO₄ and generation of Fe-bound PO₄ in the upper ~2 – 3 cm of the sediment column. Changes in the gradient of the porewater profiles also reflect PO₄ sorbing/desorbing onto the Fe-bearing minerals in the region around the redox-cline. This provides a strong qualitative case for the formation of the proposed “trapping” Fe-binding layer.

The fluxes of solid-phase PO₄ mixed down into the sediments by bioturbation and the flux of porewater PO₄ heading up into the overlying bottom waters are presented in Table 13 and Figure 23. As illustrated in Table 7, Fe-bound PO₄ makes up the majority of the solid-phase PO₄ mixed down into the sediment column. This links the P and Fe cycles, by establishing the Fe³⁺ returned through downcore mixing as a carrier for PO₄ input back into the sediments. This downward mixing of Fe-bound PO₄ appears to be the source for the very large PO₄ enrichment of the porewaters (see Figure 14).

Potential difficulty in further constraining this portion of the overall cycle stems from the apparent imbalance between the diffusive PO₄ flux into the oxic layer, and the calculated transport of Fe-bound PO₄ down into the sediments via bioturbation. (Table 13) A major concern is the validity of a steady-state assumption for the system. In

steady-state, the downward flux of Fe-bound PO₄ and the upward flux of dissolved PO₄ should balance. The apparent discrepancy may be explained by the fact that the bioturbation mixing coefficients calculated for a study site are prone to a wide range of spatial and temporal variability. In addition, as the time scales of mixing are on the order of a month ($t_{1/2}$ of ²³⁴Th = 24 days), while the equilibrium time for the Fe-bound/porewater PO₄ exchange is likely far shorter, a discrepancy between the solid-phase and porewater fluxes seems likely to occur. Further temporal sampling may help resolve such apparent discrepancies between the solid and dissolved phase fluxes. Nevertheless, there is a clear link between porewater removal and Fe-uptake of PO₄ occurring at the redox-cline.

Another indication of Fe-P coupling is the simultaneous removal of both PO₄ and Fe²⁺ from the porewaters at around the OPD. (Figures 11 - 15) The shape of the profiles exhibits similar removal features, including going to a near-zero value below the SWI. This is observable in all the cores with data available, but is most easily seen in Buzzard Bay, March 2003. (Figure 14)

At every site, there are peaks in both the “Fe-bound” and “Loosely Sorbed” PO₄ profiles that mimic the shape of the Total PO₄ profile almost precisely (Figures 1 – 3). Since the loosely-sorbed phase is much smaller than the Fe-bound phase, this establishes the Fe-bound phase as the primary driver of porewater/mineral interactions in P cycling. Porewater O₂ profiles, where available, show that the maximum value of the Fe-bound phase concentration occurs at the apparent OPD. There is an accompanying enrichment of Fe in this layer in Buzzards Bay and Mass Bay, as well as a much smaller enrichment in Hingham Bay. (Figure 4, 5) Buzzards Bay and Mass Bay both exhibit higher weight

percentages of Fe in the core tops than in the downcore regions, while Hingham Bay has the highest overall Fe content, but only a small shift in weight % as observed in the other cores (Table 6).

The Fe/Al ratios (Figure 10) from the cores reveal functions of the Fe cycle that explain these features. In Mass Bay, Fe is enriched relative to Al in the upper 2 cm of the core and depleted relative to the surface Fe/Al ratio downcore, where an asymptotic value of both Fe and Fe/Al is reached. However, Hingham Bay's Fe/Al profiles show a much smaller enrichment value at the surface and a large enrichment downcore, between 6 cm during January and 9 cm during September. A reason for this is suggested by the relationship between porewater Fe and H₂S in the Hingham Bay cores (Figure 11). There is a very shallow incidence of H₂S in the porewaters, beginning at around 5 cm. This is coupled with a drawdown of the porewater Fe, which exhibits a removal feature in its profile between 2.5 cm and 5 cm, eventually going to zero concentration by 6 cm. This indicates that Fe and S are reacting in the 5-6cm range to form ferrous sulfides, which are insoluble and precipitate out. The formation of these sulfides removes Fe from the porewaters before it has a chance to diffuse upwards to the OPD. This short-circuits the Fe cycle, by removing mobilized Fe away from the regions where oxidation and precipitation, with associated PO₄ scavenging, may occur. Such processes may account for the lack of a seasonal variation in the Fe-bound P phase at Hingham Bay.

Therefore, I propose that the short-circuited Fe cycle present in Hingham Bay limits the trapping effect of the "iron curtain" at the oxic/anoxic boundary layer. One way to support this is to look at the porewater fluxes in the sediments. During the late summer/fall time periods, every core exhibits higher fluxes upward from porewater

concentration maxima of both PO_4 and tCO_2 (Tables 8 and 9). This is an obvious consequence of increased respiration occurring in situ. Barring the dissolution of carbonate minerals, the tCO_2 value can be taken as a rough estimate of C_{org} remineralization in the sediment column. Every core shows an enrichment of its PO_4 fluxes upwards into the oxic/anoxic boundary layer relative to this tCO_2 and its expected Redfield value. Indeed, as noted in the porewater section of the results, C_{org} is likely to possess even higher C:P ratios in coastal settings. A comparison of tCO_2 : PO_4 ratios of the fluxes across the sites reveals an immediate distinction between Hingham Bay and the other sites. In Buzzards Bay and Mass Bay, the ratio remains relatively constant, whereas in Hingham Bay, the ratio value decreases from 18 in January to 7 in September, a 2.5x increase in the amount of PO_4 flux relative to the tCO_2 produced. (Table 11) All of this occurs without any major shifts in the solid-phase PO_4 concentrations. Therefore, the dampened Fe cycle in Hingham Bay leads to a greater release of PO_4 back into the overlying waters compared to the other two sites, where dissolved H_2S was not found.. However, the flux across the SWI is still much lower than the flux calculated from the maximum gradient (20x in winter, 3x in summer), indicating that the Fe-associated trapping effect is still an important factor (Table 15).

Impact on Benthic Fluxes

The return of PO_4 to the overlying waters is perhaps the most crucial part of this investigation because this is what ultimately affects coastal nutrient availability and production rates. The different conditions at each site alter the expected return of P to the water column. Table 14 presents all the returning PO_4 fluxes, from both porewater SWI and benthic flux chamber (BFC) measurements (Figure 24). With the exception of the

BFC flux in Buzzards Bay, there is a very consistent trend of fluxes depleted or nearly depleted in PO_4 relative to tCO_2 in the cold months, countered with moderate enrichment of PO_4 during the summer/fall, relative to Redfield values. (Tables 11 and 12, Figure 25) This trend may be a result of increased Fe-reducing remineralization during the summer and a thinner oxic layer at the SWI. As respiration ramps up during the warmer months, increased bioturbation supplies more Fe-OOH* PO_4 downcore, which is then consumed as the electron acceptor for microbial diagenetic processes. This removal of Fe-OOH from the surface and the subsequent reduction downcore release PO_4 that had been previously bound on these Fe-bearing minerals. This enriches the porewaters in P relative to tCO_2 and drives accelerated fluxes out of the sediments, which in turn reflect this PO_4 enrichment with steep decreases in the tCO_2 : PO_4 ratio (Tables 10 - 12). The thinner oxic layer reduces the volume of unsaturated Fe-OOH-containing sediment that the upwards fluxing PO_4 must pass through before returning to the overlying waters. As the temperature and incoming C_{org} drop, the remineralization rate in the sediment column decreases. This means that less Fe^{3+} from the Fe-OOH minerals is reduced for respiration, releasing a lesser amount of Fe-bound PO_4 in the process. Also, decreased bioturbation lowers the rate at which Fe-bound PO_4 is mixed downcore. Decreased respiration moves the OPD downcore, increasing the thickness of the oxic layer, in turn increasing the volume of Fe-OOH-containing sediments that PO_4 fluxing upwards must pass through. Thus, porewater PO_4 once again becomes trapped on the Fe-bearing minerals at the oxic/anoxic boundary layer, but with a diminished return rate back downcore. This retards the return rates of PO_4 back into the overlying waters, and leads to a depletion of PO_4 relative to tCO_2 in the fluxes into the water column, as seen in

Table 11 and 12. In the next year, this cycle will start back up again, with cycling down of Fe-bound PO_4 to be released, raising porewater PO_4 concentrations and producing an enriched PO_4 flux to the overlying waters. Additional evidence for this behavior is illustrated in Figure 13, which shows that for the majority of the sites, the Fe-bound PO_4 inventory during the winter is higher than in the summer/fall, indicating increased storage of PO_4 during that time.

This picture is supported by examining the reduction in the porewater PO_4 calculated fluxes as the PO_4 passes through the oxic/anoxic boundary layer on its way out of the sediments. Table 15 shows that, generally, over 90% of the porewater PO_4 does not make it past the oxic/anoxic boundary layer during winter. In the summer/fall, this amount is reduced to between ~40 – 60%. This provides evidence for the increased trapping effect formed by the thickening of the oxic layer during winter. Note also that Hingham Bay has the highest trapping efficiency of the three sites, with 95% trapped in winter and 61% trapped during the summer. This is connected to its relatively high amount of total Fe in the sediment column.

Potential Consequences: Mineral Precipitation and P Burial

Another potential removal mechanism for pore-water phosphate is the precipitation of phosphate-type minerals. This is a proposed consequence of the elevated P ratios in the pore-waters formed by the solid-phase cycling trap. Such minerals include struvite ($\text{MnNH}_4\text{PO}_4 \cdot 6\text{H}_2\text{O}$), apatite ($\text{Ca}_{10}(\text{PO}_4)_6(\text{F-OH})_2$) and vivianite ($\text{Fe}_3(\text{PO}_4)_2 \cdot 8\text{H}_2\text{O}$). Several studies have indicated that coastal sediment settings have the potential to act as generation sites for these mineral types (Martens, Berner et al. 1978; Jahnke, Emerson et

al. 1983; Van Der Zee, Slomp et al. 2002). Apatite formation is not indicated by the Ca-bound profiles from the SEDEX analysis.

The porewater profiles of PO_4 at Buzzards Bay site exhibit loss of phosphate at depth. During March, this removal seems to be located in the region between 2 cm and 20 cm (Figure 14). A shift in this region occurs during the transition to August (Figure 15). During that time, the return to previous PO_4 levels is not observed. Instead, phosphate decrease begins at 3.8 cm and approaches a value of $\sim 15 \mu\text{M}$ around 21.9 cm. Porewater Fe^{2+} levels exhibit removal at depth at these times as well, with a very strong removal during August. These regions of removal do not correspond with known areas of FeS formation and must be the work of a different removal process.

To investigate whether the precipitation of phosphate-type minerals is occurring, we determine if the pore-waters achieve super-saturation with respect to the solid phases in question. Here we examine vivianite, the only mineral with sufficient data for solubility analysis.

Porewater concentrations from the August 2004 Buzzards Bay cores were analyzed by the algorithm described in *Martens* and laid out in the Appendix. (Martens, Berner et al. 1978) The generated solubility profiles (Figure 26) indicate that the pore-waters are supersaturated with respect to vivianite from a depth range of ~ 0.4 cm to around ~ 8.0 cm. When plotted as Ion Activity Product (IAP) vs. Depth against Concentration vs. Depth, the peak of the IAP values appears to coincide with the peak of the dissolved iron curve and with the maximum of the pore-water phosphate profile. (Figure 26) The IAP curve then passes back under the solubility curve at 8.2 cm, which is very close to the inflection point for the pore-water Fe curve, indicating that the switch

from addition to removal of Fe^{2+} occurs at a similar depth. To test the model's reaction and sensitivity to changes in pH, potentially resulting from inaccurate back calculation values from dissolved porewater concentrations, a lower value of $\text{pH} = 7.5$ was also fed into the calculations. The lower pH shifted the super-saturation region into a tighter range, but still a significant region of the core exhibited the potential for vivianite precipitation.

Vivianite formation would represent a solid-phase removal process for P in the lower reaches of the sediment column, made possible by the simultaneous occurrence of porewaters enriched in both PO_4 and Fe. The PO_4 profiles at the other research sites do not exhibit nearly as strong a PO_4 removal at depth (Figures 9, 12, 16, 17). This may be attributed to the porewater Fe removal caused by sulfide. As iron sulfide minerals precipitate, they may lower the concentration of Fe below that required for vivianite precipitation. This is true at deep points in the sediments, and may limit potential areas of mineral precipitation to areas immediately adjacent to the dissolved Fe and PO_4 maxima.

Conclusion:

The phosphorus in Massachusetts coastal sediments exhibits a dynamic exchange cycle between porewaters, solid phases, and the overlying waters. By use of the SEDEX Method, P-associated phases were separated and analyzed. Fe-bound PO_4 is the primary actively cycling solid phase in this environment, which constitutes the majority of porewater/solid-phase P exchange. Fe-bound PO_4 primarily determines the shape of the solid-phase PO_4 profile, but was not necessarily its chief component. The Fe-bound phase is shown to peak in the near-surface sediments, near the depth of removal of porewater Fe^{2+} , with a decline downcore where Fe is reduced by microbial diagenesis. While never going to a zero value, Fe-bound PO_4 is not the principle solid-phase associated with burial at these sites.

Fe-associated minerals that form along the oxic/anoxic boundary layer act as a trapping feature, enriching porewaters greatly in PO_4 with respect to the expected values from organic matter decomposition. It appears that the PO_4 flux to the bottom waters is enhanced by a weakening of the Fe cycle due to precipitation and burial of Fe as iron sulfides, such as was observed in Hingham Bay. Benthic flux observations bear this out, displaying depletion in P relative to C in fluxes in the winter, and a significant enrichment of P in the summer/fall.

The principal phases that are associated with PO_4 burial downcore in our sites are Ca-bound, "refractory" and $\text{P}_{(\text{org})}$. Thus, Fe-bound PO_4 does not appear to be a major sink for P in these sediments. Also, in contrast to other studies of similar regions, this study does not note any significant formation of authigenic Ca-type minerals (apatite, chloro-fluoro-apatite, etc.).

In coastal environments, the sedimentary cycles of both Fe and P are intertwined. In order to properly constrain the return of bioavailable P to the overlying waters, the particulars of the local Fe cycle must be known, especially the availability of sulfide in the sediments to form non-exchanging Fe(S) compounds. Exchanging Fe-bearing minerals, such as the ferric oxy-hydroxides (Fe-OOH), are capable of generating a substantial enrichment of PO₄ in the porewaters, altering both the return rate of P to the water column and potentially generating conditions conducive to phosphate mineral precipitation capable of long-term burial in the sediments, depending on the rates of remineralization ongoing in the sediment column. The variations in the Fe cycle resonate in the P cycle, generating a seasonal pattern of depletion/enrichment of P in benthic fluxes. Future studies of nutrient cycling in this type of region must be sure to take these Fe-based interactions into account, if an accurate picture of P in a redox-driven world is to be created. This clarified picture of the coastal P cycle has numerous implications for other fields of inquiry, including coastal productivity models, “red-tide” predictions and general oceanic nutrient sinks.

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Appendix:

Flux Calculations:

Fick's First Law describes fluxes of chemical species by diffusion:

$$F = -\phi D_i \frac{\partial C_i}{\partial x} \quad (\text{Fick's First Law}) \quad (1.1)$$

where F = flux across a plane of reference, ϕ = porosity of the sediments, D_i = diffusion coefficient of species (i), and the derivative = concentration gradient of species (i). (Berner 1980) This equation was adapted for porous sediments via several modifications. First, the diffusion coefficient was calculated using data compiled by Boudreau for tortuosity and other sediment effects. (Boudreau 1997) This accounts for both temperature and salinity effects on the standard $D_{(sw)}$ of the dissolved species. Then, the porosity of the sediments was taken into account by the following relationship (Morford, Martin et al. 2007):

$$D_{(sed)} = \frac{D_{(sw)}}{\phi^F} \quad (\text{where } \phi = \text{porosity, } F = \text{formation factor}) \quad (1.2)$$

where:

$$F = c\phi^{-v} \quad (\text{where } c \text{ and } v = \text{constants}) \quad (1.3)$$

Then (1.3) can be arranged into:

$$\log\left(\frac{1}{F}\right) = c + v\log(\phi) \quad (1.4)$$

Incorporating (1.4) into (1.2), one arrives at:

$$D_{(sed)} = D_{(sw)}\phi^{(v-1)} \quad (1.5)$$

where $D_{(sed)}$ is the diffusive coefficient in the sediment column. From the pore-water profiles, regions for flux calculations were selected to both reflect the hypothesized

removal mechanism in the iron precipitation regions and the area of the curve with the steepest gradient. By using the region of highest gradient (ΔPO_4 vs. Δz), a realistic value of the phosphate flux in the sediment column can be calculated. These regions were determined by visual examination of the profiles and a trial-and-error determination of the gradient location.

Porewater Flux Calculation Parameters:

The following are the depth ranges in the porewater profiles where the maximum downcore gradients were determined to be. In Buzzards Bay, the range for March 2003 was from 1.08 to 2.65 cm. For August 2004, the range was from 0.16 to 1.27 cm. Second, in Hingham Bay, the January 2002 range was from 1.23 to 1.88 cm. In September 2002, the range was from 1.68 to 2.09 cm. Finally, in Massachusetts Bay, the range for February 2002 was from 0.95 to 1.91 cm. The range for October 2002 was from 1.83 to 2.23 cm.

Tables 16 and 17 contain the sets of parameters utilized in flux calculations. All data was provided by W. Martin via personal communication.

Solid Phase Fluxes:

Table 18 contains the parameters utilized in calculating the solid phase fluxes.

The following equation was used for this calculation:

$$F_{(solid-phase)} = -\rho(1-\phi)D_{bio} \left(\frac{dC}{dz} \right) \quad (1.6)$$

where ρ = density of sediments (g/cm^3), ϕ = porosity of sediments and D_{bio} = bioturbation mixing coefficient (cm^2/yr).

If additional terms were required, the porewater parameters for the corresponding site were employed. Density of coastal sediments was taken to be $\sim 2.6 \text{ g/cm}^3$ in all locations.

Vivianite Solubility Calculations:

Following the method outlined in *Martens et al*, 1978, I calculated the saturation state of the pore waters by substituting in the K_a values adjusted for the seawater conditions and setting up a series of two equations. From the total SRP, the following equation is used to determine the PO_4^{3-} ion concentration:

$$m_{\text{PO}_4^{3-}} = \frac{\Sigma P}{1 + a_{\text{H}^+} / K_3' + a_{\text{H}^+}^2 / K_2' K_3' + a_{\text{H}^+}^3 / K_1' K_2' K_3'}$$

where K_1' , K_2' and K_3' refer to the 1st, 2nd and 3rd disassociation constants for phosphoric acid adjusted for a seawater solution and the activity of hydrogen ions in the solution (ie. the pH). The pH of the solution was back-calculated from the measured values of tCO₂, Alkalinity and other measured dissolved charged species. The values for the phosphoric constants were obtained from Kester and Pytkowicz 1968 for values indicated for the salinity and temperature averages from the Buzzards Bay site. (Kester and Pytkowicz 1967) The values utilized were 2.96×10^2 for K_1' , 0.74×10^6 for K_2' and 0.82×10^9 for K_3' . The initial values for the K_a constants were taken from *Principles and Applications of Aquatic Chemistry* and are as follows: $K_1 = 10^{-2.15}$, $K_2 = 10^{-7.2}$, $K_3 = 10^{-12.35}$. (Morel and Hering 1993) From there, I calculated the Ion Activity Product (IAP) for vivianite using the following formula:

$$IAP = \gamma_{\text{Fe}^{+3}}^3 \gamma_{\text{PO}_4^{3-}}^2 m_{\text{Fe}^{+3}}^3 m_{\text{PO}_4^{3-}}^2$$

Gamma is the Total ion activity coefficient and m is the molality of the dissolved species. Following per *Martens* 1978, I took a total ion activity (TIA) of 0.25 for iron. The total ion activity for phosphate was determined per Martens with the following formula:

$$\gamma_{T_{PO_4}} = \frac{K_1 K_2 K_3}{K_1' K_2' K_3'} \gamma_{H_3PO_4}$$

where TIA of H_3PO_4 is assumed to equal 1. After the IAP is calculated, the value is compared to the precipitation constant for vivianite, which is taken to be 10^{-36} , again, per Martens. If the IAP is found to be larger than the solubility constant, then the solution can be considered supersaturated, making precipitation of vivianite a realistic possibility.

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TABLES:

SEDEX-generated Solid Phase PO₄ Values

PO ₄ Concentration (umol/g sed)			Loosely-sorbed	Fe-Bound	Ca-Bound	Refractory	Organic P	Total PO ₄
Buzzards	Mar-03	SWI	3.08	11.86	6.38	2.97	6.94	31.22
Bay		20 cm	2.16	3.64	6.80	2.79	6.99	22.37
	Aug-04	SWI	3.90	15.39	7.81	3.18	7.84	38.12
		22 cm	1.18	3.27	7.75	3.49	5.90	21.58
Hingham	Jan-03	SWI	4.88	11.16	9.83	6.97	6.09	38.93
Bay		15.5 cm	0.70	4.69	17.41	8.19	6.74	37.73
	Sep-03	SWI	4.20	17.21	6.27	7.91	8.83	44.42
		10.2 cm	0.44	7.31	6.54	7.66	9.15	31.10
Mass.	Feb-02	SWI	3.94	12.04	10.39	8.64	6.20	41.20
Bay		21 cm	0.61	2.59	9.67	10.26	2.94	26.07
	Oct-02	SWI	4.33	8.31	11.44	8.46	5.15	37.68
		9.2 cm	0.69	1.90	9.93	10.44	2.51	25.47

(units of umol PO₄/g sediment)

(Table 1)

Percent Totals of SEDEX PO₄ Solid Phases

Phase Fractions (%total)			Loosely-sorbed	Fe-Bound	Ca-Bound	Refractory	Organic P	Total PO ₄
Buzzards	Mar-03	SWI	9.86	37.98	20.42	9.50	22.23	100.00
Bay		20 cm	9.64	16.28	30.37	12.45	31.26	100.00
	Aug-04	SWI	10.24	40.38	20.49	8.33	20.57	100.00
		22 cm	5.45	15.15	35.91	16.16	27.34	100.00
Hingham	Jan-03	SWI	12.54	28.67	25.24	17.91	15.64	100.00
Bay		15.5 cm	1.85	12.44	46.13	21.71	17.86	100.00
	Sep-03	SWI	9.45	38.74	14.10	17.81	19.89	100.00
		10.2 cm	1.41	23.51	21.02	24.62	29.44	100.00
Mass.	Feb-02	SWI	9.56	29.22	25.20	20.96	15.06	100.00
Bay		21 cm	2.33	9.93	37.10	39.37	11.26	100.00
	Oct-02	SWI	11.49	22.06	30.35	22.44	13.65	100.00
		9.2 cm	2.71	7.46	38.98	40.99	9.85	100.00

(Phase fraction = (phase/Total PO₄)*100)

(Table 2)

Determination of Variability in the SEDEX Method

Buzzards Bay (March 2003)						
# 2-2 Depth = 0.23cm	Step I	Step II	Step III	Step IV	Step V	Total PO ₄
Average	1.765	11.37	6.263333	2.835	7.4825	29.07333
Std. Dev	0.421861	1.708469	0.737111	0.130767	0.122848	1.287646
% CV	23.90147	15.02611	11.76868	4.612591	1.641806	4.42896
#2-23 Depth = 20.03cm						
Average	1.373333	3.17	6.183333	2.786667	6.933333	20.49
Std. Dev	0.880814	0.527162	1.245325	0.130128	0.152753	1.738649
% CV	64.13694	16.62972	20.14002	4.66967	2.203161	8.485353

*(where %CV = Std. Dev/Average * 100) (#2-2 = quadruplicate, #2-23 = triplicate)*

(Table 3)

Inventories of Solid-Phase PO₄ in Sediments

PO ₄ Inventories (umol/cm ²)			Loosely-sorbed	Fe-Bound	Ca-Bound	Refractory	Organic P	Total PO ₄
Buzzards Bay	Mar-03	SWI	0.14	0.53	0.29	0.13	0.31	1.41
		21 cm	26.69	55.17	85.76	35.25	90.48	293.35
	Aug-04	SWI	0.10	0.38	0.19	0.08	0.19	0.95
		22 cm	14.52	44.72	94.93	40.99	74.78	269.94
Hingham Bay	Jan-03	SWI	0.15	0.34	0.30	0.21	0.19	1.19
		15.5 cm	8.03	44.85	145.24	65.19	55.66	318.96
	Sep-03	SWI	0.19	0.77	0.28	0.35	0.39	1.98
		10.2 cm	5.08	49.23	36.06	45.67	46.43	182.47
Mass. Bay	Feb-02	SWI	0.25	0.77	0.66	0.55	0.40	2.63
		21.3 cm	17.98	67.85	189.66	198.21	67.26	540.96
	Oct-02	SWI	0.27	0.52	0.71	0.53	0.32	2.35
		9.2 cm	9.14	23.43	87.37	83.46	24.73	228.13

(inventory in units of (umol PO₄/cm² sediment)

(Table 4)

Organic Carbon and PO₄ in Sediments

Organic C:P		Expected C:P from [C _{org}]	Measured P _(org) (SEDEX)	Expected C:P/ SEDEX P	(C:P) _{org} from [C _{org}], [P _{org}]
Buzzards	SWI	20	7.5	2.7	277
Bay	Deep	13	6.5	2	218
Hingham	SWI	25	8	3.3	333
Bay	Deep	22	7	3.1	332
Mass.	SWI	19	5.2 - 6.5	3.2	341
Bay	Deep	10	2.5 - 3.0	3.6	403

(where Expected C:P = ([C_{org}]/106), and P(org) = SEDEX Step V, average values for both sampling times)

(Table 5)

Solid Phase Iron Amounts (% weight) and Fe:Al Ratios

		% Fe	Fe:Al
Buzzards	SWI	3.3	0.52
Bay	z > 8cm	2.8 - 3.2	0.45 - 0.48
Hingham	SWI	4.0 - 4.5	0.55 - 0.60
Bay	z > 8cm	4.2 - 4.6	0.55
Mass.	SWI	(2/02) 3.7	0.5 - 0.56
Bay		(10/02) 3.4	
	z > 5 cm	(2/02) 2.8 - 3.5	0.38 - 0.42
		(9/02) 2.4 - 2.5	

(where (% Fe) is by weight)

(Table 6)

Solid Phase PO₄ Downcore Fluxes from Oxidic/Anoxic Boundary Layer

Solid Phase Fluxes		Downward	Downward	
Buzzards		Total PO ₄ Flux	Fe-bound Flux	% of Flux as Fe-Bound Fraction
Bay	Mar-03	-37	-36	97
Fluxes	Aug-04	-20	-16	82
Hingham				
Bay	Jan-03	-10	-9	94
Fluxes	Sep-03	-125	-88	70
Mass				
Bay	Feb-02	-16	-14	85
Fluxes	Oct-02	-33	-13	39

(flux in umol PO₄/ cm² yr, "-" sign indicates downward flux)

(Table 7)

Porewater PO₄ Fluxes (derived from maximum porewater gradient)

PO ₄ Fluxes (max gradient)			
Buzzards		Flux Up	Flux Down
Bay	Mar-03	1.50	0.01
Fluxes	Aug-04	8.78	0.84
Hingham		Flux Up	
Bay	Jan-03	7.45	N/A
Fluxes	Sep-03	38.24	N/A
Mass		Flux Up	
Bay	Feb-02	2.21	N/A
Fluxes	Oct-02	5.32	N/A

(all fluxes in units of $\mu\text{mol}/\text{cm}^2 \text{ yr}$)

(Table 8)

Porewater tCO₂ Fluxes (derived from maximum porewater gradient)

tCO ₂ Fluxes (max gradient)			
Buzzards		Flux Up	Flux Down
Bay	Mar-03	23.00	-6.07
Fluxes	Aug-04	160.68	-10.18
Hingham		Flux Up	
Bay	Jan-03	134.44	N/A
Fluxes	Sep-03	270.93	N/A
Mass		Flux Up	Flux Down
Bay	Feb-02	27.84	-18.63
Fluxes	Oct-02	110.39	-10.71

(all fluxes in units of $\mu\text{mol}/\text{cm}^2 \text{ yr}$)

(Table 9)

Porewater PO₄ and tCO₂ Fluxes (derived from SWI gradient) and C:P Ratios

SWI Porewater Fluxes				
Buzzards		PO ₄ (SWI)	tCO ₂ (SWI)	tCO ₂ :PO ₄ (SWI)
Bay	Mar-03	0.05	270.81	5416.21
Fluxes	Aug-04	5.63	334.77	59.46
Hingham				
Bay	Jan-03	0.33	742.11	2248.80
Fluxes	Sep-03	14.7	803.27	54.64
Mass				
Bay	Feb-02	0.92	286.51	311.42
Fluxes	Oct-02	3.22	316.31	98.23

(fluxes in units of $\mu\text{mol}/\text{cm}^2 \text{ yr}$)

(Table 10)

Ratio of C:P in Porewater Fluxes (derived from maximum porewater gradient)

tCO ₂ :PO ₄		Redfield Ratio
		106.00
Buzzards Bay		
March		August
19.23		17.75
Hingham Bay		
January		September
18.04		7.08
Mass Bay		
February		October
21.03		22.77

(Table 11)

Benthic Flux Chamber PO₄ and tCO₂ Fluxes (derived from chamber time series) and Ratios

Benthic Flux Chamber		PO ₄ Flux	tCO ₂ Flux	tCO ₂ :PO ₄ (Chamber)
Buzzards	Mar-03	0.67	307.97	459
Bay	Aug-04	0.9	471.65	524
Hingham	Jan-03	11.55	1186.32	102
Bay	Sep-03	12.55	928.28	73
Mass.	Feb-02	3.14	460.63	146
Bay	Oct-02	14.14	1056.8	74

(flux is in units of umol/cm² yr)

(Table 12)

Comparison of SWI and Benthic Chamber Fluxes Out and Solid-phase Fluxes Downcore

Solid Phase Fluxes		Total PO ₄	Fe-bound	Porewater PO ₄		PO ₄ Flux
Buzzards		Flux Down	Flux Down	Flux Up	Flux Down/Flux Up	Benthic Chamber
Bay	Mar-03	-37.33	-36.31	1.51	24.69	0.67
Fluxes	Aug-04	-20.54	-16.83	9.63	2.13	0.9
Hingham		Flux In				
Bay	Jan-03	-10.53	-9.84	7.45	1.41	11.55
Fluxes	Sep-03	-125.65	-87.59	38.25	3.29	12.15
Mass		Flux In				
Bay	Feb-02	-16.37	-13.98	2.21	7.41	3.14
Fluxes	Oct-02	-32.71	-12.90	5.32	6.15	14.14

(fluxes in units of umol/cm² yr)

(Table 13)

Comparison of SWI and BFC Fluxes w/ C:P Ratios

	PO ₄		tCO ₂			
Buzzards Bay	Porewater SWI	BFC	Porewater SWI	BFC	tCO ₂ :PO ₄ (SWI)	tCO ₂ :PO ₄ (BFC)
Mar-03	0.05	0.67	270.81	307.97	5416.21	459.66
Aug-04	5.63	0.90	334.77	471.65	59.46	524.06
Hingham Bay						
Jan-02	0.33	11.55	742.11	1186.32	2248.80	102.71
Sep-02	14.70	12.15	803.27	928.28	54.64	76.40
Mass Bay						
Feb-02	0.92	3.41	286.51	460.63	311.42	135.08
Oct-02	3.22	14.41	316.31	1056.80	98.23	73.34

(fluxes in units of $\mu\text{mol PO}_4/\text{cm}^2 \text{ yr}$)

(Table 14)

Trapping Effects of Oxidic/Anoxic Boundary Layer

PO ₄ Trapping					
Buzzards Bay		Maximum Gradient Flux	SWI Flux	"Trapped" PO ₄	% of Maximum Gradient Trapped
Fluxes	Mar-03	1.50	0.05	1.45	96
	Aug-04	8.78	5.63	3.15	35
Hingham Bay					
Fluxes	Jan-03	7.45	0.33	7.12	96
	Sep-03	38.24	14.7	23.54	62
Mass Bay					
Fluxes	Feb-02	2.21	0.92	1.29	58
	Oct-02	5.32	3.22	2.10	39

(units of flux in $\mu\text{mol PO}_4/\text{cm}^2 \text{ yr}$, "trapped PO₄" = difference between fluxes)

(Table 15)

Flux Calculation Parameters: PO₄

Pore-Water Fluxes						
Upward Fluxes						
		T (deg C)	D(sw)	v	Average Porosity	D(sed)
Buzzards	Mar-03	8	165	3.48	0.841477706	107.54
Bay	Aug-04	20	240.6	2.6	0.865529222	190
Hingham	Jan-02	5	146.2	1.86	0.786983777	118.9811
Bay	Sep-02	15	209.15	2.95	0.805594807	137.2099
Mass.	Feb-02	5	146.2	1.75	0.817714286	125.7183
Bay	Oct-02	10	177.6	1.75	0.8165	152.5493

Downward Fluxes						
		T (deg C)	D(sw)	v	Average Porosity	D(sed)
Buzzards	Mar-03	8	165	3.48	0.805315938	96.44514
Bay	Aug-04	20	240.6	2.6	0.783948454	162.988

(Table 16)

Flux Calculation Parameters: tCO₂

		T (deg C)	Range(cm)	D(sw)	v	Average Porosity	D(sed)
Buzzards	Mar-03	8	0.11 - 1.07	267.7	3.48	0.829566	168.4218
Bay	Aug-04	20	0.16 - 0.81	332.6	2.6	0.888896	275.4759
Hingham	Jan-03	5	0.13 - 1.22	202.7	1.86	0.890554	183.4686
Bay	Sep-03	15	0.15 -1.06	289.3	2.95	0.814731	194.0107
Mass.	Feb-02	5	0.13 -0.68	202.7	1.75	0.8218	174.9558
Bay	Oct-02	10	0.0 -0.5	246	1.75	0.84975	217.7226

(Table 17)

Flux Calculation Parameters: Solid Phase PO₄

			Average		
Buzzards		Range	Porosity	Gradient	Dbio
Bay	Mar-03	0.23 - 0.7	0.849916	-9.47232	10.1
Fluxes	Aug-04	0.13 - 0.4	0.90599	19.93631	3.3
Hingham					
Bay	Jan-03	0.415 - 1.23	0.873033	-6.01361	3
Fluxes	Sep-03	0.14 - 0.99	0.851363	-13.5471	24
Mass					
Bay	Feb-02	0.68 - 1.25	0.834	-7.58373	5
Fluxes	Oct-02	1.17 -1.83	0.8296	-10.6987	6.9

(units of gradient in delta(umol PO₄/g sediment)/delta(cm))

(Table 18)

Figure 1: SEDEX Profiles – Buzzards Bay

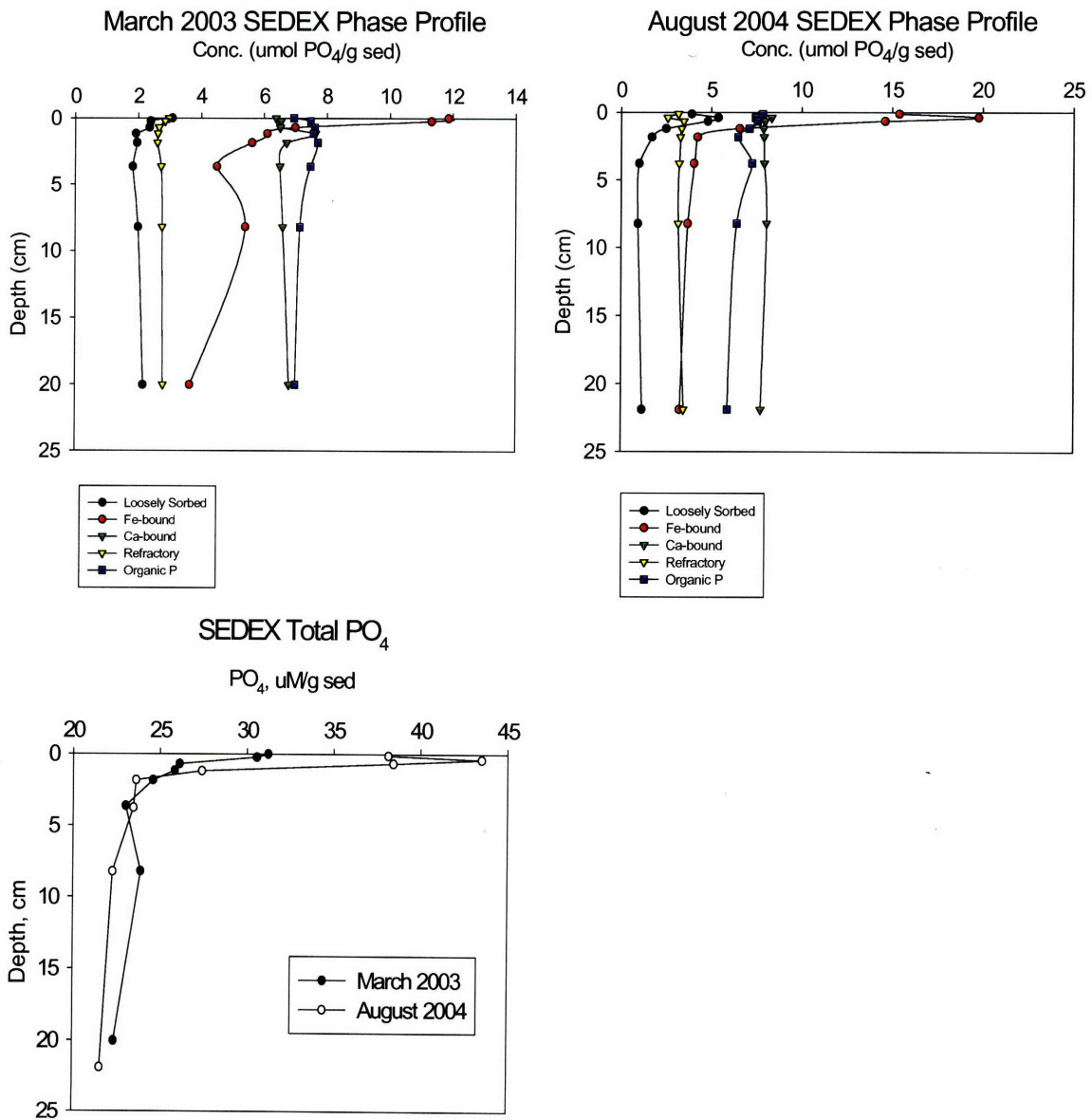


Figure 2: SEDEX Profiles – Hingham Bay

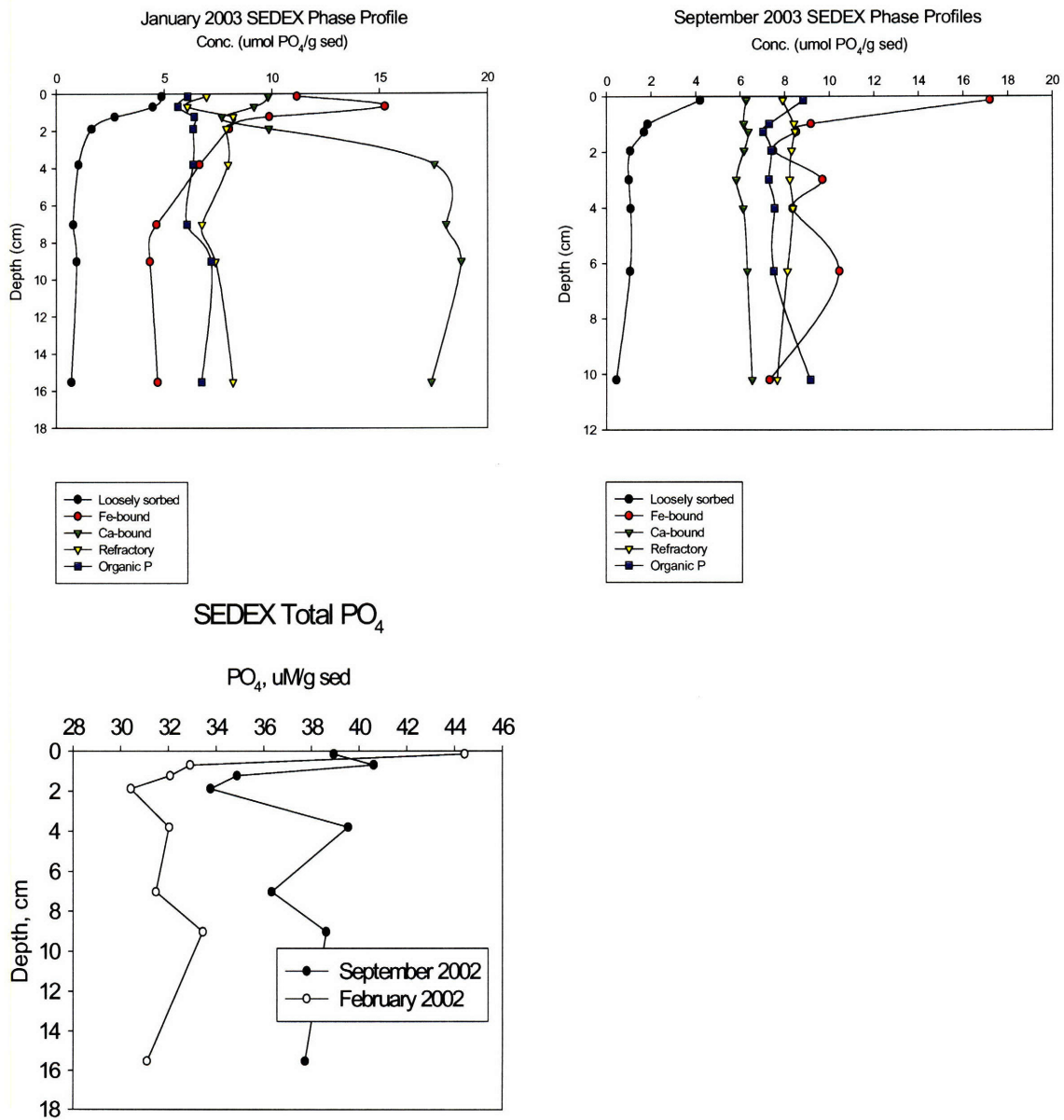


Figure 3: SEDEX Profiles – Massachusetts Bay

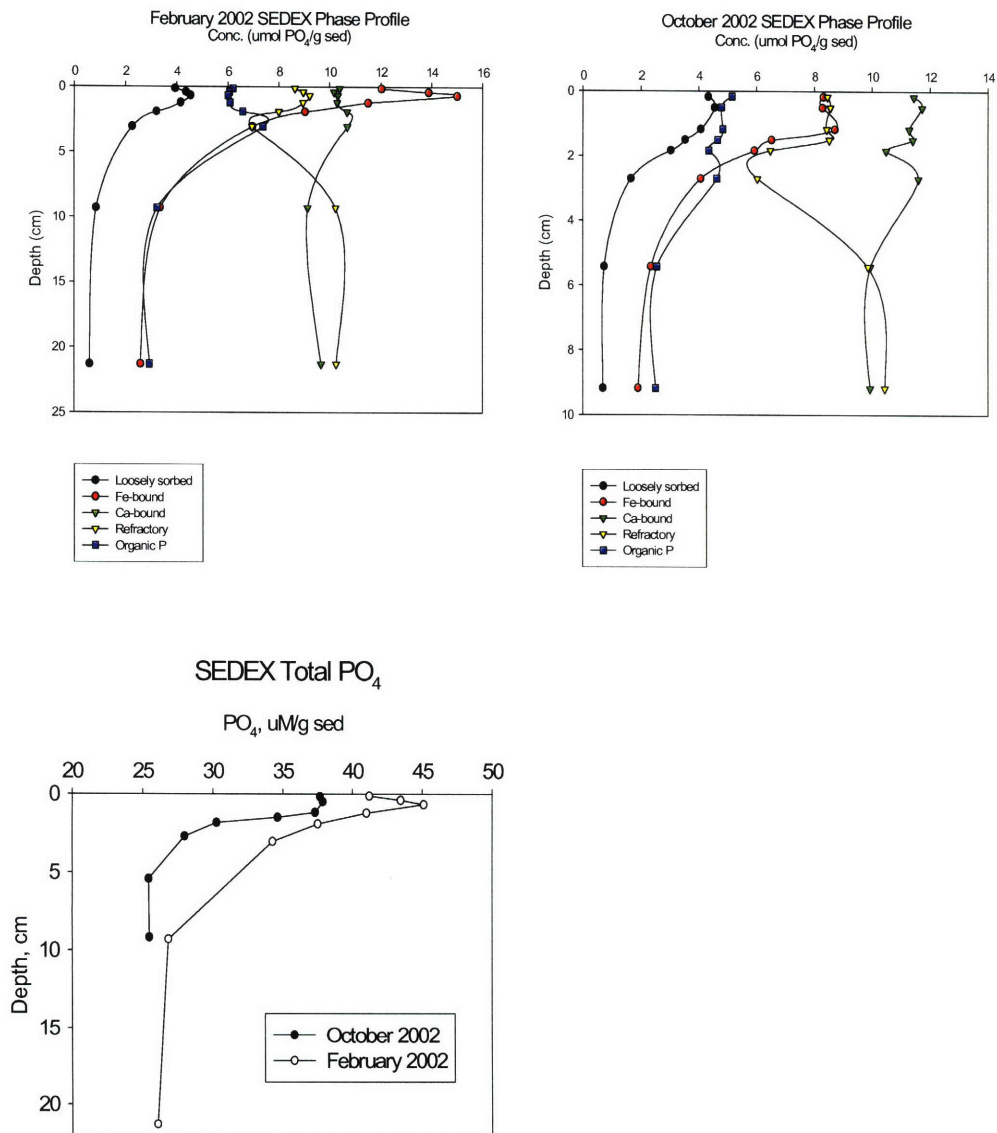


Figure 4: Hingham Bay - Additional Profiles

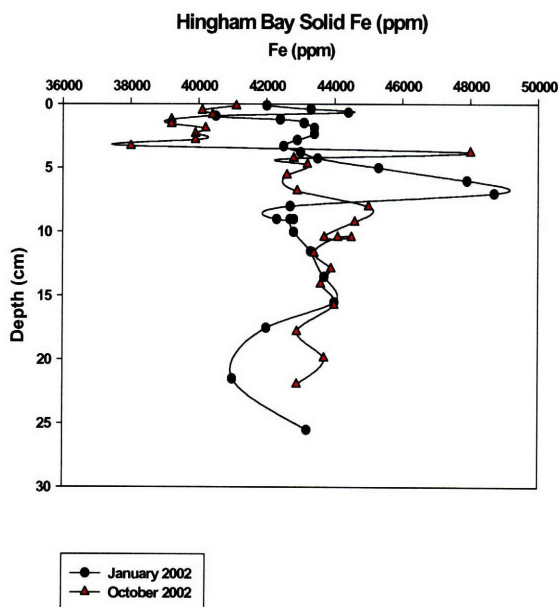
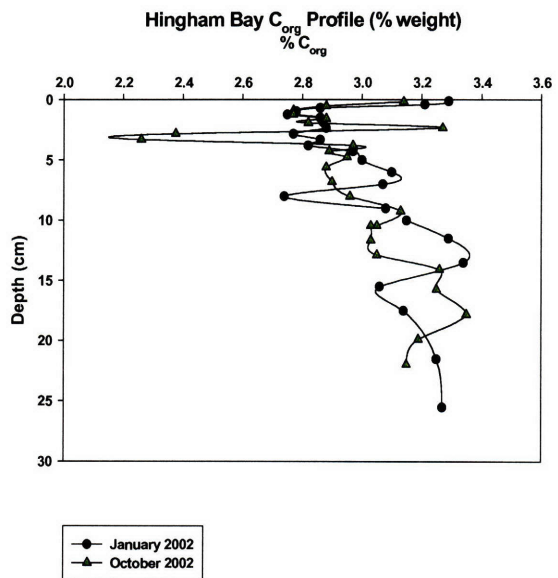
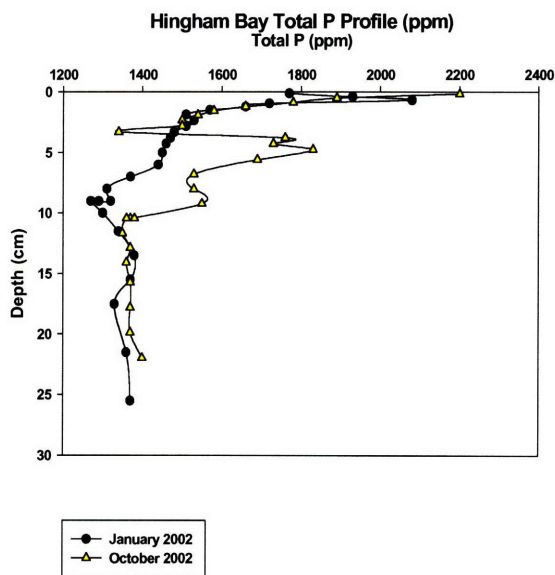


Figure 5: Massachusetts Bay – Additional Profiles

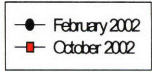
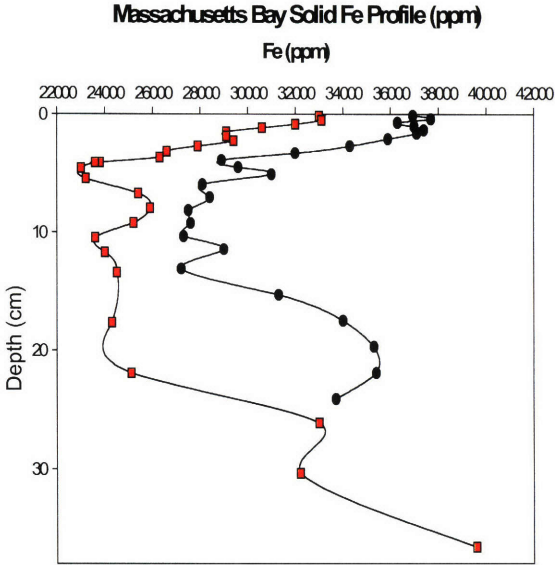
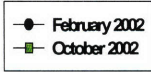
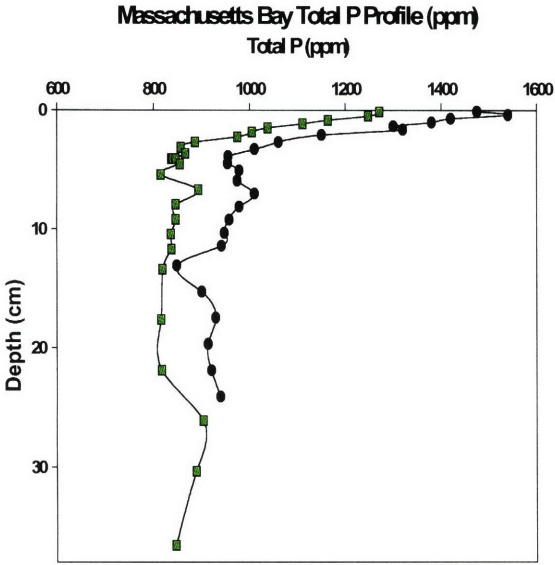


Figure 6: Hingham and Mass Bay: Total P vs. SEDEX Total PO₄

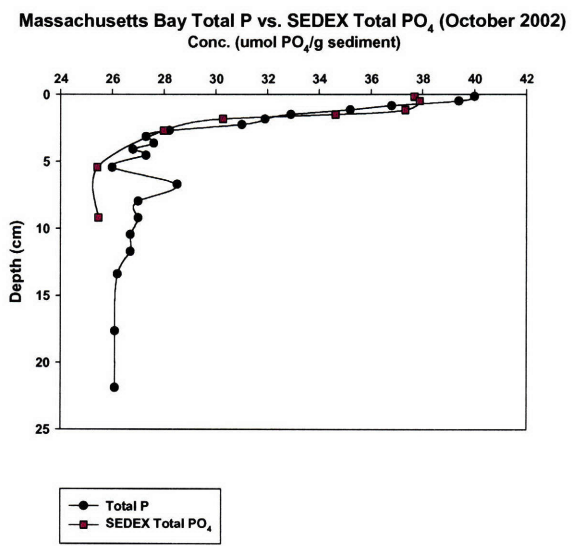
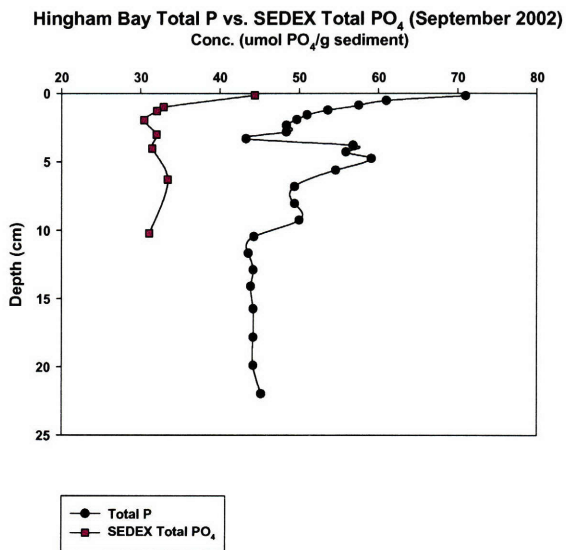
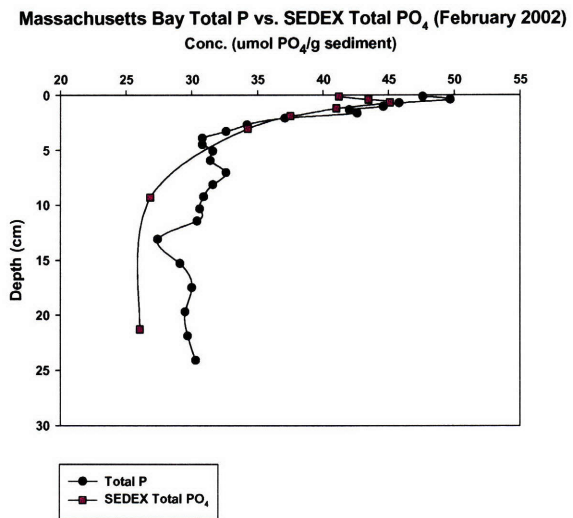
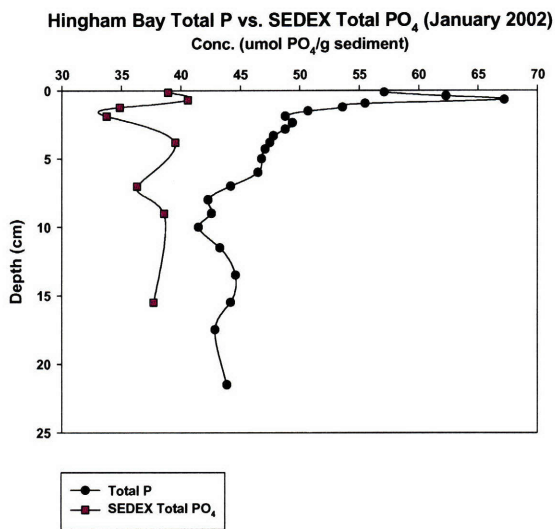


Figure 7: Buzzards Bay Phase Fractions

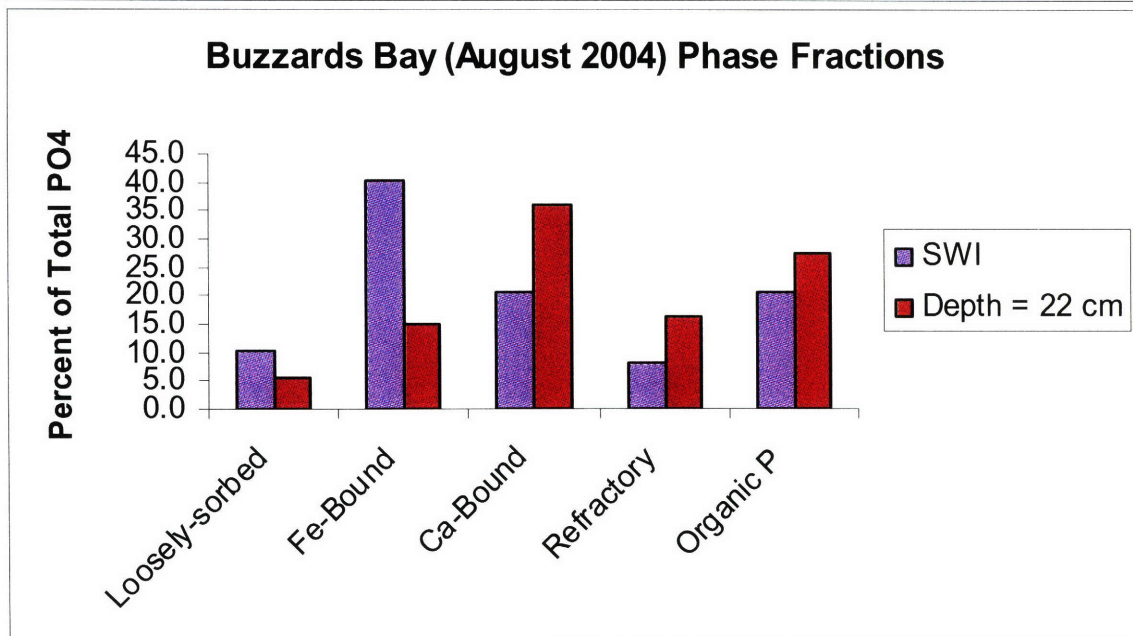
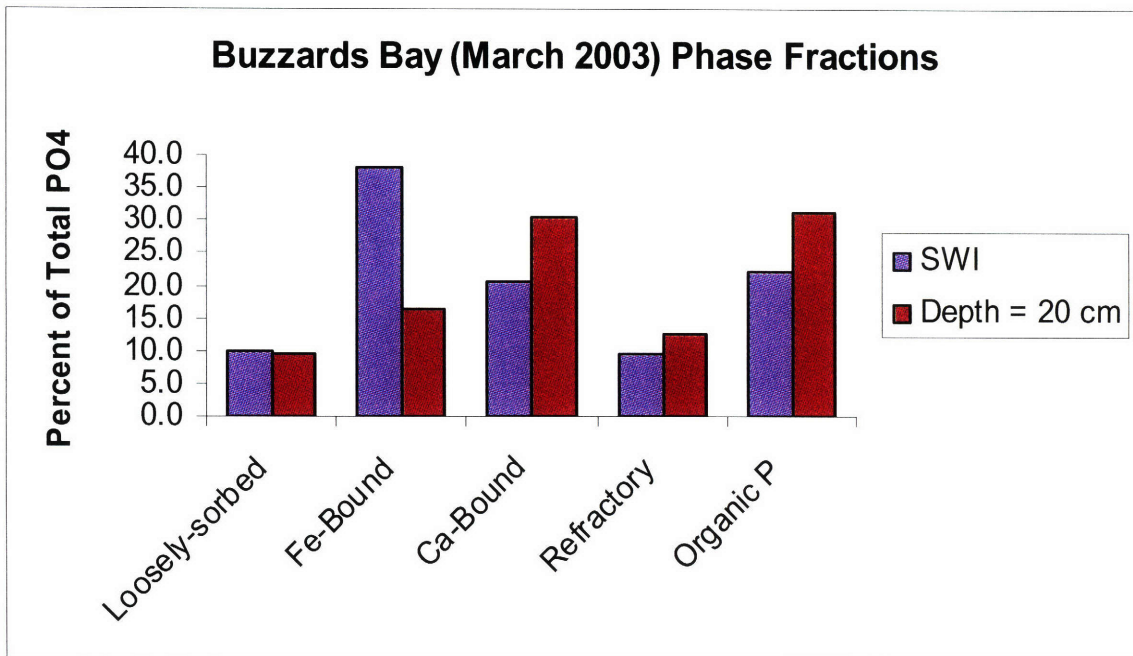


Figure 8: Massachusetts Bay Phase Fractions

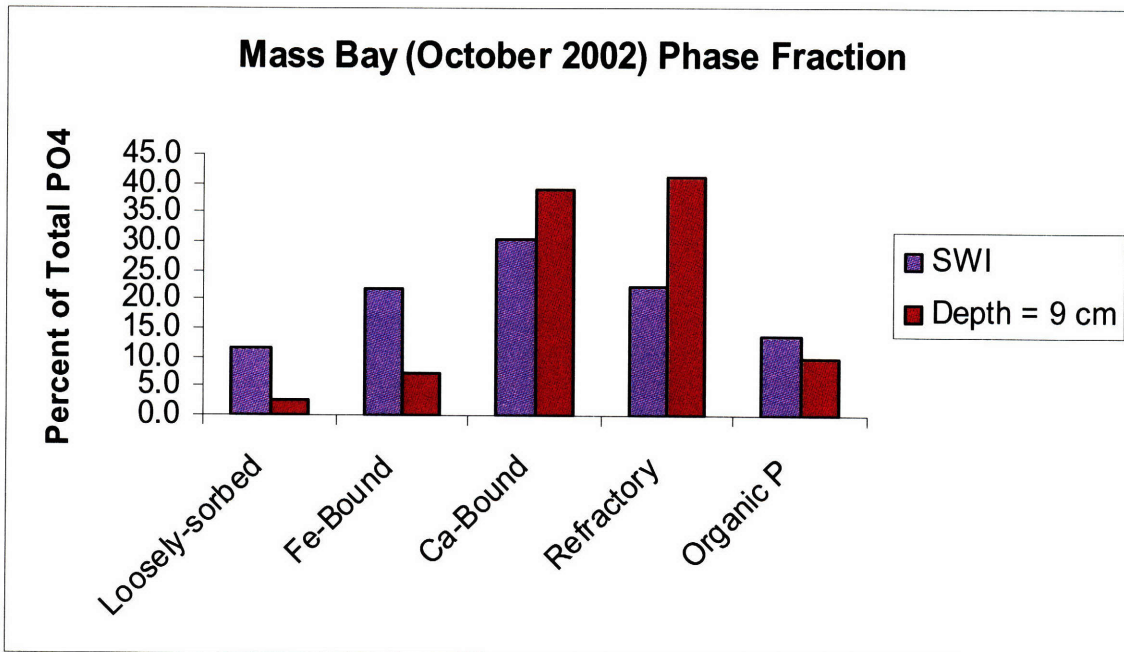
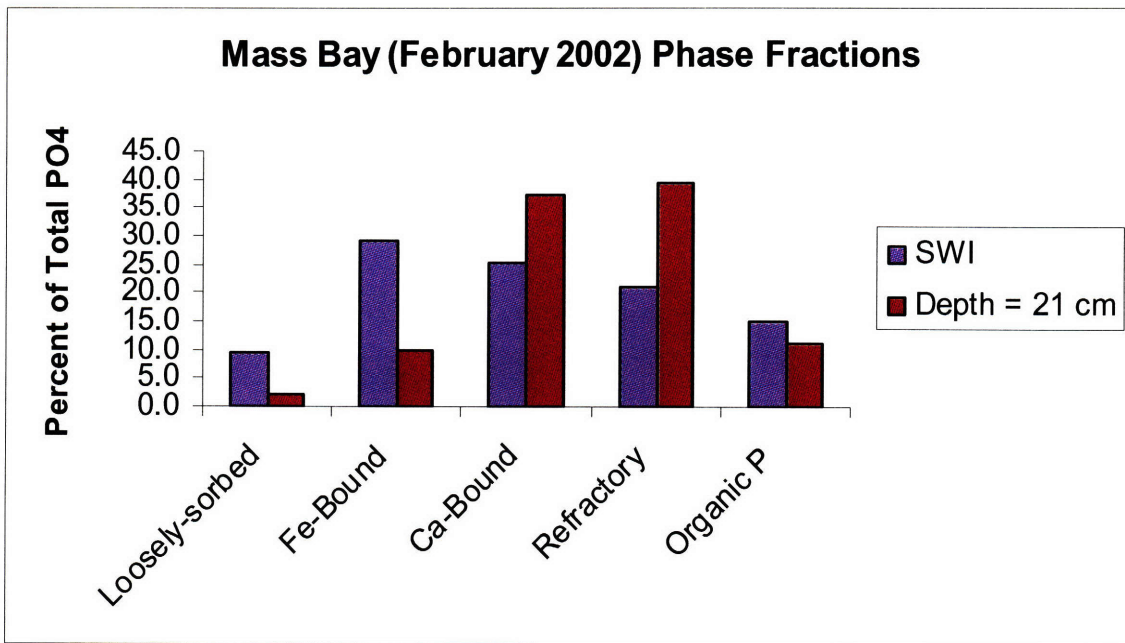


Figure 9: Hingham Bay Phase Fractions

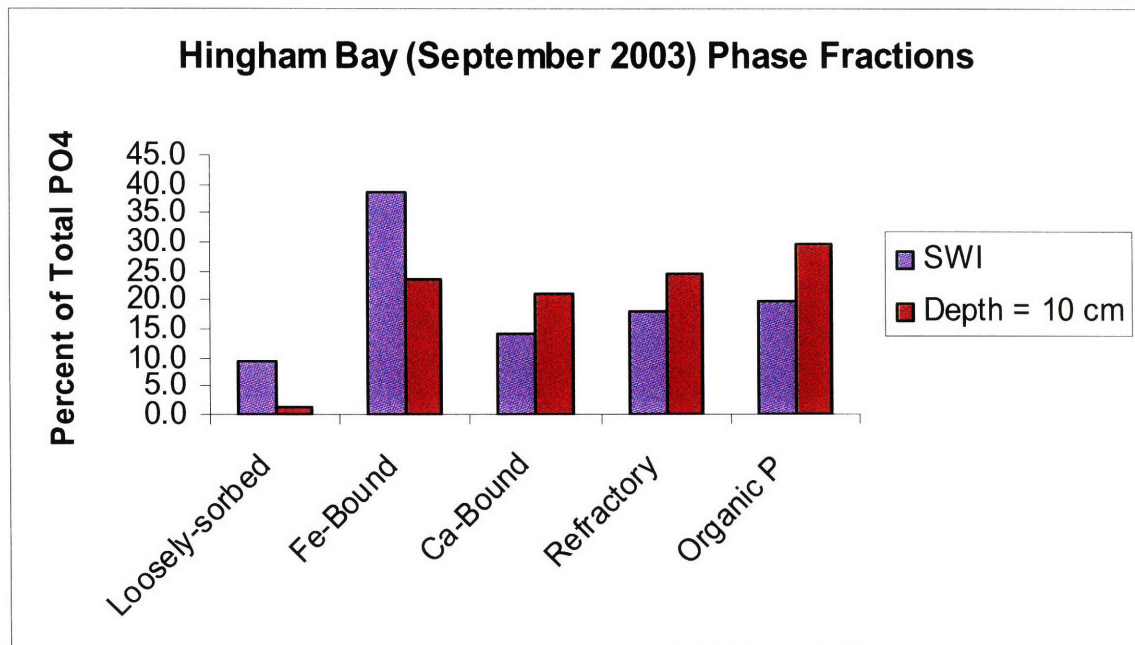
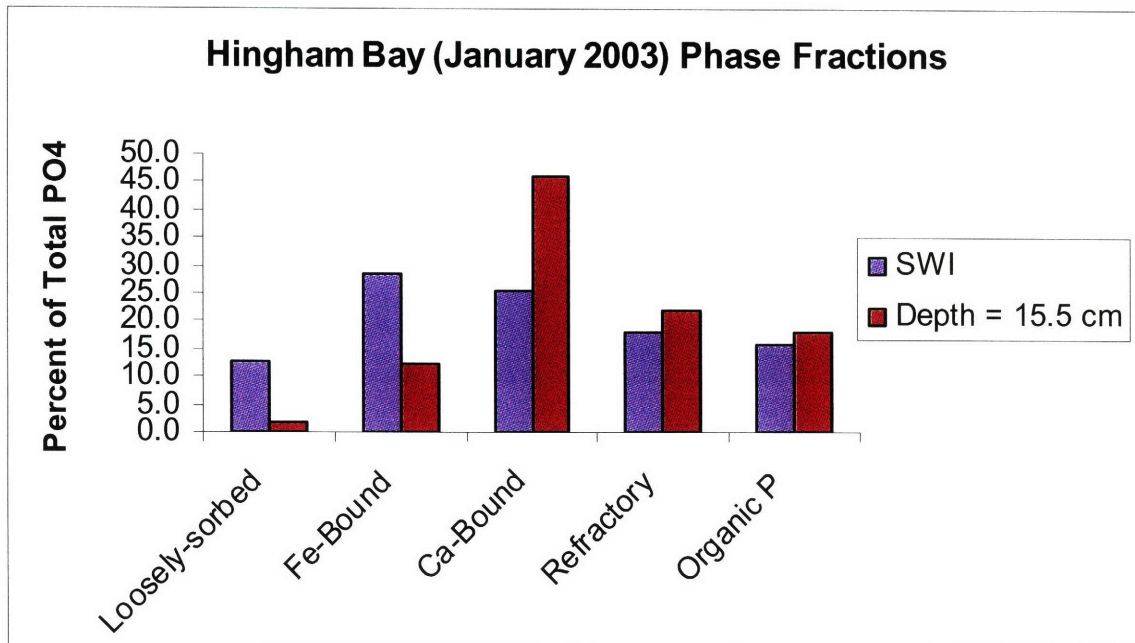


Figure 10: Fe/Al and P/Al ratio profiles

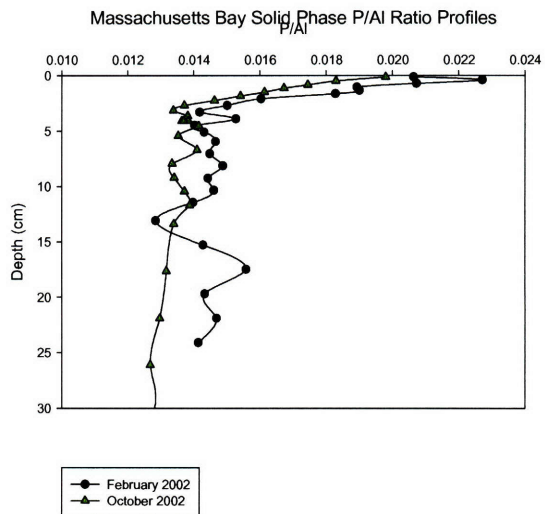
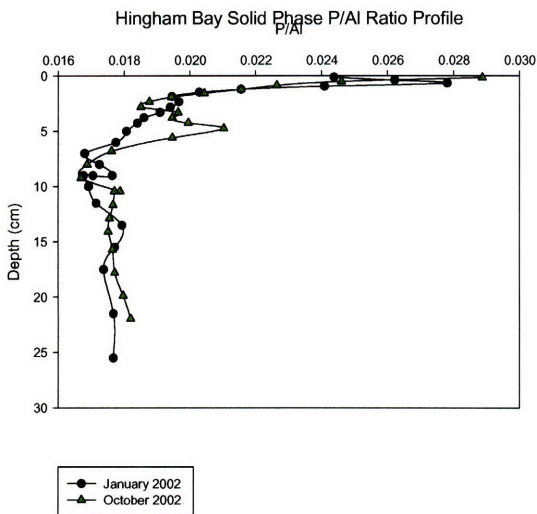
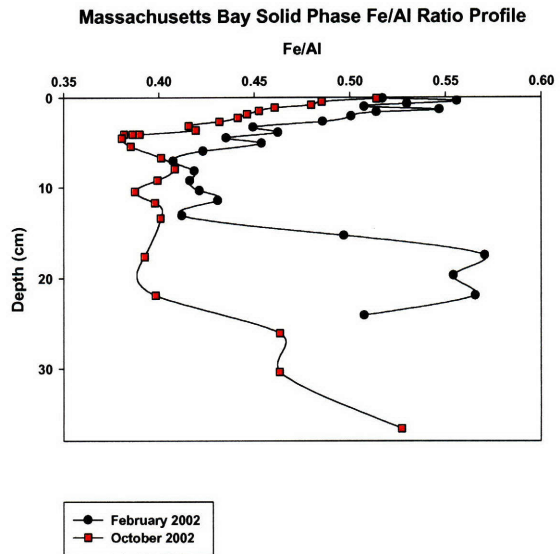
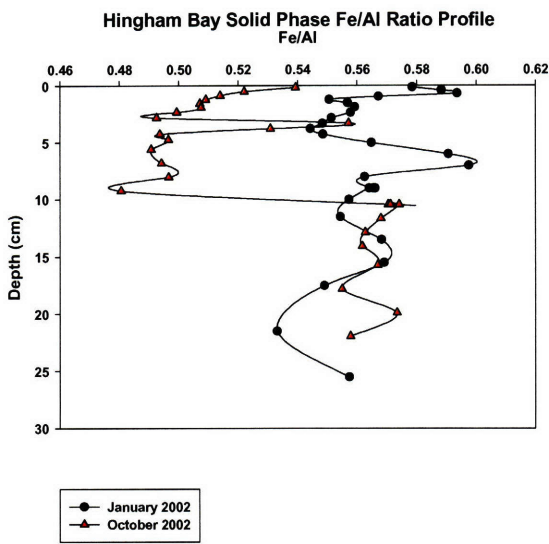


Figure 11: Hingham Bay (January 2003) Porewater Profiles

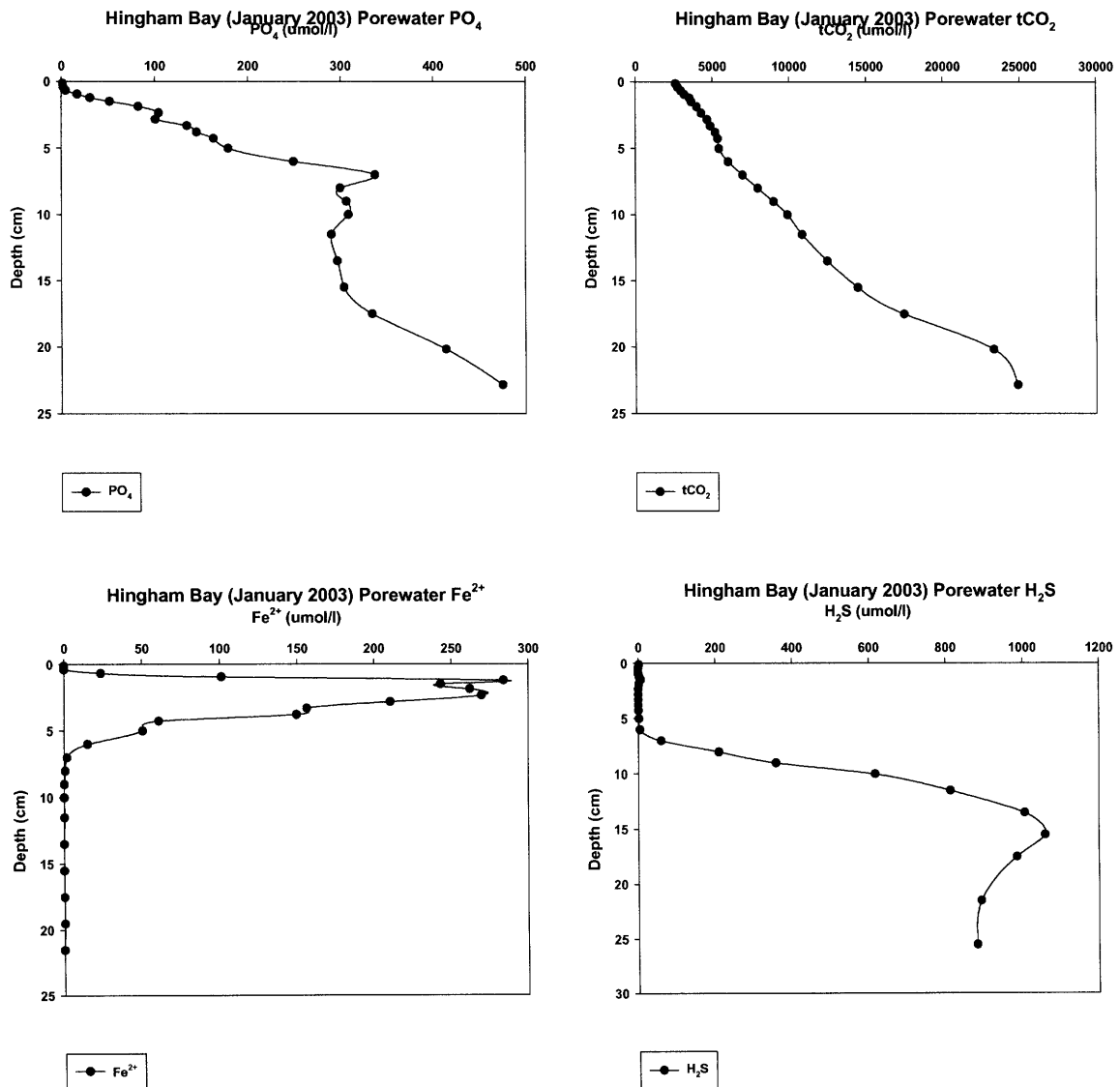


Figure 12: Hingham Bay (September 2002) Porewater Profiles

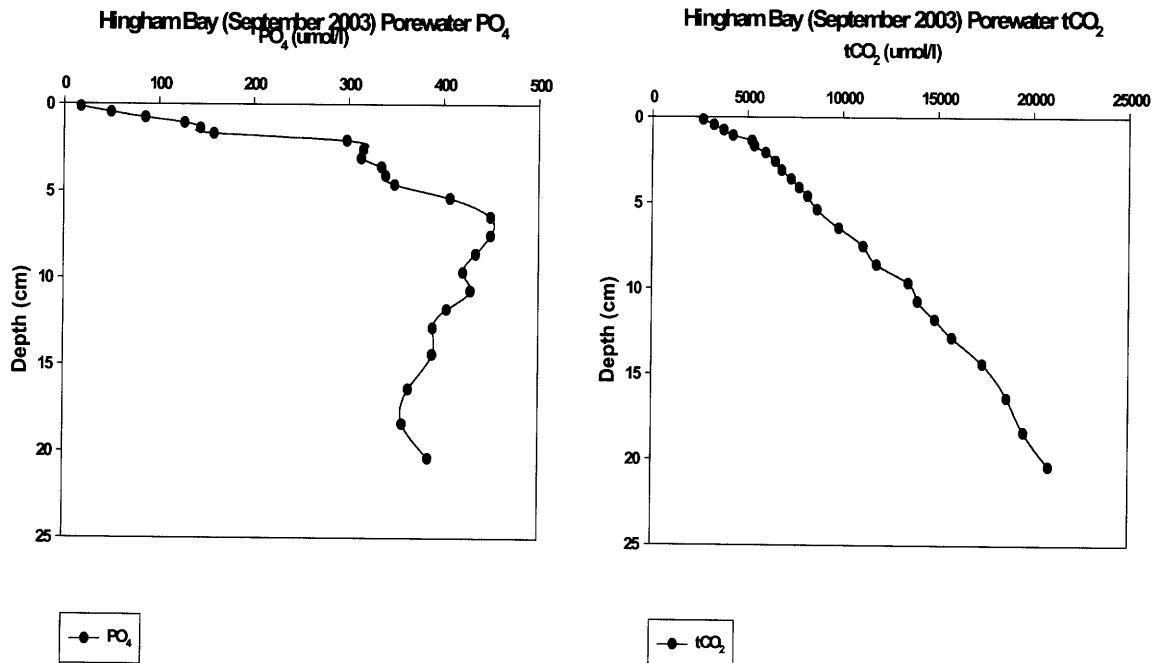


Figure 13: Surface Fe-Bound PO_4 Inventories

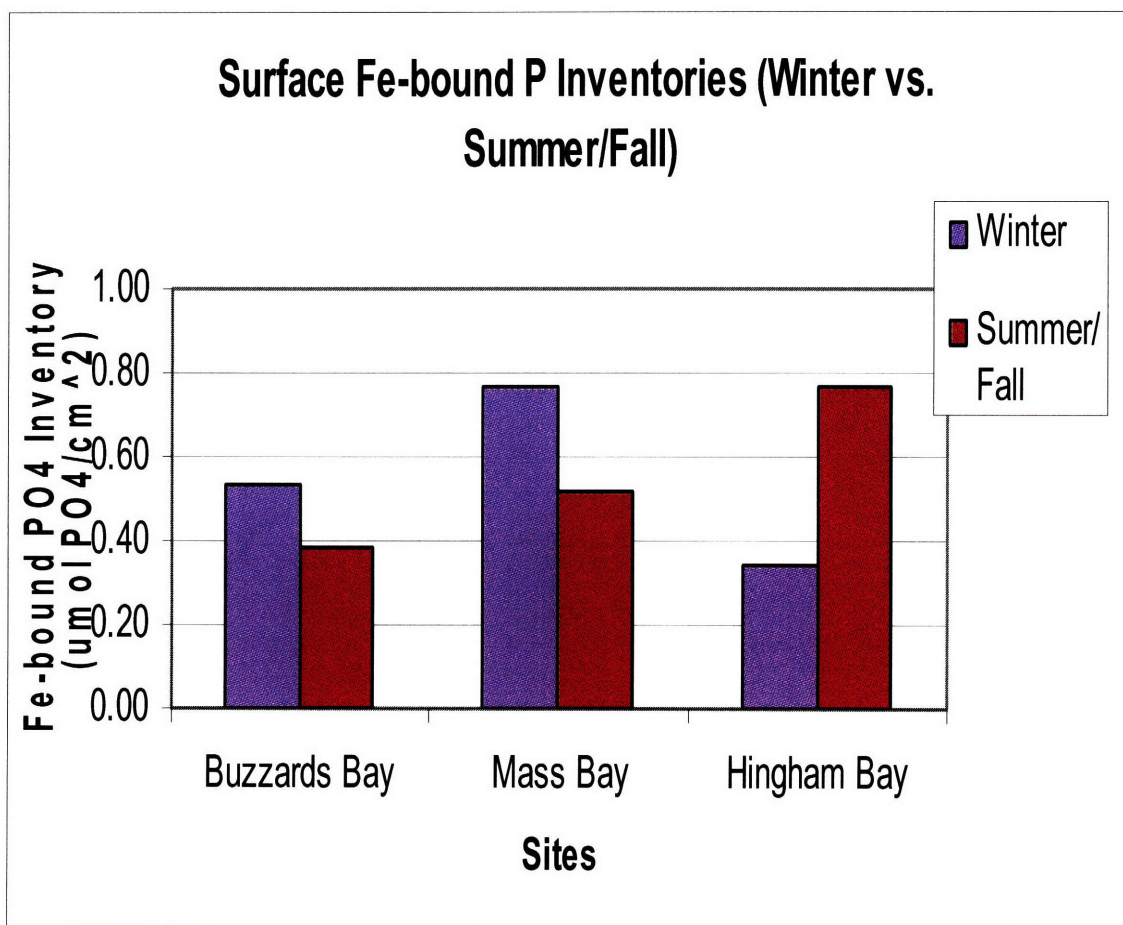


Figure 14: Buzzards Bay (March 2003) Porewater Profiles

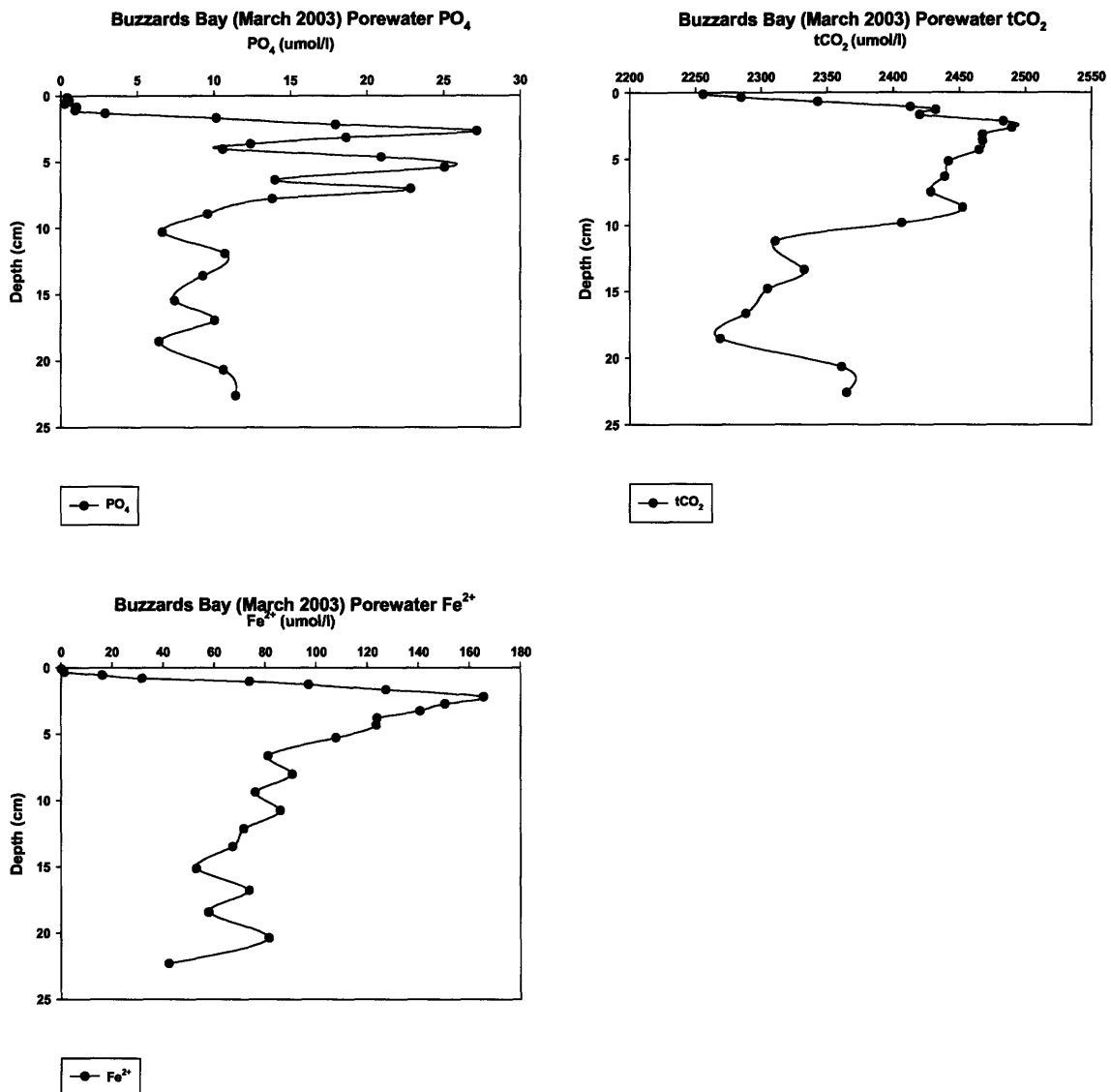


Figure 15: Buzzards Bay (August 2004) Porewater Profiles

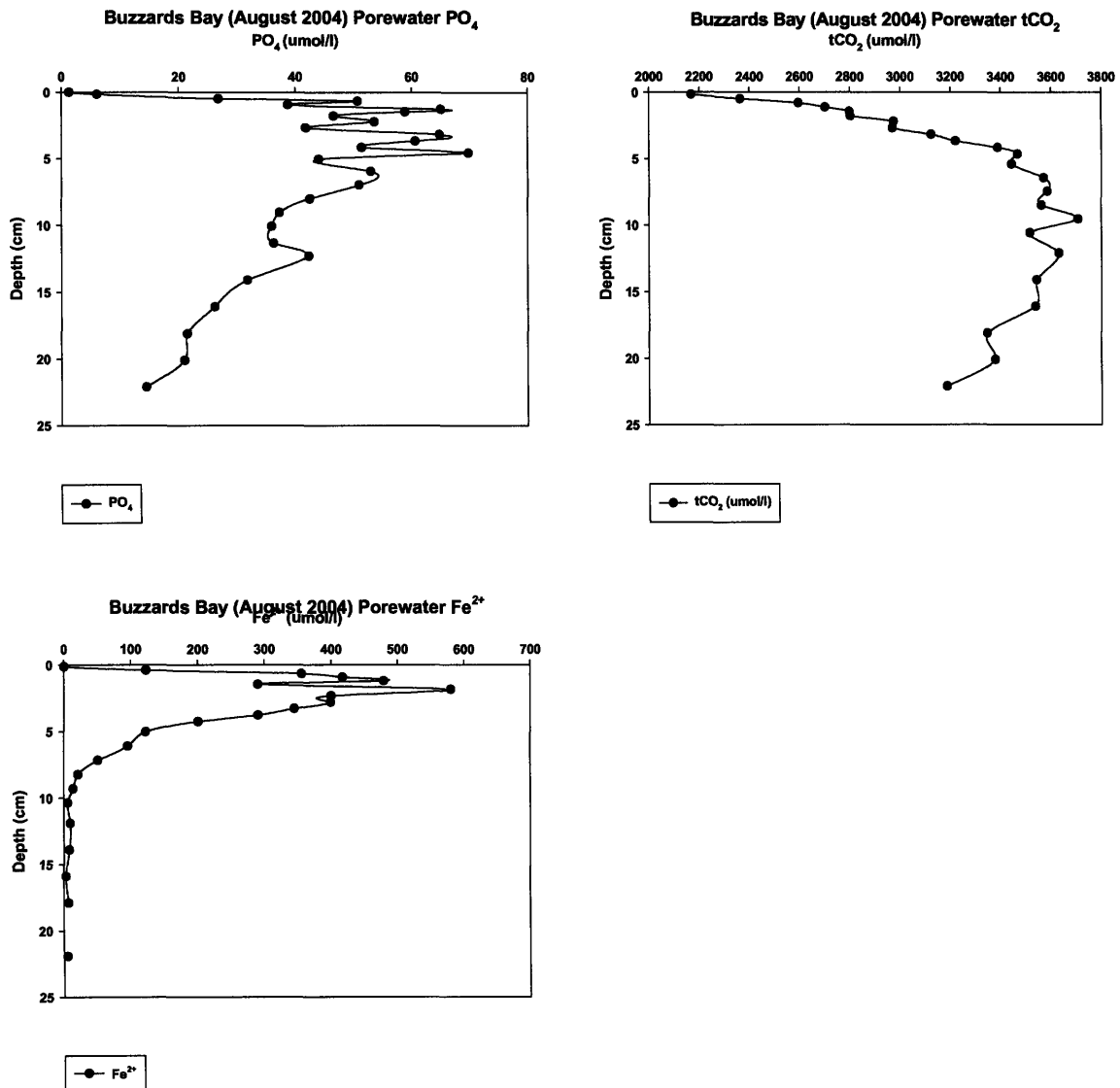


Figure 16: Massachusetts Bay (February 2002) Porewater Profiles

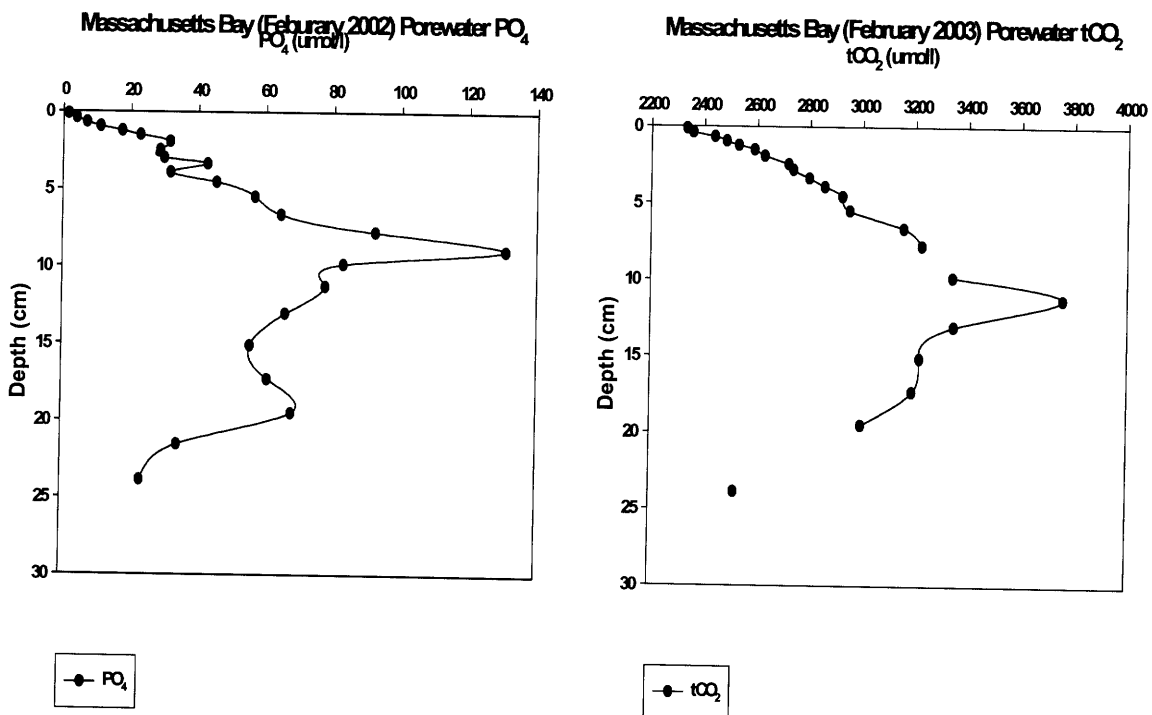


Figure 17: Massachusetts Bay (October 2002) Porewater Profiles

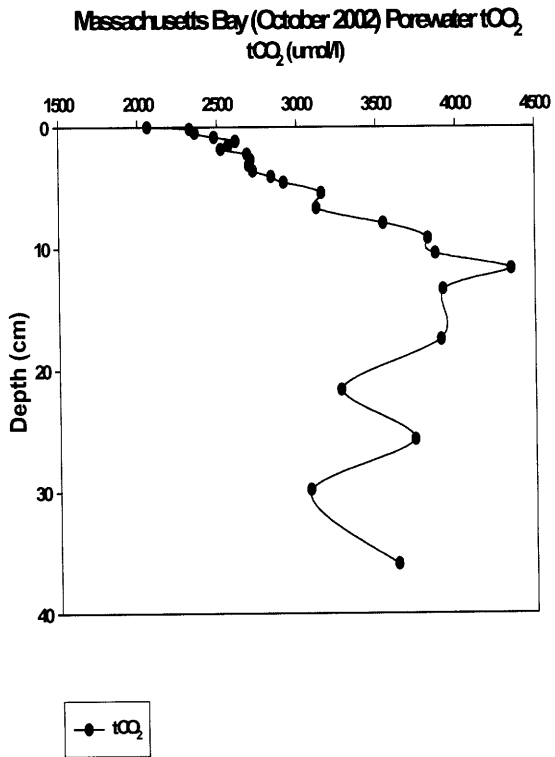
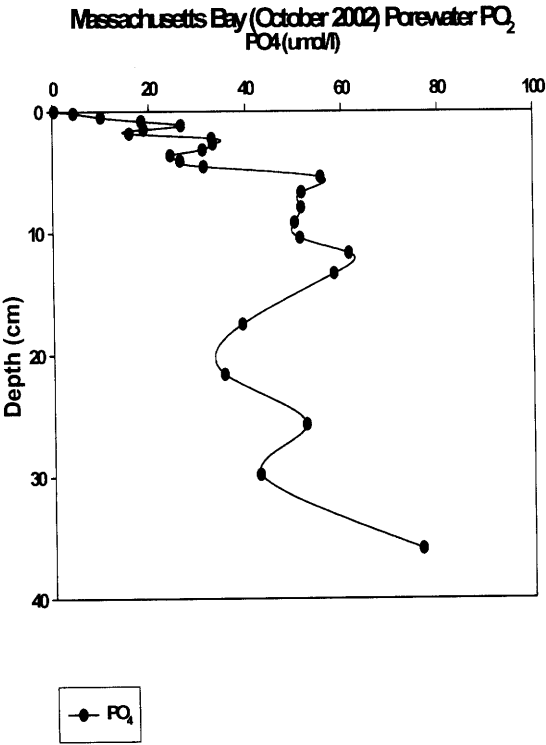
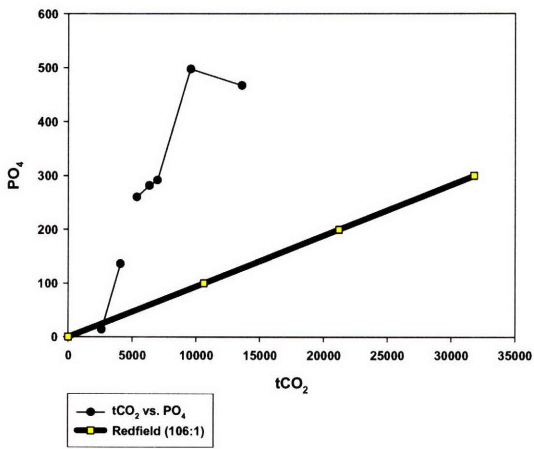
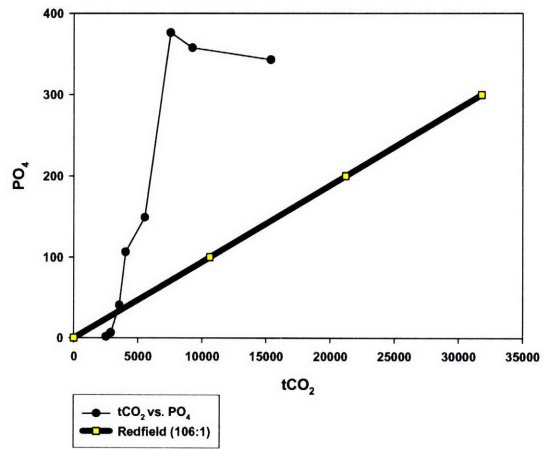


Figure 18: Hingham Bay and Mass Bay PO_4 vs tCO_2

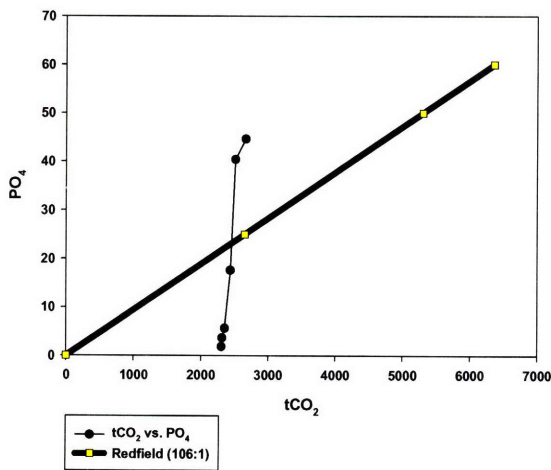
Hingham Bay (September 2002) Porewater tCO_2 vs. PO_4



Hingham Bay (January 2002) Porewater tCO_2 vs. PO_4



Massachusetts Bay (February 2002) Porewater tCO_2 vs. PO_4



Massachusetts Bay (October 2002) Porewater tCO_2 vs. PO_4

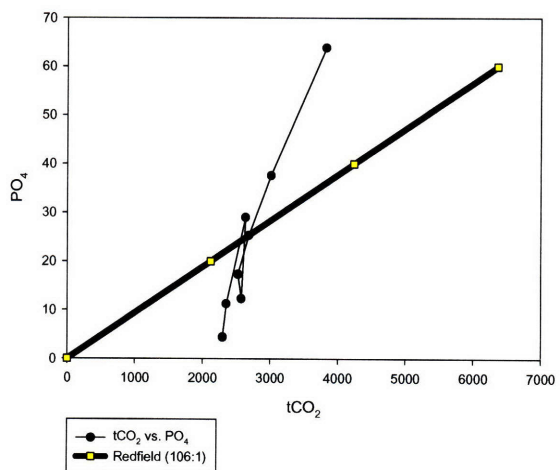
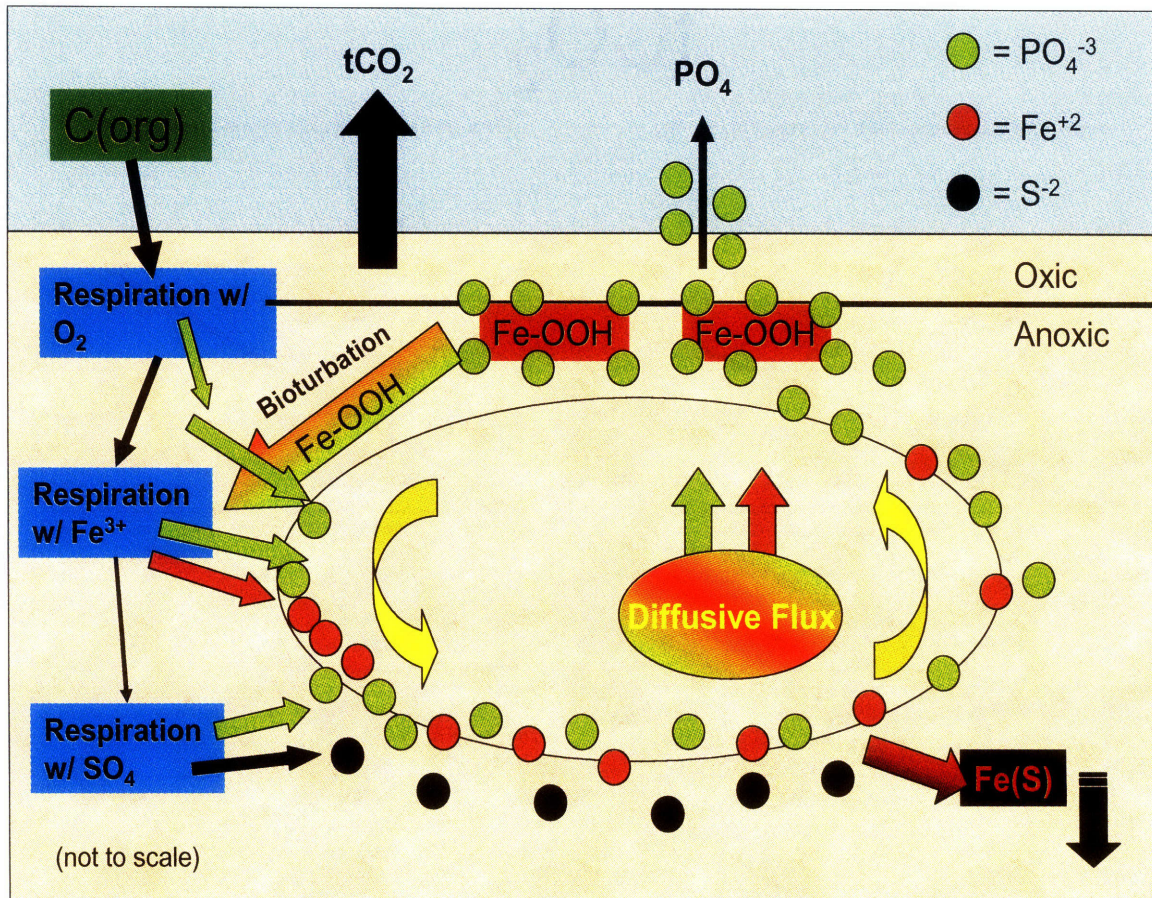


Figure 19: The Sedimentary Fe-P Cycle



- C_{org} respiration releases PO_4 into porewaters.
- PO_4 diffuses upwards until oxic/anoxic layer.
- PO_4 sorbs onto Fe-OOH
- Fe-OOH*P is mixed downwards and/or oxic layer thickness decreases
- Fe-OOH reduced \Rightarrow PO_4 released
- In certain sites, S^{-2} removes Fe^{2+} from depth, weakening Fe-trapping ability

Figure 20: Buzzards Bay Porewater and Fe-bound PO₄

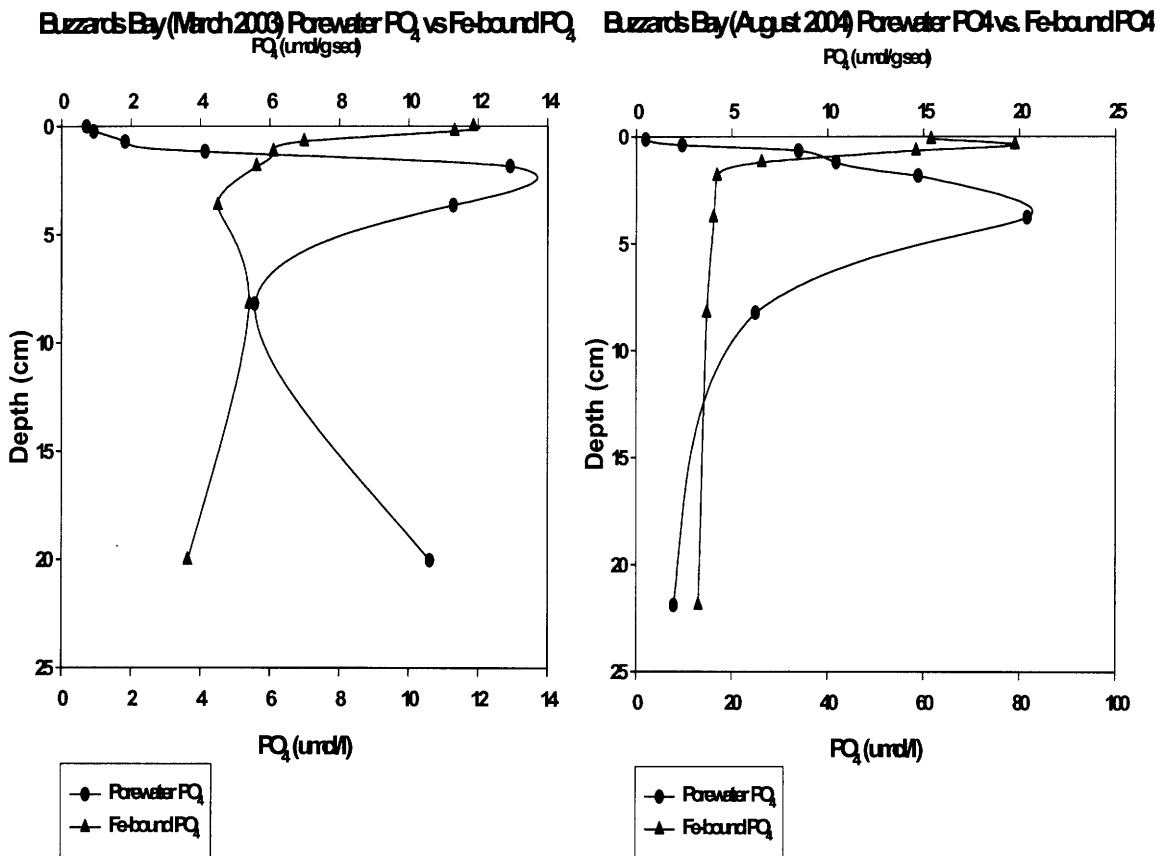
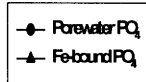
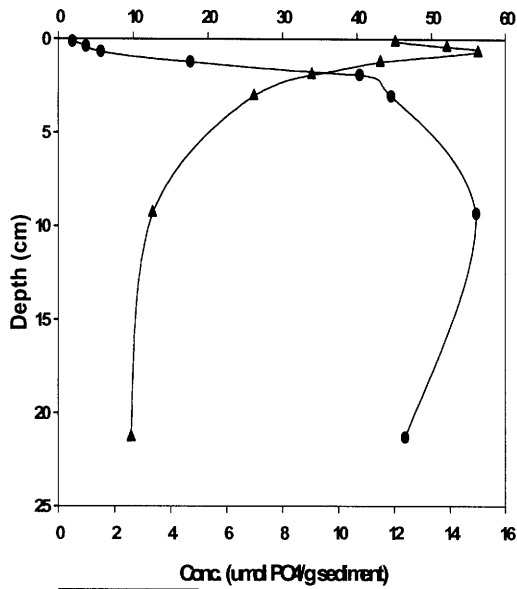


Figure 21: Massachusetts Bay Porewater and Fe-Bound PO_4

Massachusetts Bay (February 2002) Porewater PO_4 w/ Fe-bound PO_4
Conc. ($\mu\text{mol/l}$)



Massachusetts Bay (October 2002) Porewater PO_4 w/ Fe-bound PO_4
Conc. ($\mu\text{mol/l}$)

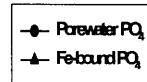
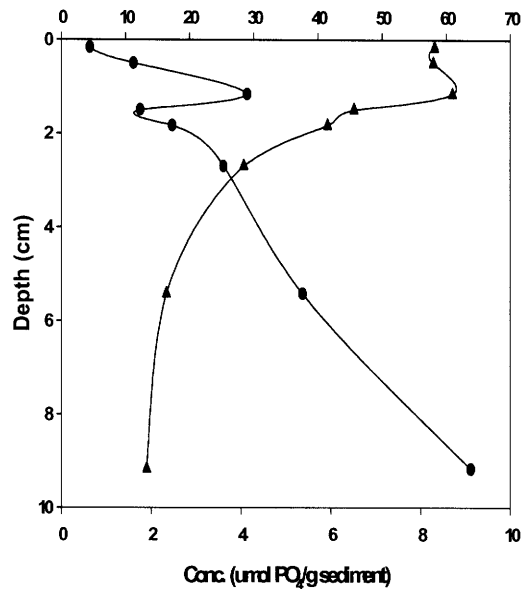


Figure 22: Hingham Bay Porewater and Fe-bound PO₄

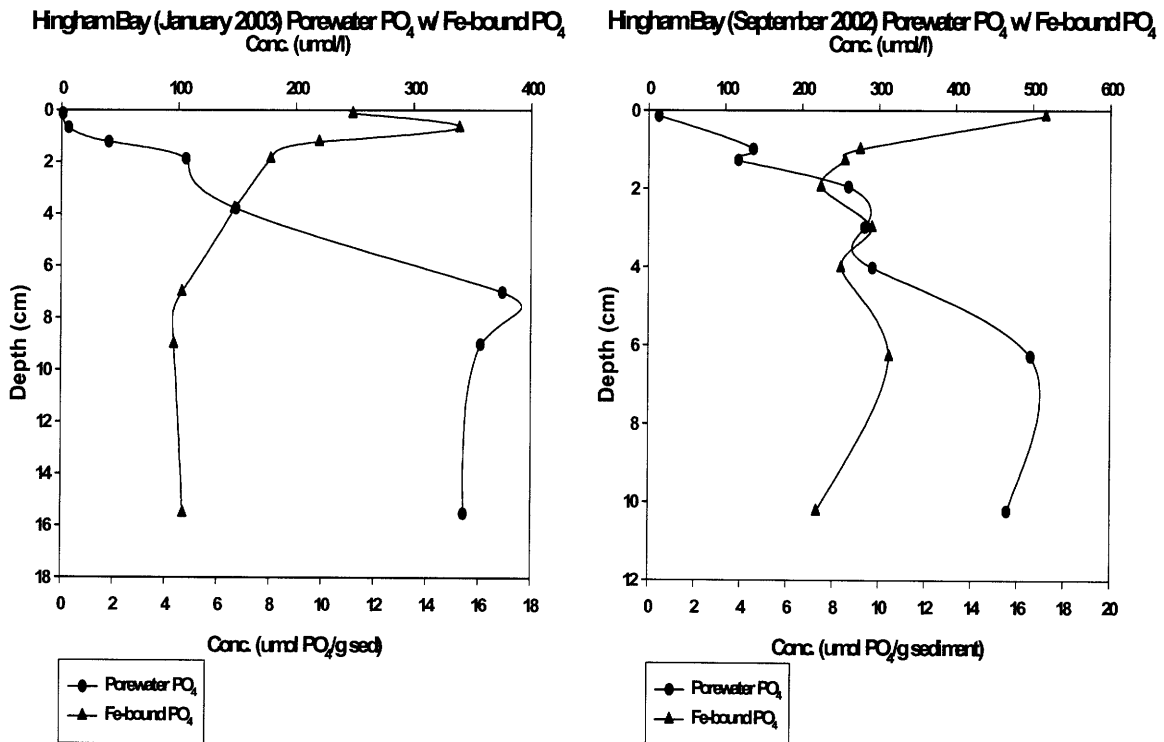


Figure 23: Solid-bound PO_4 Down Into Sediments

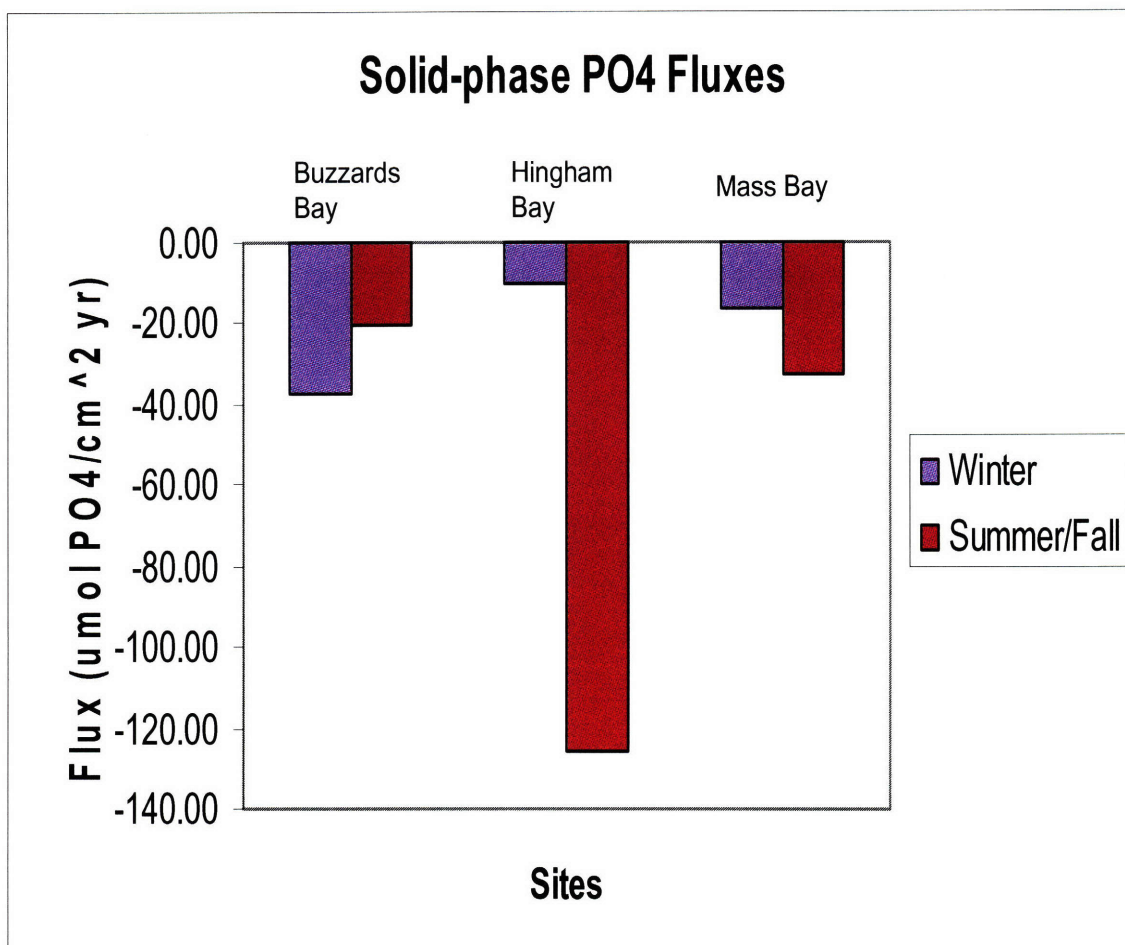


Figure 24: Porewater Fluxes Out of Sediments

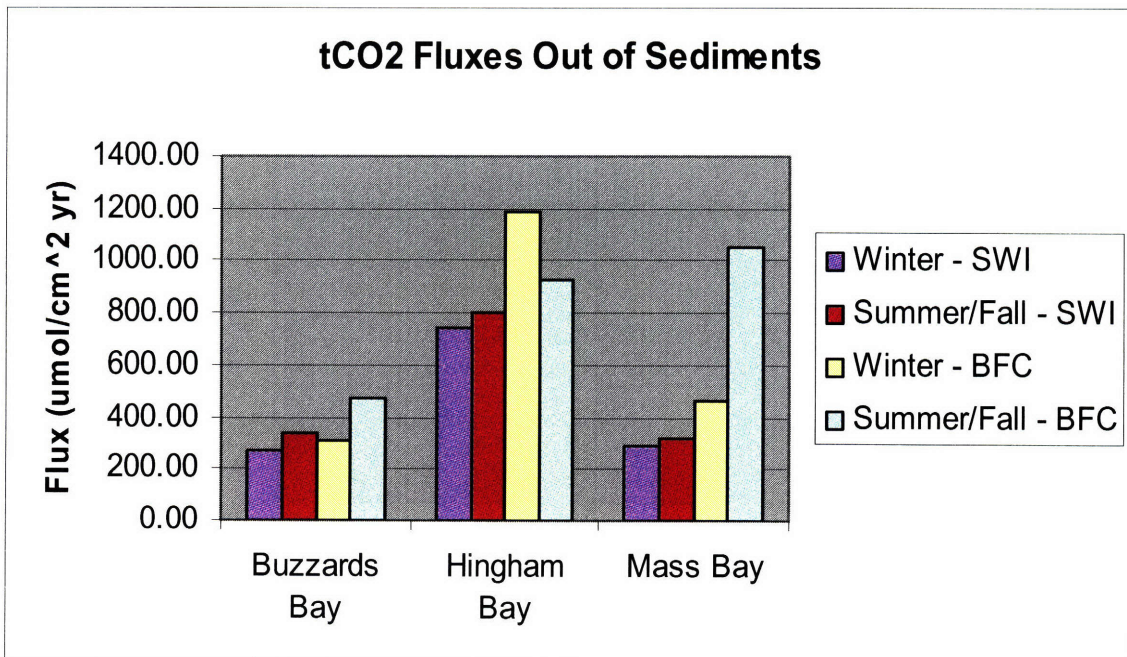
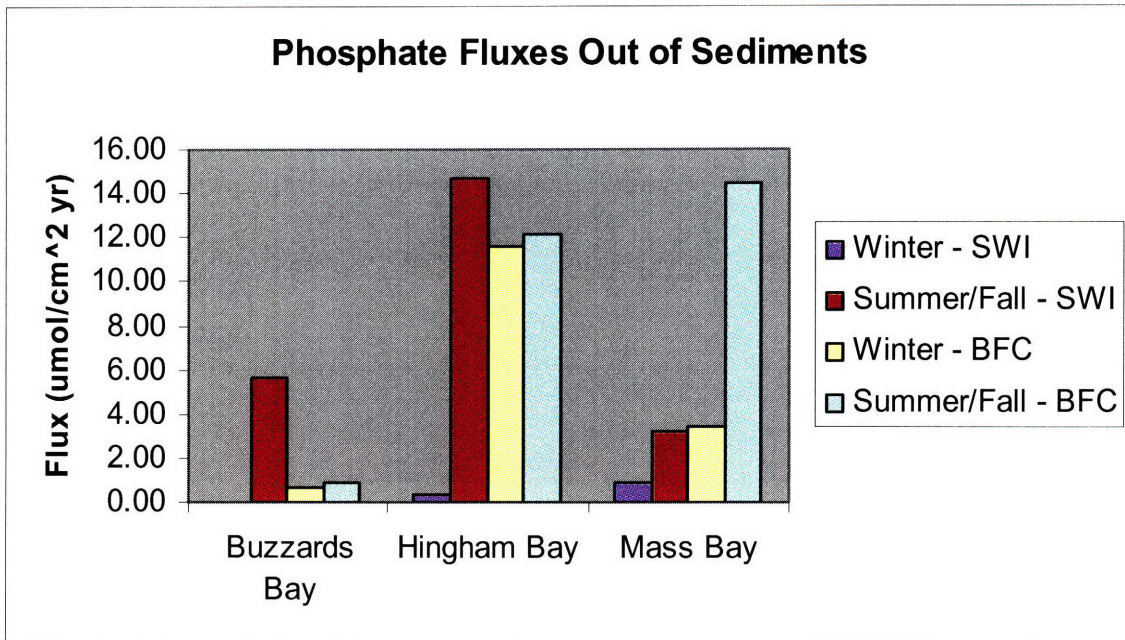


Figure 25: C:P Ratios of SWI and BFC Fluxes

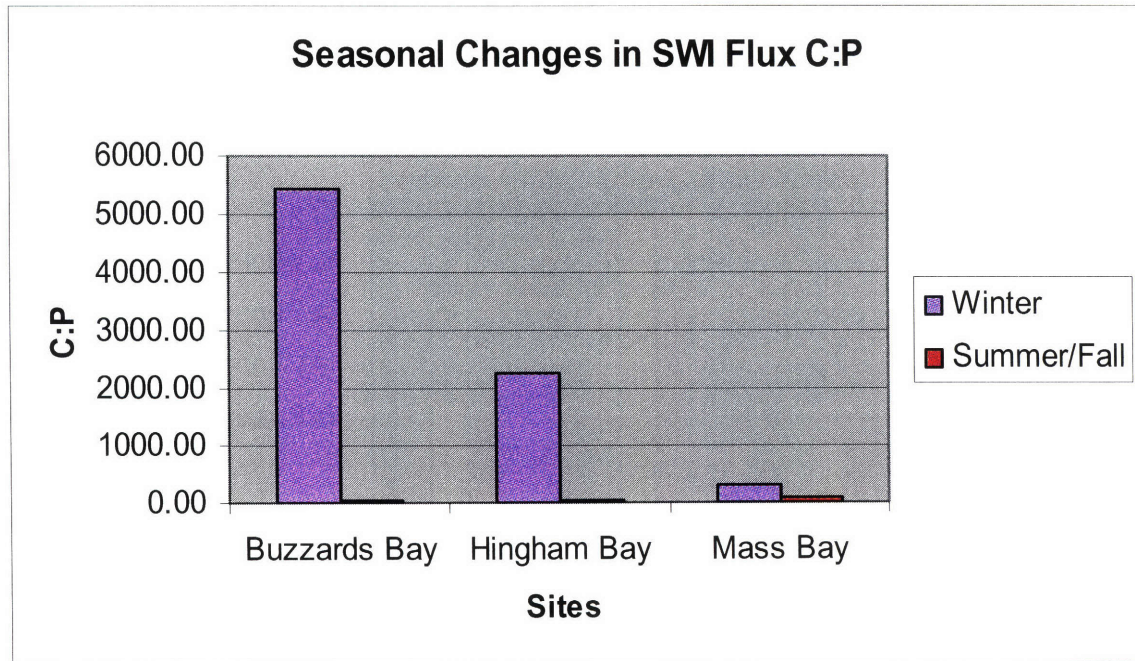
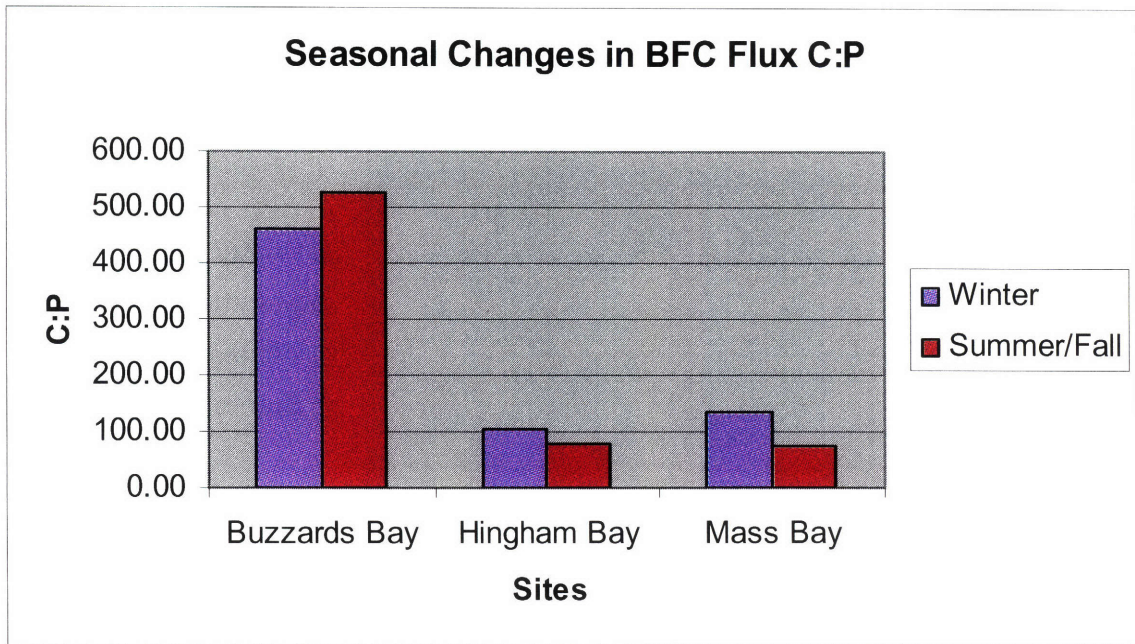


Figure 26: Vivianite Precipitation Curves

